A MULTI-COMPONENT HYDROGEOLOGIC EVALUATION OF A SHALLOW GROUNDWATER FLOW SYSTEM IN GLACIAL DRIFT

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ABSTRACT

A multi-component analysis of a shallow groundwater flow system in glacial drift at the Whiteshell Nuclear Research Establishment, Manitoba is described. The investigation included geologic test drilling, definition of the hydraulic potential distribution, hydrochemical analysis, studies of the temperature distribution in the groundwater zone, radioactive tracer studies of groundwater flow, observations of secondary permeability characteristics and mathematical modelling of the geohydrologic regime.

The stratigraphic section consists of seven glacial deposits which are: a lacustrine sand and gravel unit, a lacustrine silt unit, a lacustrine clay unit, a clay-loam till, a basal sandy drift, an undifferentiated sandy till, and a lacustrine sand unit. The glacial deposits rest on Precambrian metamorphic and intrusive rock.

The groundwater flow system is bounded on the east by a topographic high formed by the lacustrine sand, on the west by the Winnipeg River and in the subsurface by relatively impermeable Precambrian bedrock. Groundwater potential distribution exhibits four major geohydrologic zones. From east to west these are: an upland recharge area consisting of the topographically high lacustrine sand unit, a strong central discharge area, a strong central recharge area and a lowland discharge area confined to the banks of the Winnipeg River. Predominantly lateral flow occurs in the more permeable basal sandy drift with lateral potential gradients existing only locally in the overlying units. Maximum horizontal tracer velocities detected at

three sites are less than 10 feet per year. Groundwater flow through the less permeable upper clay-loam till and glacio-lacustrine silt and clay units occurs via joints and sand laminae.

Variations in electrolytic conductivity and distribution of Ca⁺⁺, Mg⁺⁺, Na⁺, and SO₄⁼ exhibit distinct increases in the directions of groundwater movement. The increase in major-ion concentrations with distance along the flow path in the lacustrine sand unit and the basal sandy drift is less pronounced than in the clay-loam till and the glacio-lacustrine silt and clay units due to a more active flow system and limited availability of soluble minerals in the sandy deposits.

Results of the modelling suggest that the location of the boundary between the central recharge area and the central discharge area is a result of a pronounced thinning of the basal, highly permeable, sandy drift. Strong recharge gradients in the central recharge area could not be simulated in the model.

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INTRODUCTION

The objective of this study is to define the groundwater regime in the Environmental Control Area, Whiteshell Nuclear Research Establishment (WNRE), Pinawa, Manitoba. Since May, 1968, field and laboratory investigations have been conducted to provide hydrogeologic information as a basis for design and location of radioactive liquid and solid waste disposal facilities in the glacial drift which underlies the area. Drift thickness ranges from 30 to 75 feet. The drift is underlain by relatively impermeable Precambrian crystalline rock.

The field study included investigations of the geologic framework, subsurface permeability distribution, hydraulic potential distribution, temperature distribution in the groundwater zone, groundwater-flow velocities and major-ion distribution in the groundwater zone. A two-dimensional, steady-state model of the groundwater flow system was used to refine the field interpretation of subsurface permeability distribution.

The WNRE is located 65 miles east-northeast of Winnipeg on the Winnipeg River (Fig. 1). The Environmental Control Area (Fig. 2) occupies 4 square miles of the WNRE site. Radioactive waste is presently being disposed of in only a small portion of the area (Fig. 2).

This thesis reports on the hydrogeologic environment of the area. Specific application of the results to radioactive waste disposal will be the subject of a future report.

Four previous geotechnical studies of the plantsite have been conducted.

Shawinigan Engineering Limited (1960) summarized the results of an initial geological and geophysical investigation of the WNRE site. Charron (1964) conducted a reconnaissance survey of the radioactive waste disposal capabilities of the ground—water regime. Zwarich (1968) is continuing studies on ion exchange characteristics of the glacial drift in the Environmental Control Area. Lund (1967) investigated groundwater flow velocities at one location using a tritiated heavy—water tracer.

Apparent anomalies in the results of her study were partially responsible for prompting the present investigation.

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Chemical analyses were conducted by the Soil Science Chemistry Laboratory,
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Chemistry Branch, WNRE. The use of tritiated heavy water was supervised by the
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by Great Plains Augering Limited, Saskatoon in 1968 and 1969 and by Mobil Augers
and Research Limited, Edmonton, in 1970.

The investigation was conducted in 1968 with direct financial support of the Environmental Control Section, WNRE, at which time I was a summer research assistant on the project. The study was continued through 1969-70 by the University of Manitoba under research contract with Atomic Energy of Canada, Limited. During this period I was employed on the project as a research assistant.

GEOMORPHIC AND HYDROLOGIC SETTING

The Environmental Control Area is located on the western edge of the Canadian Shield. Sedimentary terrain in the area originated from Wisconsinan glaciations of the Pleistocene Epoch and associated episodes of Glacial Lake Agassiz (Elson, 1966, McPherson, 1968; 1970). Deglaciation of the area occurred between 11,700 and 13,000 years BP with subsequent inundations by Glacial Lake Agassiz prior to 7300 years BP (Elson, 1966).

Relief in the Environmental Control Area varies between 905 feet and 840 feet above mean sea level (Fig. 3). The regional topography (Fig. 2) is a plain with isolated highs. Aerial photographs of the WNRE site and surrounding area (Fig. 4 and Appendix A) illustrate the low relief and marshy character of the terrain.

Oblique modified infra-red photographs (Fig. 5 and Appendix A) show vegetation patterns delineating areas of both bog and sand outcrop.

In relation to worldwide climatic conditions, WNRE is within the region Dfb. (Koppen and Geiger, 1936), characterized by areas near the center of a continent with temperature extremes greater than the world average for the latitude. Recorded temperature extremes at WNRE are -53.5° F. and 108° F. (A. Reimer, personal communication). Mean daily temperature drops below 32° F. during the first part of November and rises above 32° F. near the end of March with an annual mean daily temperature of approximately 35° F.

Precipitation varies greatly with a mean of 20 inches per year, the greatest portion occurring during the summer months (Smith and Ehrlich, 1967).

Recorded extremes are a low of 12.5 inches in 1961 and a high of 26.6 inches in 1968

(A. Reimer, personal communication).

METHODS OF INVESTIGATION

During the initial stages of the study, a base map prepared by Shawinigan Engineering Limited (1960) was used in conjunction with standard 1:16,000 aerial photographs to coordinate field studies. The field studies were conducted during the April-September periods of 1968, 1969 and 1970.

Physical Character of Glacial Drift

The geologic framework in the Environmental Control Area was defined by 76 test holes drilled by a truck-mounted hydraulic drill equipped with a six-inch diameter solid-stem auger flights. Disturbed samples were collected from auger flights, with sample depths estimated to within ± 1 foot. A small portable power drill (Mobil Augers Limited, Minuteman Model) equipped with three-inch auger flight was used for two preliminary test holes but its use was discontinued due to the bouldery character of the glacial till and poor quality of samples. Geologic logs of test holes with interpretations are listed in Appendix B.

Shelby tube cores 0.25 feet inside diameter and two feet in length were taken to define in detail the texture of the glacial drift and to investigate the frequency and character of joints and fractures. Immediately after sampling the cores were sealed using plastic sheeting in the tubes. The samples were then stored at room temperature until inspection could be made. A fluorescent dye, sodium fluorescein, was injected into the glacial drift prior to coring in an attempt to study the physical character of the joints.

Inspection of cores in which sodium fluorescein was thought to be present was conducted under ultraviolet light.

Piezometer Design and Installation

To define the hydraulic potential distribution in the glacial drift, 60 observation sites consisting of 210 piezometers and water table observation wells were installed in the glacial drift. Water levels were measured periodically.

Piezometers of the standpipe design (Fig. 6) were constructed of polyvinylchloride (PVC) pipe. Piezometers installed in 1968 and 1969 had an inside diameter of 0.94 inches and those installed in 1970 had an inside diameter of 0.80 inches. The bottom two feet of the standpipes were perforated and screened with fine-mesh stainless stell screen in 1968 and 1969 and with fine-mesh fibreglass matte in 1970. The intake area was packed with coarse silica sand. Bentonite or very fine silica sand was then placed above the coarse sand pack as a seal against grout. The hole was sealed by a grout mixture of sulphate-resistant cement.

Three piezometers using mild steel standpipes were installed in 1969.

These piezometers were not entirely satisfactory because of difficulty preventing silt and sand from entering the standpipe as well as corrossiveness of the groundwater to mild steel. The PVC piezometers generally attained equilibrium within a week after installation (Fig. 7 and Appendix C) and were subsequently flushed and tested for response and return to equilibrium. The use of PVC standpipes made the installations ideal for hydrochemical purposes. Representative piezometer hydrographs are

shown in Figures 8 – 10. Hydrographs of all other piezometers installed from 1963 to 1970 are included in Appendix C. Piezometer locations and elevations were determined by transit and level with horizontal control of approximately 10 feet and vertical control of 0.1 feet.

Hydraulic conductivities of the glacial drift were obtained from piezometer response tests using the methods of Hyorslev (1951; 1954).

Hydrochemistry and Geothermometry

Groundwater samples were collected from the screened zone of piezometers using a narrow-diameter downhole sampler and by vacuum pump. Samples were isolated from the atmosphere immediately after sampling and stored in polypropylene bottles until analyses were made.

The results of analyses for Ca⁺⁺, Mg⁺⁺, Na⁺, C1⁻, and SO₄⁻ are tabulated in Appendix D. Concentrations of HCO₃⁻ and CO₃⁻ were calculated by the procedure discussed by Barnes (1964) from titration data and are listed with total dissolved solids in Appendix D. Field pH measurements were obtained using a null-balance pH meter (Radiometer Model pHM-4) following the procedure of Barnes (1964). Electrolytic conductivities were measured in the water at the intake area of piezometers using an extended-cable down-hole conductivity cell (Solubridge Model B3-338). Field pH and conductivity results are also listed in Appendix D. Groundwater temperatures were measured at intervals in piezometers with an extended-cable thermistor (Whitney Model TC-5A) and are summarized graphically in Appendix E along with meteorological summaries for 1968, 1969 and 1970 obtained from the Environmental Control Section, WNRE.

Heat transfer from the thermistor and connecting cable into the narrow-diameter standpipes resulted in difficulties in obtaining reproducible values. Conduction of heat from surface down the standpipe and convection within the standpipe are also believed to greatly reduce the reliability of the data.

Tracer Injection Experiments

Groundwater flow velocities were studied by injecting tritiated heavy water into the glacial drift at three locations in the study area (Fig. 3). The injection experiments were initiated in July, 1969 at a site immediately north of the present Active area. Injections at the other two sites were made in August, 1969. Each injection site comprises one 3.5 inch-diameter injection well and up to 40 sampling tubes concentrated predominantly in the anticipated direction of a flow. Sampling tubes were constructed and installed in the same manner as piezometers except that the perforated area was increased to 3 or 5 feet, depending on the sampling zone desired. The locations of injection wells and sample tubes are shown in Figures 11 – 13.

Samples were obtained from sample tubes by vacuum lines set in the perforated area of the sampling tube. Tritium analyses were conducted using a liquid scintillation counter supplied by the Environmental Control Branch, WNRE. Beta activity fluctuations for selected sample tubes are presented in Appendix F.

Prior to injection of tritium, hydraulic connection of all sample tubes to the injection well was confirmed by pressure injection of pure water. At injection site No. 1, a trial injection of sodium fluorescein was conducted to test injection and sampling procedures.

An injection well was also installed approximately 5 feet east of piezometer nest No. 20. Only sodium fluorescein was injected at this point, the primary purpose of the experiment being to take Shelby tubes within the injection zone as a means of investigating secondary permeability characteristics of the fine-grained deposits.

Steady-State Mathematical Modelling of the Groundwater Regime

A computer program initially prepared by Morris (1968; 1969) for IBM

Systems 360-65 using the finite-element techniques outlined by Zienkiewicz and

Cheung (1966) was used in the analysis of the groundwater flow pattern and subsurface

permeability distribution. The program solves the two-dimensional steady-state flow

equations for groundwater flow in saturated, anisotropic, heterogeneous porous media

with an irregular water table and impermeable flow boundaries or their equivalent.

The program was adapted for specific application to the flow regime in the WNRE

area with solutions for the hydraulic potential distribution being plotted by a Calcomp

plotter. Program listings and comments on the program are compiled in Appendix G.

STRATIGRAPHY

The stratigraphic section in the Environmental Control Area, WNRE, consists of seven units overlying Precambrian bedrock. These are: a lacustrine sand and gravel unit, an interbedded lacustrine silt, a lacustrine clay, a carbonate-rich, clay-loam till, a basal sandy drift rich in Precambrian-derived pebbles and boulders, a sandy till, and a unit of predominantly fine to medium-grained sand. The total thickness of these units overlying the Precambrian bedrock varies between 30 and 75 feet. The surficial geology is shown in Figure 14 and geologic cross sections are shown in Figures 15 – 17.

In the western plain. As shown in Figure 15, the northwestern portion of the upland is underlain by only one stratigraphic unit, a thick deposit of relatively uniform fine to medium-grained lacustrine sand. The unit includes a few thin interbeds of silt and clayey silt which are probably discontinuous. This deposit extends westward beneath the clay-loam till and is terminated in the northeastern portion of the upland by the unit which appears to be a sandy and in places silty till. Figure 16 shows that the southern portion of the upland is underlain by a much different stratigraphic sequence. The thick section of sand shown in Figure 15 is not present, and a lacustrine sand and gravel deposit overlies the lacustrine clay.

West of test hole 48 (Fig. 16) the surficial sand and gravel deposits pinch out and the section consists of, in descending order, lacustrine clay, clay-loam till, and a basal unit of fine to coarse-grained sand.

The stratigraphic sequence east of the upland has not been defined by test drilling. However, based on the topographic setting, the poorly drained character of the terrain, and the continuity of both the clay-loam till unit and the lacustrine clay under the upland (Fig. 16) it appears likely that the area is underlain by both of these units.

A north-south geologic cross section (Fig. 17) illustrates the relatively uniform character of the stratigraphy in the more central portions of the Environmental Control Area.

Lacustrine Sand and Gravel Unit

The lacustrine sand and gravel unit consists of up to 15 feet of fine to coarse-grained sand with minor silt beds and gravel. The deposit is restricted to an isolated topographic high in the southeastern portion of the area (Fig. 15). Grainsize characteristics suggest relatively high hydraulic conductivities, probably in the range of 1×10^{-5} to 1×10^{-2} fps.

Lacustrine Silt Unit

The lacustrine silt unit observed in the Environmental Control Area represents a late phase shallow-water deposit of Glacial Lake Agassiz (McPherson 1968). This unit consists of massive to laminated very fine sands, sandy silts, silts,

Clayey silts and silty clays with a maximum thickness of 15 feet occurring near the Winnipeg River, and wedges out in the central portion of the area. Where the unit is less than two feet thick, it is somewhat obscured by the effects of soil formation.

Easily eroded masses resembling primary bedding structures were recognized in cores of this unit (Fig. 18). The structures are generally one to three inches in length and less than 1/2 inch in thickness. They may represent scours subsequently filled with coarser silt. They were not observable unless the cores were washed with a high pressure stream of water. The lower contact of the silts to the deep-water clays appears gradational over several feet.

In the poorly drained western part of the Environmental Control Area underlain by the lacustrine silt unit, perched water table conditions frequently exist. Water-table observation wells located adjacent to areas frequently covered by water indicate that the water table lies three to five feet below ground level. This pending is thought to be due to local interbeds of silty clay near the soil zone which retard downward percolation of surface water.

Lacustrine Clay Unit

The lacustrine clay unit was deposited during a high-water phase of Glacial Lake Agassiz (McPherson, 1970). The thickness has a maximum of 25 feet near the Winnipeg River. Carbonate pebbles are present throughout this unit but become more common in the lower portions. Fragments of till-like material up to 10 inches in diameter (Fig. 19) were noted in isolated cases. Angular buff-coloured calcareous silt clasts are very common

near the bottom of the unit (Fig. 20). The pebbles, till and silt clasts are probably due to ice rafting although Weller (1960), Hills (1963) and Elson (1966) have discussed possible alternate origins for such features.

Laminae of fine to very fine sand are present throughout the unit at angles up to 90° to the bedding (Fig. 21). These laminae may cut horizontal or shallow dipping laminae of clay and silty clay or may be conformable to them. A probable origin of the sand laminae is infilling of slump or dessication cracks formed by fluctuations in the lake level.

The lower contact of the clay is marked by a stratified zone up to three feet thick of pebbly, sandy silt, and silty sandy clay (Fig. 22).

Grain-size analyses of the glacio-lacustrine clays conducted by Zwarich (1968) on five samples collected from drill holes near the present radioactive disposal area (Fig. 2) indicate a 60:30:10 ratio of clay, silt and sand, respectively. It is expected that the largest portion of the silt and sand occur as laminae within the clay unit rather than a homogeneous unit of clay, silt and sand.

Inspection of Shelby tube cores taken near piezometer nest No. 20 (Fig. 2) indicate well-defined jointing throughout this unit. The joints may appear as thin dark lines cutting laminae (Fig. 23) or as planes coated with carbonate and gypsum precipitates (Fig. 24). Rootlets follow the joint surfaces as deep as 14 feet (Fig. 20).

The results of Hvorslev response tests indicate that the lacustrine clay unit is characterized by hydraulic conductivities in the range of 1×10^{-10} to 1×10^{-7} fps at depths below ground surface greater than 14.5 feet and by conductivities in

the range of 1×10^{-10} to 1×10^{-5} fps at depths less than 14.5 feet below ground surface. Laboratory tests conducted by Templeton Engineering Company Limited (1960) on small samples obtained from the lacustrine clay near the plant (south half, section 21) yielded values in the range 1×10^{-10} to 3×10^{-9} fps. This suggests that joints, fractures or other secondary in situ characteristics of the lacustrine clay unit significantly increase the hydraulic conductivity. A more detailed analysis of the hydraulic conductivities associated with the lacustrine clay is discussed in a later section.

Clay-loam Till Unit

The lacustrine clay unit grades through a stratified zone of pebbly, sandy silt, and silty sandy clay, a few feet thick into a pebbly, carbonate-rich, clay-loam till. The deposit ranges from 0 to 30 feet in thickness. Grain-size analyses (Zwarich, 1968) of composite samples show approximately equal proportions of clay, silt and sand near the upper contact with the sand content increasing to over 50 per cent near the lower contact. Pebbles and cobbles are predominantly dolomitic with the remainder commonly being highly weathered Precambrian material. The lower contact is marked by a pavement of Precambrian cobbles and boulders. Considerable difficulty penetrating this layer was encountered at many locations. This unit correlates with the Libau Drift mapped and described by McPherson (1968) in the area west of the Winnipeg River.

Joint planes coated with carbonate precipitates (Fig. 25) were commonly observed in Shelby tube cores taken near the upper contact of this unit.

The results of Hvorslev response tests indicate that the clay-loam till is characterized by hydraulic conductivities in the range of 1×10^{-7} to 1×10^{-4} feet per second (fps). Joints appear to account for a significant secondary permeability network although only a few irregular joint-fracture features were observed in the Shelby tube samples from this unit. A more detailed analysis of the hydraulic conductivities associated with this unit is presented in a later section.

Basal Sandy Drift Unit

The basal sandy drift unit underlies the clay-loam till unit in the western third of the study area (Figs. 15 - 17). The deposit varies in thickness from 0 to 15 feet. Texturally, it varies from a silty till to unsorted gravel and sand to a well sorted fine to coarse-grained sand with few, if any, till-like features. Compositionally the unit is characterized by angular rock fragments of Precambrian Shield lithologies.

Sedimentary rock fragments were observed only in isolated cases. The upper contact of the basal sandy drift with the clay-loam till is defined by a boulder pavement.

McPherson (1970) states:

The main basis for distinguishing the Belair Drift is its occurence on surface as a sandy end moraine or outwash, its characteristic mineralogy indicating it was primarily derived from acid igneous rocks, and its stratigraphic position in the subsurface being stratigraphically lower than Lake Agassiz sediments and Libau Drift.

By this definition, the basal sandy drift unit can be stratigraphically correlated with the Belair Drift.

The results of Hvorslev response tests indicate that the basal sandy drift is characterized by hydraulic conductivities in the range of 1×10^{-7} to 1×10^{-4} fps. A detailed analysis of the hydraulic conductivities associated with this unit is presented in a later section.

Undifferentiated Sandy Till

In the eastern portion of the area the lacustrine sands are overlain by a deposit which is apparently a sandy and in places silty till. Only limited test hole data is presently available in this area, making correlation of this unity to the western portion of the area impossible.

Lacustrine Sand Unit

The lacustrine sand unit consists of up to 60 feet of fine to medium-grained sand in the northeastern portion of the area (Fig. 15). In the lower portions, several interbeds of silt and clay, often containing angular carbonate rock fragments. The lacustrine sand unit is overlain by the clay-loam till, although its stratigraphic relationships to the basal sandy drift are not clear.

The results of Hvorslev response test indicate hydraulic conductivities in the range of 1×10^{-7} to 1×10^{-5} fps, hydraulically similar to the basal sandy drift.

Precambrian Bedrock

Underlying the glacial drift in the WNRE area are granite and granodiorite of Precambrian age (Templeton, 1960). The bedrock contact defined in Figs. 15 - 17 has been assumed to occur within a foot or two of the maximum depth penetrated at

each of the test holes drilled to refusal. This assumption is supported by the results of a geophysical investigation of the Environmental Control Area by Pakiriah (1970), by the relative uniformity of penetration depths within most portions of the area. The similarity of the steady-state mathematical model and the field data also support this location for the impermeable lower bounding surface.

Templeton Engineering Company Limited (1960) concluded on the basis of four diamond-drill holes in the southern half of section 21 that a deposit of "broken and weathered granite bedrock rather than glacial till" rests on the solid bedrock surface. Results of the present investigation indicate that this is a misinterpretation of the deposit here referred to as the basal sandy drift.

Although the overall bedrock surface in the Environmental Control Area has been defined to be gently rolling, local relief of over 20 feet over a horizontal distance of 10 feet was observed at tritium injection site No. 2. McPherson (1970) noted local relief on the bedrock surface in drift covered areas near Pinawa, Manitoba in excess of 100 feet.

The bedrock is assumed to have negligible permeability. Examination of outcrops near WNRE and the steady-state mathematical model support this assumption.

FIELD STUDIES OF GROUNDWATER FLOW

The objective of this portion of the investigation was to define the geohydrologic regime in the Environmental Control Area by interrelating studies of fluid potential distribution, hydrochemistry, subsurface temperature variations, groundwater flow velocities and fracture flow characteristics.

Groundwater Flow Pattern

A piezometer network of 210 piezometers and water-table observation. wells was utilized to establish the hydraulic potential distribution (Figs. 26 - 28) and the magnitude of transient head variations generated within the system by water-table fluctuations. Representative piezometer hydrographs (Figs. 8 - 10) illustrate the response to seasonal precipitation and indicate short term reversals in hydraulic potential gradients in the central discharge area due to water-table fluctuations. A corresponding increase in recharge head is observed in the central recharge area. These transient effects would not be expected to affect the flow pattern as the condition is short-term and rapidly returns to the original gradient direction.

Data used for preparation of flow diagrams (Figs. 26 – 28) were taken from piezometer hydrographs (Figs. 8 – 10 and Appendix C) as an average value for 1969 and 1970. Values for 1968 were discarded due to abnormally high rainfalls which resulted in highly fluctuating values (Appendix C).

The flow pattern interpretations (Figs. 26 - 28) were made according to the methods of van Everdingen (1963) to account for a 10:1 vertical exaggeration and Freeze (1969) to approximate refractions of equipotential lines at permeability field boundaries defined by test drilling. The potential distribution indicates two areas of recharge referred to as the upland and the central recharge areas and two areas of discharge referred to as the central and lowland discharge areas (Fig. 29). The upland recharge area occurs within the outcrop area of the lacustrine sand and gravel. The lowland discharge area is confined to the banks of the Winnipeg River. Vertical discharge gradients in the central discharge area have a maximum of 0.3 ft/ft with a head avove the average water-table level of up to 12 feet and vertical recharge gradients in the central recharge area have a maximum of 0.2 ft/ft. Horizontal potential gradients are restricted to the basal sandy drift and the lacustrine sand with a local horizontal gradient existing in a narrow zone along the boundary between the central discharge and central recharge areas.

A flow pattern interpretation (Fig. 28) along a north-south cross section along the western boundary of Section 28 indicates recharge to the basal sandy drift except near the northern edge of the study area. The sandy drift unit pinches out and small discharge gradients are present.

The hydraulic potential distribution suggests two scales of flow regimes in the Environmental Control Area. The smallest scale 'local' flow systems consists of recharge in the upland recharge area and discharge in the central discharge area, and recharge in the central discharge area discharing into the Winnipeg River (Fig. 29).

The larger scale 'regional' flow system consists of recharge in the upland recharge area and discharge into the Winnipeg River with lateral ground-water flow continuing under the central discharge area. This co-existence of flow systems has been described by Meyboom (1962) and T6th (1966), as well as others.

The seasonal stability of recharge-discharge boundaries was studied in detail by installing a line of piezometers at closely spaced intervals across the central discharge-central recharge area boundary in section 28 (Fig. 29). During the period of study from summer, 1968 to fall, 1970, this boundary was not observed to vary significantly. On the basis of these data, it has been assumed that the recharge-discharge boundaries are stable with respect to seasonal fluctuations of precipitation, evapotransipiration, and freezing.

Hydrochemistry

The major-ion chemistry of groundwater in a shallow flow system results from numerous geochemical processes. The results of studies by such investigators as Meyboom (1962), T6th (1966), Parsons (1967), Rozkowski (1967), Charron (1969), van Everdingen (1970), and Cherry, et al (1971), suggest that the most important geochemical influences are: (a) generation of gases in the soil zones of recharge areas, (b) solution of porous media minerals, (c) precipitation of mineral phases within the pore network, (d) ion exchange between the pore solution and clay minerals in the porous medium, and (e) oxidation-reduction reactions.

These processes are interdependent and are affected by mineral associations in the porous media, the relative order of geologic units along the groundwater flow path

rate of groundwater flow and the length of the groundwater flow path. Using these guidelines, the major-ion distributions and electrolytic conductivity in the groundwater regime in the Environmental Control Area were investigated as a means of refining the flow system interpretation based on the hydraulic potential distribution.

Hydrochemical cross sections (Figs. 30 and 31) were prepared for Ca⁺⁺, Mg^{++} , Na⁺, Cl⁻, HCO₃⁻, CO₂⁻ and SO₄⁻ and electrical conductivity and exhibit hydrochemical patterns which correlate very closely with the flow pattern interpretations obtained from hydraulic head data (Figs. 26 – 28). In recharge areas the conductivity increases downwards to a maximum of 600 millimhos/cm. In the central discharge area conductivity increases from a minimum of 10 millimhos/cm near the Precambrian bedrock to greater than 680 millimhos/cm near the water table. A complex hydrochemical pattern is developed in a zone along the boundary of the central recharge and central discharge boundary (Figs. 30 – 32). This is probably due to a sluggish flow regime with long groundwater residence times resulting from low hydraulic potential gradients coupled with low permeability material, as well as short-term reversals of groundwater potential distribution west of this zone.

Groundwater in the lacustrine sand deposit and the basal sandy drift up to the central recharge-central discharge area boundary has very low concentrations of dissolved solids. Samples taken from shallow piezometers in the upland recharge area indicate that essentially all the dissolved solids are picked up in or near the soil zone. The subsequent lack of major variations in the dissolved mineral load in these deposits is due to the low remaining H_2CO_3 agressivity as well as the lack of

relatively soluble salts in the predominantly quartzo-feldspathic sand. The more permeable deposits would also tend to allow a more active flow regime with resulting shorter residence time and therefore fresher water.

West of the central discharge-central recharge area boundary in the basal sandy drift, conductivity increases over several hundred yards to a maximum of 450 millimhos/cm. This rapid increase results from downward influx of highly mineralized water from the upper units, possibly enhanced by increased availability of fine-grained soluble materials in the drift unit.

In the lacustrine clay and clay-loam till units in the central recharge area, downward seepage is accompanied by increases in Na⁺, Ca⁺⁺, Mg⁺⁺, Cl⁻, and SO₄⁼. In the central discharge area, upward seepage is accompanied by similar increases. This is attributed to the availability of soluble sulphate and chloride salts in the fine grained units.

The hydrochemical patterns (Figs. 30 and 32) indicate that precipitation infiltrates in the lacustrine sand and moves downward and laterally with dissolved solids remaining below 50 mg/l. A portion of the water in the central discharge area is forced upwards into the overlying upper till and lacustrine deposits with total dissolved solids reaching 200 mg/l near the water table.

Water infiltrating in the central recharge area rapidly picks up a maximum of 6000 mg/l dissolved solids as it descends to the basal sandy drift unit where it is somewhat diluted by the fresher westward lateral flow from the upland recharge area. The combined flows travel over the Precambrian surface to discharge in and near the Winnipeg River.

Field pH data have been obtained (Appendix D) and will be incorporated in a future report dealing more intensively with the hydrogeochemical regime in the WNRE area.

Groundwater Temperature Distribution

Groundwater temperature variations with depth and time were studied during the spring and summer of 1969. Typical variations (Fig. 32 and Appendix E) indicate that seasonal temperature variations affect the groundwater temperature through the full section of glacial drift at WNRE.

There is an apparent correlation of the range of temperatures and discharge-recharge zones. From graphical results (Appendix E) of 13 piezometer nests in discharge areas and 8 nests in recharge zones it was found that the mean of groundwater temperatures were 1.1 C° and 1.6 C° respectively for the first week of July, 1969. At this time the approximate mean groundwater temperature was 5.2° C. Temperature data from water-table observation wells was not included in this test.

Tracer Velocity Studies

To obtain real values of horizontal groundwater flow velocity, tritiated heavy water was injected into the flow system at three selected locations (Figs. 11 - 13) in the Environmental Control Area. Injection site 1 (Fig. 11) was installed in the basal sandy drift in the central discharge area. During the injection process the tracer was observed to arrive at sample points 13 and 15 located approximately 8 feet from the injection point. Arrivals of the tracer at other sample points located

8 to 48 feet from the point of injection were not observed during the period July, 1969 to August, 1970 although fluctuations in tritium concentration were noted in the active sample points. The movement of the tritium is confirmed by the gradual decrease in tritium concentration (Appendix F) at the injection point. Hydraulic connection of all sample tubes to the injection point was confirmed prior to injection of the tracer by pressure testing.

The tracer is known to have been successfully injected into the ground-water flow regime in the basal sandy drift. The data show that the lateral rates of groundwater movement have not been sufficient to move the tracer completely away from the area of injection. Hydraulic potential data require that the tracer front move westward and it is very unlikely that the tracer front could move undetected between the sample tubes to the west. Since sample tubes are present from 8 to 15 feet west of the injection site, it can be safely assumed that west-lateral groundwater velocities are considerably less than 10 feet per year at this location.

Injection site 2 (Fig. 12) was installed in the basal sandy drift in the central recharge area. Arrivals of the tracer during injection was observed in sample tubes 26 - 43, 1 and 4. Arrivals of the tracer at other sample tubes located 8 to 58 feet from the point of injection were not observed during the period August 1969 to August, 1970, although decreases in tritium concentrations were observed in the active sample tubes and the injection well. The procedures and results were very similar to those observed at injection site 1. This indicates a maximum horizontal flow velocity of less than 10 feet/year.

Injection site 3 (Fig. 13) was installed adjacent to injection site 2 in the central recharge area and bracketed the contact of the lacustrine clay and the clay-loam till units. Arrival of the tracer was observed in sample tube 6 during injection. Arrivals of tritium at sample tubes 8 and 48 feet from the injection point were not observed during the period August, 1969 to August, 1970. However, very low tritium levels did occur sporadically in various sample tubes, probably due to contamination during sampling. Similarly to injection sites 1 and 2, this indicates maximum westward horizontal flow velocities of less than 10 feet/year.

The experiment was designed to detect horizontal flow velocities and was not capable of detecting vertical flow. The sampling and injection methods resulted in a large mass of tritiated water (Figs. 11 - 13) which, if moving horizontally, would necessarily pass through at least one sample point.

The possibility of hydraulic fracturing due to the pressure injection method would increase the flow velocity because of increased permeability and would not have caused an anomalously low value for the flow rate. Movement of the tracer was confirmed by fluctuating tritium levels in active holes (Appendix F).

A previous tracer study in 1967 by B. Lund (1967) gave a horizontal flow velocity of 430 feet per year. The study was conducted from surface to a depth of 20 feet in the centre of section 28 (Fig. 3) in an area of extremely low lateral hydraulic potential gradients. The velocity determined by here is now thought to be due to flow in the unsaturated zone as a result of a groundwater 'mound' built by the injection of tritiated heavy water rather than flow through the saturated zone.

Tracer Investigations of Secondary Permeability Characteristics

Joints in the lacustrine clay and clay-loam-till units were observed in Shelby tube cores and disturbed samples. In an attempt to define the effects of the icints on permeability in the fine-grained deposits, a fluorescent dye, sodium fluorescein was injected into these units and then Shelby tube cores were taken of the injected material.

Examination of cores under ultraviolet radiation showed no visible dye on joint surfaces or in the bulk of the material. The most permeable joints are coated with carbonate accumulations and the coating may mask the presence of the dye. Intergranular flow of groundwater in the clay unit and upper carbonate till unit is probably negligible with respect to flow through joints. Photographs of fractures (Figs. 19 - 25 and Appendix A) support this view.

Further study on the problem of quantitative determination of fracture flow is being considered. This aspect is very important to radioactive waste disposal as ion exchange capacities of a geologic medium will be drastically reduced if waste solutions do not have access to intergranular pore spaces.

Hvorslev Response Tests

Of the 210 piezometers installed in the area of investigation, 170 were subjected to response tests based on the principles and methods described by Hvorslev (1951; 1954). The ratios of vertical versus horizontal permeabilities were assumed unity for all the hydrostratigraphic units. Data summarized by Davis and DeWiest (1966) and Bakhtiari (1970) suggest that this assumption is normally acceptable

within 50 per cent for sandy unconsolidated deposits not containing fine-grained interbeds or oriented platy mineral orientations. The $K_{\rm v}/K_{\rm h}$ ratios for the fine grained deposits were assumed to be unity due to the relatively orthogonal orientation of joints and laminae prominent in these deposits, although the actual value of $K_{\rm v}/K_{\rm h}$ was not determined.

Hydraulic conductivities obtained from water-level response tests can be affected by: (1) water-intake geometry, (2) textural variations within the intake zone, (3) disturbances of the formation during piezometer installation (4) nature and efficiency of the grout seal above the intake zone, (5) gas accumulations in the intake zone, (6) any artificial restriction or enhancement of the hydraulic conductivity of the formation or intake zone resulting from piezometer installation. Because of these and other factors involved in water-level response tests, the resulting hydraulic conductivities are likely not reliable to more than one order of magnitude.

Figures 33 and 34 summarize the permeabilities calculated from the water-level response tests. Data which appear to have been significantly affected by one or more of the above factors have been omitted. The exclusion of results was usually based on obvious evidence of sedimentation or clogging or poor graphical correlations of water-level recovery with time.

All of the 7 piezometers in the lacustrine sand unit from which reliable water-level response data were obtained yielded hydraulic conductivities in the range 1×10^{-7} to 1×10^{-4} fps. Of the 34 piezometers in the basal sandy drift from

which reliable data were obtained, 22 yielded values in the range 1×10^{-7} to 1×10^{-5} . The remainder of the conductivities are in the range 1×10^{-7} to 5×10^{-10} . Auger samples from these units have been compared visually to glacial sands from southwestern Manitoba. The visual textural correlations support the conclusion that conductivities in the range 1×10^{-7} to 1×10^{-5} are characteristic of most zones within the two units.

Of the 36 piezometers in the clay-loam till from which reliable water-level response data were obtained, 20 yielded hydraulic conductivities in the range of 1×10^{-9} to 1×10^{-7} pfs. Of the remainder, 9 are in the range 1×10^{-7} to 9×10^{-6} and 9 in the range 1×10^{-9} to 1×10^{-10} fps. There is little doubt that joints in the clay-loam till account for most, if not all, of the hydraulic conductivities larger than about 1×10^{-8} fps and that the wide range of the data reflects heterogeneous characteristics in the permeability system. Variability in the character and frequency of joints in till at numerous locations in western Canada has been observed by A. Vonhof (personal communication, 1970).

Hydraulic conductivities in the lacustrine clay unit vary within the relatively wide range of 1×10^{-10} to 1×10^{-5} fps. Fig. 35 indicates that 10 of the 31 piezometers located in this unit at depths between 5 and 14.5 feet below ground surface yield hydraulic conductivities greater than 1×10^{-7} fps. Of the 14 piezometers in this unit at depths below 14.5 feet, only one yielded a hydraulic conductivity greater than 1×10^{-7} fps. The hydraulic conductivity data suggest that the lower portion of the lacustrine clay and the clay-loam till have rather similar hydraulic properties.

The significantly higher frequency of larger hydraulic conductivities nearer the ground surface can be accounted for by the increased occurrence of joints, dessication cracks, and channels cuased by plant roots and burrowing animals.

The seasonal water-table fluctuations in the area underlain by the lacustrine clay is between 2 and 7 feet below ground surface. Decline of the water table below 7 feet probably occurs during years of exceptionally low precipitation. The clay fraction of sediments in the Lake Agassiz basin generally contain abundant montmorillonite (Kushnir, 1970) which characteristically undergoes major volume changes on wetting and drying. The zone of water-table fluctuation is subject to numerous cycles of wetting and drying and is therefore commonly a zone in which dessication cracks and joints develop.

STEADY-STATE MATHEMATICAL MODEL OF THE FLOW REGIME

The steady-state mathematical model (Morris, 1969) was used to: (a) refine the interpretation of the groundwater flow pattern, (b) test and refine the field interpretation of the subsurface permeability distribution, (c) test the sensitivity of the flow regime to water-table fluctuations, (d) test the assumption of an impermeable bedrock contact and (e) determine if this mathematical model was useful in predicting flow patterns as a method of initial stage planning of hydrogeologic models.

The boundary conditions defining the limits of the model (Fig. 35) are:

(a) the bedrock surface is a lower bounding impermeable surface, (b) the lateral bounding surfaces are vertical and coincide with the topographic high of the lacustrine sand to the east and the Winnipeg River to the west, (c) the water table is constant and acts as a free-surface and (d) steady-state conditions exist.

The models tested include isotropic and anisotropic homogeneous, isotropic and anisotropic 2-layer and isotropic and anisotropic 2-layer with variations in permeability within the basal sandy drift and the lacustrine sand and gravel. Most combinations of permeability ratios from one to 10,000 were modelled. In all cases the lacustrine silt unit, lacustrine clay and clay-loam till were considered to have unit vertical permeability.

By using a water-table configuration determined from the observation well network (Fig. 26) a model was developed which very closely approximated the

conditions observed in the field. This model (Fig. 35) was developed using unit permeability for the clay-loam till, clay and silt deposits and a permeability of 1,000 for the sandy drift and lacustrine sand deposits. This model does not, however, simulate the magnitude of recharge gradients existing in the central recharge area, although all other features are very similar.

Studies currently in progress are attempting to determine whether improved correlations between the model and the flow system can be obtained. The modelling technique is also being used in an analysis of cross section B-B' (Fig. 27). The results described above indicate that (a) although from a small-scale viewpoint each hydrostratigraphic unit is heterogeneous in texture and permeability, it is only the large scale permeability contrasts between the units which are the dominant internal factor in the development of the flow pattern, (b) the initial assumption that the bedrock is relatively impermeable is acceptable and (c) the location of the bedrock contact as indicated by the drilling data is reasonable. Although many different conductivities have been assigned to the basal sandy drift and the lacustrine sand, the results obtained from the model using a uniform conductivity continue to be the closest representation of the head distribution.

SUMMARY OF CONCLUSIONS

The stratigraphic section found in the WNRE Environmental Control Area conforms generally to the Pleistocene stratigraphic succession defined by McPherson (1968, 1970).

The glacial drift sequence includes a lacustrine sand and gravel deposit, a late-phase deposit of Glacial Lake Agassiz silts, a deep-water deposit of clay and silty clay, a clay-loam till, a basal sandy drift rich in Precambrian-derived pebbles and boulders, an undifferentiated sandy till, and a uniform lacustrine sand. The glacial drift is underlain by relatively impermeable Precambrian granite and granodiorite.

The groundwater flow pattern is characterized by a westward flow to the Winnipeg River with four zones of recharge and discharge. From east to west these are: an upland recharge area consisting of an outcrop of the lacustrine sand and gravel which act, as a groundwater divide, a central discharge area, a central recharge area and a lowland discharge area confined to the banks of the Winnipeg River which is the groundwater divide to the west. Tracer studies at three locations suggest maximum groundwater flow velocities to be less than 10 feet per year.

 Ca^{++} , Mg^{++} , Na^{+} , SO_4^{--} and electrolytic conductivity increased in the directions of apparent groundwater movement and correlated with flow directions interpreted from the hydraulic potential distribution. Values increased laterally in the more permeable basal units from the upland recharge area in the east to the Winnipeg River. Values in the till and lacustrine units increase upwards in the

hydrochemical pattern has developed along a zone between the central discharge and central recharge area boundary due to stagnant flow conditions. The excellent correlations of hydrochemistry and groundwater flow pattern are due to the relatively active flow regime in the Environmental Control Area, lack of solubility constraint for most major ions, and highly contrasting hydrochemical facies in each geologic unit. Hydrochemistry confirmed flow pattern interpretations which would otherwise have required considerably more detailed instrumentation.

Laboratory studies of cores taken from the lacustrine and upper till units indicate that groundwater flow is occurring mainly along joint surfaces and sand laminae. This is supported by the accumulation of carbonate precipitates on joint surfaces and oxidized zones. Attempts to quantitatively describe joint-fracture flow by means of fluorescent dye tracing was not successful.

A steady-state mathematical model of the groundwater flow regime using boundary conditions defined by test drilling and the hydraulic potential distribution suggests a permeability ratio between the lacustrine clays and clay-loam till deposits and the basal sandy drift and lacustrine sand and gravel units of 1:1000. The model supports the conclusion that the water-table conditions and the discharge-recharge boundaries are relatively stable. The use of this model with preliminary hydrostratigraphic data may allow a more economical approach to simple hydrogeologic investigations.

The complementary use of several hydrogeologic techniques has allowed a relatively detailed interpretation of the shallow groundwater flow system in glacial drift at WNRE. Some of the more general conclusions resulting from the investigation are summarized as follows:

- 1) A comprehensive understanding of the stratigraphic framework is a prerequisite to the design of a piezometer network and the interpretation of the resulting potential distribution.
- 2) Hydrochemical patterns are a result of flow patterns and stratigraphy and may be used in some cases to aid in groundwater flow interpretations.
- 3) Influence of joints and small sand laminae increase hydraulic conductivities of otherwise low-permeable materials.
- 4) A finite-element steady-state model of groundwater flow may refine interpretations of field data and reduce the cost of detailed field study.

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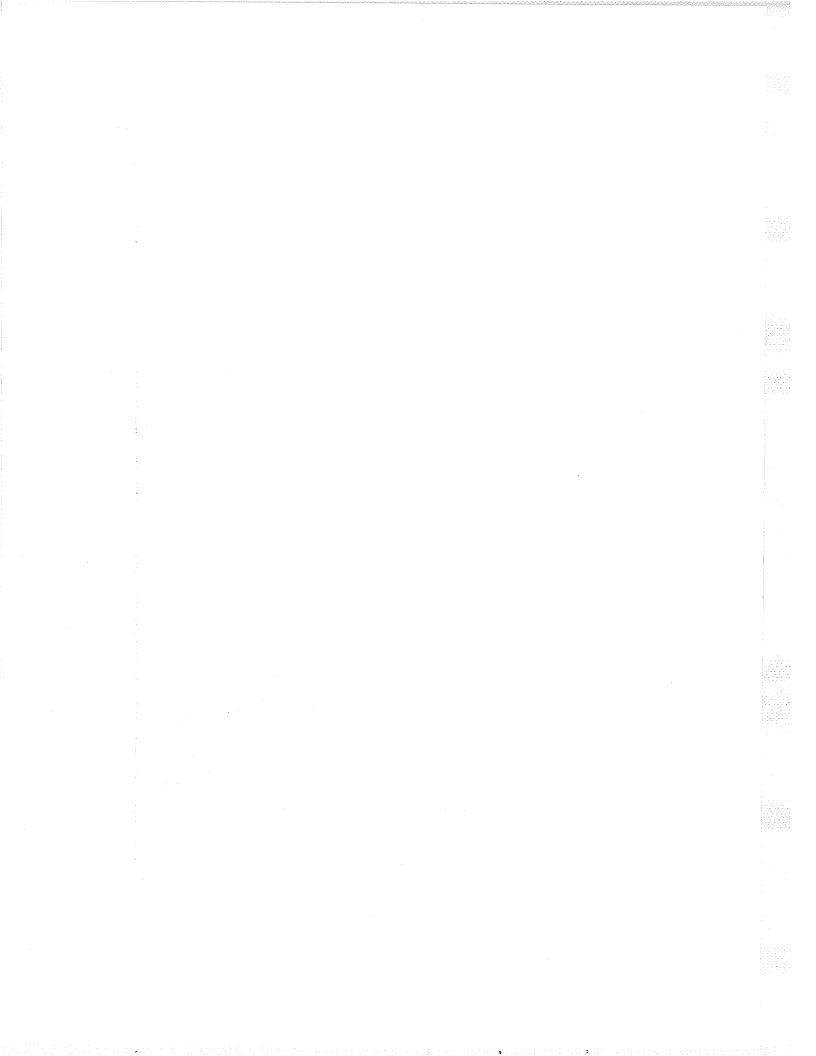
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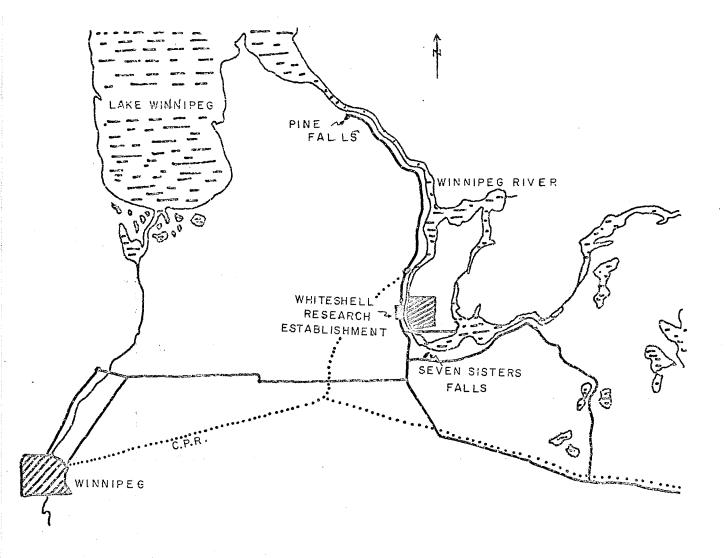
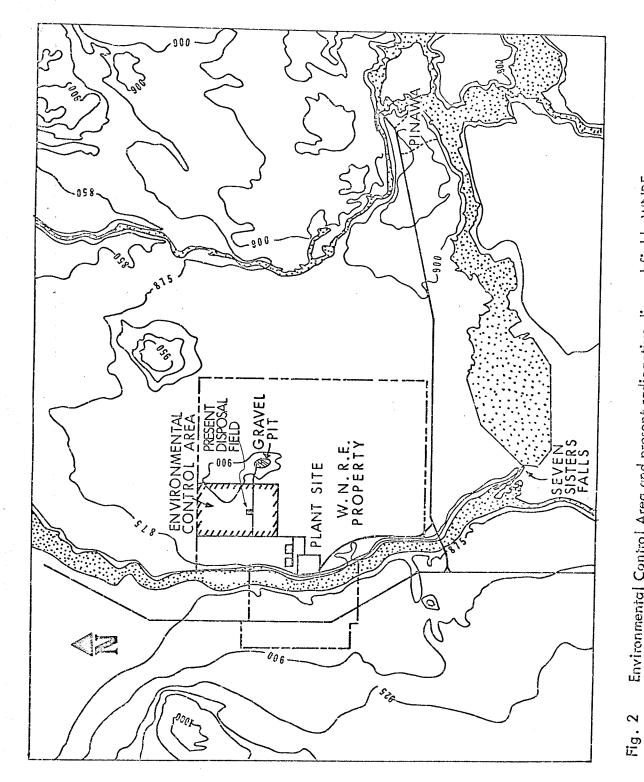


FIG. 1. Vicinty Map-Whiteshell Nuclear Research Establishment



Environmental Control Area and present radioactive disposal field, WNRE.

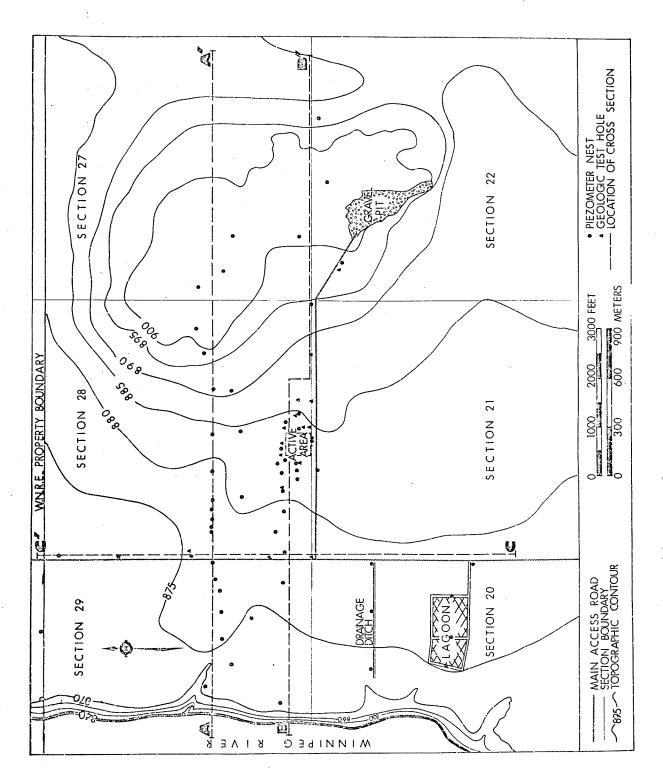
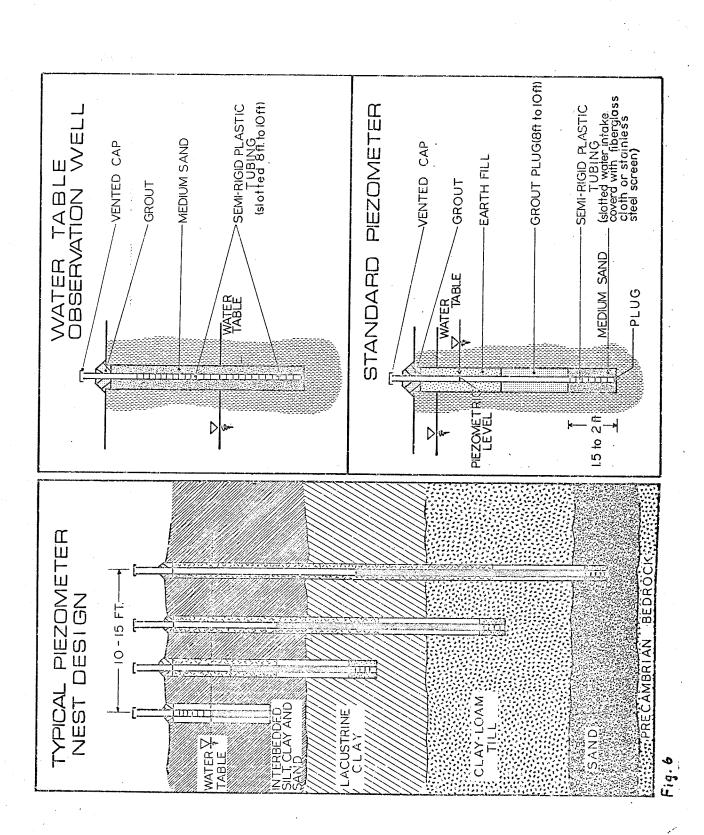


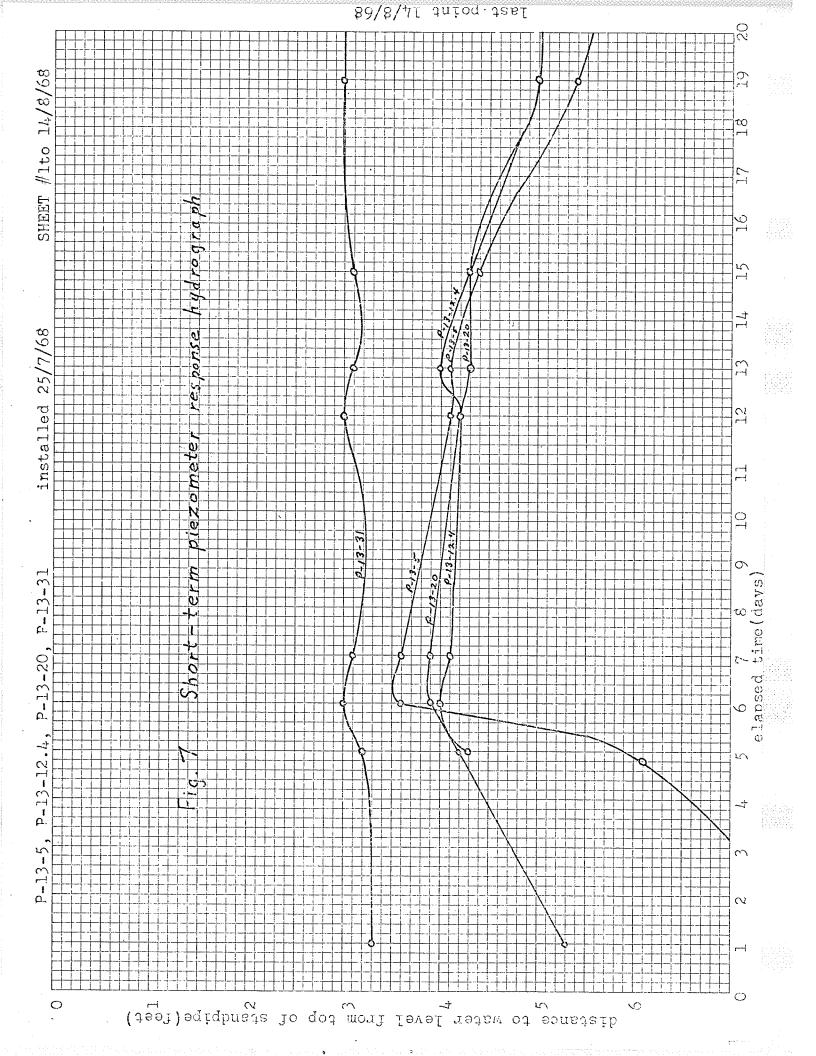
Fig. 3 Topography and location of instrumentation.

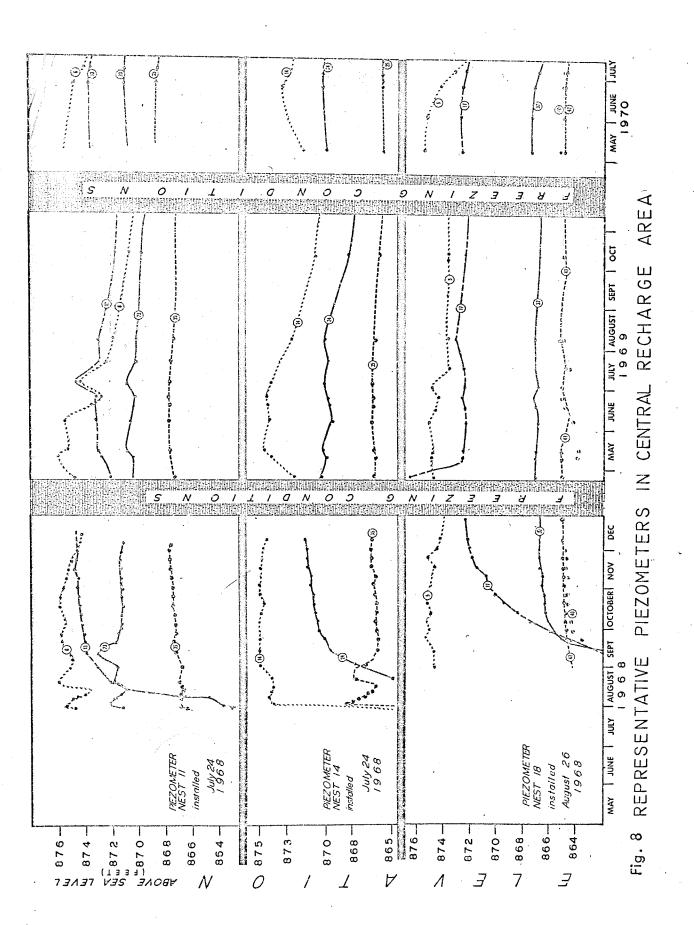


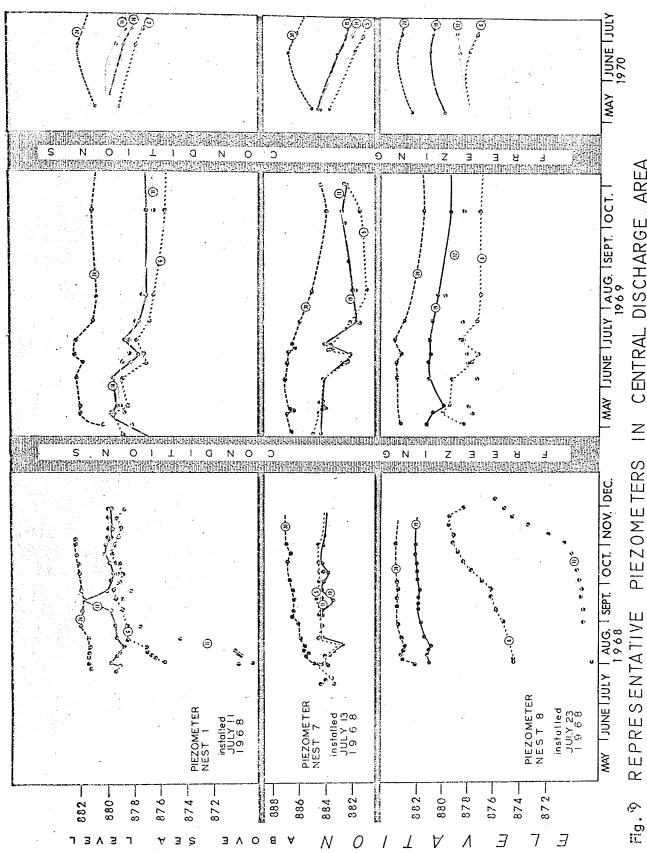
Fig. 4 Looking west towards Winnipeg River over portion of Environmental Control Area.

Fig. 5 Modified infra-red photograph of the upland area of recharge.

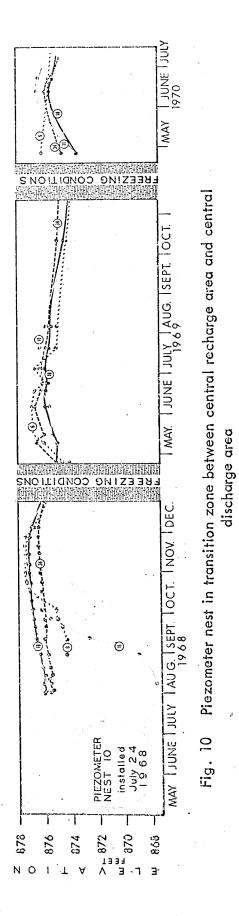


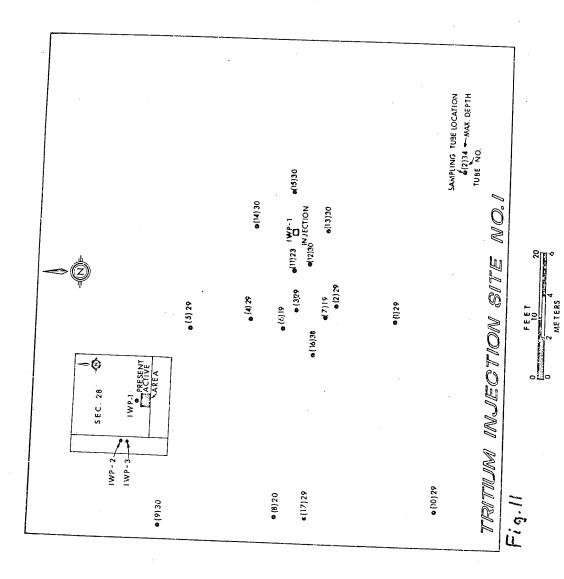






CENTRAL DISCHARGE AREA PIEZOMETERS IN





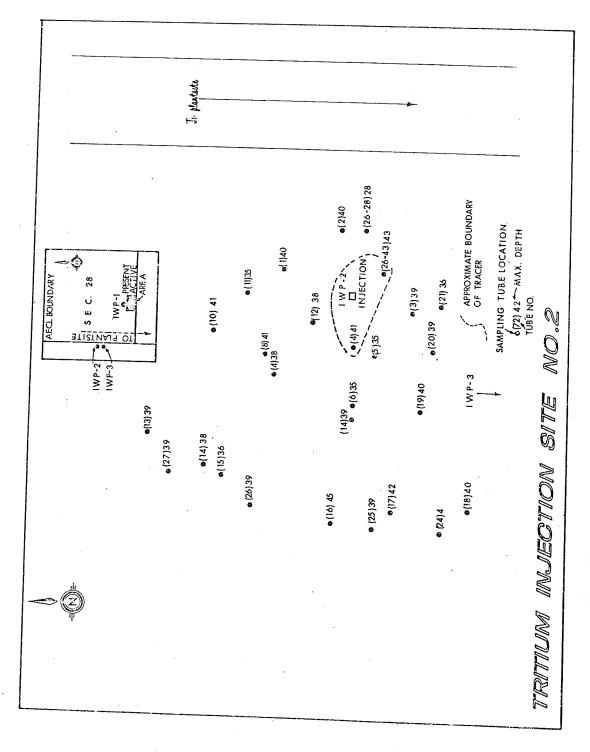


Fig. 12 Layout of tritium injection site No. 2.

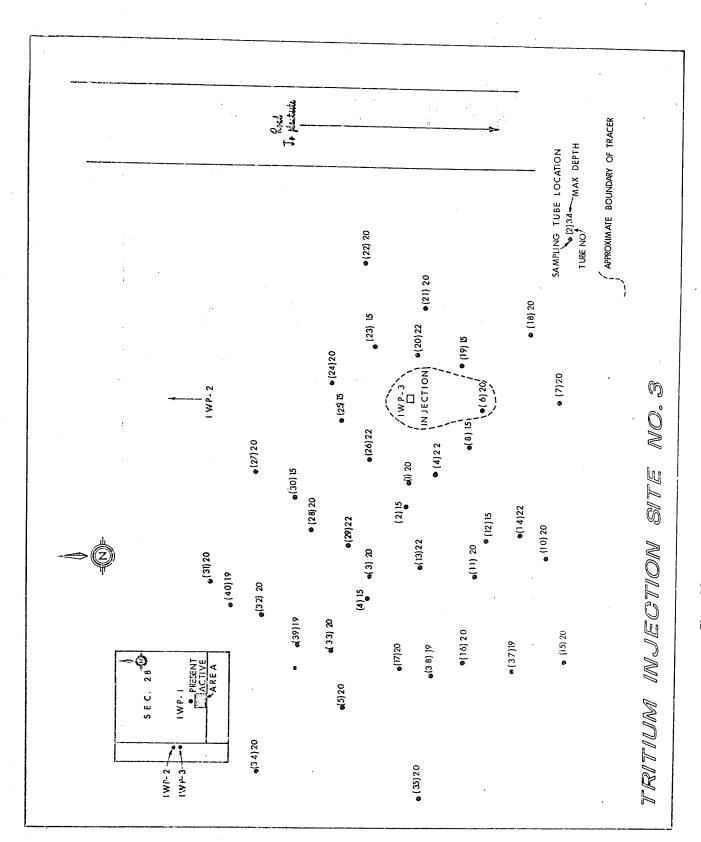


Fig. 13 Layout of tritium injection site No. 3.

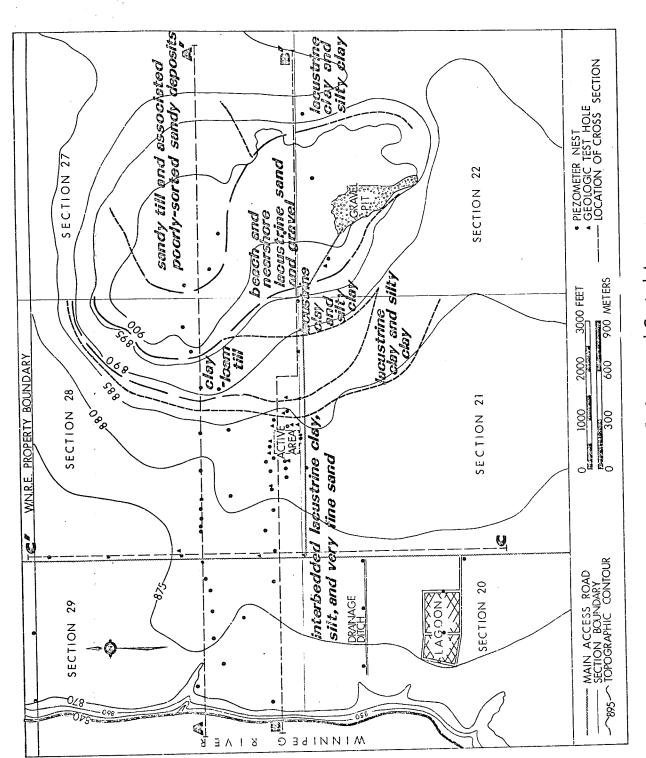


Fig. 14 Surficial geology - Environmental Control Area.

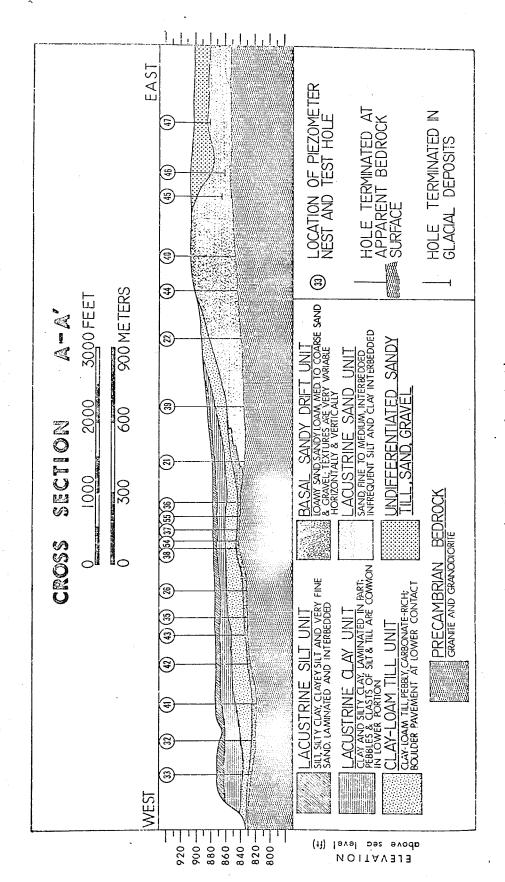


Fig. 15 Geologic cross section A-A'

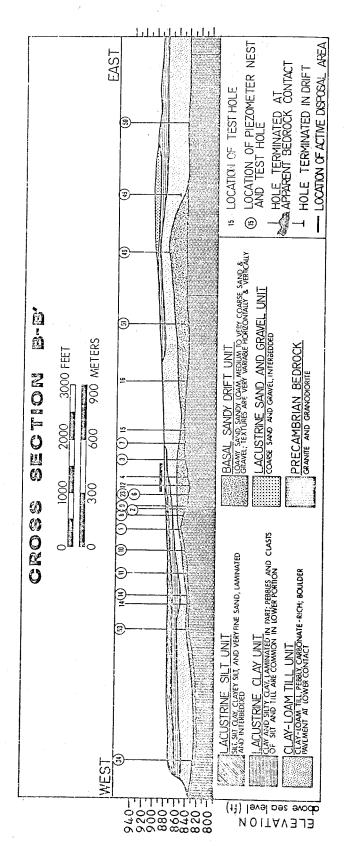


Fig. 16 Geologic cross section B-B'

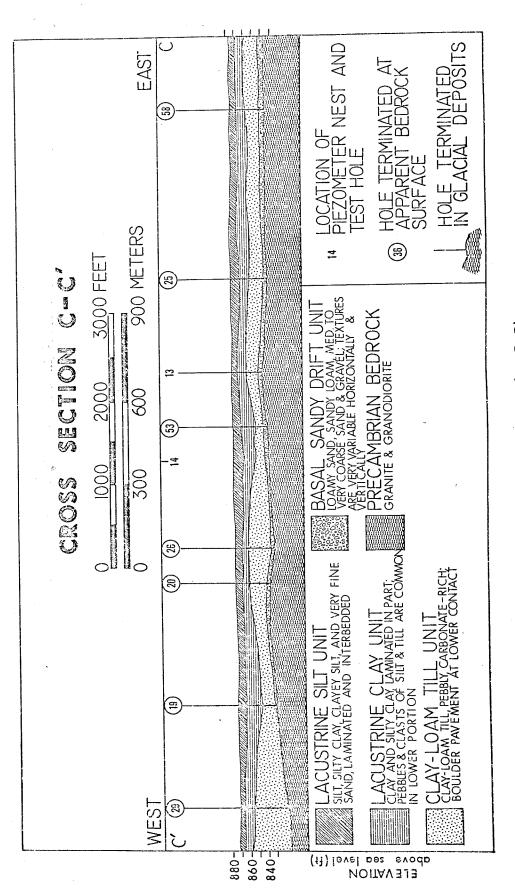


Fig. 17 Geologic cross section C-C'



Fig. 18 Soft pockets of easily eroded silt in the upper unit – possibly primary bedding structure. Sample taken from a depth of 5 feet.



Fig. 19 Ice rafted till fragment in lacustrine clay unit. This photo taken at bottom of a 9 foot cut.



Fig. 20 Angular calcareous silt fragments embedded in lower part of the clay unit – also note precipitates and rootlets on joint surfaces. This sample taken from a depth of 14 feet.



Fig. 21 Steeply dipping laminae of very fine sand in clay unit – carbonate precipitates give a whitish tone to the laminae.

Bedding is essentially horizontal.



Fig. 22 Stratified transition zone between lacustrine clay unit and the clay-loam till.

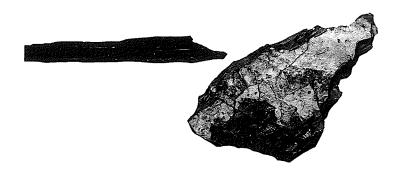


Fig. 23 Joints cutting laminae in clay unit. Joints have been accentuated artificially for illustrative purposes.

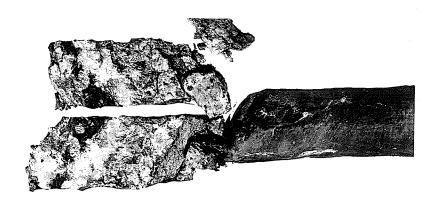


Fig. 24 Carbonate and gypsum precipitates on vertical fracture in lacustrine clay unit.

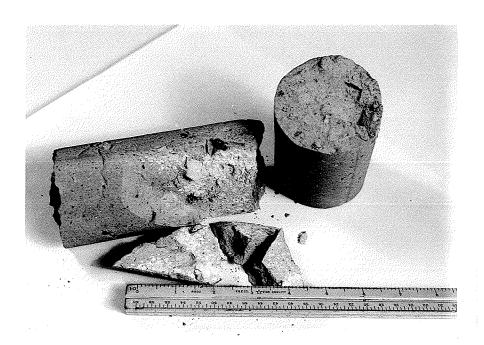
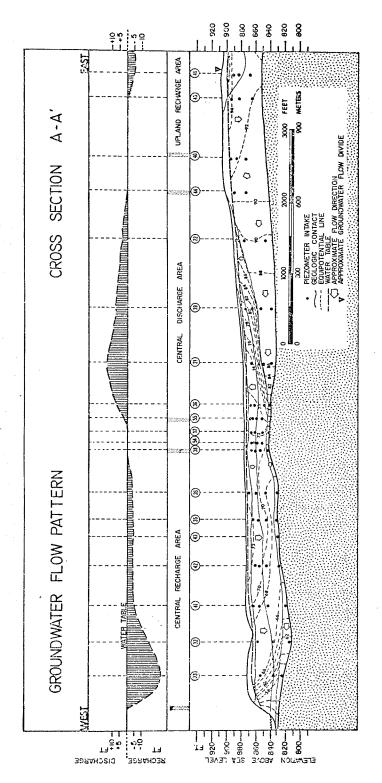
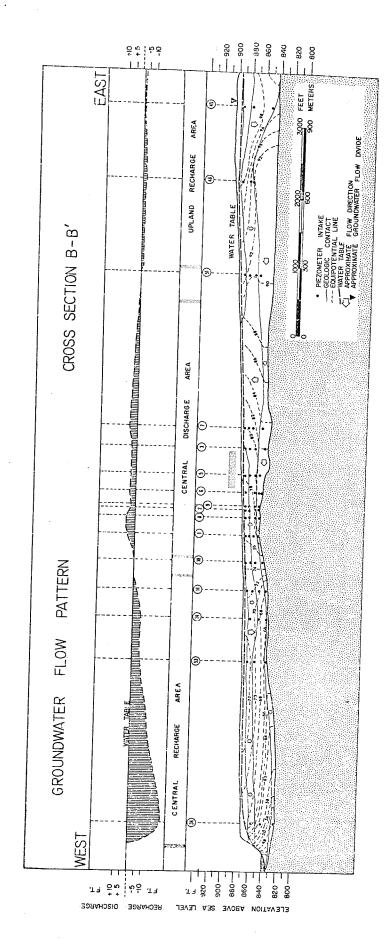


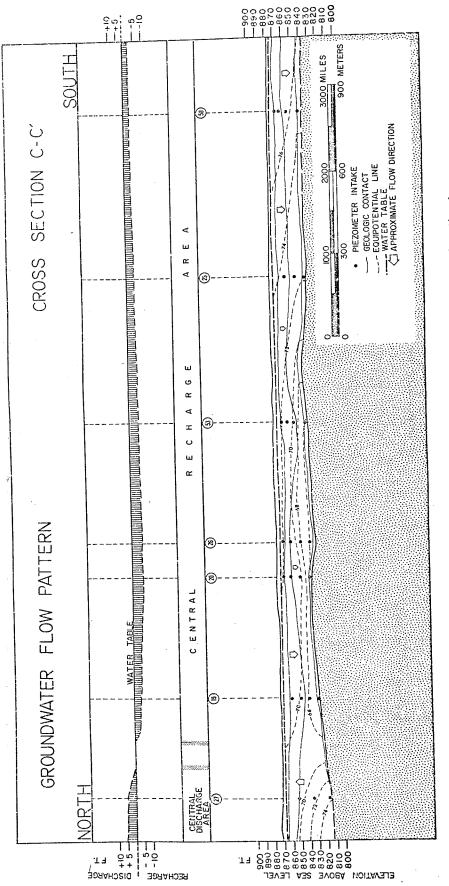
Fig. 25 Carbonate precipitates on joint surface in clay-loam till unit. Sample taken from a depth of 20 feet.



Flow pattern interpretation based on the hydraulic potential distribution Cross section A-A' Fig. 26



Flow pattern interpretation based on the hydraulic potential distribution Cross section B–B' Fig. 27



Flow pattern interpretation based on the hydraulic potential distribution Cross section C-C' Fig. 28

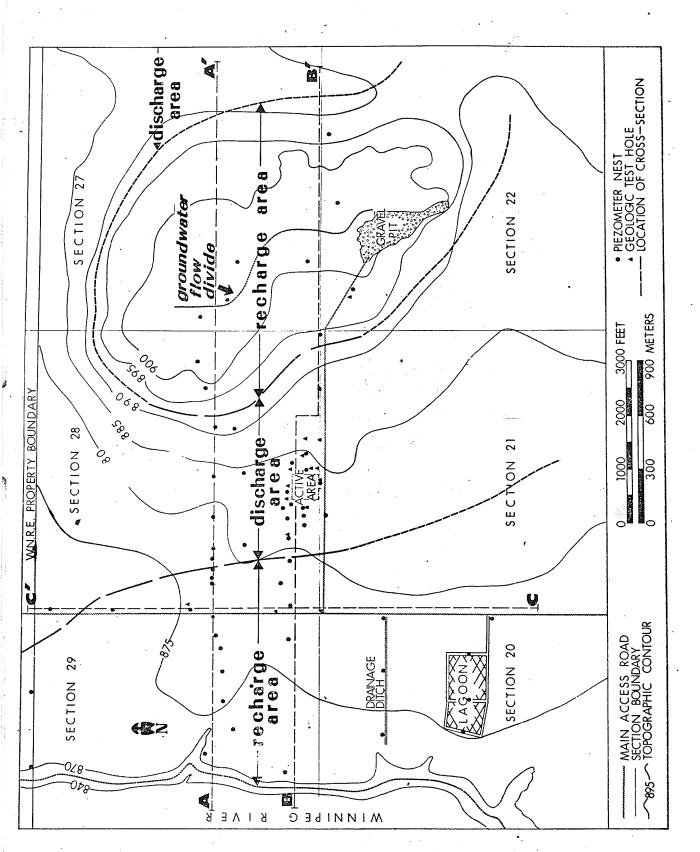
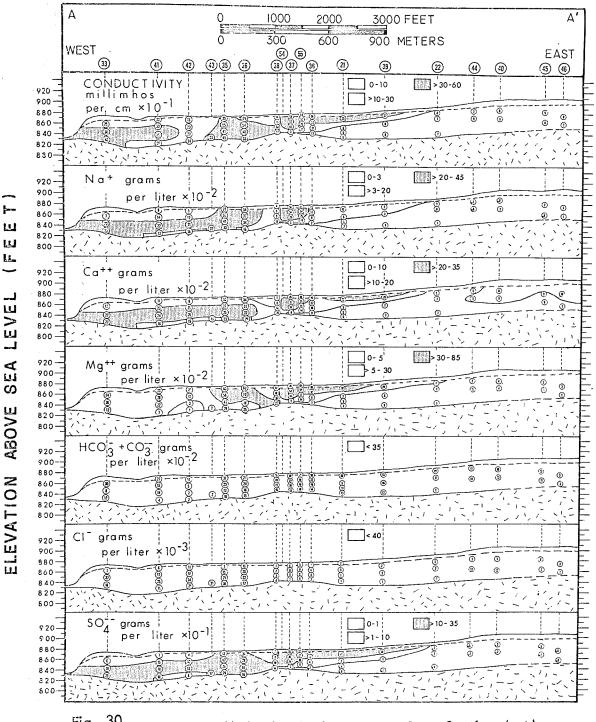


Fig. 29. Areal distribution of recharge-discharge areas.



- Fig. 30 Hydrochemical patterns - Cross Section A-A'

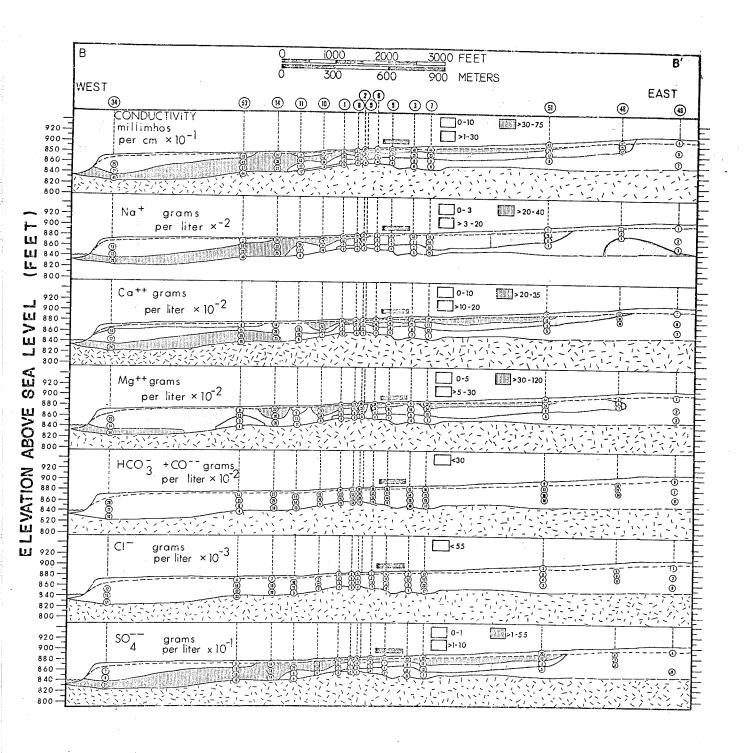
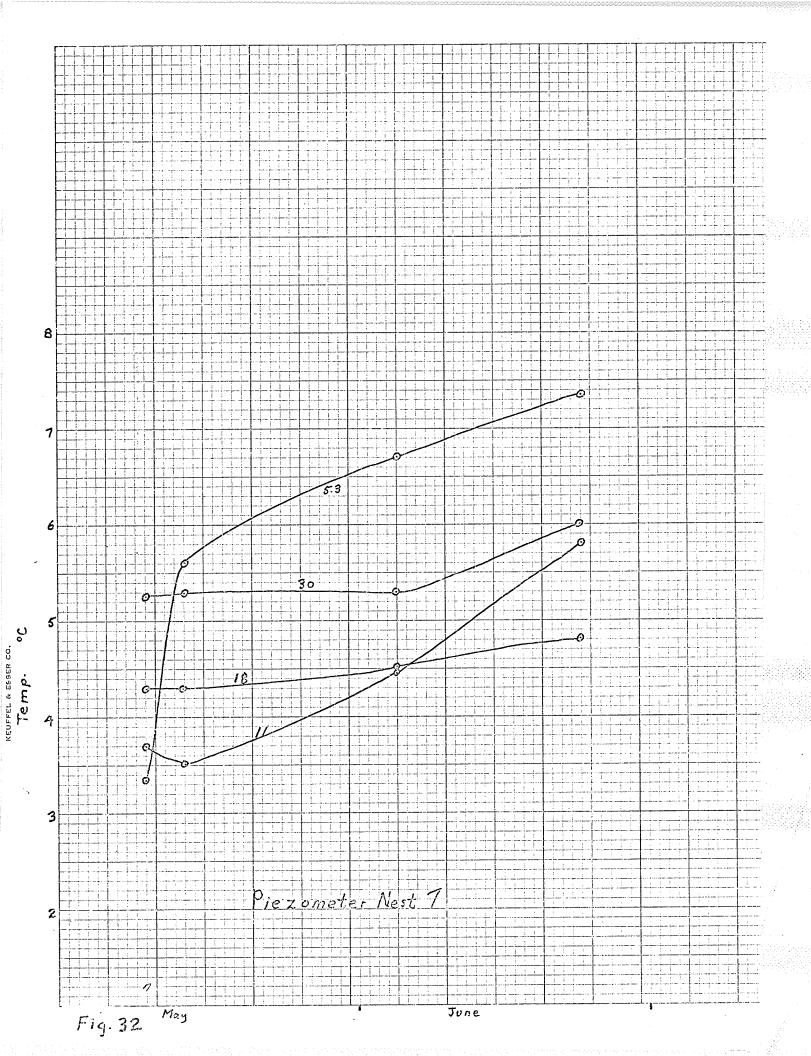
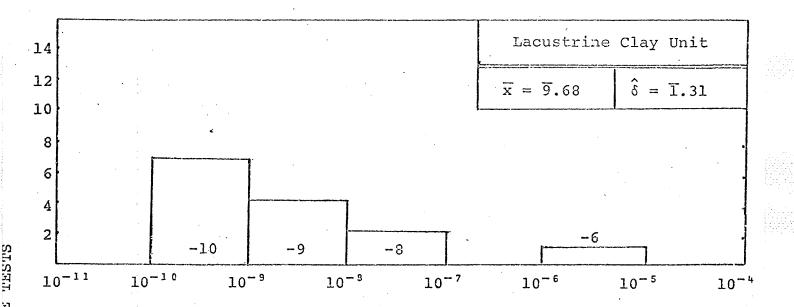
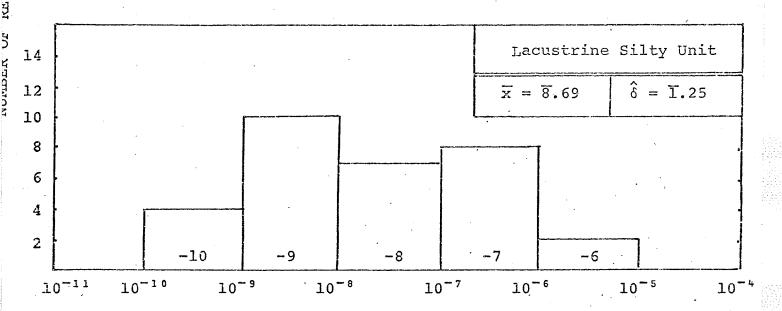


Fig. 31 Hydrochemical patterns - cross section B-B'



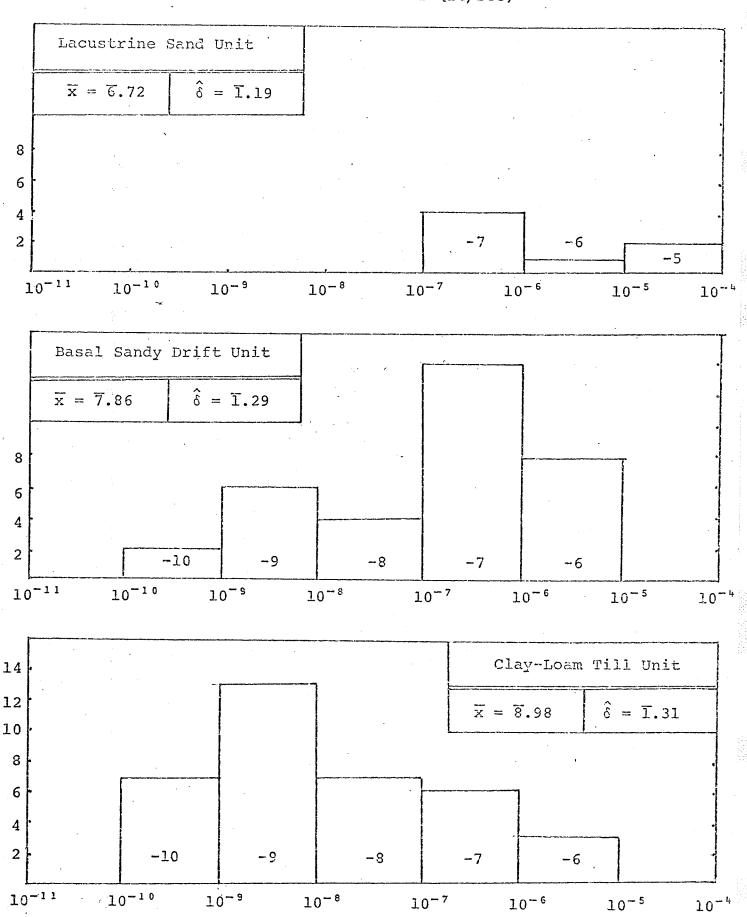




 \overline{x} = Unbiased estimate of logarithmic mean.

 $\boldsymbol{\hat{\delta}}$ = Population logarithmic standard deviation.

Fig. 33 Statistical summary of the Hvorslev hydraulic conductivity data from the lacustrine clay unit grouped according to depth intervals.



NUMBER OF RESPONSE TESTS

Fig. 3 Statistical summary of Hvorslev hydraulic conductivity data grouped according to hydrostratigraphic units.

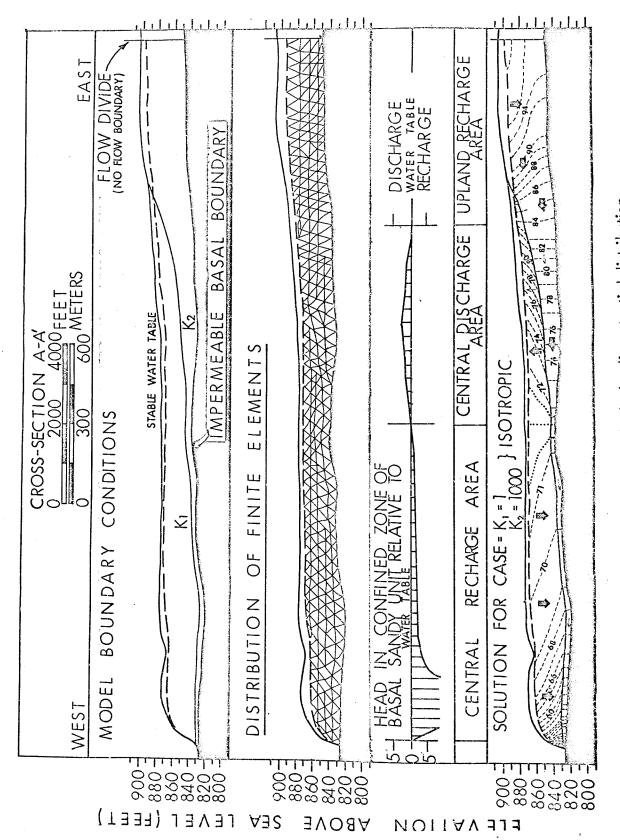


Fig. 35 'Best fit' model of the hydraulic potential distribution.

APPENDIX A

COMPILATION OF PHOTOGRAPHS



Fig. 1A Modified infra-red photograph looking east over Environmental Control Area. Outcrop of lacustrine sand and gravel unit is on the left. Present waste disposal area is in top centre.



Fig. 2A Caissons for subsurface disposal of radioactive wastes.



Fig. 3A Winter conditions in Environmental Control Area



Fig. 4A Parting along bedding in lacustrine silty unit.



Fig. 5A Jointing in lacustrine clay unit with precipitates on joint surfaces.



Fig. 6A Jointing in lacustrine clay unit without visible precipitates.



Fig. 7A | Ice-rafted clayey-silt fragment embedded in lacustrine clay unit.



Fig. 8A Clay-loam till unit.



Fig. 9A Installation of piezometers in central discharge area.



Fig. 10A Installation of piezometers in upland recharge area.



Fig. 11A PVC piezometer with intake zone wrapped with fine mesh screen.



Fig. 12A Reading water levels in piezometers in late fall after water-table freezing begins.



Fig. 13A Injection of tritium tracer at injection site No. 1.



Fig. 14A Injection vessel with level gauge.



Fig. 15A Sampling for tritium at injection site No. 1. Colouring of the water sample due to sodium fluorescein remaining in the groundwater from a trial injection.



Fig. 16A Injection site No. 2 in late fall.

APPENDIX B

GEOLOGIC LOGS OF TEST HOLES

TEST HOLE #1

ILUI IIOLL :	-	·
Depth (ft.)	Characteristics	Interpretation
0-1	Top soil.	
1-15	Finely laminated clay and silty clay; gypsum flakes and iron concretions present throughout; colour 2.5Y browns and olive browns.	lacustrine clay unit
15-16	Interlayered sandy, pebbly clay, sandy silts.	
16-36	Sandy, silty, clayey till; carbonate pebbles; colour 2.5Y light clive brown and yellowish brown near 16 feet becoming 2.5Y4/2 below.	clay-loam till
36-39	Till and interbedded medium to coarse sand; igneous pebbles and quartzo-feldspathic sand.	basal sandy drift
39	End of hole on rock.	
TEST HOLE #2	-	
Depth (ft.)	Characteristics	Interpretation
0-1	Top soil.	
1-12	Complexly laminated clay with carbonate pebbles and sand grains common in upper few feet; gypsum flakes common in upper five feet; colour 2.5Y4/2 dark grey brown with lighter grey laminae, mottled 2.5Y shades in upper 5 feet.	lacustrine clay unit
12-13.5	Layered silty sands; gradational from clay to	•

till-like material.

2.5Y4/4 near 28 feet.

Silty sandy clayey till with carbonate rock

fragments; silt content increases slightly downward; colour 2.5Y4/2 becoming

clay-loam till

13.5-28

TEST HOLE #2	! Cont'd.	
28-40	Very poorly sorted gravelly sands with till- like interbeds; rock fragments and sand of igneous origin; colour 5Y4/2 olive grey.	basal sandy drift
40	End of hole due to auger breakage.	
TEST HOLE #3		
Depth (fr.)	Characteristics	Interpretation
0-2	Top soil.	
2-12	Laminated clay with isolated pebbles and sand grains in upper portion; occasional laminae of very fine sand between 5 and 10 feet; silt content increased slightly near 12 feet; colour 2.5Y/3/2.	lacustrine clay unit
12-21	Sandy silty clayey till with carbonate pebbles; colour 2.5Y4/4.	clay-loam till
21-21.5	Coarse gravel and cobbles.)
21 .5-24	Medium to coarse pebbly sand, colour 5YN3.) basal sandy drift
24-30	Very sandy slightly silty till; colour 5YN3.)
30	End of hole on rock.	
TEST HOLE #4		
Depth (ft.)	Characteristics	Interpretation
0-1	Top soil.	
1-7	Laminated silty clay; pebble content increases with depth; till-like zones; gypsum flakes common; colour 2.5Y greys.	lacustrine clay unit

Highly variable sand-sill-clay layered complex.

7-10

TEST HOLE #4	Cont'd.	
10-23	Sandy silty clayey till with carbonate pebble colour 2.5Y3/2.	s; clay-loam till
23-4 2	Quartzo-feldspathic sand, medium to fine, with clay interbeds a few inches thick in lower 15 feet; colour 5YN/3.	basal sandy drift
42	End of hole on rock.	
TEST HOLE #5		
Depth (ft.)	Characteristics	Interpretation
0-1	Top soil.	
1-3	Clay; laminated; pebbles common; colour 2.5Y5/4.	lacustrine clay unit
3-21	Till; carbonate pebbles predominate; colour 2.5Y5/4 grading downward to 5Y5/4.	clay-loam till
21-46	Sand; medium to fine poorly sorted gravelly; igneous origin; sandy till-like layer near the upper contact and clay interbed believed present near bottom; colour 5YN3.	basal sandy drift
46	End of hole on rock.	
TEST HOLE #6		
Depth (ft.)	Characteristics	Interpretation
0-1	Top soil.	
1-6	Clay; laminated; silty, pebbly; colour 2.5Y4/	2. lacustrine clay unit
6-28	Till; carbonate pebbles predominate; becomes progressively more sandy downwards; pebble content decreases downwards; colour 2.5Y4/4 becoming 5YN/3 around 15 feet.	clay-loam till

TEST HOLE #6 Cont'd.

28-47

Very poorly sorted medium, gravelly sand, igneous crigin; below 40 feet driller reported till-like drilling condition but samples were

basal sandy drift

not obtained due to boulders.

47

End of hole on rock.

TEST HOLE #7

Depth (it)	Characteristics	Interpretation
0-1	Top soil.	
1-31	Sandy silty clayey till with carbonate pebbles; sand and silt content increasing downward; colour 2.5Y4/2.	clay-loam till
31-37	Silt, sand, pebbles, boulders; colour 5YN/3.	basal sandy drift
37	End of hole on rock.	

TEST HOLE #8

Depth (ft.)	Characteristics	Interpretation
0-1	Top soil.	
1-5	Laminated, mottled, clay; carbonate pebbles present; colour 5Y4/3.	lacustrine clay unit
5-23	Normal; till with carbonate pebbles; colour 5Y4/3 grading to olive grey.	clay-loam till
23-36	Sand; medium very poorly sorted; many cobbles; lower few feet become very fine silty sand; colour 5YN/4.	basal sandy drift
3 6 _.	End of hole.	

TEST	HOLE	#9

Depth (ft.)	Characteristics	Interpretation
0-2	Top soil.	v. v
2-5	Laminated, mottled, clay; gypsum flakes common; colour 5Y4/2.	lacustrine clay unit
5-26	Normal till with carbonate pebbles; becomes more silty in lower portions; colour 5Y5/4.	clay-loam till
26	End of hole on rock. NOTE: second hole drilled five feet from TEST HOLE #9 also ended at 26 feet on rock.	
TEST HOLE #	10	
Depth (ft.)	Characteristics	Interpretation
0-1	Top soil.	
1-5	Laminated, mottled, clay; carbonate pebbles present; colour 2.5Y light browns.	lacustrine clay unit
5-23	Till; normal; colour 2.5Y4/4.	clay-loam till
23	End of hole on rock. NOTE: second hole drilled five feet away from TEST HOLE #10 also ended at 23 feet on rock.	
TEST HOLE #1	<u>11</u>	
Depth (ft.)	Characteristics	Interpretation
0-1	Top soil.	
1-11	Clay; massive, mottled; pebbles absent; colour 2.5Y3/2.	lacustrine clay unit
11-17	Till; normal; colour 2.5Y4/4.	cloy-loam till
17	End of hole on rock.	

TEST HOLE #11 Cont'd.

New hole started 4 feet away.

13-22 Till; normal; carbonate pebbles predominate. clay-loam till

22-30 Sand; fine to very fine well sorted, no basal sandy drift

pebbles; colour 5YN/4.

30 End of hole on rock.

TEST HOLE #12

Depth (ft.)	Characteristics	Interpretation
0-2	Road fill.	
2-11	Clay; laminated, mottled, light grey silt balls (angular fragments) in dark clay.	lacustrine clay unit
11-22	Till; normal; colour 5Y4/3 grading to 5YN/4.	clay-loam till
22-40	Sand; upper three feet is coarse poorly sorted silty; 25 to 40 is fine well sorted dense clean sand; colour 5YN/4.	basal sandy drift

End of hole, no more auger available.

TEST HOLE #13

40

Depth (ft.)	Characteristics	Interpretation
0-4	Road fill.	
4-7	Very fine sand and silt.	lacustrine silty unit
7-16	Clay; massive; no pebbles noticed; many gypsum flakes; colour 5Y4/3.	lacustrine clay unit
16-18	Transition zone.	
18-29	Till; normal; carbonate pebbles predominate; colour 5Y5/3.	clay-loam till

TEST	HOLE	#13	Cont'd.

29-35 Sand; very fine silty sand interlayered with basal sandy drift black silty clay; becomes gravelly in last foot.

35 End of hole on rock.

TEST HOLE #14

Depth (ft.)	Characteristics	Interpretation
0-3	Road fill.	
3-29	Clay; layered massive and laminated zones; pebbled content increasing with depth; angular silt fragments increase with depth; gypsum flakes concentrated around 20 feet; laminae of very fine sand; colour 5Y5/4 and 5Y4/3, sand laminae 5Y7/1.	lacustrine clay unit
29-30	Clay; massive; very compact; slightly sandy; colour 5Y4/3.	lacustrine clay unit
30-38	Till; hard, very silty; carbonate and granitic pebbles; very bouldery in the last foot.	clay-loam till
38	End of hole on rock.	

TEST HOLE #15

Depth (ft.)	Characteristics	Interpretation
0-4	Road fill.	
4-11	Clay and silty clay; fine sand laminae present, pebbly.	lacustrine clay unit
≈11 -25	Till; upper portion has deficiency of sand but approaches equal proportions near lower boundary	clay-loam till Ty:
25-30	Till; slightly less clay than above and very bouldery; some portions drill like sand.	basal sandy drift
30	End of hole on rock.	

TEST HOLE #16

No drilling log available.

TEST HOLE #17

Depth (ft.)	Characteristics *	Interpretation
0-2	Top soil.	
2-8	Sand and gravel, interbedded.	lacustrine sand and gravel unit
8-16	Silty clay and clay; laminated, pebbly; colour 5Y greys and clives.	lacustrine clay unit
16-42	Till; carbonate pebbles predominate; colour 5Y4/1 dark grey.	clay-loam till
42-45	Till; slightly more silty than above; drills very bouldery.	basal sandy drift
TEST HOLE P	_1	

Depth (ft.)	Characteristics	Interpretation
0-1	Top soil.	
1-2	Silty sand.	lacustrine silt unit
2-15	Clay; laminated; many gypsum flakes and much iron staining; pebbles absent near surface but become common near 15 feet.	lacustrine clay unit
15-15.5	Gravelly layer.	
15.5-23	Till; mottled, with carbonate pebbles; upper two feet silty and gravelly, lower material normal till composition; dark grey fine sand layers present in upper 4 feet.	clay-loam till
23-30	Sand; clean fine grained with minor gravel; becomes silty and clayey near 27 and resembles a sandy till near 30 feet.	basal sandy drift
30	End of hole for piezometer installation.	

Depth (ft.)	Characteristics	Interpretation
0-1	Top soil.	
1-8 4	Clay; laminated; many gypsum flakes and iron staining; carbonate pebbles predominant over highly weathered igneous pebbles; colour 2.5Y3/2.	lacustrine clay unit
8-21	Till; normal with dark grey fine sand layers less than 0.1 inches thick; near 8 feet till oppears banded with sand and gravel; colour 2.5Y4/2.	clay-loam till
21-29	Sand; poorly sorted, gravelly; igneous fragments.	basal sandy drift
29-30	Sandy till.	

TEST HOLE P-3

Same as Test Hole #8.

TEST HOLE P-4

Same as Test Hole #10

TEST HOLE P-5

Same as Test Hole #4.

TEST HOLE P-6

Same as Test Hole #11.

TEST HOLE P-7

Same as Test Hale #6.

and samples.

than above.

boulders.

Till; normal; olive brown.

Sand; very coarse gravelly dirty; in bottom)

Sand; gravelly, bouldery; less silt and clay)

three feet approaches a sandy till.

End of hole due to difficult drilling in

15-23

23-30

30-41

41

Depth (ft.)	Characteristics	Interpretation
0-2	Top soil.	
2-14	Clay; laminated; gypsum flakes common; dark grey brown.	lacustrine clay unit
14-21	Till; silty; igneous and carbonate pebbles.	clay~loam till
21-27	Sand; very poorly sorted, gravelly, bouldery, silty; possibly sandy till.	basal sandy drift
27-29	Till; normal; mainly carbonate pebbles.	
29	End of hole on rock.	
TEST HOLE P-	<u>-9</u>	
Depth (ft.)	Characteristics	Interpretation
0-2	Top soil.	
2-9	Clay; mottled, laminated; dark grey brown.	lacustrine clay unit
9-15	Clay-till transition; based on drilling feel	

clay-loam till

basal sandy drift

Depth (ft.)	Characteristics	Interpretation
0-2	Top soil.	
2-13	Clay; generally non-mottled massive; no silt fragments; slightly silty; minor gypsum flakes.	lacustrine clay unit
13-15	Transition to till.	
15-23	Till; normal; olive brown.	clay loam till
23-32	Sand; silty, clayey; gravelly; possibly a sandy till; upper three feet are very till-like.	basal sandy drift
32	End of hole on rock.	·

Depth (ft.)	Characteristics	Interpretation
0-2	Top soil.	
2-18	Clay; laminated; lacks usual mottling and silt fragments; patches of white and grey material; colour 5Y4/3.	lacustrine clay unit
18-27	Till; normal; carbonate pebbles predominate; colour 5Y4/3.	clay-loam till
27-30	Till; becomes very sandy near 30 feet; colour) 5Y5/4.	basal sandy drift
30-36	Sand; very coarse to coarse, gravelly,) slightly silty; colour 5Y5/2.	
36	End of hole.	

Depth (ft.)	Characteristics	Interpretation	
0-2	Top soil.		
- 2-6	Clay; laminated, mottled; many gypsum flakes and angular silt fragments; colour olive grey 5Y4/3 and 5Y4/4.	lacustrine clay unit	
6-21	Till; normal; carbonate and igneous pebbles; colour 5Y4/3 grading downward to 5Y4/1.	clay-loam till	
21-25	Sand and gravel; very poorly sorted; angular) pebbles and sand grains; granitic.		
25-37	Sand; very fine to fine well sorted; a dense) silt or very fine sand layer encountered at) 36 feet.	basal sandy drift	
37-42	Gravel; based on drilling feel; no samples) obtained.		
42	End of hole on rock.		
TEST HOLE P-	13		
Depth (ft.)	Characteristics	Interpretation	
0-2	Top soil.		
2-11	Clay; massive, non-mottled; occasional pebbles; fissile; colour 5Y4/2.	lacustrine clay unit	
11-13	Clay-till transition.		
13-25	Till; more silty than normal; carbonate and igneous pebbles; colour 5Y4/3.	clay-loam till	
25 - 30	Till; slightly more sandy and silty than normal; colour 5YN4.	clay-loam till	
30-33	Sand; silty gravelly, resembling sandy till.	basal sandy drift	
33	End of hole due to auger breakage.		

Depth (ft.)	Characteristics	Interpretation
0-2	Top soil.	
2-22	Clay; laminated in some portions; fine sand layers at angles to layering; some silt fragments in lower portion; pebble content increases with depth; colour 5Y4/2 to 5Y4/3.	lacustrine clay unit
2-35	Till; normal; seems abnormally soft in places but no visible variation in grain size ratios; colour 5YN4.	clay-loam till
35	End of hole on rock.	

TEST-HOLE P-15

Same as Test Hole #12

Depth (ft.)	Characteristics	Interpretation
0-2	Top soil.	
2-16	Till; igneous and carbonate pebbles; occasional slightly silty zone; colour 2.5Y4/4 grading to 5Y3/1 near 16 feet.	clay-loam till
16-49	Sand; very fine to fine well sorted; drills as) if layered possibly with clayey interbeds;) colour 5YN4.	
49-52	Silt and clay; interbedded; colour 5YN4.	lacustrine sand unit
52-55	Sand; granitic, very coarse, poorly sorted.	
55,	End of hole on rock.	

48

End of hole on rock.

Depth (ft.)	Characteristics	Interpretation	
0-1	Top soil.		
1-13	Clay; laminated; many gypsum flakes; silt fragments increase near 13 feet; colour 2.5Y3/2.	lacustrine clay unit	
13-21	Till; normal; colour 2.5Y5/4.	clay-loam till	
21-34	Sand; very fine; well sorted; silt and clay interbeds; colour 5YN4; sand becomes coarse and poorly sorted from 31-34 feet.	basal sandy drift	
34	End of hole on rock.		
TEST HOLE P-	18		
Depth (ft.)	Characteristics	Interpretation	
0-4	Top soil and fill.		
4-11	Clay; laminated to massive; colour 5Y6/2,) 5Y5/4 and 5Y4/3.		
11-27	Silty clay; very soft; appears bedded; colour) dark grey 5Y4/1.	lacustrine clay unit	
27-36	Till; more silty than normal; very plastic; colour 5Y4/1.	clay-loam till	
36-48	Sand; from 36 to 41 feet very fine silty sand, well sorted from 41 to 48 feet medium poorly sorted silty pebbly sand with boulders; definitely not a till.	basal sandy drift	

Depth (ft.)	Characteristics	Interpretation	
0-5	Top soil and road fill.		
5-27	Silty clay and clay; upper 10 feet somewhat till-like; colour 5Y5/3; below 10 feet generally massive; pebbles common; colour 5Y5/3 and 5Y4/1.	lacustrine clay unit	
27-42	Till; less clay than normal; carbonate and igneous pebbles common; colour 5Y5/1.	clay loam till	
42-44	Sand?; driller reported sand but no sand sample was obtained.	basal sandy drift	
44	End of hole on rock.		
TEST HOLE P-	20		
Depth (ft.)	Characteristics	Interpretation	
0-3	Top soil.		
3-17	Clay; massive to slightly laminated; silt fragments common.	lacustrine clay unit	
17-36	Till; normal; igneous and corbonate pebbles; colour 5Y4/1.	clay-loam till	
36-39	Till; very bouldery and sandy; sand interbeds suspected.	basal sandy drift	
39	End of hole on rock.		
TEST HOLE P-21			
Depth (ft.)	Characteristics	Interpretation	
0-2	Top soil.		
2-17	Clay; many selenite crystal concentrations; laminated grading to more massive in lower portions; occasional pebble; colour 5Y5/4 and 5Y4/1.	lacustrine clay unit	

TEST	HOLE	P-21	Cont'd.

	A.A.B.CAD	
17-25	Till; normal; igneous and carbonate pebbles; colour 5Y4/1 dark grey.	clay-loam till
25-45	Sand; silty coarse very poorly sorted; layered; colour 5YN5.	basal sandy drift
45	End of hole on rock.	
TEST HOLE P-	<u>22</u>	
Depth (ft.)	Characteristics	Interpretation
0-1	Top soil.	
1-9	Till; normal; igneous and carbonate pebbles; colour 5Y5/3.	clay-loam till
9-10	Boulders.	
10-43	Sand; fine, well sorted; colour 5Y5/1.	lacustrine sand
43-53	Sand; granitic; coarse pebbly; no silt or clay.)	unit
53	End of hole on rock.	
TEST HOLE P-	23	
Depth (ft.)	Characteristics	Interpretation
0-1	Top soil.	
1-9	Clay.	lacustrine clay unit
9-11	Clay-till transition.	
11-21	Till; normal; colour 5Y4/3.	clay-loam till
21-24	Bouldery.	
24-37	Sand; coarse very poorly sorted; gravelly, bouldery; possible slightly silty; igneous origin.	basal sandy drift

TEST HOLE P-23 Cont'd.

37-41 Till; very bouldery; more sandy and silty than basal sandy drift normal; colour 5YN4.

41 End of hole on rock.

TEST HOLE P-24

Depth (ft.)	Characteristics	Interpretation
0-3	Fill.	
3-10	Sili and clayey silt; laminated; occasional pebble; colour - greys and olives 5Y.	lacustrine silt unit
10-24	Clay and silty clay; no pebbles apparent; laminated.	lacustrine clay unit
24-51	Till; silty; less pebbles than usual; colour - dark grey.	clay-loam till
51-56	Sand; isolated pebbles and cobbles; slightly silty; has appearance of till in places in terms of pebble distribution.	basal sandy drift
5 6	End of hole on rock.	

Depth (ft.)	Characteristics	In terpretation
0-3	Fill.	
3-10	Silts and clayey silts; laminated; greys and olives 5Y.	lacustrine silt unit
10-23	Clays and silty clays; grades into above; laminated.	lacustrine clay unit
23-38	Till; slightly more silty and less clayey than usual; colour 5Y4/1.	clay-loam till
38-43	Till; very silty and sandy; igneous pebbles; bouldery; difficult drilling.	basal sandy drift
43	End of hole on rock.	

Depth (ft.)	Characterist	ics	Interpretation
0-2	Top soil.		
2-10	Silty and clayey silt; lami pebbles and silt fragments, and olives 5Y.		lacustrine silt unit
10-15	Clay and silty clay; grade calcite concentrations (?)		lacustrine clay unit
15-42	Till; normal; igneous and colour 5Y4/1.	carbonate pebbles;	clay-loam till
42-45	Sand; medium, pebbly; me sorted; clean to slightly si or coarse gravel.		basal sandy drift
45	End of hole on rock.		

Depth (ft.)	Characteristics	Interpretation
0-2	Top soil.	
2-27	Clay and silty clay; laminated; colour 5Y4/2 and 4?4 becoming 5Y4/1 near base; upper few feet are very silty.	lacustrine clay unit
27-56	Till; clayey silty sandy; carbonate pebbles predominant; in places clay content drops; colour 5Y4/1.	clay-loam till
56-59	Till; very pebbly, sandy, bouldery; igneous material.	basal sandy drift
59	End of hole on rock.	

Depth (ft.)	Characteristics	Interpretation
0-2	Top soil.	
2-8	Silt and silty clay; laminated; olive 5Y colouring.	iacustrine silt unit
8-24	Clay and silty clay; laminated in places; oxidized near 8 feet grading rapidly to 5Y4/1 unoxidized.	lacustrine clay unit
24-75	Till; generally normal with minor variations in silt-sand-clay ratios as well as pebble content; pebbles mainly carbonate; no sign of sand or gravel interbeds.	clay-loam till
75-77	Clay; laminated; occasional pebble; gypsum concentrations.	
77	End of hole on rock.	

Depth (ft.)	Characteristics	Interpretation
0-4	Road Fill.	
4-19	Clay and silty clay; laminated to massive; occasional pebble; gypsum flakes; olive 5Y colouring.	lacustrine clay unit
19-56	Till; carbonate pebbles predominate; colouring 5Y4/1.	clay-loam till
56	End of hole on rock.	
TEST HOLE P-	31	

Depth (ft.)	Characteristics	Interpretation
0-3	Top soil and peat.	

TEST	HOLE	P-31	Cont'd.
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TEST HOLE P	TEST HOLE P-3! Cont'd.				
3-15	Silt and clayey silt; olives and yellow browns; massive.	lacustrine silt unit			
15-37	Clay; soft, slightly silty; grades into the above; dark grey 5Y.	lacustrine clay unit			
37-42	Till; normal; soft; dark grey 5Y.	clay-loam till			
42-45	Bouldery; no sample, auger broke in hole.	basal sandy drift			
45	End of hole due to auger breakage.				
TEST HOLE P-32					
Depth (ft.)	Characteristics	Interpretation			
0-2	Top soil.				

2-10	Silt and sand; very fine sands and silts; laminated; iron oxide staining; olive, yellow, and yellow brown colour.	lacustrine silt unit
10-32	Clay; slightly silty; silt fragments; generally massive.	lacustrine clay unit
32-46	Till; normal; mainly carbonate pebbles; soft; dark grey 5Y.	clay-loam till

46-59	Till; sandy, silty, becoming very sandy near bottom; igneous pebbles and sand grains pre-	basal sandy drift
	dominate; dark grey 5Y; very bouldery.	

59 End of hole on rock.

Depth (ft.)	Characteristics	Interpretation
0-2	Top soil.	
2-10	Silt and clayey silt; very fine sand laminations; yellow and yellow brown colouration.	lacustrine silt unit

TEST	HOLE	P-33	Cont'd.
-			

10-35	Clay; slightly silty; dark grey 5Y.	lacustrine clay unit
35-42	Till; normal; mainly carbonate pebbles; soft; dark grey 5Y.	clay-loam till
42-48	Till; sandy, silty becoming very sandy near bottom; igneous pebbles, bouldery; dark grey 5Y.	basal sandy drift
48	End of hole on rock.	

Depth (ft.)	Characteristics	Interpretation
0-2	Top soil.	
2-13	Silt and very fine sand; olives, yellows, and yellow brown.	lacustrine silt unit
13-24	-Clay; slightly silty.	lacustrine clay unit
24-46	Till; silty, sandy slightly clayey; very soft; dark grey.	clay-loam till
46-49	Till; sandy, silty; igneous pebbles; contact) noted due to sudden increase in boulder) content.	basal sandy drift
49-51	Till; as above but very pebbly.	

Depth (ft.)	Characteristics	Interpretation
0-2	Top soil.	
2-8	Silts and very fine sand; laminated; yellows, browns, olives.	lacustrine silt unit
8-15	Clay; slightly silty; many silt fragments and gypsum flakes.	lacustrine clay unit

TEST	HOLE	P-35	Cont'd.

	· · · · · · · · · · · · · · · · · · ·		
15-17	Transition to till.	Ţ.	
17-38	Till; clayey, silty, sandy; dark grey unoxidized; soft.	clay-loam till	
38-43	Till; sandy (medium to fine) slightly silty, pebbly, bouldery; igneous pebbles.	basal sandy drift	
43	End of hole on rock.		
TEST HOLE P	-36		
Depth (ft.)	Characteristics	Interpretation	
0-2	Top soil.		
2-6	Silts and very fine sand.	lacustrine silt unit	
6-19	Clay; laminated; oxidized olives 5Y; many silt fragments; selenite concentrations; occasional pebble; becoming oxidized dark grey 5Y near bottom.	lacustrine clay unit	
19-22	Clay-till transition.		
22-31	Till; clayey silty, sandy; carbonate pebbles predominate; dark grey 5Y.	clay-loam till	
31-37	Till; sandy, silty, pebbly; silt content decreased downwards; igneous pebbles.	basal sandy drift	
37	End of hole on rock.		
TEST HOLE P-37			
Depth (ft.)	Characteristics	Interpretation	
0-2	Top soil.		
2-8	Silt; laminations of very fine sand and clay.	lacustrine silt unit	

TEST HOLE P-37 Cont'd.			
8-16	Clay; slightly silty; occasional pebble; silt fragments; selenite concentrations; oxidized olives 5Y.	lacustrine clay unit	
16-17	Transition to till.		
17-34	Till; clayey silty sandy; very soft; mainly carbonate pebbles.	clay-loam till	
34-36	Till; silty, pebbly, bouldery; igneous pebbles.	basal sandy drift	
36	End of hole on rock.		
TEST HOLE P-	38	•	
Depth (ft.)	Characteristics	Interpretation	
0-2	Top soil.		
2-4	Silt and very fine sand; laminated.	lacustrine silt, unit	
4-14	Clay; slightly silty; occasional pebble; silt fragments.	lacustrine clay unit	
14-16	Transition to till.		
16-31	Till; normal; mainly carbonate pebbles;	clay-loam till	
31-33	Till?; very bouldery; no samples as auger broke in hole.	basal sandy drift	
33	End of hole in boulders.		
TEST HOLE P-	39		

Depth (ft.)	Characteristics	Interpretation
0-2	Top soil.	
2-4	Silts and very fine sand.	lacustrine silt unit

TEST I	OF	LE	P-39	Cont'd.
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4-8	Clay; silty; silt fragments common; oxidized olive 5Y.	lacustrine clay unit
8-24	Till; normal; carbonate pebbles predominate.	clay-loam till
24-42	Sand; layered very fine silty sand and medium) sand; no pebbles; occasional thin clay) interbed.	lacustrine sand
42-46	Medium to coarse well-sorted sand.	unit
46	End of hole in boulders.	·

Depth (ft.)	Characteristics	Interpretation
0-1	Top soil.	
1-56	Sand; clean, well sorted fine to medium sand; no pebbles; oxidized to 18 feet.) lacustrine sand
56-57	Clay; layered, very silty, black; angular carbonate pebbles.) unit))
57	End of hole on rock	

Depth (ft.)	Characteristics	Interpretation
0-1	Top soil.	
1-7	Silty clay, clayey silt and very fine sand; interbedded, laminated; oxidized.	lacustrine silt; unit
7-20	Clay, lacustrine massive.	lacustrine clay unit
20-50	Till, clayey, silty, sandy, with carbonate pebbles; 5Y5/1 grey.	clay-loam till
50-54.5	Till, very sandy silt; igneous pebbles.	basal sandy drift
54.5	End of hole on solid rock.	

Depth (ft.)	Characteristics	Interpretation
0-1	Top soil.	
1-6	Silt, silty clay and very fine sand; interbedded, laminated; gradational lower contact.	lacustrine silt unit
6-18	Clay, lacustrine; olive grey 5Y becoming dark grey; occasional pebble and silt ball.	lacustrine clay unit
18-22	Transition to till.	•
22-43	Till; olive grey 5Y4/2 grading to dark grey 5Y4/1; clayey, silty, sandy; very plastic.	clay-loam till
43-45.5	Till; silty sandy; igneous pebbles.	basal sandy drift
45.5	End of hole in boulders.	
TEST HOLE P-	43	
Depth (ft.)	Characteristics	Interpretation
0-1	Topsoil.	
1-6	Silt, clayey silt and very fine sand; interbedded, laminated; oxidized.	lacustrine silt unit
6-18	Clay, lacustrine.	lacustrine clay unit
1823	Transition to till.	•
23-40	Till; silty, sandy, slightly clayey; carbonate pebbles.	clay-loam till
40-44	Sand, medium, poorly sorted; upper foot bouldery; slightly pebbly.	basal sandy drift
44	End on solid rock.	

Depth (ft.)	Characteristics	Interpretation
0-53	Sand, very fine to fine; massive, poorly sorted; no pebbles, few fines.	lacustrine sand unit
TEST HOLE P-	-45	
Depth (ft.)	Characteristics	Interpretation
0-1	Top soil.	
1-6	Till; silty, sandy, slightly clayey; igneous pebbles.	
6-43	Sand; fine to very fine; poorly sorted, no fines; oxidized to 20 feet; no pebbles; no clay interbeds observed.	lacustrine sand unit
TEST HOLE P-	46	
Depth (ft.)	Characteristics	Interpretation
0-1	Top soil.	
1-4	Sand; coarse to medium with occasional) pebble; oxidized.	
4-8		
. •	Sand; coarse; very poorly sorted; pebbly,) slightly silty; oxidized.	
8-17		undifferentiated sandy till
	slightly silty; oxidized.) Silt; sandy, pebbly; and sand; fine, silty;)	undifferentiated sandy till
8-17	slightly silty; oxidized.) Silt; sandy, pebbly; and sand; fine, silty;) oxidized olive 5Y; pebbly.) Silt and sand; interlayered; unoxidized	undifferentiated sandy till

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1601	HOLE	1 7/

Depth (ft.)	Characteristics	Interpretation
0-1	Top soil.	
1-10°	Till; silty, sandy, slightly clayey; oxidized olive 5Y; gypsum and carbonate concentrations) -)
10-18	Till; as above but more sandy and bouldery; very little clay.) Undifferentiated sandy till)
18-21	Sand; very coarse, pebbly, bouldery; oxidized.)
21	End of hole in boulders.	
TEST HOLE P-	-48	
Depth (ft.)	Characteristics	Interpretation
0-1	Top soil.	
1-4	Sand; medium to coarse; moderately sorted; no fines.	
4-7	Clay; lacustrine; occasional pebble; massive to slightly laminated; silt balls.	lacustrine clay unit
7-24	Till; silty, sandy, slightly clayey;	clay-loam ti!!
24-26	Sand; silty, pebbly, bouldery; dark grey 5Y	basal sandy drift
TEST HOLE P-	49	
Depth (ft.)	Characteristics	Interpretation
0-1	Top soil.	
*-1-10	Sand; upper 5 feet medium grained moderately sorted, 2.5YR3/6; lower 4 feet slightly silty with occasional pebble.	lacustrine sand and gravel unit
10-23	Clay, lacustrine; massive; dary grey 5Y3.5/1.	lacustrine clay unit
23-53	Till; clayey, silty, sandy; very plastic.	clay-loam till

IEST	HO	LE	P-50

Depth (ft.)	Characteristics	Interpretation	
0-2	Top soil.		
2-8	Clay; lacustrine; massive to laminated; abundant silt balls; occasional pebble.	lacustrine clay unit	
8-15	Transition to till.		
15-47	Till; clayey, silty, sandy; carbonate and igneous pebbles; olive 5Y to 20 feet and dark grey 5Y below.	clay-loam till	
47	End of hole on boulder or bedrock.		
TEST HOLE P-	51		
Depth (ft.)	Characteristics	Interpretation	
0-3	Road fill.		
3-7	Clay; lacustrine; silt balls and occasional pebble; gypsum concentrations; massive to slightly laminated.	lacustrine clay unit	
7-10	Transition to till.		
10-25	Till; clayey, silty, sandy; igneous and carbonate pebbles.	clay-loam till	
25-27	Boulders.		
27-50	Sand; medium grading to coarse near bottom; pebbly; poorly sorted near bottom; no fines.	basal sandy drift	
50	End of hole in boulders.		
TEST HOLE P-52			
Depth (fr.)	Characteristics	Interpretation	
0-2	Road fill.		

TEST HOLE P-52 Cont'd.			
2-12	Silty clay and clay; iaminae of very fine sand; pebbly with silt balls.	lacustrine silt unit	
~12 - 30	Clay; lacustrine; silt balls and occasional pebble; massive to slightly laminated.	lacustrine clay unit	
30-40	Till; clayey, silty, sandy; plastic and soft.	clay-loam till	
40	End of hole on boulder.		
TEST HOLE P-	-53		
Depth (ft.)	Characteristics	Interpretation	
0-2	Top soil.	•	
2-7	Clayey silt and silty clay; laminae of very fine sand.	lacustrine silt ynit	
7-12	Transition to well sorted clay.		
12-27	Clay; laminated to massive; silt balls and occasional pebble.	lacustrine clay unit	
27-35	Till; clayey, silty, sandy; very soft.	clay-loam till	
35-37	Till and boulders; very sandy.	basal sandy drift	
37	End of hole in boulders.		
TEST HOLE P-	-54		
Depth (ft.)	Characteristics	Interpretation	
0-2	Top soil.		
2-5	Clayey silt with laminations of very fine sand.	lacustrine silt unit	
5-14	Clay; lacustrine; silt balls and occasional pebble; massive to slightly laminated.	lacusirine clay unit	
14-26	Till; clayey, silty, sandy.	clay-loam till	

26

End of hole on boulder.

TEST: HO) LE	P-	55
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19-30

30-41

	Contract Con		
Depth (ft.)	Characteristic	s	Interpretation
0-2	Top soil.		
2-3	Clayey silt.	16 .	lacustrine silt u
3-15	Clay; lacustrine; silt balls and occasional pebble; secondary gypsum concentrations.		lacustrine clay u
15-27	Till; clayey, silty; soft, oxid	dized.	basal sandy drift
TEST HOLE P	-56		
Depth (ft.)	Characteristics	5	Interpretation
0-7	Silty clay and clayey silt; lo fine sand; lower portion has occasional pebble and silt bo	less silt,	lacustrine clay u
7-23	Till; clayey, silty, sandy.		clay-loam till
23-26	Boulders and sand.		basal sandy drift
26	End of hole on rock.		•
TEST HOLE P	-57		
Depth (ft.)	Characteristics	5	Interpretation
0-4	Silty clay and clayey silt; lovery fine sand.	uminated with	lacustrine silt ur
4-10	Clay; massive with occasions	al pebble.	lacustrine clay ur
10-19	Till; clayey, silty, sandy.	· · · · · · · · · · · · · · · · · · ·)
70.00	▼ 253 1 11. • 1	1.1.1) clay-loam till

Till; very sandy silty; mixed pebble composition; moist; soft.

Sand; possibly silty, very sandy till.

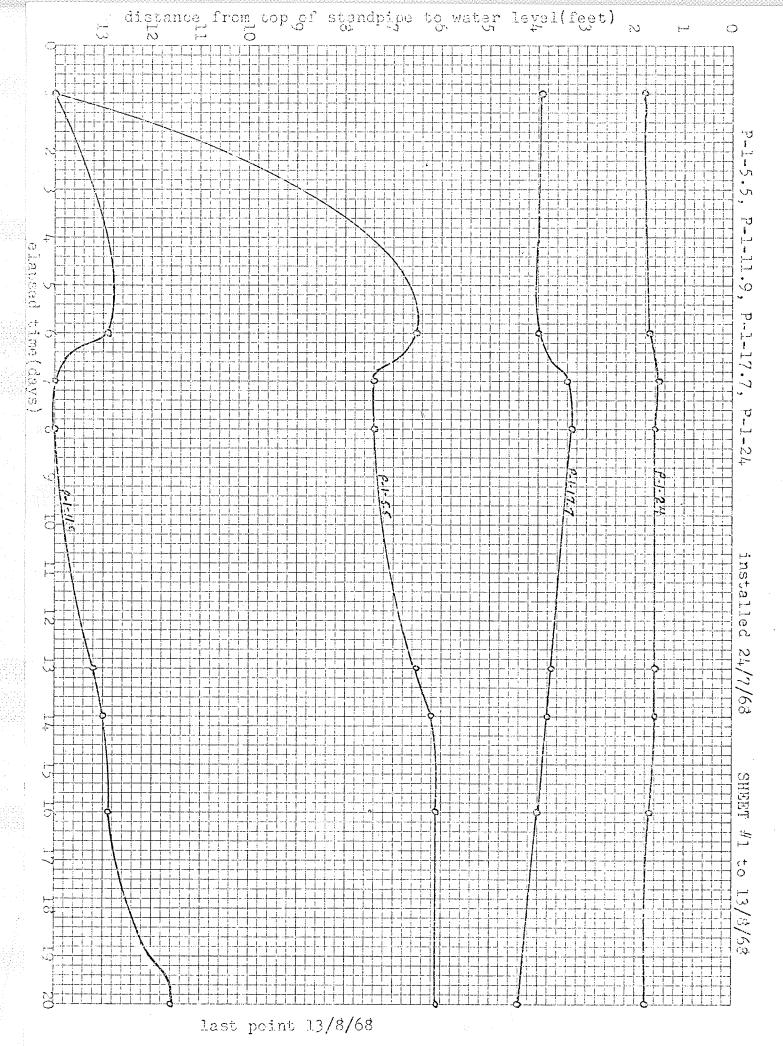
basal sandy drift

Depth (ft.)	Characteristics		Interpretation
0-1	Top soil.		
1-8	Silty clay.		lacustrine silt unit
·8-17	Clay; lacustrine; silt balls and gypsum concentrations; laminated.		lacustrine clay unit
17-35	Till; silty, sandy.		clay-loam till
TEST HOLE P-	-59		
Depth (ft.)	Characteristics		Interpretation
0-2	Road fill.		·
2-10	Clay; slightly silty.	.)	
10-25	Clay; lacustrine; silt balls; high gypsum content.)	lacustrine clay unit
25-59	Till; very clayey, silty, sandy.		clay-loam till
TEST HOLE P-	60		
Depth (ft.)	Characteristics		Interpretation
0-2.5	Road fill.		
2.5-13	Clay; silty; laminated with silt balls.	,	
13-23	Clay; lacustrine; silt balls and gypsum concentrations.) lacustrine clay uni))	
23-40	Till; very clayey, slightly silty, sandy.		clay-loam till

APPENDIX C PIEZOMETER HYDROGRAPHS

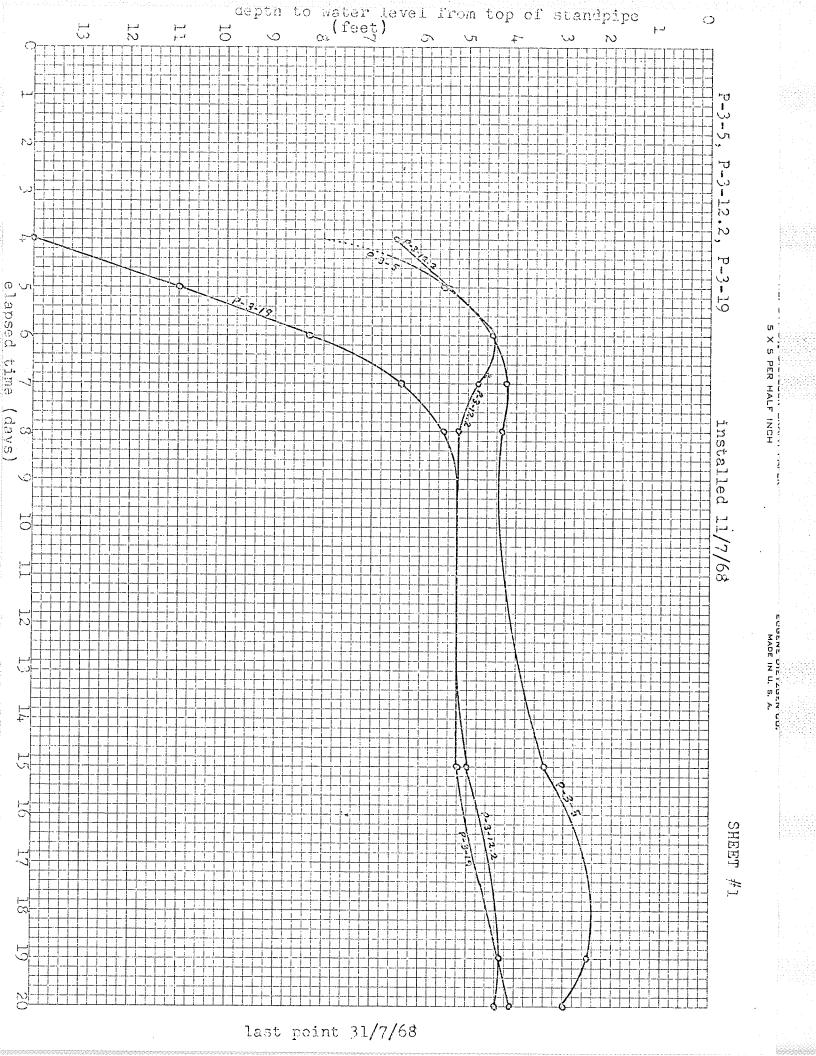
The first group of hydrographs is arranged so that, for each nest, a 20-day hydrograph is followed by a 100-day hydrograph. This allows a more convenient comparison of the initial stabilization period to a seasonal fluctuation. These hydrographs were prepared only for piezometer nests I through 14 installed during the summer of 1968. The 20-day hydrographs were prepared before survey data was available and list water levels from the top of the standpipe rather than an elevation. However, they are useful in illustrating short-term response characteristics.

Two-year hydrographs of piezometer nests 1 through 40 were prepared to illustrate seasonal and annual fluctuations. Precipitation bar graphs indicate the relative sensitivity of each hydrograph to short-term and long-term variations in precipitation.

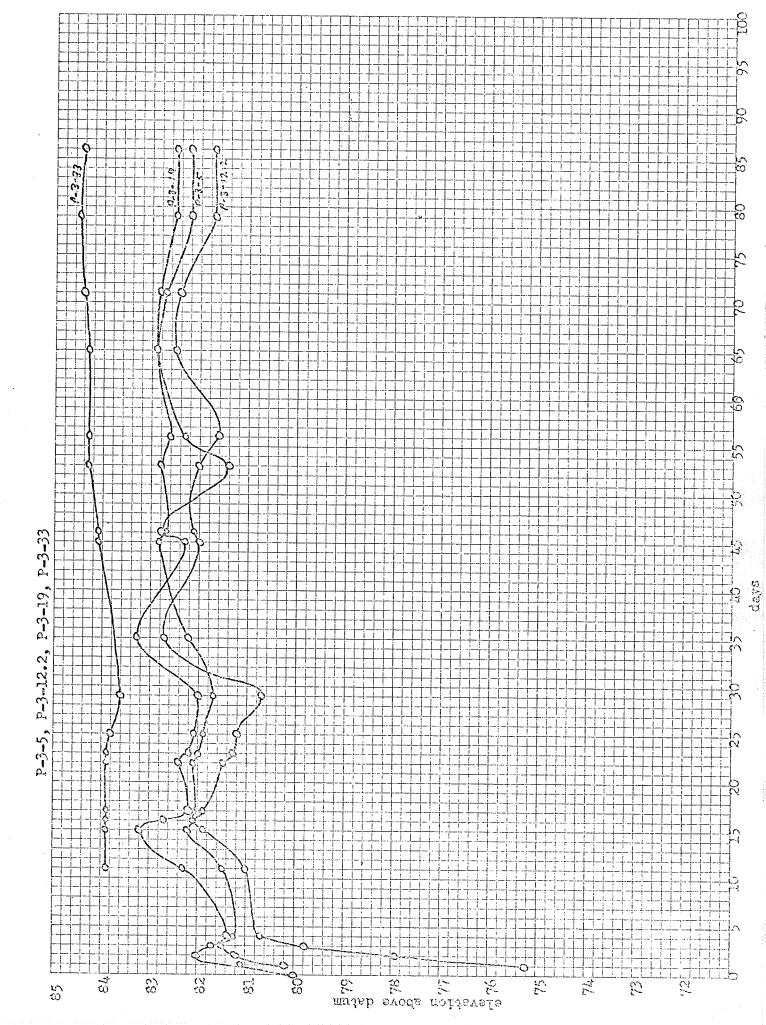


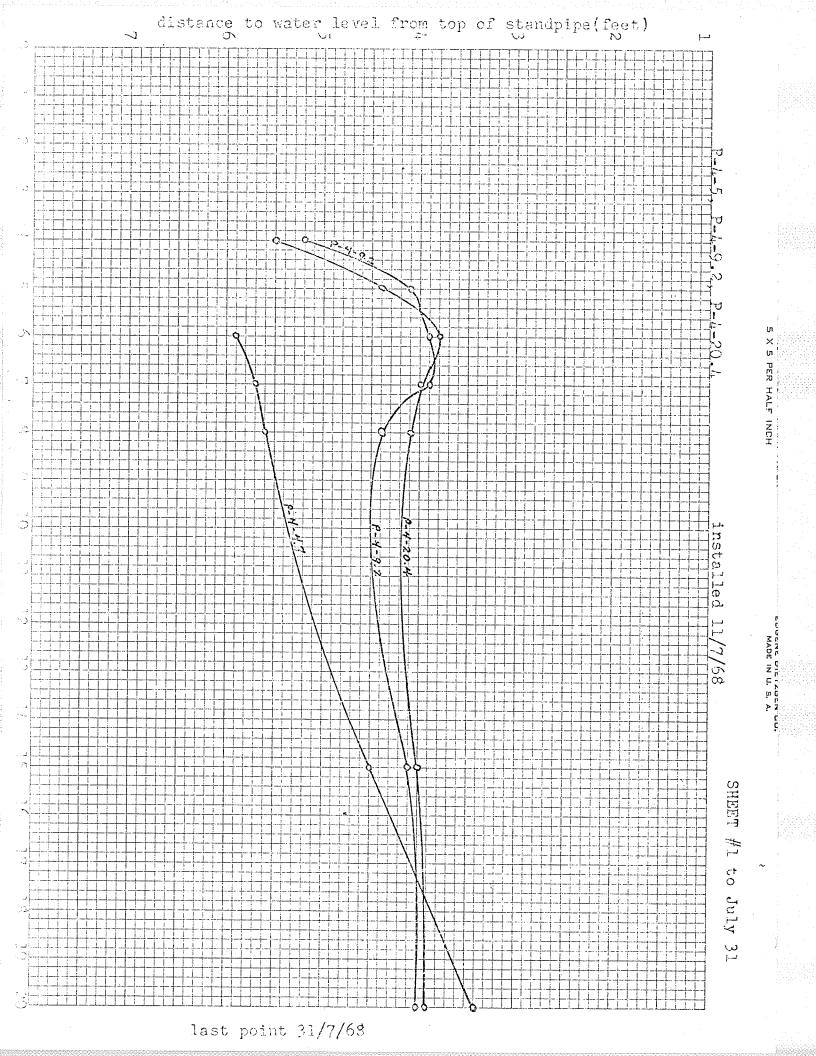
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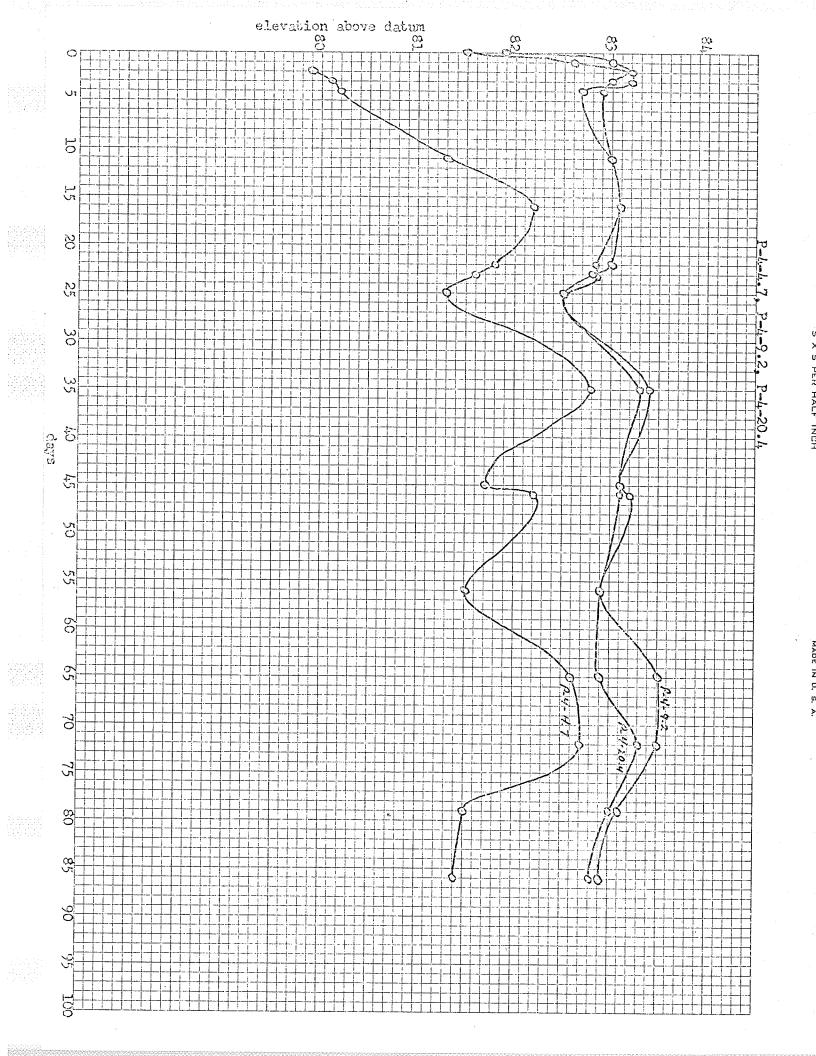
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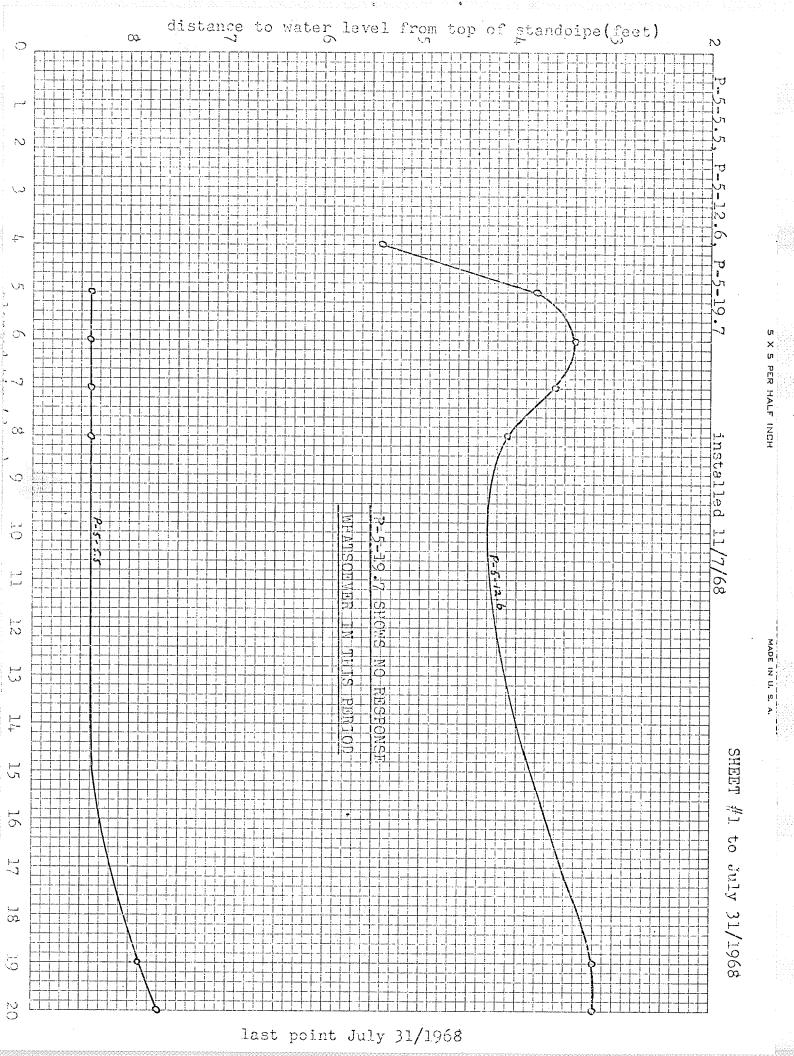


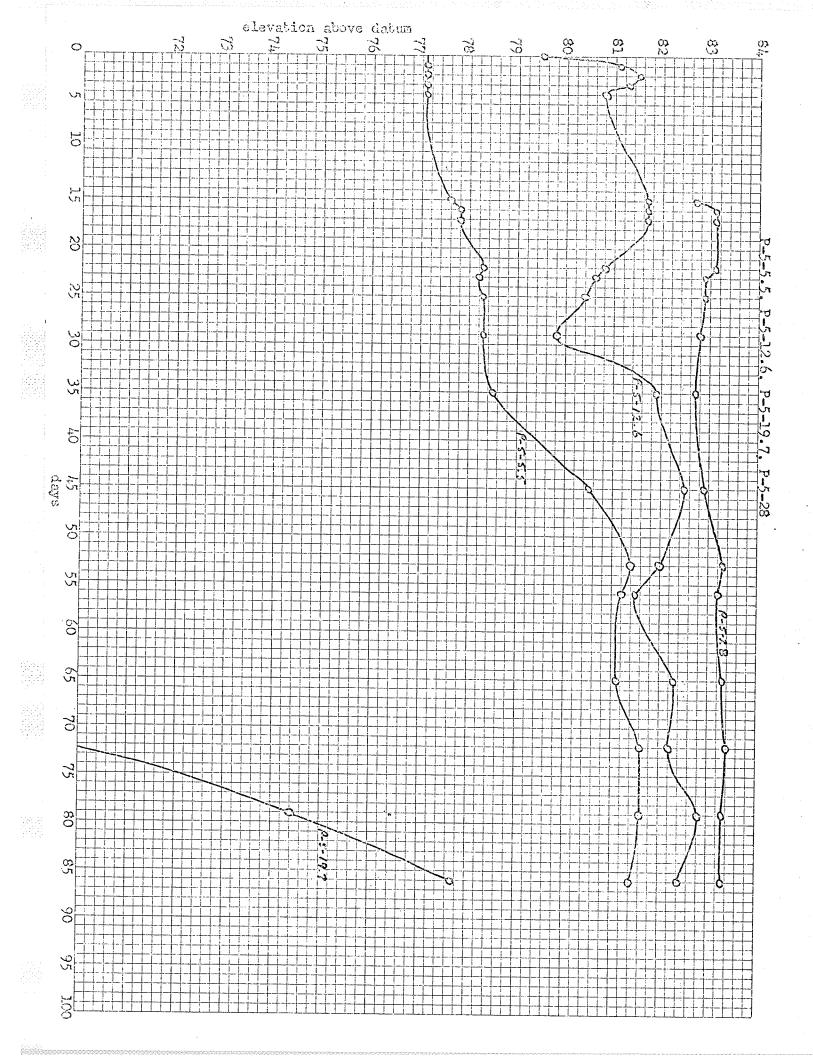
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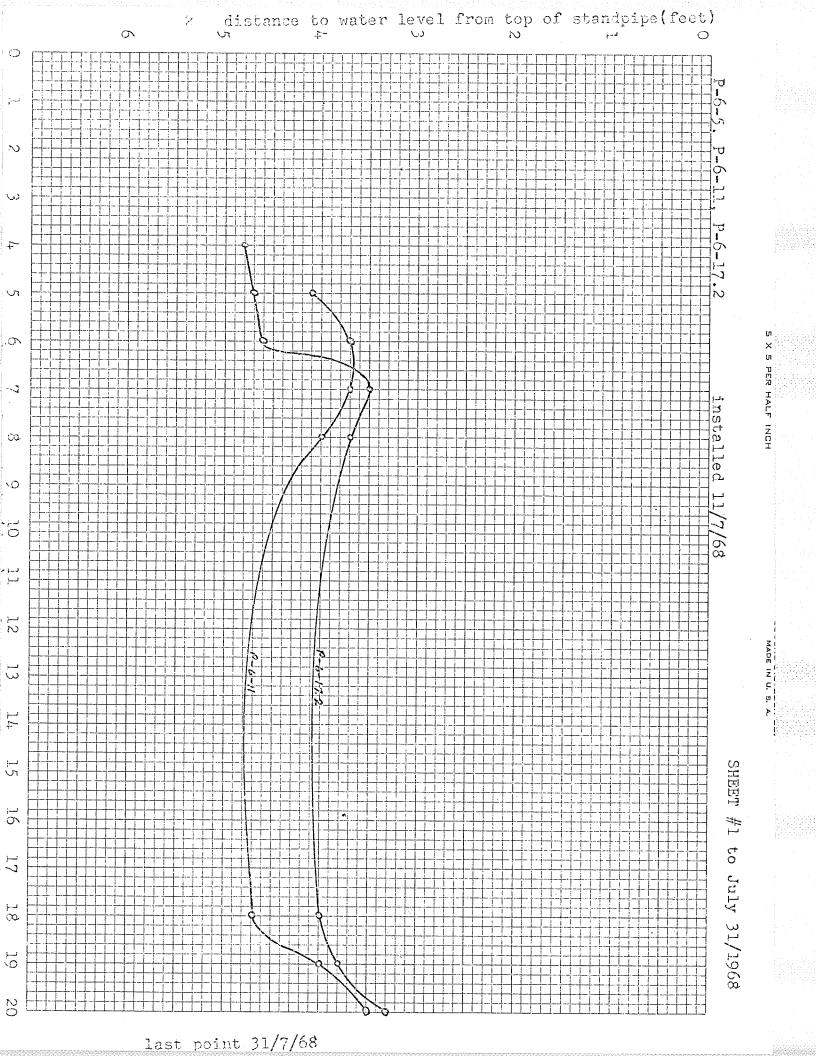




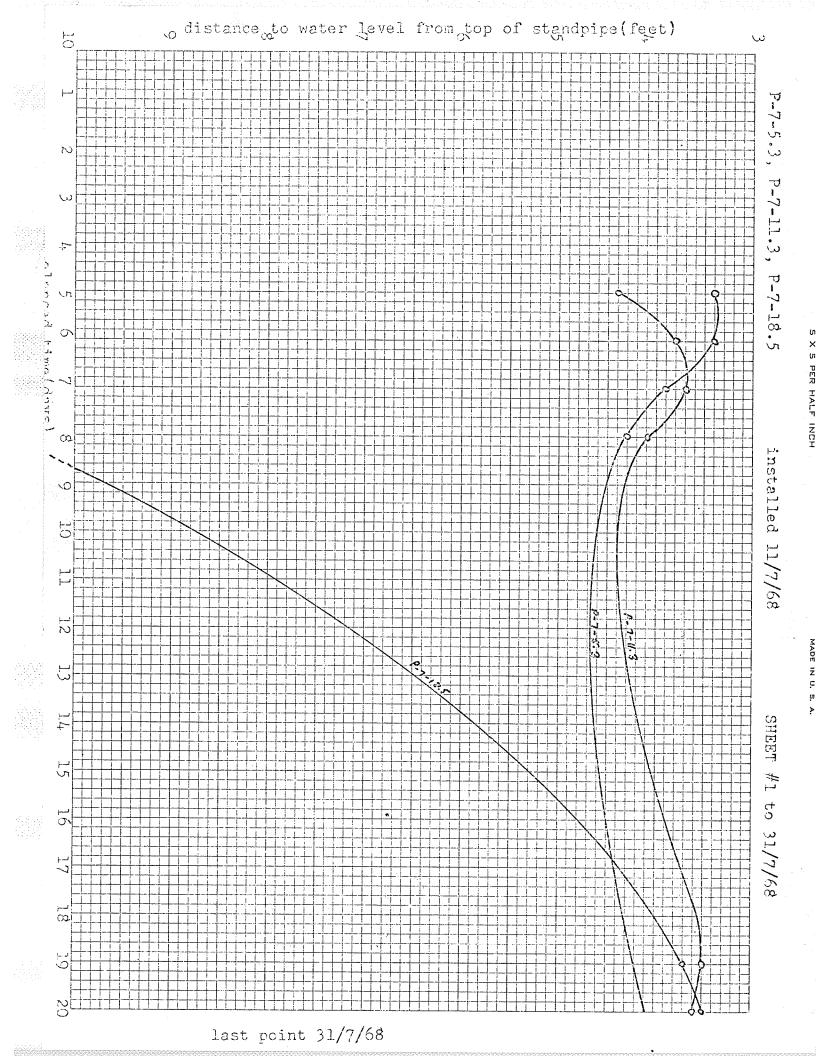




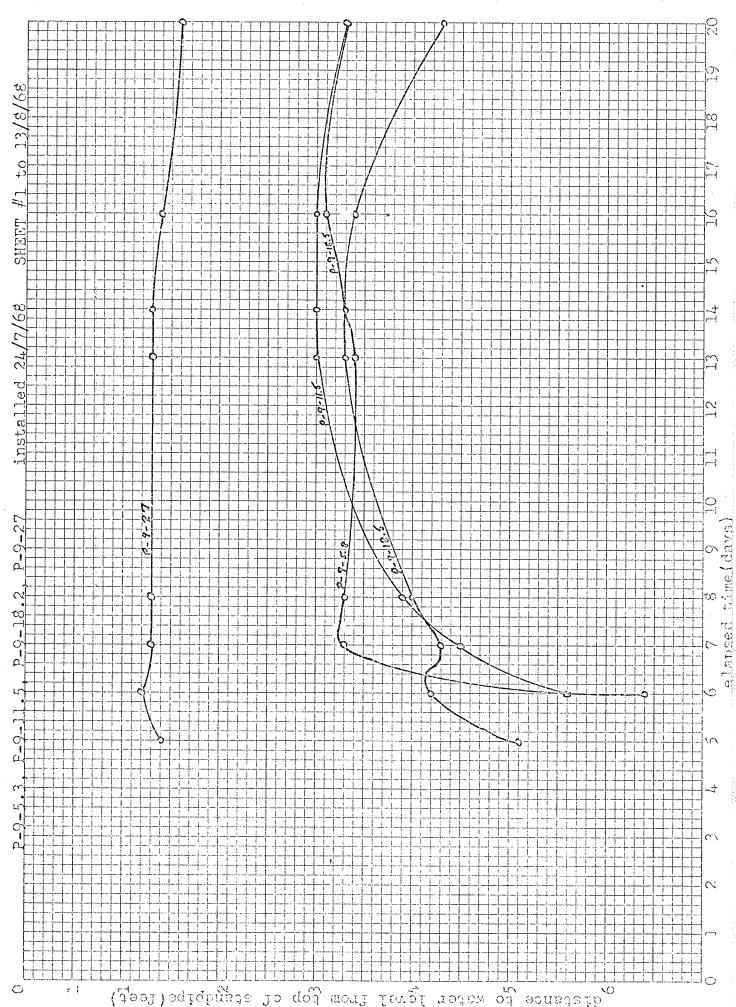


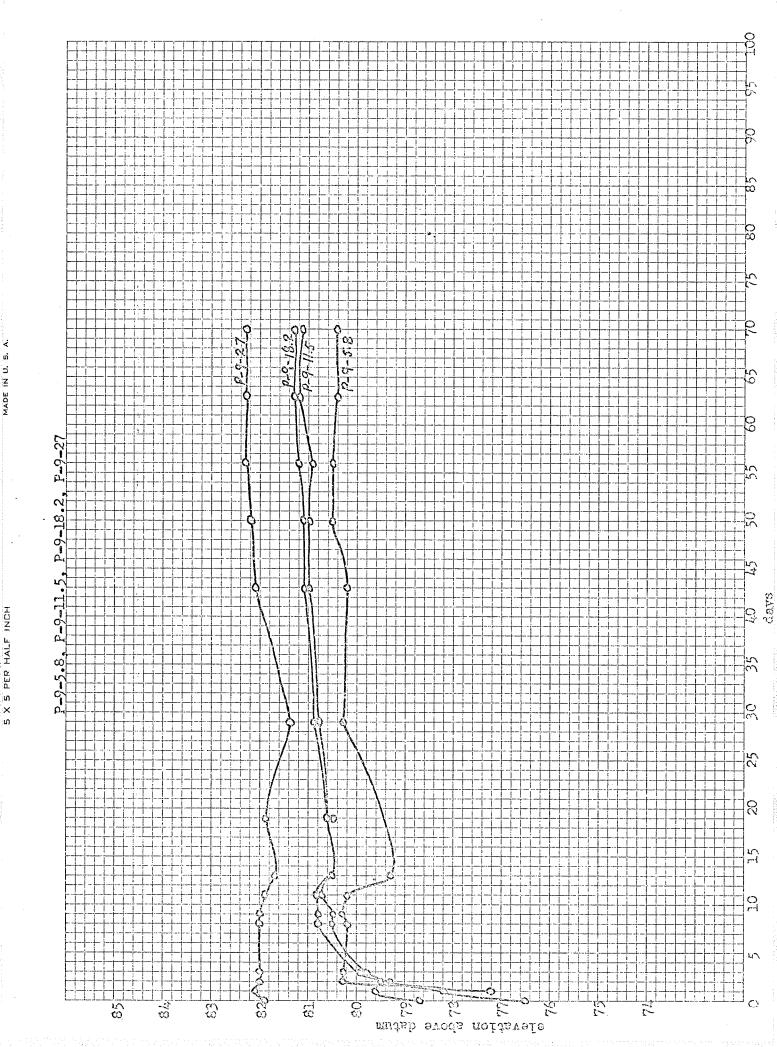


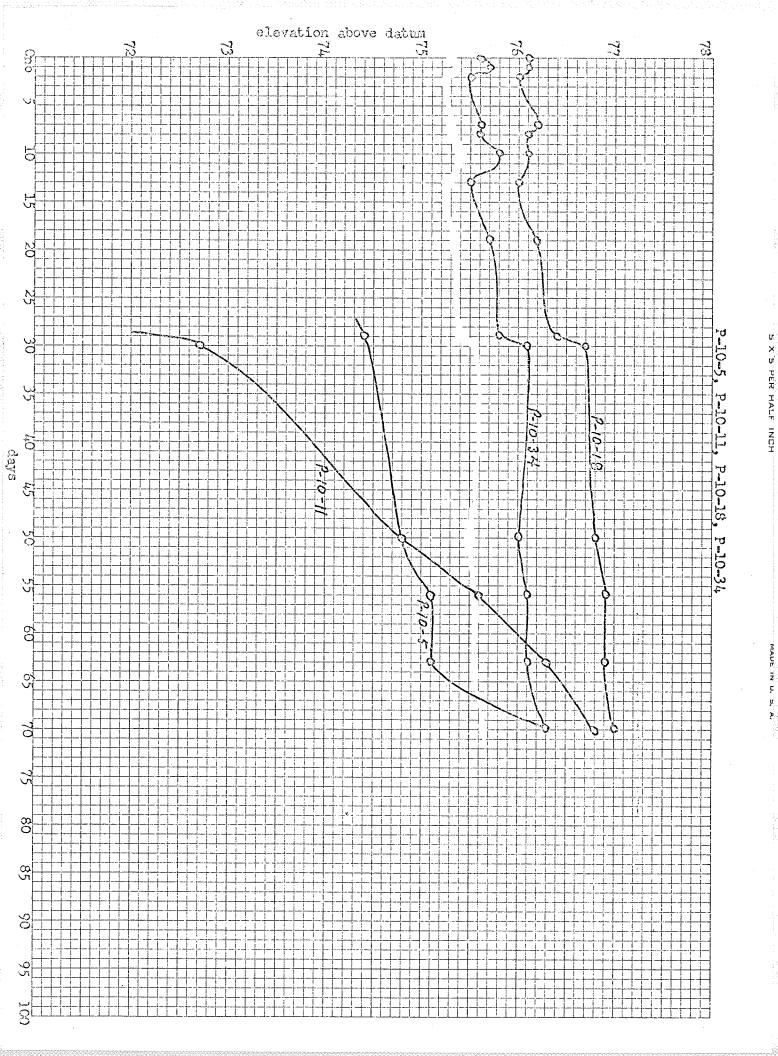
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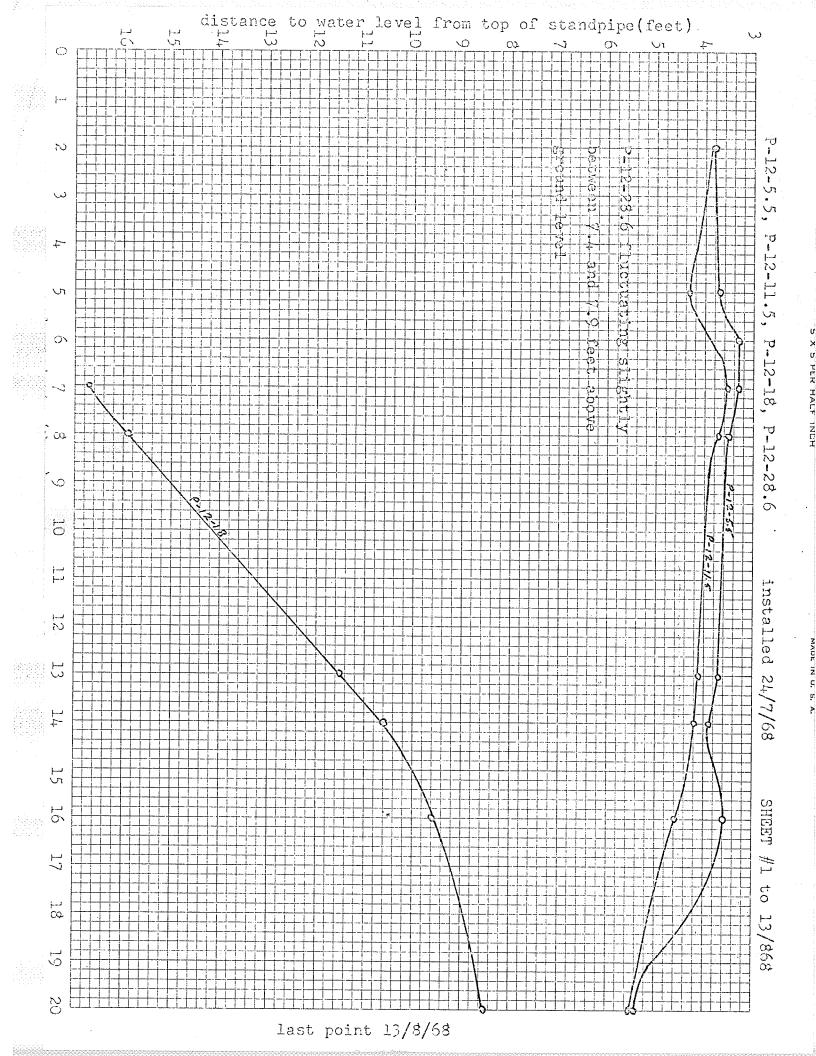


last point 13/8/68

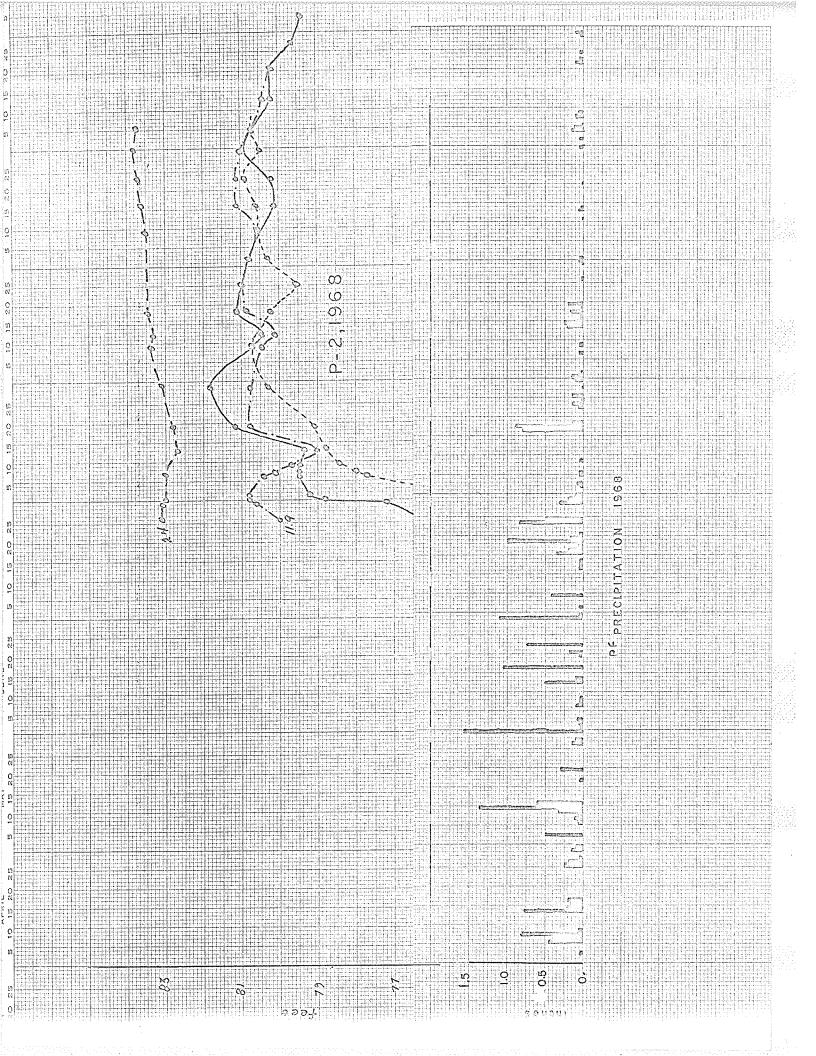


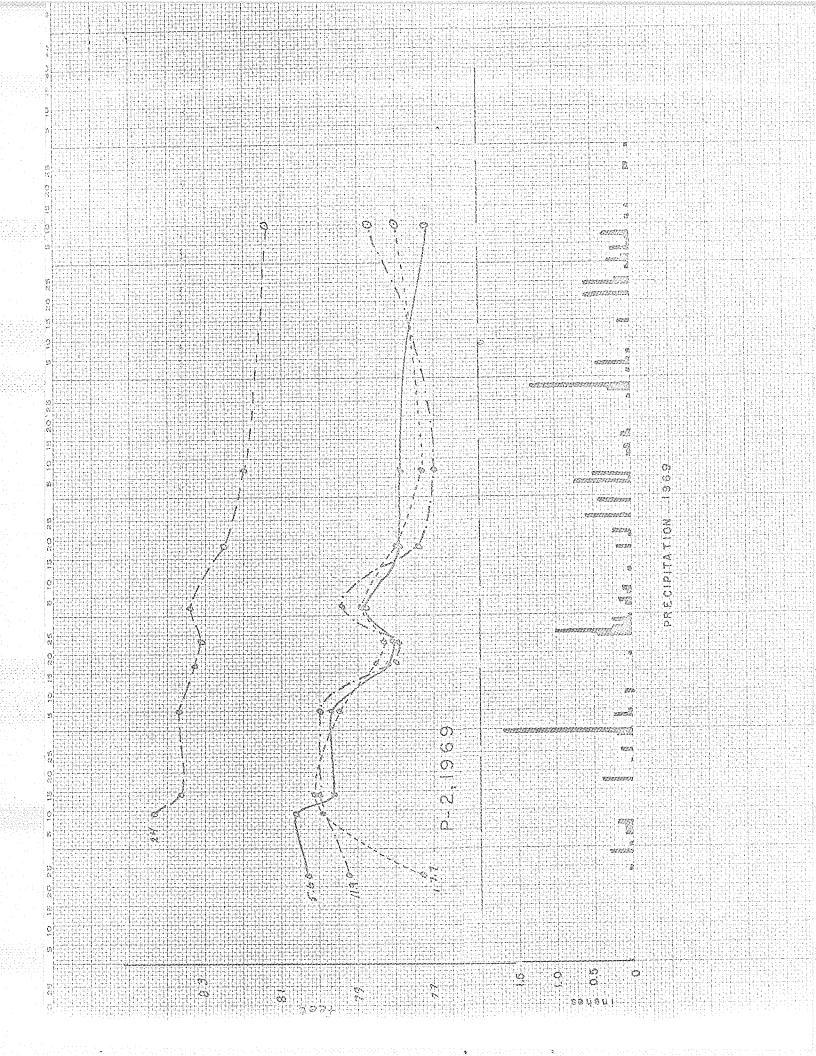


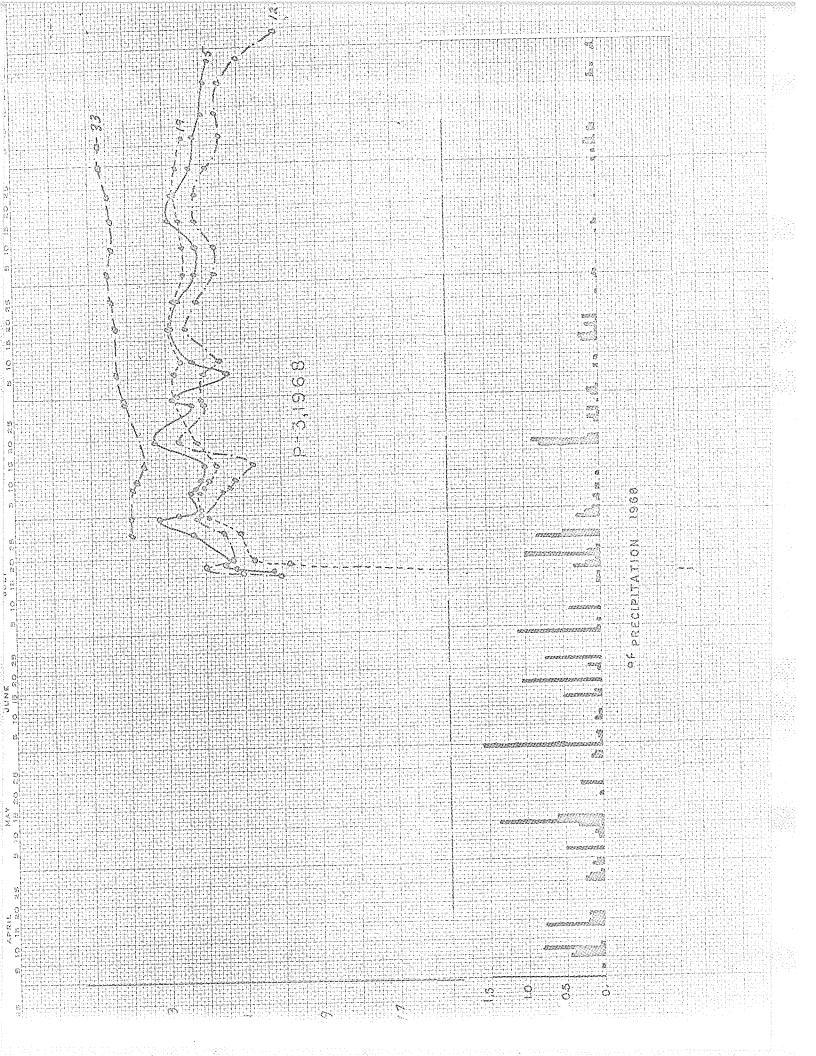


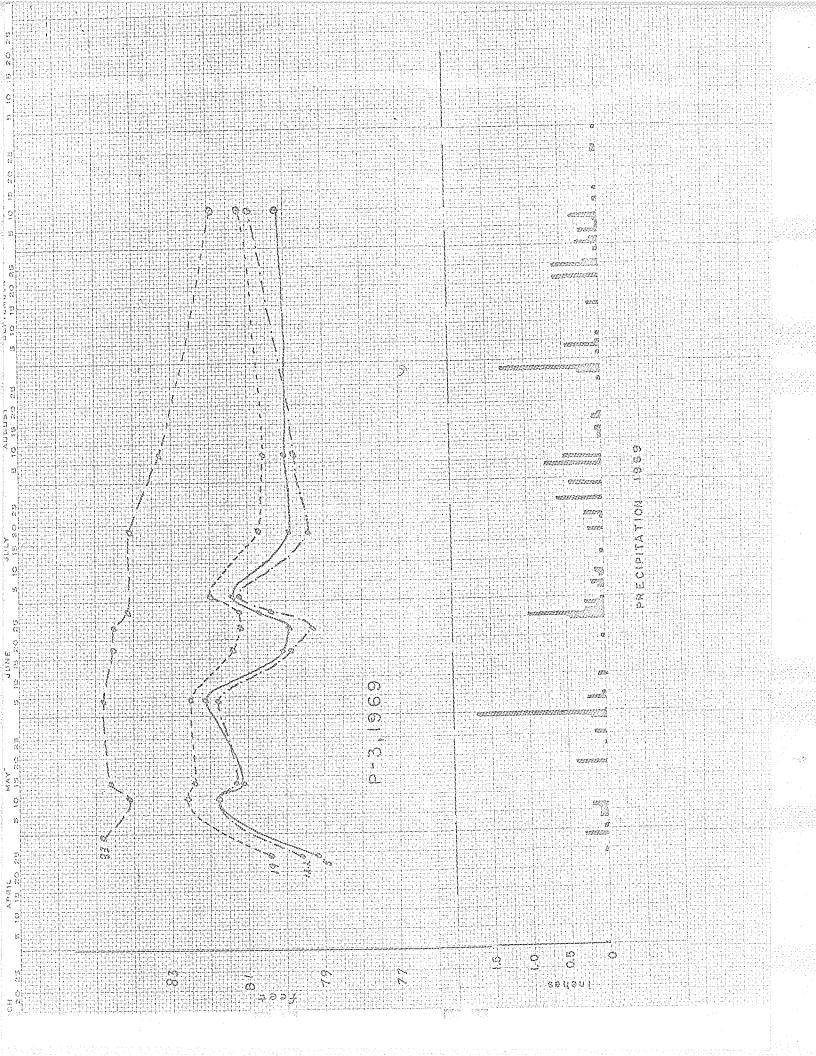


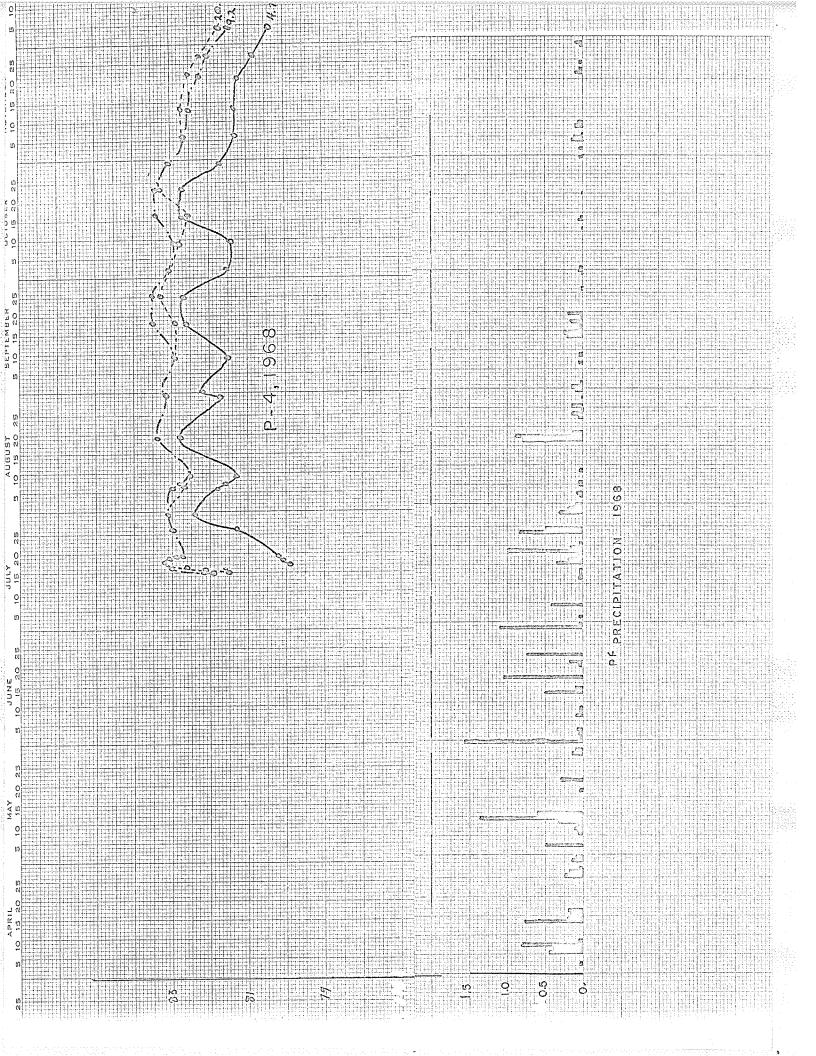


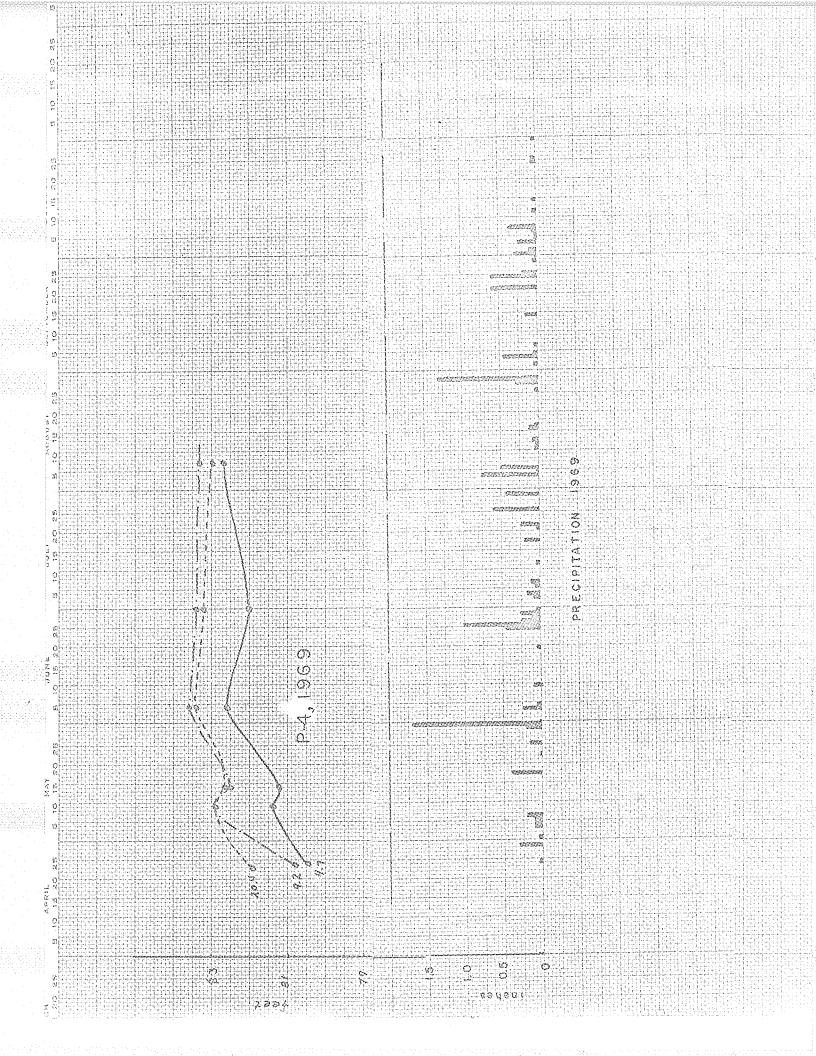






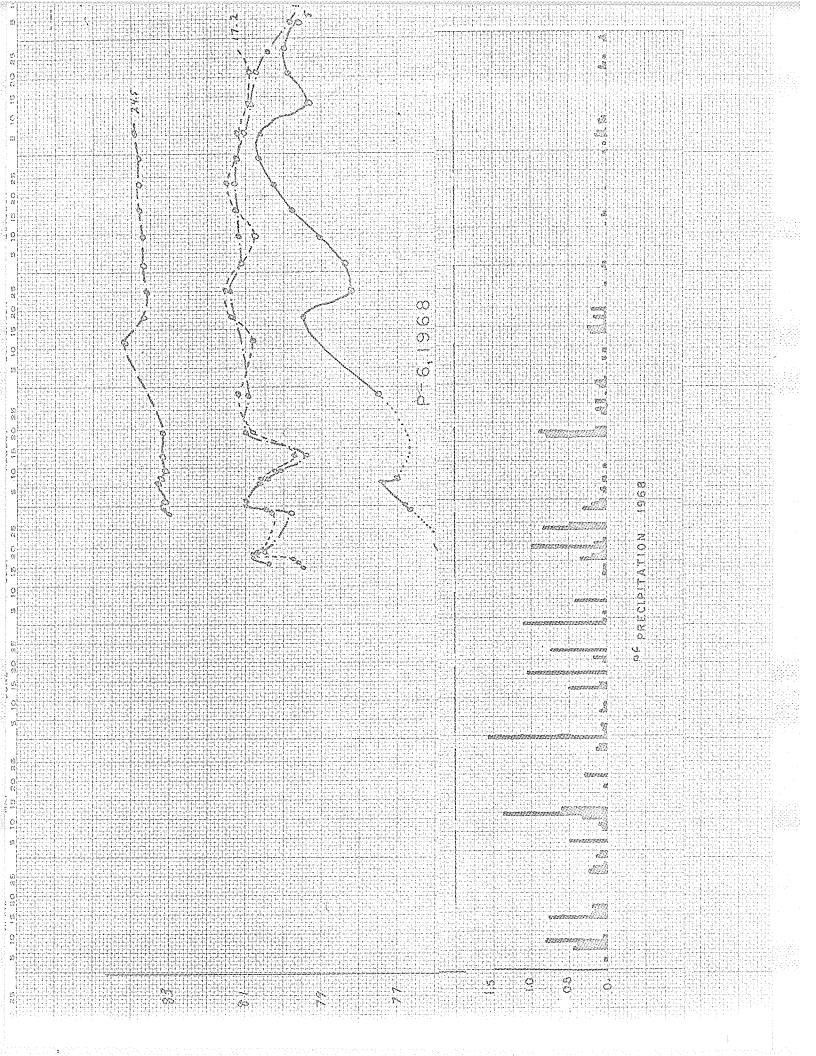


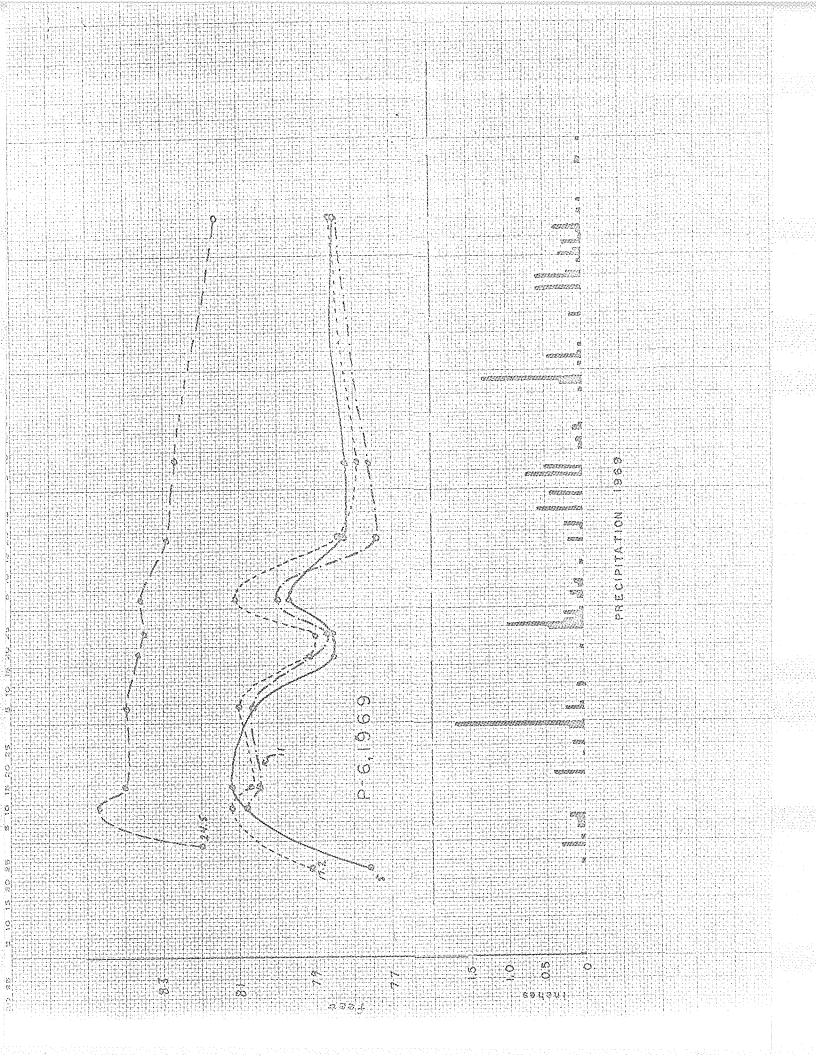


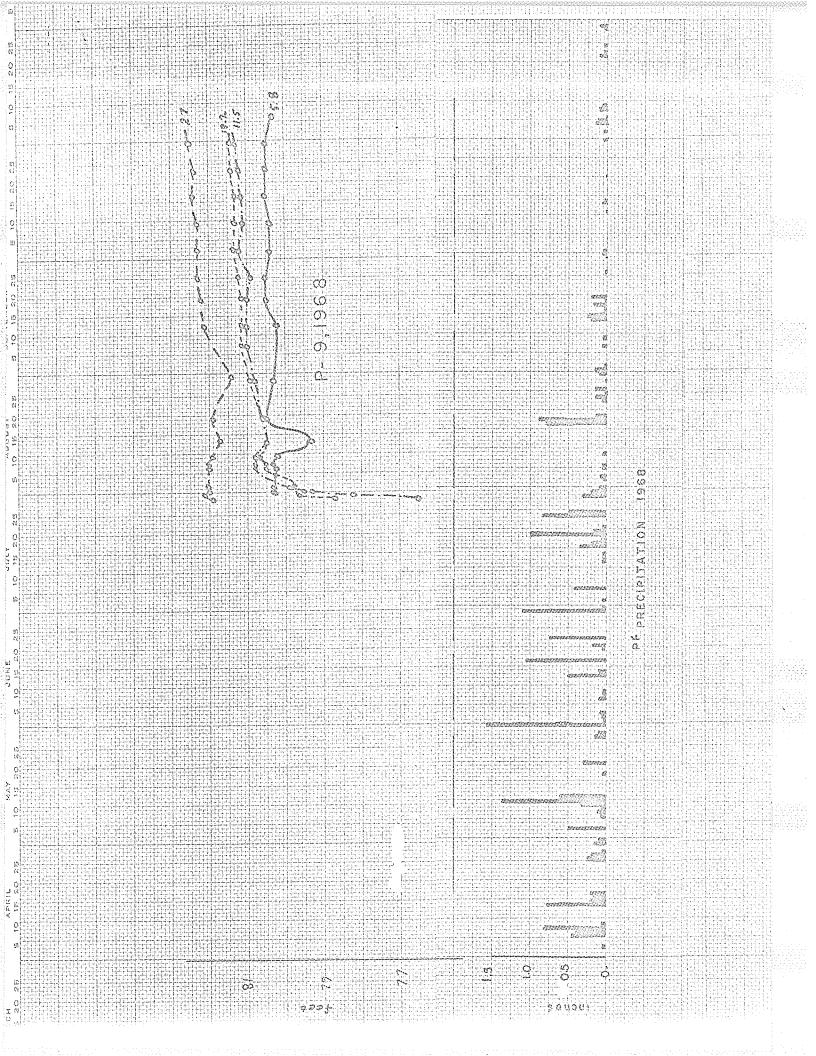


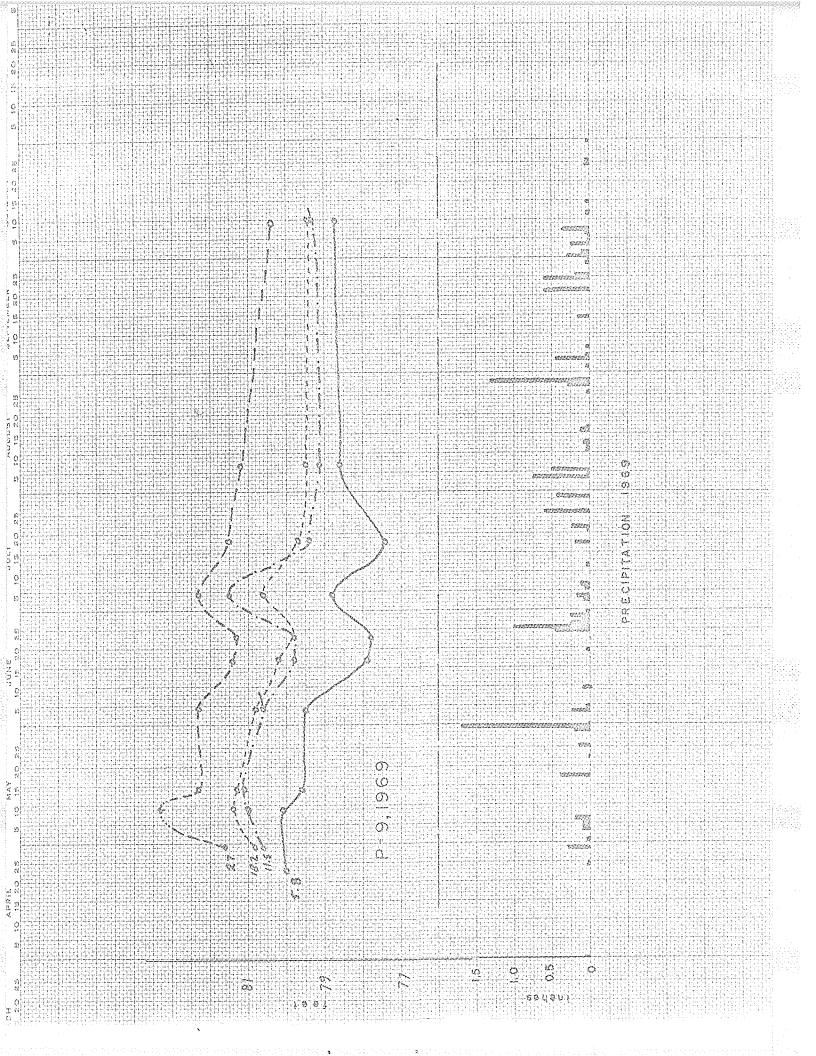


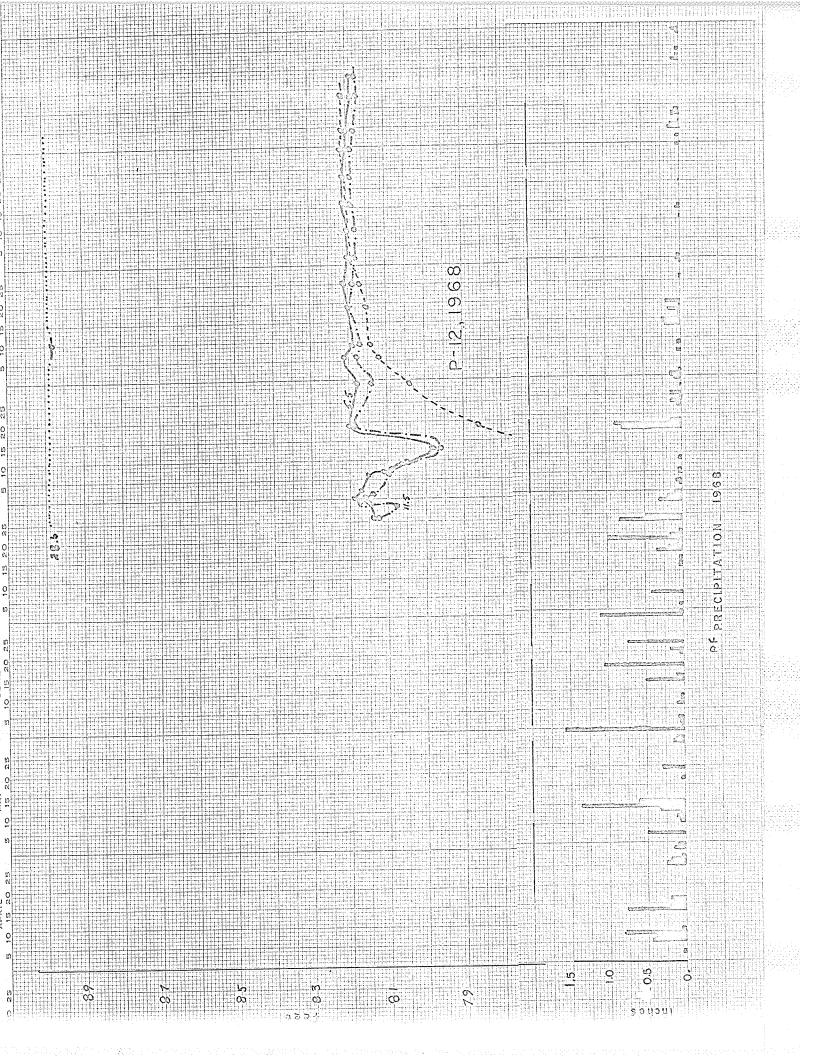


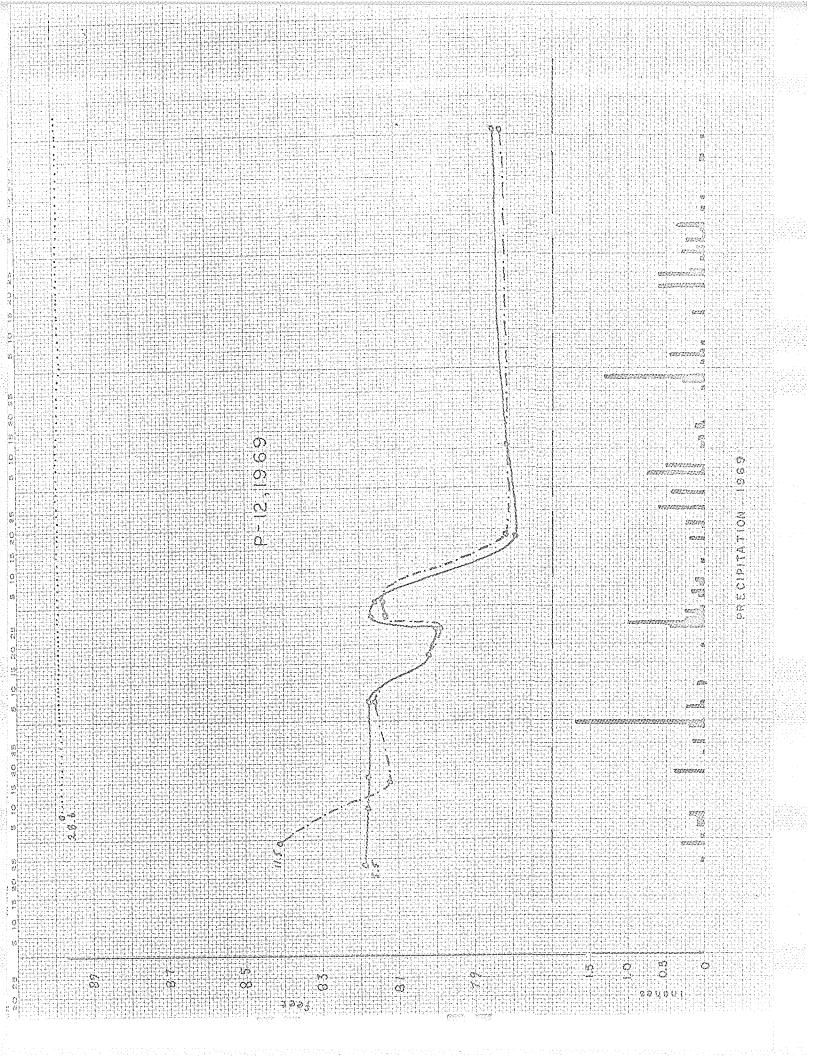


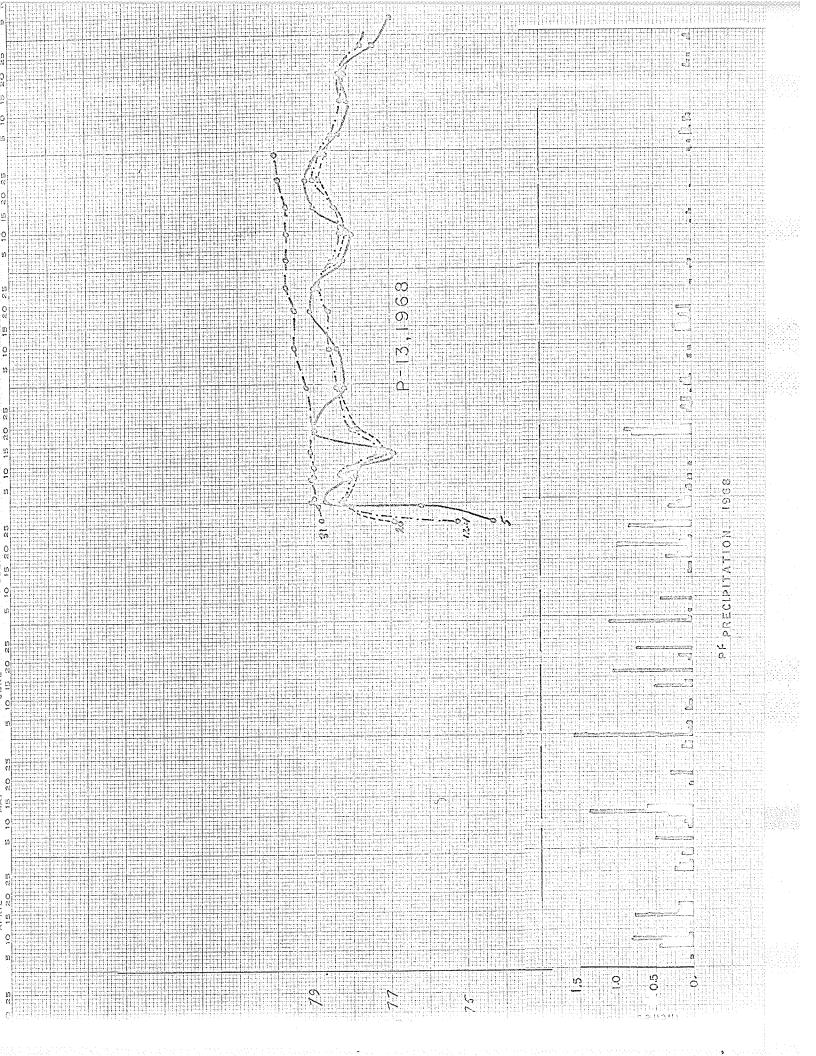


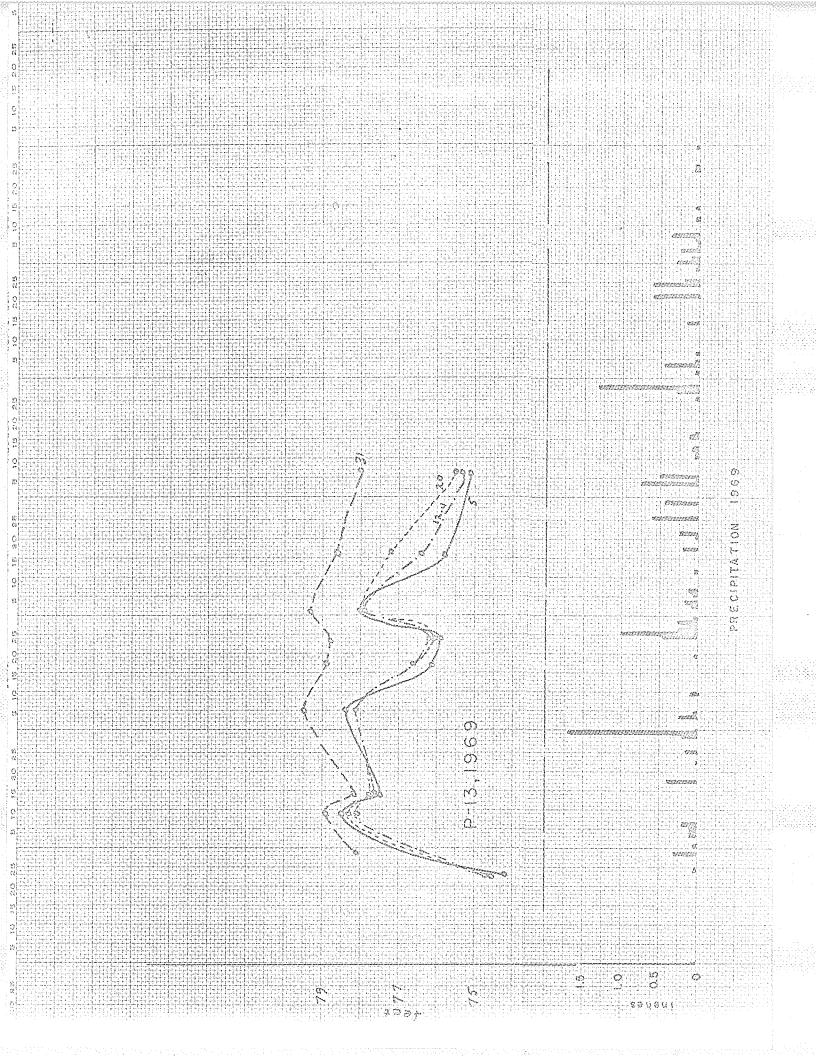


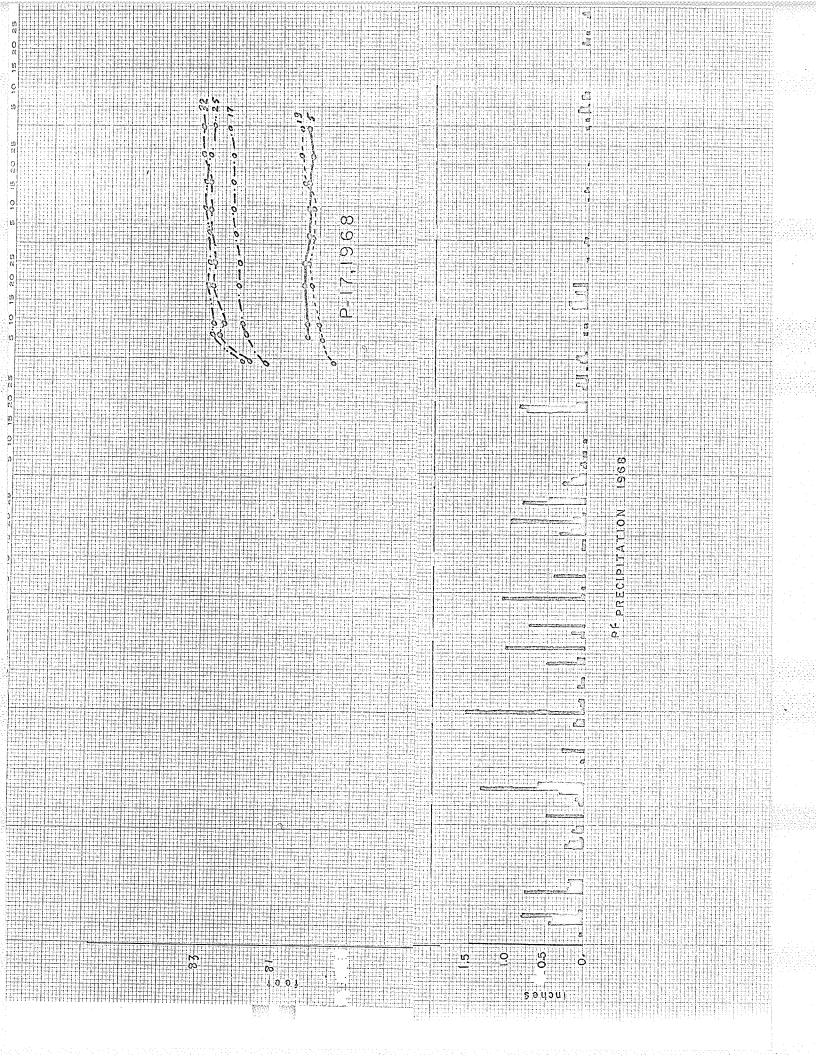


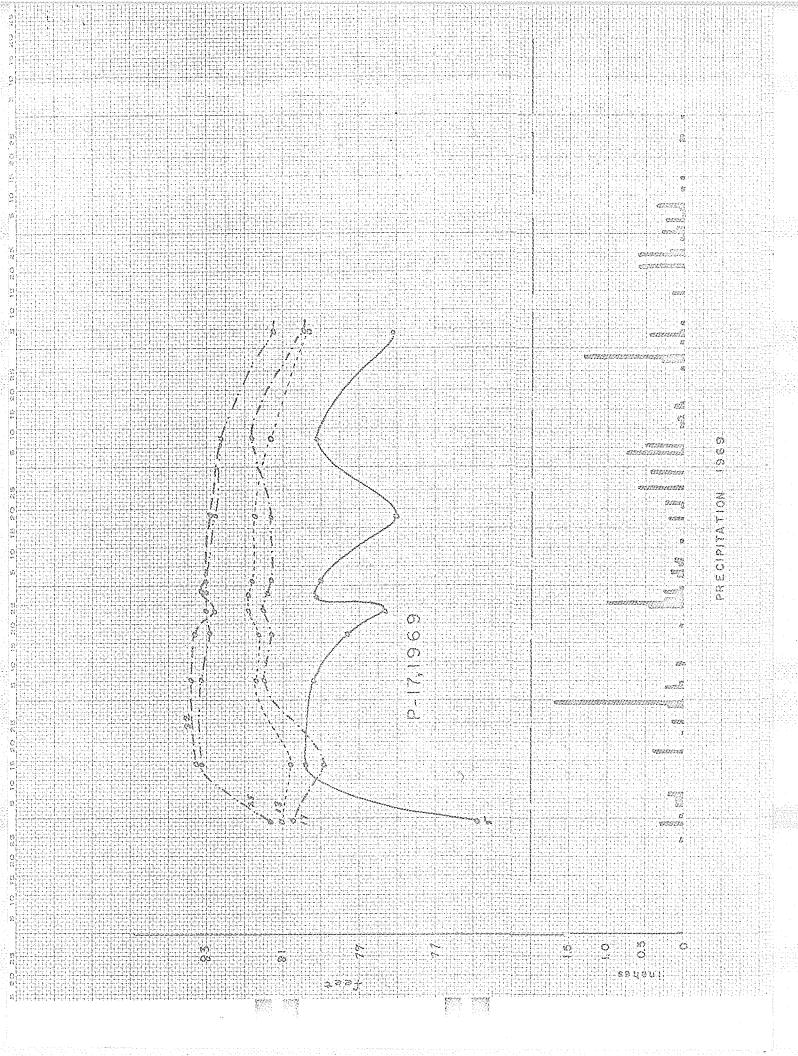


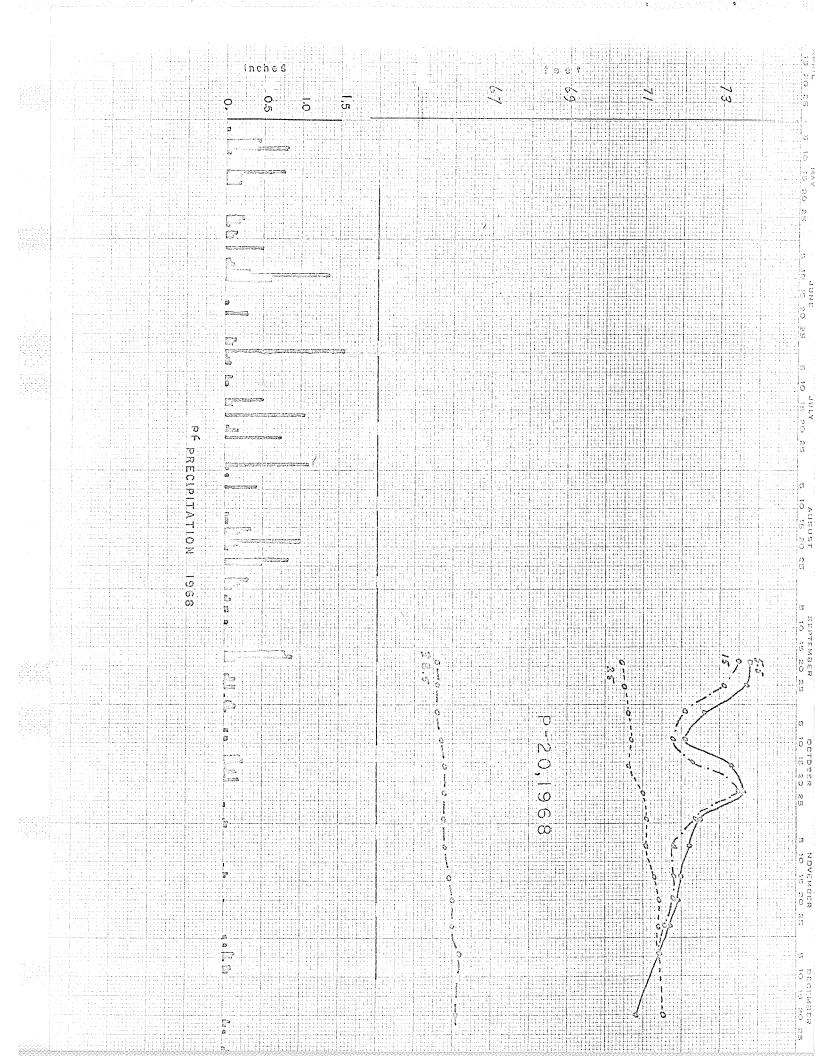












APPENDIX D

HYDROCHEMICAL ANALYSES

The hydrochemical analyses listed in Table 1 were conducted by the Whiteshell Nuclear Research Establishment, Research and Development Branch, in 1968 on samples obtained from relatively newly installed piezometers. The analyses listed in Table 2 were obtained in 1969 after piezometers had been stabilized, flushed, and drained, allowing true formation waters to be sampled with a minimum of contamination. Analyses for anions of the 1968 samples were conducted by the University of Manitoba, Department of Agriculture under the supervision of Dr. M. Zwarich, and the results are available although not listed in this appendix.

In Table 2, the samples listed under 'Special Samples', were isolated under low pressure from the atmosphere immediately after sampling.

Table 3 lists the field conductivities obtained by downhole methods during 1970.

Ca ⁺⁺	226.0	0.96	51.7	268.0	220.0	210.0	79.0	492.0	197.0	128.5	92.0	354.0	122.0	78.0	31.0	304.0	62.5	92.0	139.0	726.0	137.0	850.0	17.4	139.0	45.5	49.0	62.5
																	**						Ĵ.				
Mg++	850.0 182.5	100.5	50.0	530.0	345.0	215.0	55.0	650.0	210.5	70.0	79.0	650.0	130.0	0.99	35.0	366.0	25.0	49.0	149.0				24.0	155.0	46.5	48.5	48.5
Κ ⁺	3.7	5.7	6.3	5.25	8.4	12.3	5.8	7.0	16.65	48.0	5.9	5.5	5.5	5.25	7.3	5.8	5.85	15.6	5.85	595.0	817.5	162.5	9.35	117.0	10.4	8.85	0.6
Na+	291.0	61.0	31.5	205.0	144.0	149.5	23.5	118.5	71.0	65.0	22.5	223.5	138.5	127.0	18.0	75.0	34.0	75.0	44.0	271.0	181.5	71.0	14.5	109.5	21.0	22.0	20.0
Sample Number	P- 1- 5	2-1-17	P- 1-24	P-2-5		P- 2-17		P- 6- 5	P- 6-11	P- 6-17	9		∞	∞	က	0	P- 9-11	P- 9-18	P- 9-27	P-21-12	P-21-22	\sim	P-21-39	P-23-19	P-23-28	P-23,-32	P-232-32

Table 1. Analyses by AECL, 1968.

Lab Conductivity (umhos/cm)											ı			1. 1.	1-														
·		. +					<i>3</i> .					4											•						
TDS mg/	4881	804 1352	2487	928	712	4016	1574	519	331	3319	3322	610	2872	1524	730	543	4036	1705	692	62.1	3883	3247	18504	831	4748	1648	1223	539	
CI _ mg/l	3.0	2.0	3,5	7.5	3.5	4.0	5.0	8.0	3.0	4.0	4.5	5.0	ය	5.5	7.5	3.5	3.0	4.5	5.0	3.5	6.0	8.0	10.01	0.6	4.0	5.5	5.0	2.0	
HCO 3	844.2	007.00	392.8	361	373	241.6	307	334	210	359	268		366	405	368	364	293	346	425	427	381	401		447	439	632	339	295	
CO3 mg /																	•												
SO4=	2800 485 45	1180	1490 1440	50	225	2830	840	65	40	2180	2250	610	1820	770	185	70	2800	950	30	50	2540	1960	1260	30	3650	685	80	110	RECL, 1969
Ca++	180 70 85	\$ 59	135 130	185	40	235	135	40	52	270	335	105	265	105	2	2	235	40	100	89	385	265	195	09	370	20	140	0%	yses by
Mg++	670 160 85	86	300 280	185	20	535	145	45	15	370	320	140	305	165	09	35	520	180	70	45	455	440	215	25	40	165	625		2. Analy
K+	7 7 W	י אט י	ဂ လ	6	יט	S)	<u> </u>	× 6	33) -	9	o\	_	_	ထ	٥.	1	25	6	27			∞ ;	<u>.</u>	25	ιO	2	4	_	Table
Na+ mg/	380 75 55	200	160 160	130	15	165	35	07 -	٠ ر	130	135	125	105	65	30	Ŋ	160	175	35	20	105	165	155	235	240	105	30	5	
Sample Number	P-1-5.5 -11.4 -17.7	-24	r-z-5.6 -11.9	-17.7	-24	7-6-5	7.71	۲ <u>۱</u>	, i do	7-4-4.	7.6.	-20.4	P-5-5.5	-12.6	7.61-	-23	P-6-5	!	7./1	-24.5	P-7-5.3	- C	ر د د	-30	P-8-5.9	-12	6[1	-28	

Lab Conductivity (umhos/cm)	÷.																													
TDS mg/l	3255	55 l 760	560	1869	5211	3276	808	1249	4202	223	2383	363	6550	3793	2732	099	2767	2991	2412	727	1567	2678	1546	615	564	1860	527	202	602	481
C1 - mg/1	4.0	0.0	4.5	2.5	7.5	5.5	3.5	2.5	7.0	> 0	0	1	14.0	11.0	5.0	2.5	16.0	19.0	19.0	3.0	4.0	5.0	5.0	2.5	3.0	4.5	7,0	7.0	5.5	3.0
HCO 3	322	332 332	342	639	564	625	405	444	364 364	298	229	107	561	307	532	327	805	566	503	493	337	244	229	415	327	712	361	346	425	366
CO3	•													•													-			
504= mg/1	2130	215	135	745	3390	1850	250	550 550	1300	1330	1630	150	4220	2490	1730	20	1210	1640	1380	30	098	1770	955	55	100	710	40	30	40	9
Ca++ mg/l	390	95	6	8	375	290	<u>0</u> ¦	ر ر در در	145	215	190	2	205	340	225	92	240	270	75	145	145	310	155	2	08	12	20	55	9	90
Mg++ mg/1	335	04	20	190	009	330	52	001 707	140	200	255	30	1110	400	30	220	290	260	200	40	170	235	165	20	40	275	50	45	50	40
K+ mg/1	4 4	7	4.	7	6 ;	0	I C	۸ <u>د</u>	6	14	က	9	2	15	0	2		16	15	9	%	6 I	_	_	6	က	4	ر د	0 1	^
Na+	20 20	55	15	210	265	165	C 4	5 C	375	165	75	5	435	230	200	70	195	220	220	2	45		39 30	15	52	140	5	15	70	ر ان
Sample Number	P- 9-5.8	13.2	-27	7-10-5	9	∞ <u>₹</u>	45-1-10 7-1-10	5.5	-22.8	-35	P-12-5.5	-11.7	P-13-5	-12.4	-20.0	- C2	P-14-13.5	-23.5	-35	P-15-4	P-16-6	0 7	-14.5	-20		Ç=/ =d	<u>n</u> ;	<u></u>	c7 -	-32

Table 2, continued

Lab Conductivity (umhos/cm)																. 022	1850	1500	540	. 3810	470	2400	2990	2550
TDS mg/l	1349	2904	1714	2023	1040	57.16 5567	5549	3667	491	550	391	1145	545	543		77.1	1778	1472	57.1	4192	489	2462	3173	3051
C1 -	9.5	34.0	12.0	18.0	0,0	20.0	32.0	37.0	2.5	4.5	3.0	7.0	0.9	5.0		7.0	7.0	5.0	0	5.0	2.0	17.0	17.0	18.0
HCO3-	854 630 444	464 512	732	205 205	454	522	590	434	376	376	273	307	342	288		429	351	332	434	459	337	581	537	420
CO3 mg/l																								
504= mg/1	195 705 1190	1680	560	1150	265 3450	3690	3560	2230	2	40	9	540	09	.115		110	930	760	5	2700	30	1210	1790	1800
Ca++	125 170 165	270	115	155	35 25	285	350	310	52	99	92	95	70	65		30	150	140	62	325	65	190	340	320
Mg++ mg/l	150 125 140	230	160	115	65 850	560	009	370	50	55	25	115	40	40		15	110	125	30	450	30	210	270	210
+ X + X + X	9 I I	36 16	35	130	120	. 09	17	9 '	7	4	5		· /	2		10	15	. 15	 	က	0	-4.	77	81
Na+	10 145 230	190	100	210	95 475	430	400	270	(C)	01	0.	R :	20	25	•.	170	215	92	2	250	15	240	205	265
Sample Number	P=18-5 -17	-39.5	P-19-12 22	-33	p-20-5	-15	-25	38.5	P-22-9	ထု	9].	P-23-17	87	-32.5	Special Samples	P- 7-30	P-11-22.8	P-11-34	P-12-23.6	P-13-12.9	P-13-31	P-14-13.5	P-14-23.5	P-14-35

Table 2, continued

	+ _σ Ζ	+	#+ Wa+	‡°C	=, 05			į	ς 2	Lab
Sample Number	l/gm	mg/l	mg/l	1/Bm	mg/1	mg/l	mg/l	mg/l	mg/	Conductivity (umhos/cm)
P-16-14	35		165	350	905				1000	1240
P-16-20	5	Ω.	50	95	40) C	1022	040
P-16-34	ιΩ	9	20	100	90) W	713	0/0
P-20-15	350	17	65	335	3670			0, 10	0.17	040
P-21-22	09	23	15	06	310			0.72	4952	5100
P-21-28	0	က	25	? ?	פי ע			o -	\ \ \ \ \ \ \ \ \ \ \ \ \ \ \ \ \ \ \	04/
P-21-39	5	ි ආ -	25	2 2) ער) u	440	460 .
P-22-16	15	17	35	06	15			- v	44/ 707	0.7
P-26-28	365	355	245	290	2340			25.0 0.0	3000	040
P-26-43	80 180	14	200	270	2320			7.0	2320	9520
P-27~58	115	80	4	520	1250			ο α	2450	2000
P-28-18	175	34	380	390	2290			20.00	4062	37.50
P-23-40	225	275	4	200	1080			20.0	100Z	3030 2250
P-28-71	110	1	70	135	820			37.0	1338	1630
P-29-18	130	10	300	495	2140			8.0	3512	3340
P-29-35	140	7	275	475	2050			0.6	3319	3000
P-29-54	135	13	245	490	1980		•	7.5	32.13	3000
P-31-10	25	က	90	140	200			5.0	931	3350
P-31-25	145	35	145	345	1290			18.0	2378	2250
P-31-44	255	36	190	285	1400			33.0	2633	2800
IWP-3-20	435	15	360	530	2590			21.0	458.5	4000
IWP-3-(11-15)	380	8	505	4.55	3590			29.0	5416	5000
IWP-3-(15-20)	350	7	470	200	3270			29.0	5331	4700
1WP~3-10	245	12	345	550	2600			16.0	4383	3900
7-33-18	20	4	200	105	1644	-		20.0	1644	1600
P-33-35	305	26	80	300	1600			37.0	2446	2800
7-33-48	305	39	250	370	1990			44.0	3286	3200
IWP-2-(20-36)	320	15	495	410	3190			29.0	5006	4500
IWP-2-8	325	29	0.1	190	.575			7.5	1913	3400
P-35-9 5 65 35	160	4	190	120	885			5.0	1769	1800
P-35-17	390	70	550	230	3360			20.0	5360	5000

Lab Conductivity (umhos/cm)	4200	4000	920	510	370	2800	4390	0.75	450	5300	4600	4050	7007	7,00	1400	2500	3000
TDS mg/l	4205	3969	349	454	331	2685	3766	107	- 25	/189	4928	4337	200	767	· · · · · · · · · · · · · · · · · · ·	2542	2584
CI _	21.0	34.0	7.0	2.5	2.5	8.0	10.0) · C	٠, د د	သ ပ	0.0	11.5	r r) t) (o. B	14.0
HCO3-	254	205	405	32.7	220	700	229	330	200	717	732	644	268	850) () (C7/	215
CO3 mg/l														٠			
SO ₄ =	2800	2350	220	15	25	1360	595	5	3540	000	3010	2550	260	1416	0101	017	1670
Ca++	250	275	<u>8</u> :	80	9	210	7.5	55	300) \ \ \ \ \ \ \ \ \ \ \ \ \ \ \ \ \ \ \	040	340	45	125	000	0 / 4	235
Mg++ mg/l	425	240	65	25	20	260	69	25	200	200) .	405	55	125	225) [82
K+ 1/6m	45	480 1	` '	4 (m	<u></u>	7	7	9) <u>-</u>	- c	ე (2	12	0	, L	C7
Tag +	410	, 585 285	က်	n t	ဂ ႏ	<u> </u>	2	5	345	310	270	0/7	6 0	45	80		740
Sample Number	P-35-27	7-55-47 7-36-10	0-40.30	0 40 AE		07-00-10 20-70 c	47-00-7	P-36-36	P-37-10	·p-37-15	5-37-05	2-10 c	05-10-	P-38-15	P-38-22	0-38-33	5 0-00-

APPENDIX E

GRAPHICAL PRESENTATION

OF GROUNDWATER TEMPERATURE

VARIATIONS

ANNUAL METEOROLOGICAL SUMMARY

53,1966

1968

WNRE, Pinawa, Manitoba

Health and Safety Branch

Environmental Control Section

			Temperat	ure (°	F)	,		Heating	Wir	rd
nth 	Mean Max.	Mean Min.	Mean	Ex- treme Max.	Date	Ex- treme Min.	1	Degree Days		Pre- vailing Directio
n. b. ch xil y ne ly st. t. v.	8.1 12.4 35.1 48.8 58.6 67.6 73.3 67.7 65.2 49.6 32.9 14.5	-10.1 -11.4 13.5 26.7 37.7 48.8 52.5 51.1 46.1 35.1 21.4 0.0	- 1.0 0.5 24.3 37.8 48.2 58.2 62.9 59.4 55.7 42.3 27.2 7.3	76 78 84 88 80 82 66 50	Jan. 21 Feb. 3, 29 Mar. 26 Apr. 11 May 1 June 3 July 6 Aug. 5 Sept. 14 Oct. 13 Nov. 3 Dec. 2	-34 -18 8 17 37 39 35 34	Jan. 6 Feb. 22 Mar. 11 Apr. 5 May 5 June 14 July 9 Aug. 17 Sept. 9 Oct. 24 Nov. 29 Dec. 25	2046 1870 1261 817 522 210 102 190 290 703 1135 1791	6.4 6.3 6.1 8.7 6.3 5.3 8.4 5.6	SSE NNW NNW NNW NNW NNW WNW SSE NNW SSE
ar	44.5	26.0	35.2	88	July 6	-42	Jan. 6	10,937	6.6	

Precipitation (Inches) Total Greatest 24 Hour Precipitation Precipi-Snow tation* Snow Date onth Rain Rain Date Jan. 11 0.10 16.5 1.75 .09 Jan. 24 5.1 nuary 2.6 0.27 Feb. 26 1.2 Feb. .OI 0.01 bruary March 30 3.9 Mar. 19 0.92 .10 0.15 7.7 rch I .39 2.5 Apr. Apr. 20 3.5 1:13 ril 0.78 3.27 3.69 .78 May 8 2.2 May 4.2 У 1.54 5.23 June 30 5.23 ine 1.11 July 29 4.40 4.40 ily 18 1.01 Aug. 4.31 4.31 igust 0.89 Sept. 16 2.46 ptember 2.46 0.2 26 15 0.26 1.40 0.3 1.43 Oct. tober Nov. 2.2 0.25 0.03 Nov. 1.1 vember 0.03 Dec. 1.7 7.9 0.79 cember Year 1.54 June 22.14 30 44.90 26.63 5.1. Jan. 11 * Sam of rainfall plus

ANNUAL METEOROLOGICAL SUMMARY

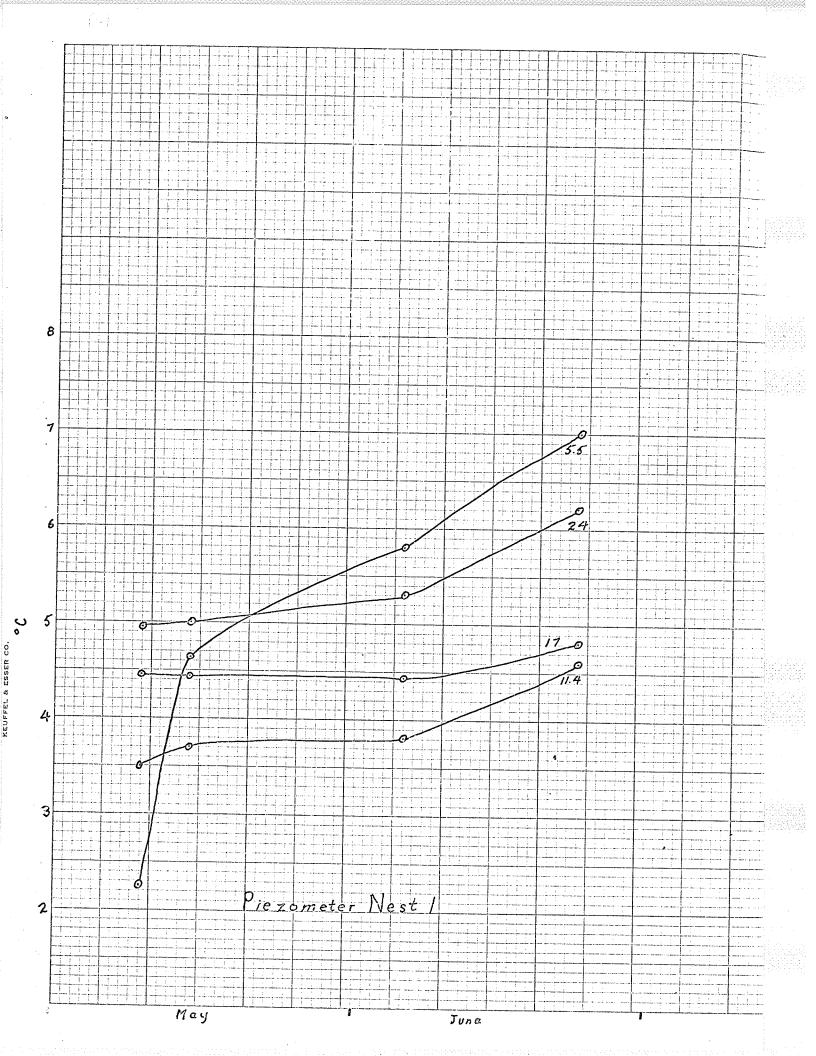
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WNRE, Pinawa, Manitoba

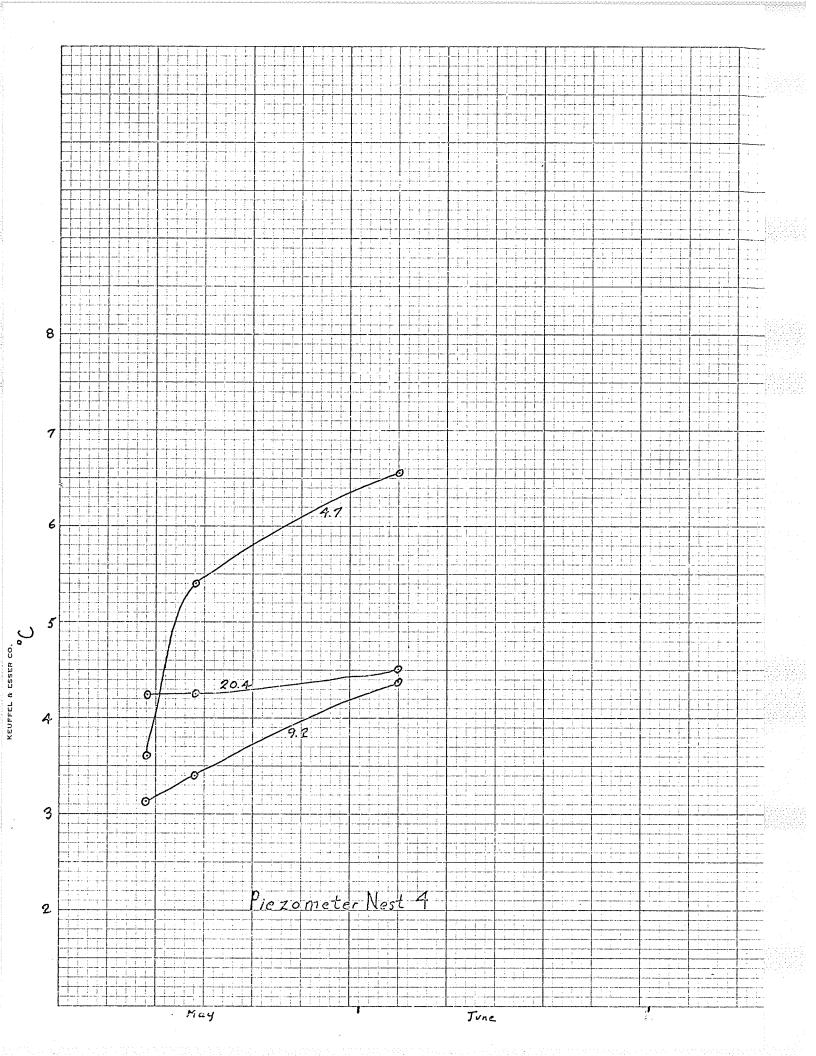
Health and Safety Branch

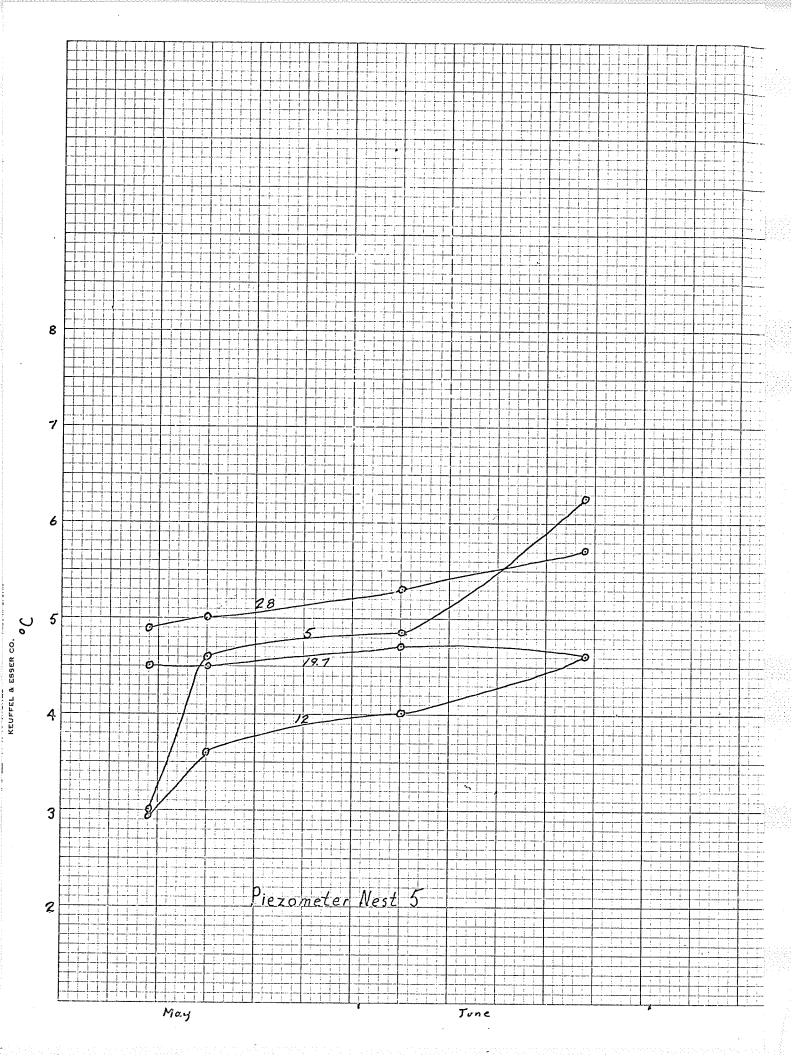
Environmental Control Section

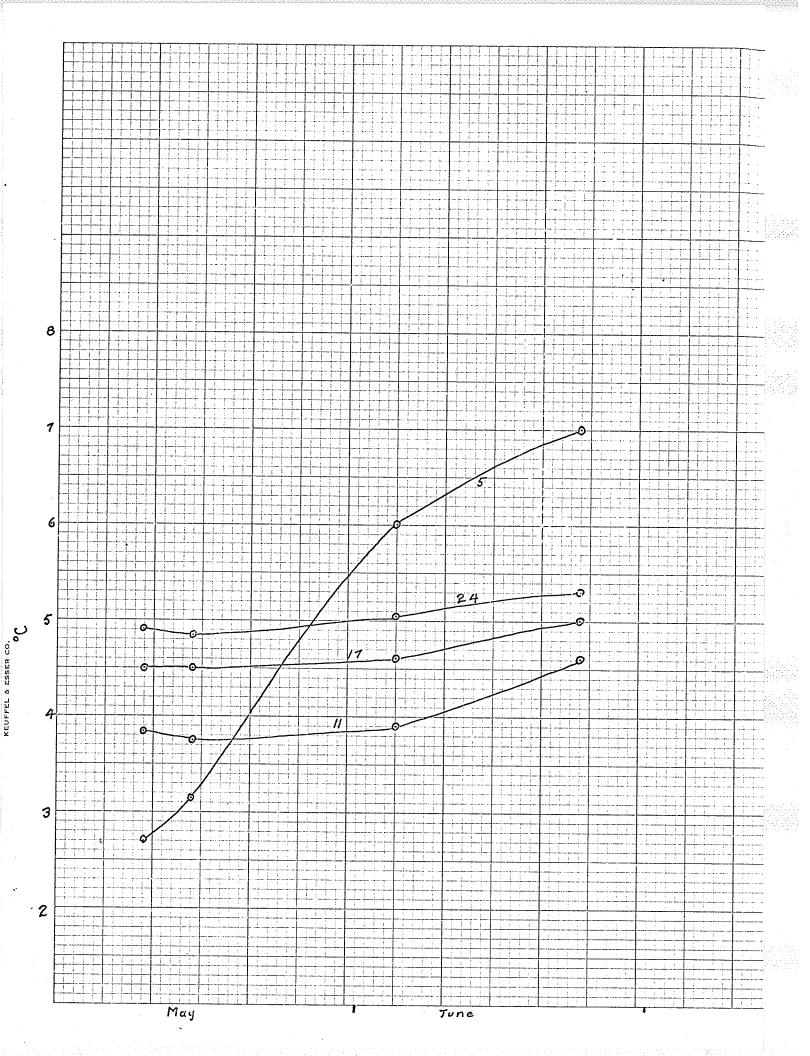
			Temperat	ure			•		Heating	r	Wi	nd
nth	Mear Max		7,	Ex- treme Max.	1	Ex- treme Min.)	Degree Days		age	Pre- vailing Direction
an. eb. ar. by ine ily ig. ept ct. ov. ar	4.0 20.0 26.0 54.6 60.8 62.9 74.5 64.3 41.2 20.3	0 - 1.6 4.0 29.4 39.4 41.1 53.5 56.6 45.5 30.0 18.8 8.7	9.2 15.0 42.0 50.1 52.0 64.0 67.9 54.8 35.6 26.1 14.5	27 36 41 70 85 77 90 84 59 50	Jan. 15 Feb. 23 Mar. 21 Apr. 25 May 26 June 9 July 12 Aug. 22 Sept13 Oct. 3 Nov. 6 Dec. 1	-40 -15 - 2 21 30 38 46 28 15	Feb. Mar Apr May June July Aug. Sept. Oct.	3 29 2 19 13 6 1 28 23 20	2147 1562 1549 691 474 390 87 35 331 913 1167 1566	556 76 55 55 7	.5 .0 .6 .1 .8 .6 .0 .4 .5 .9 .1	NNW SSE N S NNW N SSE SSE SSE SE SSE SSE SSE SSE
	45.2	2 26.0	35.6	93	July 12	-40	Feb.	3	10,912	6	.0	SSE
	,	·			Precipit	ation	·		,		· · · · · · · · · · · · · · · · · · ·	
				1	Total ecipi-	Grea	test :	24_	Hour Pr	ecipit.	ation	
onth		Rain :	Snow	ta	tion*	Rai	n		Date	Snow		Date
nuar brua rch ril y ne ly gust pteml tobe cembe	ber r er	0.01 0.35 2.72 2.60 2.06 3.29 2.46 .91 .01	27.7 3.1 1.9 1.5		2.78 0.31 0.19 0.35 2.87 2.60 2.06 3.29 2.46 1.04 0.38 1.19	• .	28 54 98 62		Jan.14 April30 May 31 June 26 July 26 Aug. 29 Sept.24 Oct. 7 Nov. 11	5.6 0.7 0.8 1.5		Jan. 15 Feb. 10 March 7 May 31 Oct.24-2 Nov. 23 Dec 7
ear		14.41	51.10	,	19.52	J., 6	54	M	ay 31	5.6		Jan. 15
	* 50	un of rain	Pall plus	s one		_					15, 15,	

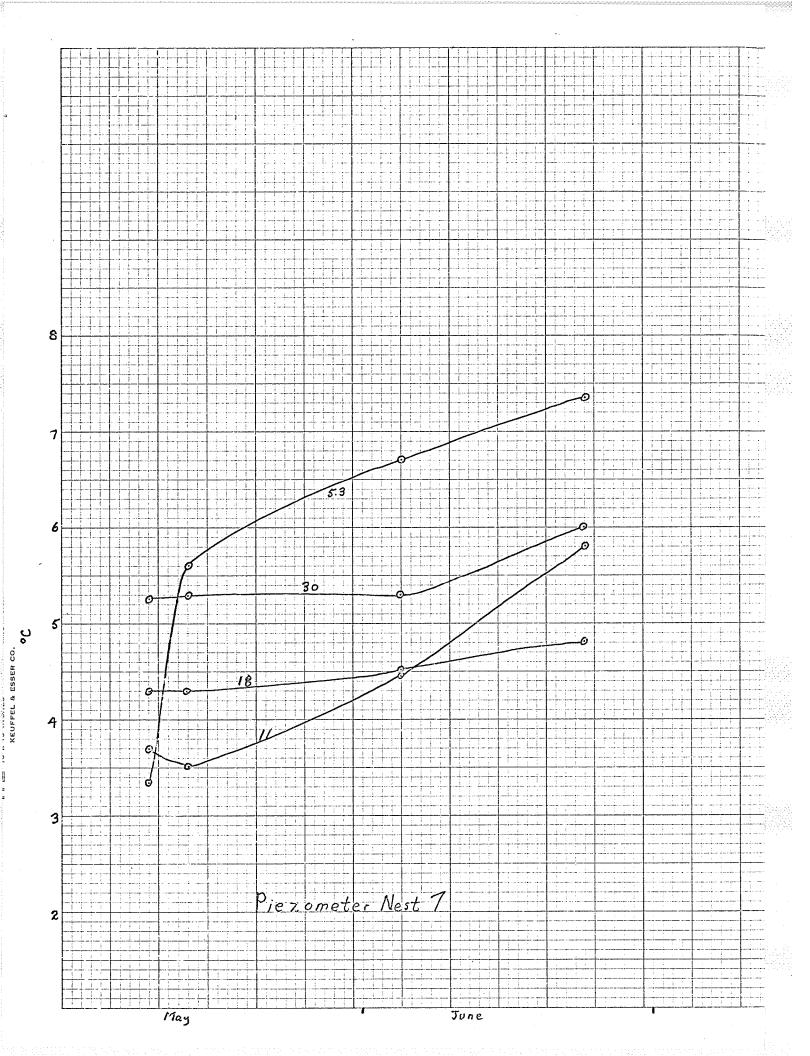


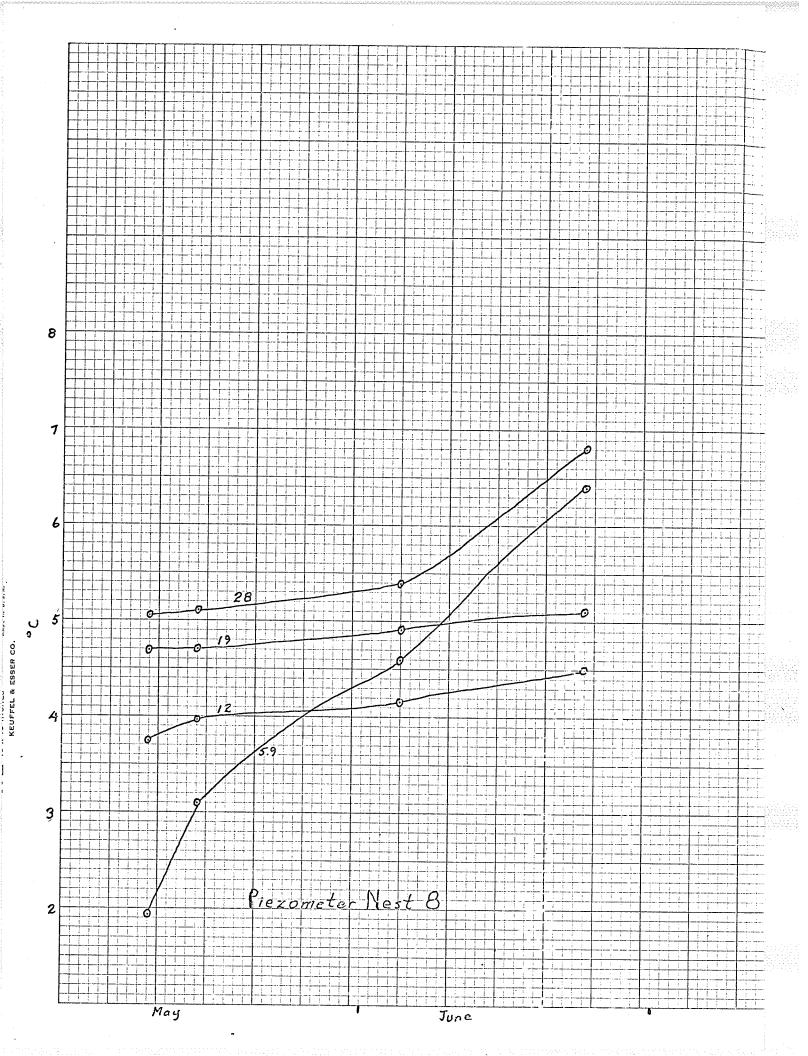
KEUFFEL & ESSER CO.

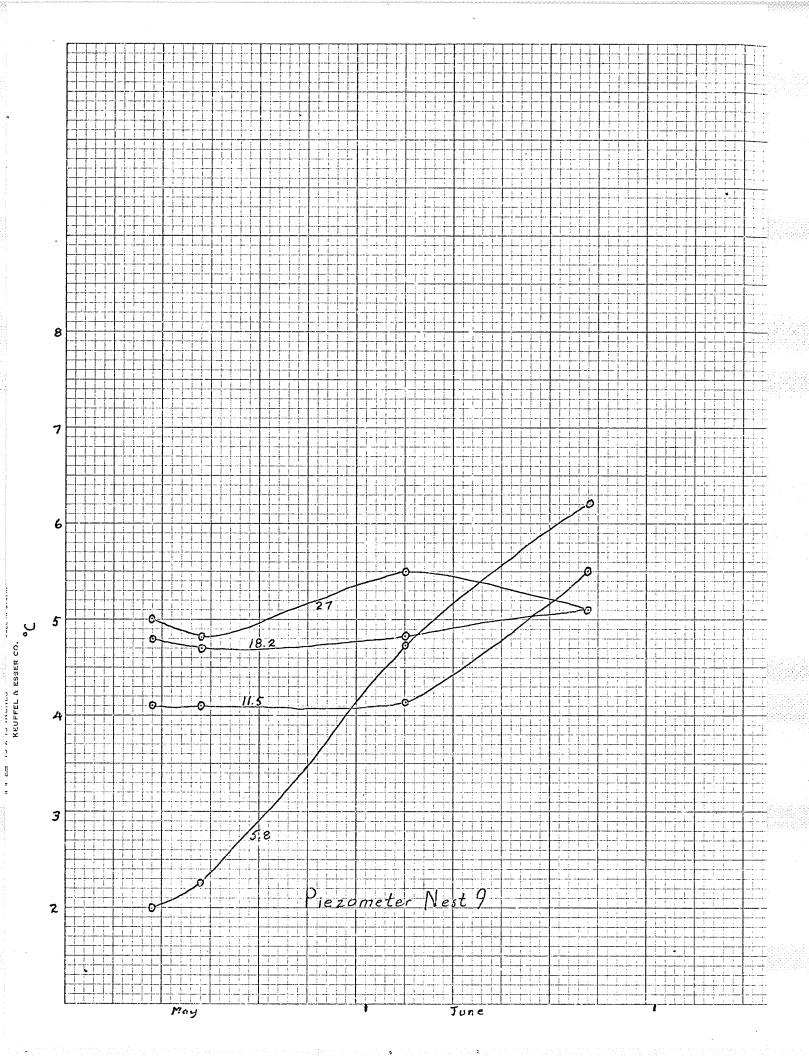


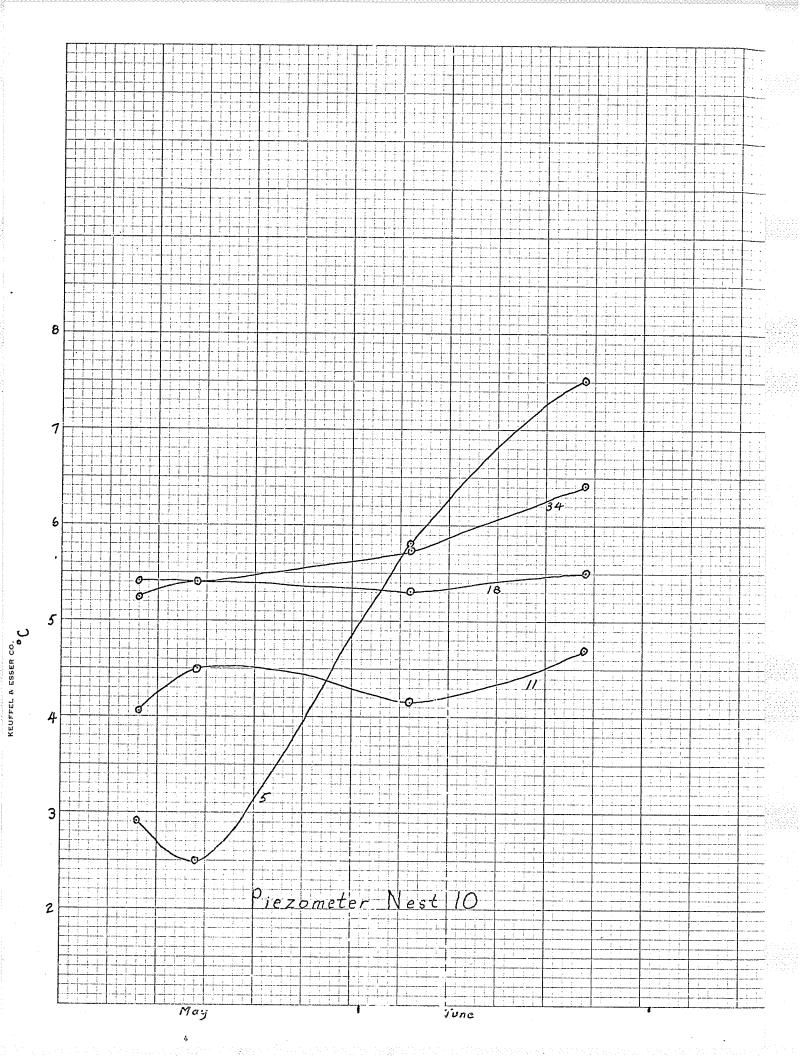


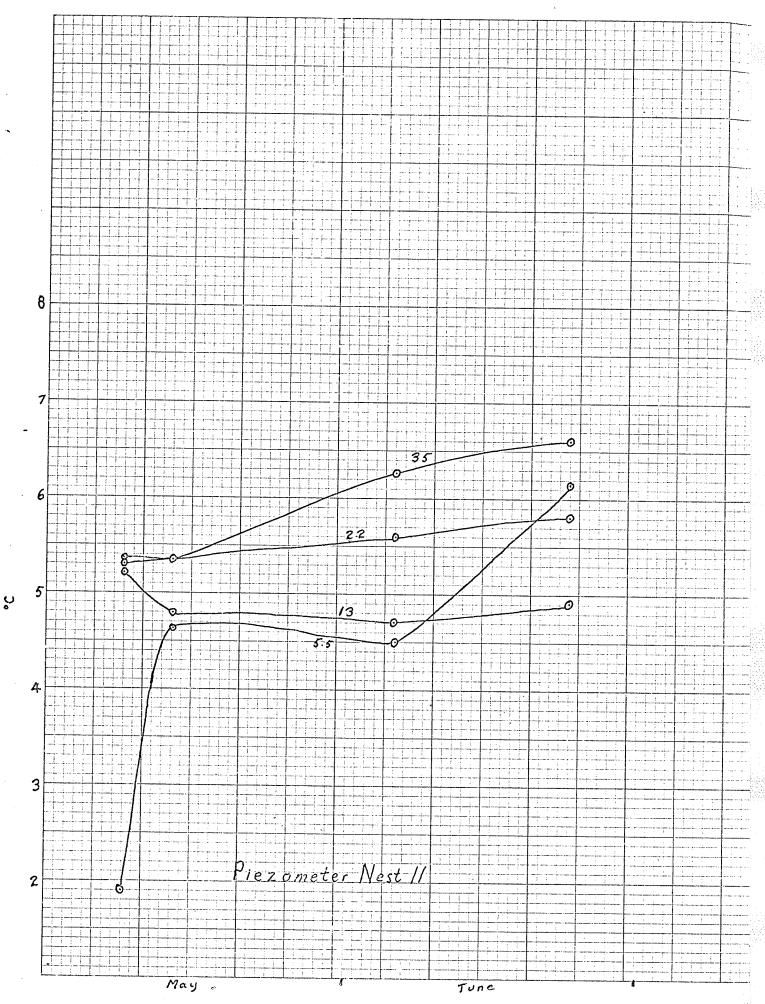








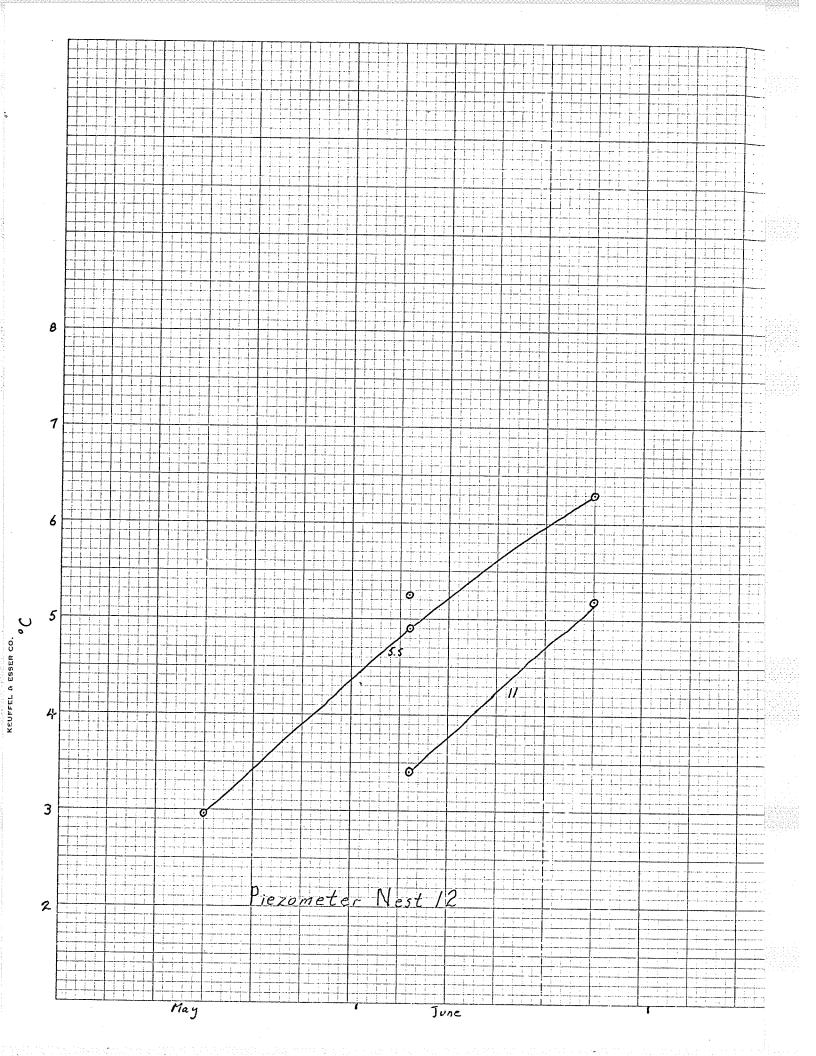


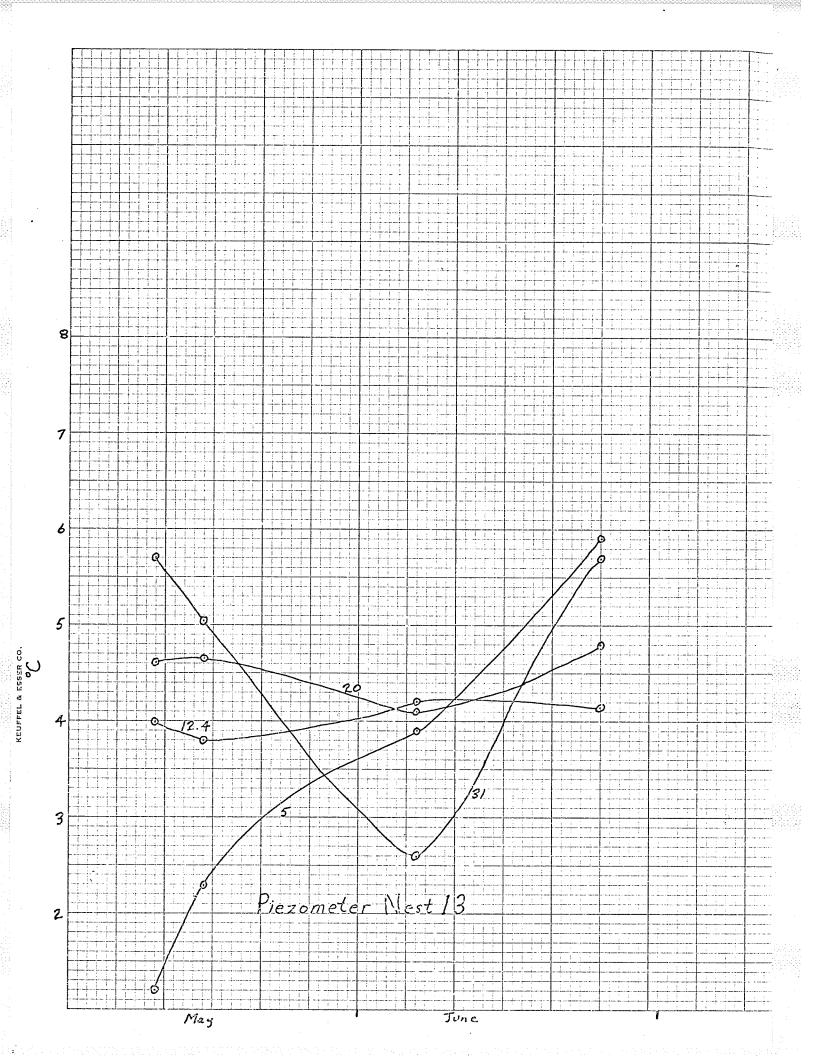


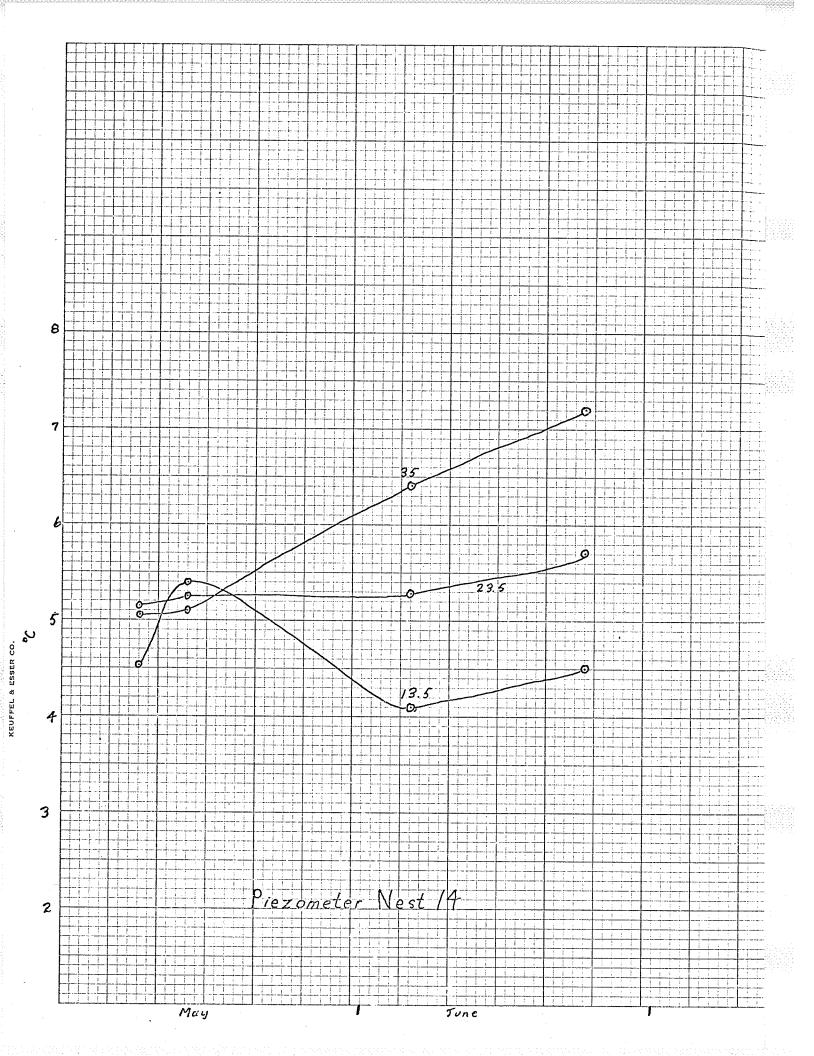
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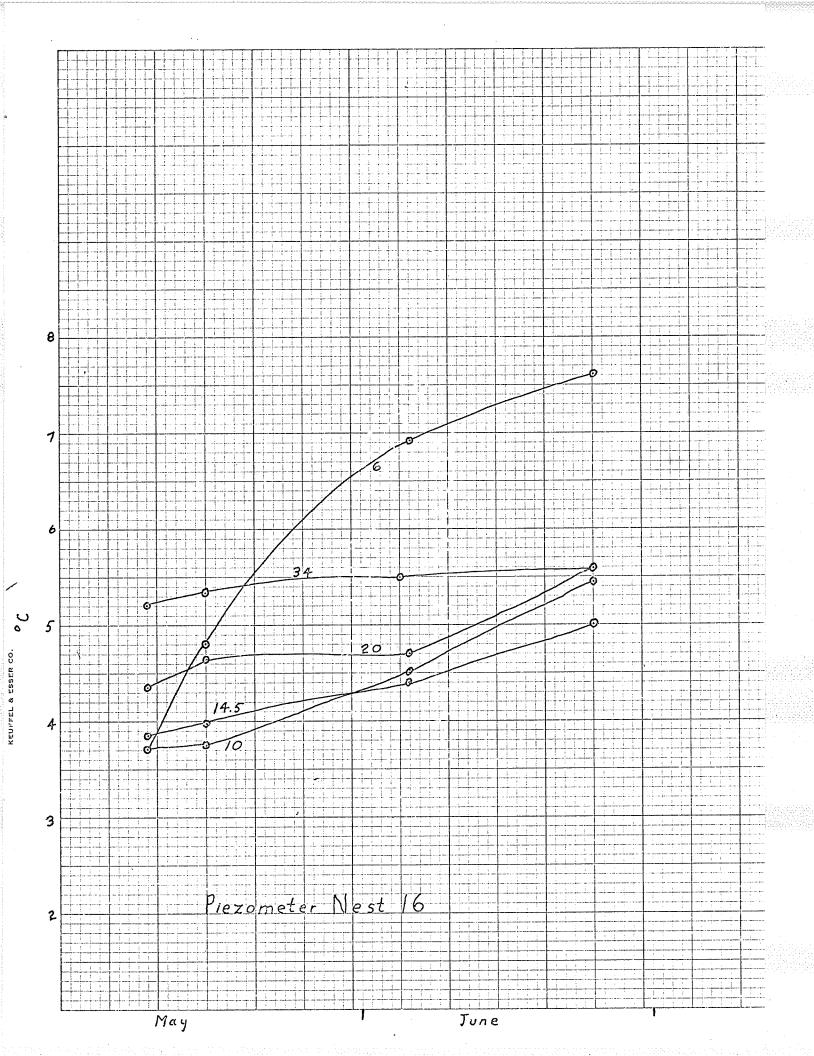
EL & ESSEN CO

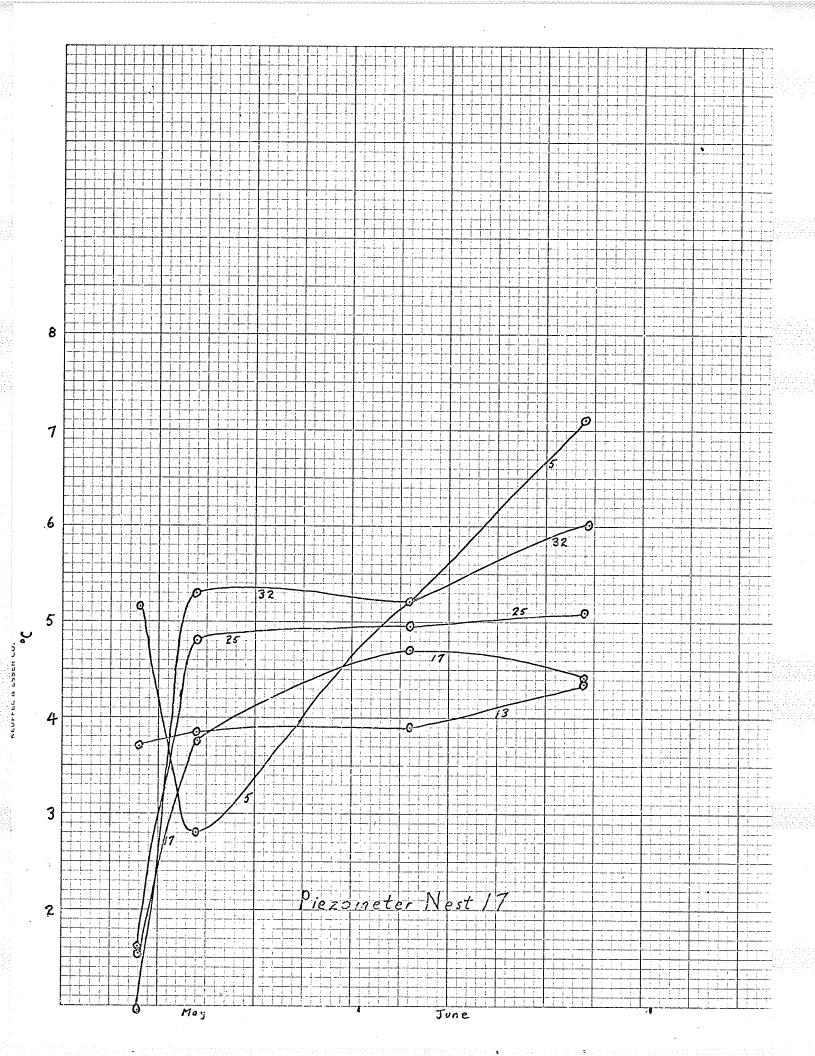
Y

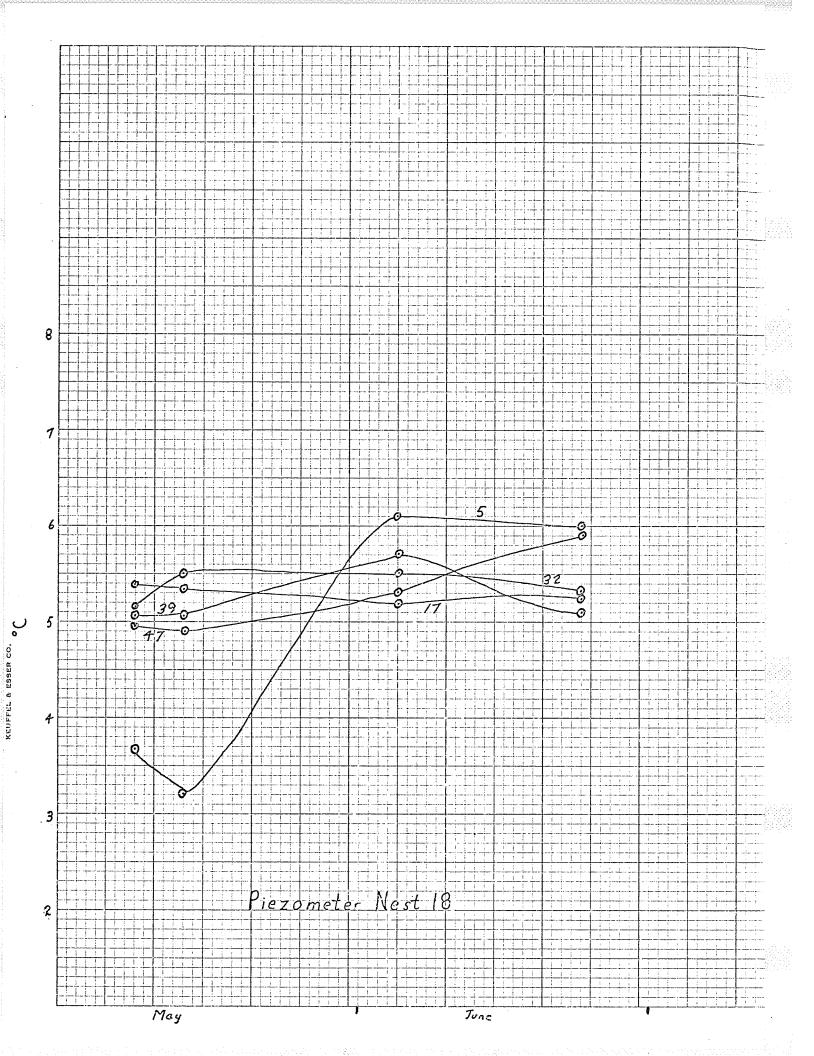


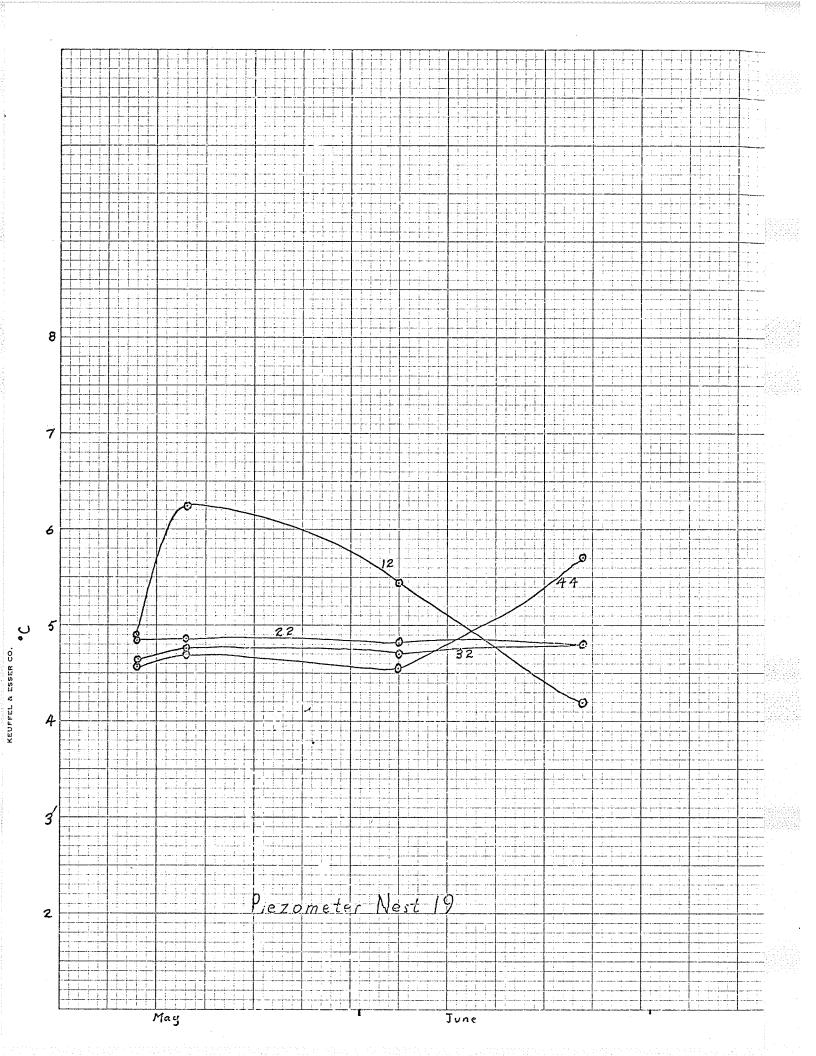


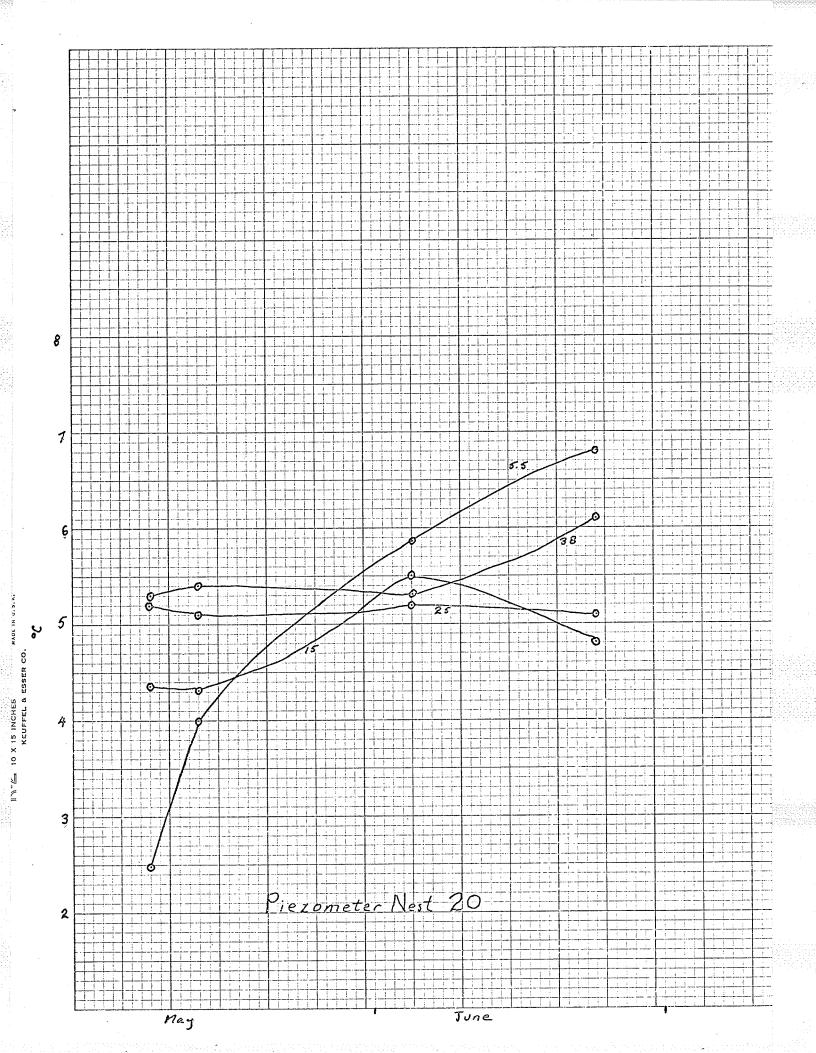


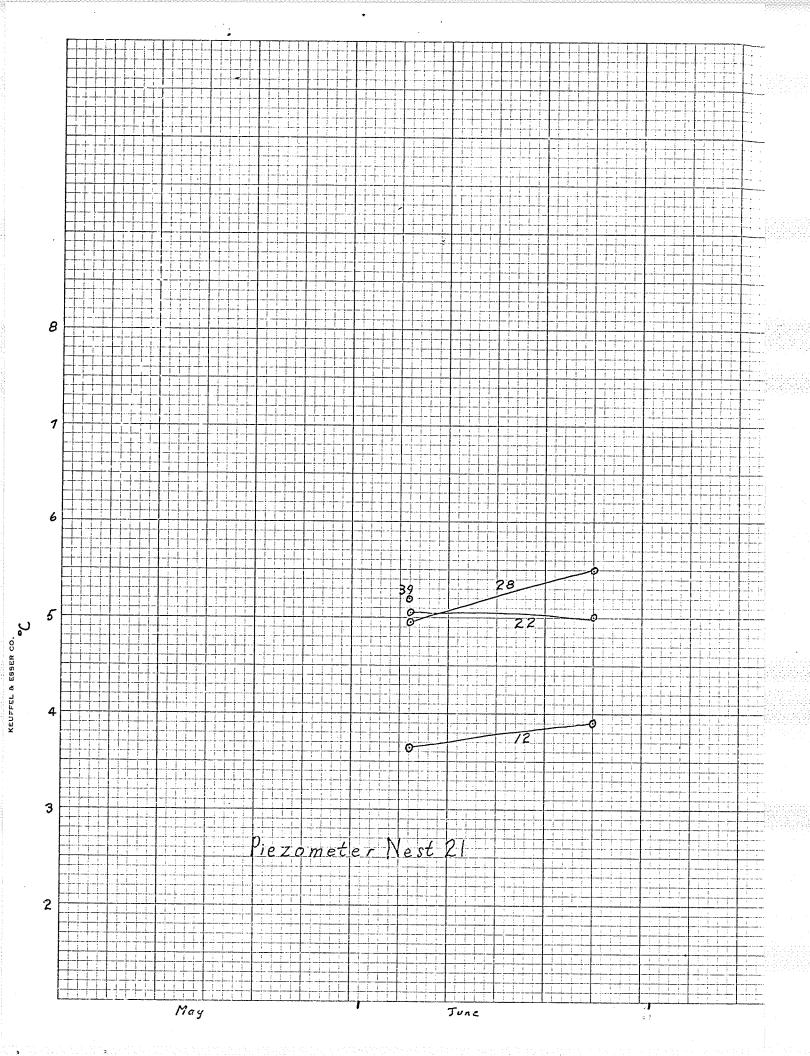


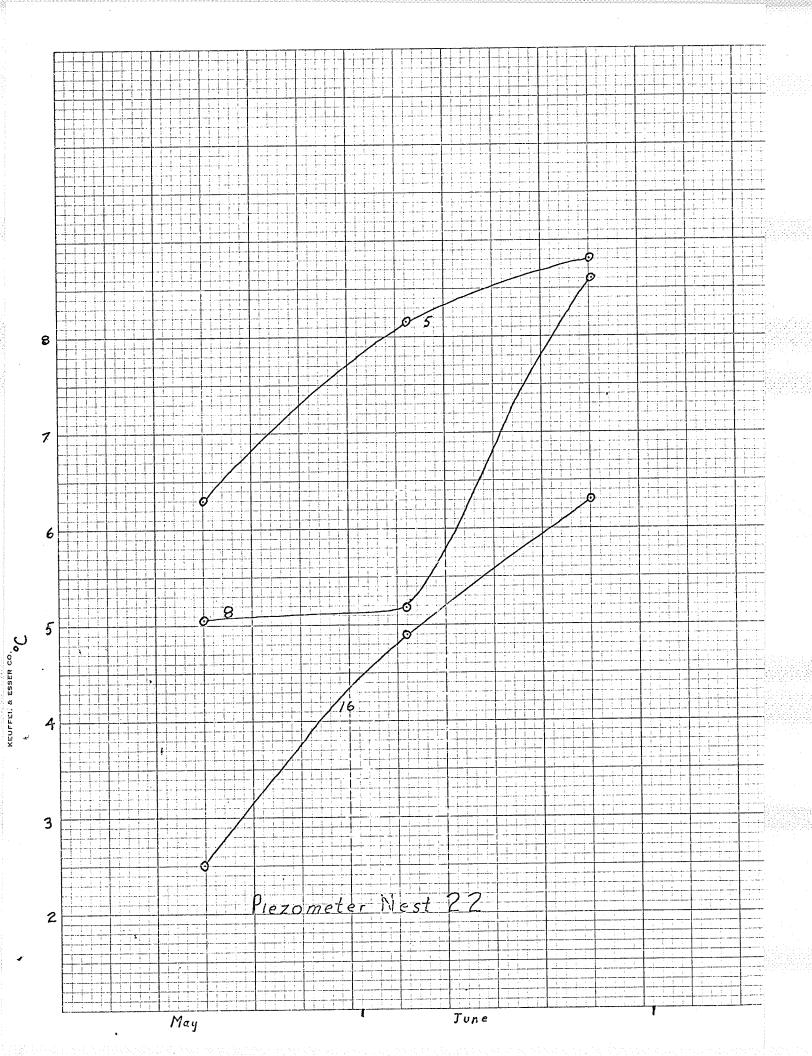












APPENDIX F TRACER VELOCITIES IN GROUNDWATER

Tritiated heavy water was injected into the flow regime under pressures of less than 10 psi. A schematic of the injection system is shown in figure 13A.

Samples were taken with a vacuum pump connected to flasks in which a 20 - 50 ml sample was collected. To ensure a fresh sample for every sampling period, a 100 ml aliquat was taken and discarded prior to final sampling. A downhole sampler identical to that used for the geochemical samples (Fig. 11) was used for the initial few samples but the difficulty of preventing contamination resulted in a modification of sampling technique. Each sample point was given individual vacuum lines and a vacuum flask was used for sampling (Fig. 15A).

For analysis, an aliquot of sample was placed in 20 ml of Chicago Nuclear NE220 liquid scintillator in 25 ml scintillation counter (low sodium) glass vials and counted by a Chicago Nuclear 3-channel scintillometer for beta activity.

Results were printed and stored on paper tape and counts repeated to ensure reliable values. Background radiation was not considered because of its relatively stable 18-25 cpm during the period of study. Decay of the tritium was not considered because of the relatively short period of study relative to the 12.26 year half-life of tritium.

Tabulated activity counts are given only for sample tubes considered active or for those indicating contamination.

INJECTION SITE #1

- July 2, 1969 injection of 20 gallons solution of 20.3 gms fluorescein at 10 psi to ensure adequate safety precautions for tritium injection.

 Immediate response of water level in all sample tubes noted when pressure applied. Injection rate @ 10 psi approximately 2 gpm.
- July 9, 1969 1 curie tritium diluted in 20 gallons water and injected under less than 10 psi.followed by 20 gallons water to flush all tritium into the formation. All sample tubes were corked during injection to prevent overflow due to pressure. Fluorescein noted in sample tube 15 immediately after injection. Samples taken immediately after injection showed IWP-1-15 to have count of 210 cpm.

 Samples taken periodically over the next 96 hours showed a rapid increase to 148,000 cpm/ml. A peak count in sample tube 15 of 826,445 cpm/ml was recorded on August 15, 1969. By November 5, the count had fallen to 675,675 cpm/ml.

The injection well samples indicated counts of 4.78×10^3 cpm/ml (top) and 3.15×10^5 cpm/ml (bottom) on July 21. The difference was a result of the injection of pure water following the tritium. A peak of 3.65×10^6 cpm/ml was reached on August 22, 1969 and by May 15, 1970 the count had fallen to 3.9×10^5 .

Sample tube 13 began to show above background beta activity 8 hours after injection. Values fluctuated but reached a peak of 4,470 on August 17, 1969. A value of 12,300 cpm/ml on July 19 is felt to be in error. By November 5, 1969 the value fell off to 460 cpm/ml.

No other sample tubes indicated the presence of tritium.

Date	Activity (cpm/ml)	Date	Activity (cpm/ml)
21/7	4.780×10^3_5	6/8	mon mon
21/7	3.150×10^{3}	6/8	1.462×10^{6}
22/7	4.149×10^{3}	7/8	2.084×10^{6}
23/7	6.493×10^{5}	8/8	1.429×10^{6}
24/7	7.751×10^{5}	10/8	1.916 × 10 ⁶
25/7	9.708 x 10 ³	11/8	$2.080 \times 10^{\circ}$
28/7		13/8	$1.770 \times 10^{\circ}$
28/7	$1.405 \times 10^{\circ}$	14/8	$2.220 \times 10^{\circ}$
29/7	1.456×10^{6}	15/8	$1.740 \times 10^{\circ}$
30/7	7.619×10^{5}	17/8	$1.870 \times 10^{\circ}$
31/7	9.709×10^{5}	18/8	$1.585 \times 10^{\circ}_{5}$
1/8	9.346×10^{5}	22/8	5.55×10^{3}
3/8	$1.0 \times 10^{\circ}$	29/8	3.65×10^{6}
4/8	-	7/9	1.136×10^{6}
4/8	1.550×10^{6}	17/9	7.368×10^{5}
5/8	/	8/10	8.620×10^{5}
5/8	1.417 × 10 ⁶	5/11	5.848×10^{9}

Date	Activity (cpm/ml)	Date	Activity (cpm/ml)
9/7	210	28/7	529,100
10/7	51,400	29/7	549,450
10/7	70,600	30/ 7	5 52,485
10/7	95,000	31/7	537,633
10/7	119,000	1/8	502,500
11/7	102,000	3/8	507,613
12/7	100,000	4/8	526,315
13/7	148,000	5/8	518,134
14/7	168,500	6/8	490, 195
15/7	170,000	7/8	5 55,555
15/7	177,600	8/8	588,234
16/7	290,000	10/8	609 <i>,7</i> 55
16/7	293,000	11/8	653,594
17/7	288,000	13/8	704,224
18/7	357,000	14/8	781,249
19/7	367,000	15/8	826,445
20/7	385,000	22/8	734,607
21/7	426,000	17/8	793,65 0
22/7	438,000	18/8	793,650
23/7	448,430	29/8	757,120
24/7	462,962	7/9	<i>7</i> 75, 193
25/7	487,804	17/9	813,674
27/7	490,195	8/10	746,268

Date	Activity (cpm/ml)	Date	Activity (cpm/ml)
9/7	7	29/7	910
10/7	3 6	30/7	775
10/7	91	31/7	940
10/7	134	1/8	850
10/7	303	3/8	850
11/7	453	4/8	1,325
12/7	535	5/8	950
13/7	540	6/8	1,565
14/7	508	7/8	2,060
15/7	435	8/8	1,755
15/7	235	10/8	2,095
16 <i>/</i> 7	375	11/8	3,171
16/7	625	13/8	2,899
17/7	625	14/8	2,436
18/7	2,225	15/8	2,650
19/7	12,300	22/8	3,200
20/7	4,700	17/8	4,467
21/7	2, 750	18/8	2,330
22/7	1,675	29/8	1,782
23/7	875	7/9	1,494
24/7	1,175	17/9	1,289
25/7	1,025	8/10	816
27/7	1,125	5/11	460
28/7	655		

INJECTION SITES #2 & #3

The procedure used for injection #1 was followed for injections 2 and 3 with the exception that the fluorescein test was not used. In injections 2 and 3 only 10 millicuries of tritium was injected rather than the 1 curie used at 1.

Serious sample tube contamination occurred in injection site 3 in sample tube 20 and 24. It probably resulted from contamination during the initial few sampling periods before a routine sampling procedure was developed. The anomalous values are not felt to be true indicators of the subsurface movement of tritium because the values are very low relative to what would be expected from the 10 mCi injection.

A series of peaks and declines in beta activity in the active holes is felt significant, although of what is not certain.

Date	Activity (cpm/ml)	Date	Activity (cpm/ml)
20/8	495,282	3/9	226,244
20/8	6 86	5/9	204,498
20/8	248	10/9	182,546
20/8	26,404	13/9	168,960
24/8	189,907	17/9	147,547
28/8	252, 650	20/9	163,335
22/8	93,023	24/9	307,666
25/8	225,73 3	1/10	184, 162
2 7/8	264,550	8/10	112,935
30/8	2 26,244	17/10	49,799
4/9	222,222	24/10	28,778
1/9	220,264	5/11	16,446
2/9	228,310	24/11	12,492

Date	Activity (cpm/ml)	Date	Activity (cpm/ml)
20/8	20	13/9	1,948
20/8	13,298	17/9	1,846
24/8	13,872	20/9	1,267
22/8	25,203	24/9	1,444
26/8	14,464	1/10	1,046
- 27/8	9,292	8/10	1,350
30/8	8,069	17/10	6,277
4/9	4,842	24/10	6,925
5/9	3, 887	5/11	7,096
2/9	6,589	24/11	6,809
10/9	2,689	6/ 5/70	4,525

Date	Activity (cpm/ml)	Date	Activity (cpm/mi)
24/8	32	4/9	52
28/8	30	5/9	23
30/8	34	10/9	30

Date	Activity (cpm/ml)	Date	Activity (cpm/ml)
20/8	29	2/9	73
20/8	45	10/9	74
24/8	74	1/10	67
22/8	73	8/10	59
26/8	93	17/10	100
27/8	72	24/10	82
30/8	94	5/11	77
4/9	135	24/11	94

Date	Activity (cpm/ml)	Date	Activity (cpm/ml)
20/8	20	17/9	161
20/8	24	24/9	101
24/8	27	1/10	114
28/8	77	8/10	106
30/8	82	17/10	100
4/9	20	24/10	86
5/9	144	5/11	82
10/9	163	24/11	101

Date	Activity (cpm/ml)	Date	Activity (cpm/ml)
21/8	57,7 35	10/9	116,157
21/8	349,950	13/ 9	106,951
21/8	6,000	17/9	83,756
24/8	5,007	20/9	85,917
26/8	9,454	24/9	76,915
28/8	106 <i>,5</i> 93	1/10	59,245
30/8	138,979	8/10	53,248
· 1/9	137,363	17/10	41,119
3/9	132,626	24/10	3 7,825
4/9	132, 978	5/11	32,873
5/9	130,208	24/11	33,213

Date	Activity (cpm/ml)	Date	Activity (cpm/ml)
24/8	28,926	13/ 9	46,168
28/8	37,059	17/9	42,706
21/8	28	20/ 9	45,241
22/8	16,239	24/9	44,339
26/8	35,844	1/10	41,098
30/8	30,754	8/10	40,919
4/9	43,195	17/10	37,900
1/9	42,444	24/10	37,797
3/9	46,749	5/11	36,416
5/9	42,633	24/11	43,370
10/9	42,047		

Date	Activity (cpm/ml)	Date	Activity (cpm/ml)
24/8	102,499	10/9	27,120
28/8	99,087	13/9	27,901
21/8	15,836	17/9	25,177
22/8	100,678	20/9	27,865
23/8	106,461	24/9	27,282
3 0/8	98,597	1/10	22,663
4/9	52,402	17/10	19,523
1/9	77, 869	24/10	19,627
3/9	60,961	5/11	20,006
5/9	42, 890	24/11	20,538

TRITIUM ANALYSES RECORD - IWP-2-1

Date	Activity (cpm/ml)	Date	Activity (cpm/ml)
24/8	853	13/9	963
28/8	2,389	17/9	695
21/8	221	20/9	744
23/8	679	24/9	650
3 0/8	1,794	1/10	<i>5</i> 45
4/9	1,322	8/10	649
5/9	1,301	17/10	511
1/9	1,488	24/10	521
3/9	1,373	5/11	<i>5</i> 50
10/9	1,070	24/11	625

APPENDIX G

LISTING OF AND COMMENTS ON
INPUT AND OUTPUT FOR THE
STEADY-STATE MODEL

The first program includes a complete set of data input and output.

The second program includes print checks for debugging new data decks.

```
THIS PROGRAM CALCULATES TOTAL HEAD AT ALL NODES IN A REGION
    WHEN GIVEN HEADS AT SPECIFIED BOUNDARY NODES
    REAL NE(352), KX(576), KY(576)
    COMMON M, CJ (352,2), IE (576,3), L1, L2, L3, T, ST (3), KL, KX, KY, S (576)
    DIMENSION LIST(353), A(353), ISR(576), HDG(20)
                                           NEIN(353) \cdot NI(353 \cdot 8)
    DIMENSION STORE (6000),
    DIMENSION G(3,6), DISPL(6), STRESS (576)
    EQUIVALENCE (NE; NN)
    JPD=5 /
    JWT=6 3
    READ JOB TITLE
  1 READ(JRD,510,END=9999)(HDG(I),I=1,20)
    WRITE (JWT, 512) (HDG(I), I = 1, 20)
512 FORMAT (*1*//20A4)
510 FORMAT (20A4)
    WRITE(JWT,600)
600 FORMAT(/ TWO DIMENSIONAL SEEPAGE PROBLEM 1// INPUT DATA 1//)
    READ NUMBER OF ELEMENTS, NUMBER NODES, NUMBER SPECIFIED NODAL
    HEADS, THICKNESS
    ALL INPUT PRINTED OUT FOR CHECKING PURPOSES
    READ(JRD, 500) NEL, NE, NBD, T
    IF(T.EQ.O.) T=1.
    WRITE (JWT, 601) NEL, NN, NBD, T
601 FORMAT(14, * ELEMENTS*14, * NODES*14, * SPECIFIED NODAL HEADS*//*
   ILEMENT THICKNESS = *F8.3//)
500 FORMAT (315, F10.3)
    READ COORDINATES OF NODES
    READ(JRD, 501)((CJ(J, I), I=1, 2), J=1, NN)
    WRITE (JWT, 602)
602 FORMAT(///* NODAL COORDINATES (FT)*
                                                          X COORD
                                              1/1
                                                    NODE
   1 Y COORD!)
    DO 99 J=1,NN
99 WRITE(JWT,694)J,CJ(J,1),CJ(J,2)
501 FORMAT (6F10.3) -
    ZERO CONSTANT VECTOR
191 DO 39 J=1,NN
 39 NL(J) = 0.
    GENERATE NODAL INCIDENCE TABLE
    WRITE (JWT, 686)
686 FORMAT(// ELEMENT INCIDENCE TABLE 1/1 ELEMENT . NODE 1
                                                                        NOD
   1E 2 NODE 3
                     KX
                                KY
                                        SLOPE!//)
    SS=0.
    DO 17 M=1, NEL
    READ(JRD,505)(IE(M,I),I=1,3), XK, YK, SL
505 FORMAT(315,3F10.3)
    IF (XK. NE.O.) XX=XK.
    TF (YK, NE, O.) YY=YK
    IF(SL.NE.O.) SS=SL
    S(M) = SS
    KX(M) = XX
    KY(M) = YY
 17 WRITE(JWT, 687)M, (IE(M, I), I=1, 3), XX, YY, S(M)
687 FORMAT (419.3F10.3)
    DO 18 J=1,NN
    I = 0 = (U) NIBN
```

```
DO-18, Mel.6, ...
 18 MI(J_*M) = 0.5
    DO 19 M=1.NEL
    DO 19 K=1.3
    J = IE(M,K)
    NEIN(J) = NEIN(J) + 1
    L = NEIN(J)
 19 NI(J,L)=M
694 FORMAT(15,6F12.4)
    READ ANY PRESCRIBED NODAL HEADS
    DO 376 J=1.NN
376 ISR(J) = 0
    WRITE(JWT, 596)
596 FORMAT(///* PRESCRIBED NODAL HEADS*//* NODE PRESCRIBED HEAD*//)
    00 574 I=1,NBD
    READ(JRD, 694)J, NL(J)
    ISR(J)=1
574 WRITE (JWT, 694) J, NL(J)
    GENERATION: AND ELIMINATION OF NODAL EQUILIBRIUM EQUATIONS
    GENERATE I THE ROW OF STIFFNESS MATRIX AND STORE IN A TEMPORARILY
172 \text{ LIST(I)} = 1
    00 100 I=1, NE
    NON ZERO BAND OF ROW I IN STIFFNESS IS FROM KL TO KH. KL = LOWEST
     NODE NO FOR NODES INCIDENT ON ELEMENTS INCIDENT ON NODE I., KH = HIGHEST
    KL = I
    KH = I
    IM=NEIN(I)
    DO 5 J=1.IM
    M=NI(I,J)
    DO 5 K=1,3
    JF=IE(M,K)
    IF(JF-NE)10,10,5
 10 IF(JF-KH)14,14,12
 12 \text{ KH=JF}
    GO TO 5
 14 IF(JF-KL)16,5,5
 16 KL=JF
  5 CONTINUE
    ZERO ALL A MATRICES IN NON - ZERO BAND
    K = KH - KL + 1
    DO 29 J=1,K
 29 A(J)=0.
    INSERT STIFFNESS MATRICES INTO NON ZERO BAND
    DO 200 J=1,IM
    M=NI(I,J)
    D0 20 L=1,3
   -IF(I-IE(M.L))20,22,20
 22 \text{ KI=IE(M,L)}
    1.1 = L
    L2 = L + 1 - 3 * (L/3)
    KJ = IE(M, L2)
    L3=L+2-3*(L/2)
    KK = IE(M, L3)
 20 CONTINUE
    CALL STIFF
    NNI, NJ AND NK ARE NODES I, J AND K FOR ELEMENT M - (POSITION IN ROW I
    RELATIVE TO KL = 1
```

NNI=I-KL+1 . INSERT STIFFNESS MATRIX I $\Delta(NNI) = \Delta(NNI) + ST(I)$ 12.61e Water TEST WHETHER NODE NJ IS A FIXED SUPPORT IF(KJ-NE)47,47,6 INSERT STIFFNESS MATRIX J .47 NJ=KJ-KL+1 A(NJ) = A(NJ) + ST(2)TEST WHETHER NODE NK IS FIXED SUPPORT 6 IF (KK-NE)8,8,200 8: NK=KK-KL+1 INSERT STIFFNESS MATRIX K $\nabla A(NK) = A(NK) + ST(3)$ 200 CONTINUE MODIFY EQUATION IF SPECIFIED HEAD AT NODE I IF ([SR (I) . EQ . 0) GO TO 497 A(NNI)=10.**25 ← NL(I)=NL(I)*10.**25 497 LINC=KH-I FOR FIRST EQUATION, BYPASS ELIMINATION IF(I-KL)60,60,50 PERFORM ELIMINATION FOR ROW I TO ZERO BELOW MAIN DIAGONAL 50 KU=I-1 00 256 K=KL,KU IK = PIVOTAL COLUMN RELATIVE TO KL = 1IK=K+1-KL IM=LIST(K+1)-LIST(K) IJ=K+IM-I-LINC IF NON ZERO BAND FOR PIVOTAL EQ ENDS TO RIGHT OF THAT FOR EQ I, EXTEND FOR EQ I IF(IJ.LE.O.) GO TO 82 KK=LINC+I-KL+2 LINC=LINC+IJ LL = IJ + KK - IDO 83 L=KK, LL 83 A(L)=0. 82 IF (IM.LE.O.) GO TO 256 .DO 55 J=1, IM IJ = IK + JKJ = LI'ST(K)+J-1IF(ABS(STORF(KJ)).LE.10.**(-10)) STORE(KJ)=0. 55 A(IJ)=A(IJ)-A(IK)*STORE(KJ) 256 NL(I)=NL(I)-A(IK)*NL(K)NORMALIZE ROW I. JJ = PIVOTAL ELEMENT RELATIVE TO KL = 1 60 JJ = I + I - KLIJ=JJLIST(I+1)=LINC+LIST(I) IF(LINC)75,75,65 65 DO 70 J=1, LINC 576 . IJ=IJ+1 " IK=LIST(I)+J-1 / • 70 STORE(IK) = A(IJ) / A(LL) NORMALIZE HEAD AT NODE I 75 NE(I)=NL(I)/A(JJ)100 CONTINUE START BACK SUBSTITUTION

```
NZENETI
   IF(N2)210,210,220
220 DO 250 K=1,N2
    I=NE-K
    KU = LIST(I+1) - LIST(I)
    SUM=0.
   00 240 J=1.KU
    IK = LIST(I) + J - 1
    IJ=I+J
240 SUM=SUM+STORE(IK)*NL(IJ)
250 NL(I)=NL(I)-SUM
   WRITE NODAL HEADS
210 WRITE(JWT,560)(HDG(I),I=1,20)
                                      NODAL HEADS 1//1
                                                       NODE
                                                             HEAD (FT
560 FORMAT(*1*//20A4,//* RESULTS*//*
   1)*//)
   DO 397 L=1,NN
397 WRITE (JWT, 694) L, NL(L)
   CALL PLOTER (NL)
GO TO 1
```

END

```
SUBRUUTINE STIFE
THIS SUBROUTINE DEVELOPS STIFFNESS MATRICES II I AND IK FOR
CONSTANT STRAIN TRIANGLE I-J-K
REAL ME (352), KX (576), KY (576)
COMMON M,CJ(352,2),IE(576,3),L1,L2,L3, T,ST(3),KL,KX,KY,S(576)
DIMENSION A(3), B(3)
SL = ABS(S(M)/57.2958)
COSA=COS(SL)
SINA=SIN(SL)
IF(S(M).LE.O.) SINA = -SINA
I = I \in (M, 1)
J = I \in (M, 2)
K = IE(M,3)
A(1) = (CJ(K, 1) - CJ(J, 1)) *COSA + (CJ(K, 2) - CJ(J, 2)) *SINA
A(2) = (CJ(1,1) - CJ(K,1)) * COSA + (CJ(1,2) - CJ(K,2)) * SINA
A(3) = (CJ(J,1) - CJ(I,1)) * COSA + (CJ(J,2) - CJ(I,2)) * SINA
B(1) = (CJ(J,2) - CJ(K,2)) * COSA - (CJ(J,1) - CJ(K,1)) * SINA
B(2) = (CJ(K,2) - CJ(I,2)) * COSA - (CJ(K,1) - CJ(I,1)) * SINA
B(3) = (CJ(I,2)-CJ(J,2))*COSA-(CJ(I,1)-CJ(J,1))*SINA
DET=CJ(J,1)*CJ(K,2)-CJ(K,1)*CJ(J,2)-CJ(I,1)*CJ(K,2)+CJ(K,1)*CJ(I,2)
1)+CJ(I,1)*CJ(J,2)-CJ(J,1)*CJ(I,2)
ST(1) = (KY(M) * A(L1) * A(L1) + KX(M) * B(L1) * B(L1)) / 2./DET
ST(2) = (KY(M) *A(L1) *A(L2) + KX(M) *B(L1) *B(L2))/2./DET
ST(3) = (KY(M) * A(L1) * A(L3) + KX(M) * B(L1) * B(L3))/2./DET
RETURN
END
```

```
SUBROUTINE PLOTER(Z).
COMMON M, CJ (352,2), IE (576,3), L1, L2, L3, T, ST (3), KE, KX, KY, S (576)
DIMENSION Z(352), IBUF(1000)
CALL PLOTS (18UF, 1000)
CALL PLOT (0.0,-11.0,-3)
CALL PLOT(1.0,1.0,-3)
DRAW AXES
CALL AXIS(0.,0.,*JOB320*,-6,40.,0.,0.,200.)
CALL AXIS(0.,0., 1,1,6.,90.,800.,20.)
SCALING
DO 20 I=1,352
\Delta = CJ(I,1)/200.
B = (CJ(I,2) - 800)/20.
PLOT POINTS
CALL SYMBOL(A, B, 0.07, 4, 0., -1)
C = A - 0.175
0 = 8 - 0.105
E = Z(I) - 300
CALL NUMBER (C, D, 0.07, E, 0., 2)
CONTINUE
CALL PLOT (40.,0.,999)
RETURN
END
```

FINITE ELEMENT SEEPAGE PROGRAM BESWICK PINAWA

TWO DIMENSIONAL SEEPAGE PROBLEM

INPUT DATA

576 ELEMENTS 352 NODES 57 SPECIFIED NODAL HEADS

FIEMENT THICKNESS = 1.000

NODAL COORDINATES (FT)

NODE	X COORD	Y COORD
1		839.0000
2	70.0000	849.0000
3	95.0000	839,0000
4	1,0000	830.0000
5	120.0000	830.0000
6	205.0000	848.0000
7	225.0000	839.0000
8	200.0000	860.0000
9	1.0000	822,5000
10	200.0000	822.0000
11	300.0000	830.0000
12	390.0000	821.0000
13	485.0000	830.0000
14	420.0000	839.0000
15	375.0000	848.0000
16	320.0000	859.0000
17	340.0000	865.0000
18	450.0000	868.0000
19	400.0000	865.0000
20	490.0000	859.0000
21	550,0000	821.0000
22	600 10000	839.0000
23	650.0000	830.5000
24	700.0000	820,0000
. 25	570.0000	848.0000
26	695.0000	859.0000
27	570.0000	865.0000
28	610.0000	869.0000
29	790.0000	865.0010
30	830.0000	869.0000
3.1	880.0000	859.0000
32	720.0000	848.0000
. 33	760.0000	839,0000
34	810.0000	830.5000
35	900,0000	817.5000
36	910.0000	839.0000
37	890.0000	848.0000
38	1000.0000	859.0000
-39	№ -940 . 0000	865,0000
40 .	1010,0000	869.0000
41	1120.0000	869.0000
42	1080.0000	865.0000
43	1030.0000	848.0000
44	1000.0000	828.0000
45	1080.0000	813.0000
Salara, a dia di	and the second s	The second secon

47	1140.0000	839.0000		
48	1250.0000	848,0000		
49	1180.0000	859.0000		
50	1240.0000	865.0000		
51	1305.0000			
52	1510.0000	870.0000		
53	1440.0000	865.0000		
54 5 5	1390.0000	859.0000		
55 57	1440.0000	848.0000		
56	1330.0000	839.0000		
57 58	1410.0000	828.0000		
20 59	1280.0000	813.0000		
60 60	1440.0000	815.0000		
	1620.0000	817.5000		
61 62	1590.0000	829.0000		jat saksaka.
63	1510.0000 1690.0000	839.0000 839.0000		
- 64	1610.0000	848,0000		
65		859.0000		
66	1550.0000			
67	1620.0000 1700.0000	865.0000 -0000.0000		
68	1860.0000	870.5000		
:69	1800.0000	865.0000		
70	1710.0000	859.0000		
71	1790.0000	848.0000		
72	1820.0000	839,0000		
73	1740.0000	829,5000		
74	1790.0000	820.0000		
75	1880,0000	829.5000		
76	1970.0000	821.0000		
77	1990.0000			
78	1920.0000	848.0000		
79	1890.0000	859,0000		
80	1950.0000	865.0000		
81	2005.0000	870.5000		
82	2150.0000	870.5000		
83	2090.0000	865.0000		
84	2020.0000	859,0000		
85	2090.0000	848.0000		
86	2140.0000	839,0000		
87	2040.0000	830.0000		
88	2160.0000	824.0000		
89	2320.0000	824.5000		
90	2225.0000	831.0000		
91	2310.0000	839,0000	en de la production de la completa del la completa de la completa del la completa de la completa del la completa de la completa de la completa del la completa della d	
92	2250.0000	848.0000		
93	2200.0000	859.0000		
94	2250.0000	865,0000		
95	2330.0000	871.0000		
96	2515.0000	871.0000		
97	2430.0000	865.0000		
9,8	2380.0000	859,0000		- Negative
- 99	2420.0000	848.0000		
100	2470,0000	839.0000		
101	2400.0000	832.5000		
102	2490.0600	824.5000		
103	2560.0000	834.0000		
104	2650.0000	825.0000		
105	2610.0000	839.0000		i.
106	2580.0000	848.0000		
107	2540.0000	859.0090		
108	2610.0000	865.0000		
109	2695.0000	871.0000		
110	2800.0000	871.0000		
111	2730.0000	865,0000		\ \ \ \ \ \ \ \ \ \ \ \ \ \ \ \ \ \ \
	그는 눈이 반 없으면 하고 있다니다.		그리아 보는 생님, 그리고 있는 것도 되는 것은 그는 그리고 있는 그리고 있는 것은 그리고 그리고 있는 것은	

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113	2710.0000	848.0000
114	2819.0000	859.0000
115	2880.0000	848.0000
116	2730.0000	835.0000
117	2810.0000	825.0000
	2990.0000	827.0000
118		
119	2900.0000	835.0000
120	3630.0000	848.0000
121	2990.0000	.859.0000
122	2890.0000	865.0000
123	2970.0000	871.0000
124	3105,0000	871.0000
125	3030,0000	865,0000
		865.0000
126	3200.0000	
127	3280.0000	871.5000
128	3120.0000	859.0000
129	3200.0000	848.0000
130	3060.0000	835.5000
131	3115.0000	829.5000
132	3295,0000	831.0000
133	3210.0000	839.0000
134		848.0000
	3310.0000	
135	3290,0000	859.0000
136	3350.0000	865.0000
137	3410.0000	871.5000
138	3527.0000	871.5000
139	3450.0000	865.0000
140	3400.0000	859,0000
141	3500,0000	859.0000
142	3420,0000	848.0000
143	3370.0000	841.5000
144	3450.0000	834,5000
145	3615.0000	836.0000
146	3540.0000	844.0000
147	3600.0000	852.5000
148	3650.0000	859,0000
149	3560.0000	865.0000
150	3610,0000	871.5000
	3750.0000	871.5000
151		
152	3700.0000	865.0000
153	3900.0000	871.5000
154	3840.0000	865.0000
155	3800.0000	859.0000
156	3760.0000	852.5000
157	3700.0000	845.0000
158	3790.0000	835,5000
159	3980.0000	835,0000
	3840.0000	844.0000
160		
161	3910.0000	852.5000
162	3970.0000	859.0000
163	4010.0000	865.0000
164	4080.0000	871.5000
165	4020.0000	842.0000
166	4100.0000	852.5000
167	4170.0000	859,0000
168	4270.0000	871.5000
169	4210.0000	865.0000
170	4120.0000	834.5000
171	4210.0000	843.0000
172	4290.0000	852,5000
173	4330.0000	859.0000
174	4390.0000	865,0000
1.75	4440,0000	871,5000
176	4290.0000	834.5000
177	4380.0000	845.0000
n p l adist i Promision	UUVVI e VOC # .	TOURS & CHORES
and the state of the	<u>a sugue, se term a juda a ridag dilik</u>	

Regional Administration of the Control			
179 180	4500.0000		
181	4550.0000 4600.0000		
182	4770.0000		
183	4700.0000		
184	4640.0000		어른 시간의 어느는 지역의 이 어른 사이에서 이 사는 사는 이 뒤를 다 가를 맞춰
185	4600,0000		
186	4550.0000		
187	4440.0000	834.5000	
188	4580.0000	833.0000	
189	4610.0000		
190	4670.0000		
191	4720.0000		
192	4690.0000		
193	4790.0000		
194 195	4790.0000 4880.0000		
196	4820.0000		
197	4920.0000		
198	4850.0000		
199	4990.0000		
200	5150.0000	873.0000	
201	5100.0000	865.0000	
202	5030.0000		
203	4980.0000		
204	4890,0000		
205	4810.0000		
206 207	4980.0000 5030.0000		
208	5120.0000		
209	5200.0000		
210	5290.0000		
211.	5350,0000	840.0000	
212	5280,0000	849.5000	
213	5180.0000	849.5000	
214	5150.0000	855.0000	
215	5190,0000		
216	5240.0000	865.0000	
217 218	5300.0000 5480.0000	874.0000	
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220	5350.0000	858.5000	
221	5400.0000	849,5000	
222	5470.0000	830.0000	
. 223.	5500.0000	840.0000	
224	5540.0000	849.5000	
225	5460.0000	860.0000	
226	5600,0000	859.0000	
227	5610.0000	864.0000	
228 229	5490.0000	865.0000 875.5000-	
230	5680.0000	877.5000	
231	5830.0000	878.5000	
232	6000.0000	880.0000	
233	5800.0000	867.5000	
234	5760.1000	859.0000	
235	5710.0000	849.5000	
236	5680.0000	840.0000	
237	5630.0000	830.5000	
238	5800,0000	831.5000	and a first of the contract of
239 240	5820.0000 5870.0000	840.0000	
240	5900,0000	849 . 5000 859 . 0000	
242	5950.0000	870.0000	
243	6120.0000	881.5000	

246 6490,0000 870,0000 247 6010,0000 849,5000 249 5950,0000 832,5000 250 6050,0000 834,5000 251 6050,0000 844,5000 252 6170,0000 844,5000 253 6100,0000 844,5000 254 625,0000 849,5000 255 6100,0000 870,0000 256 6230,0000 870,0000 257 6280,0000 870,0000 258 6150,0000 870,0000 259 6100,0000 880,0000 250 6210,0000 870,0000 250 6210,0000 870,0000 250 6210,0000 870,0000 250 6300,0000 870,0000 250 6410,0000 870,0000 250 6410,0000 870,0000 250 6410,0000 870,0000 250 6410,0000 870,0000 250 6410,0000 870,0000 250 6410,0000 870,0000 250 6410,0000 870,0000 250 6410,0000 870,0000 250 6410,0000 870,0000 250 6410,0000 870,0000 250 6410,0000 870,0000 250 6410,0000 870,0000 250 6480,0000 889,0000 251 6580,0000 870,0000 257 6580,0000 870,0000 257 6580,0000 870,0000 257 6780,0000 882,0000 257 6780,0000 882,0000 258 6810,0000 870,0000 259 6790,0000 882,0000 259 6790,0000 882,0000 259 6790,0000 882,0000 259 6790,0000 882,0000 259 6790,0000 882,0000 259 6790,0000 882,0000 259 6700,0000 882,0000 259 7700,0000 882,0000 259 7700,0000 882,0000 259 7700,0000 882,0000 259 7700,0000 882,0000 259 7700,0000 882,0000 259 7000,0000 882,0000 250 7110,0000 882,0000 250 7110,0000 882,0000 250 7110,0000 882,0000 250 7110,0000 882,0000 250 7110,0000 882,0000 250 7110,0000 882,0000 250 7110,0000 882,0000 250 7110,0000 882,0000 250 7110,0000 882,0000 250 7110,0000 882,0000 250 7110,0000 882,0000 250 7110,0000 882,0000 250 7110,0000 882,0000 250 7110,0000 882,0000 250 7110,0000 882,0000 250 7110,0000 882,0000 250 7110,0000 882,0000 250 7110,0000 882,0000 250 7110,0000 882,0000		en plantaja da kanda kanta alkanga antaja kanda antala ja kanda ja kanda ja kanda ja kanda kanda kanda kanda d Antala kanda kanda kanda kanda antala kanda k	Toping graphic harmonic and an individual water in the last of the last contemplate of the property of the last of	and the state of t
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268 6300.0000 836.0000 269 6480.0000 837.5000 270 6510.0000 859.0000 271 6580.0000 870.0000 272 6580.0000 870.0000 273 6600.0000 882.0000 274 6620.0000 889.000 275 6780.0000 889.000 276 6770.0000 885.0000 278 6710.0000 870.0000 279 6690.0000 870.0000 280 6670.0000 839.000 281 6620.0000 839.000 282 6770.0000 849.5000 284 6810.0000 859.0000 285 6840.0000 870.0000 286 6880.0000 882.0000 287 6900.0000 882.0000 288 6900.0000 882.0000 289 7000.0000 870.0000 290 6920.0000 849.5000 293 4905.0000 849.5000 294 6890.0000 849.5000 295 <td< th=""><th>267</th><th>6340.0000</th><th>849,5000</th><th></th></td<>	267	6340.0000	849,5000	
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311	7320.00°n	870.0000	
312	7360,0000	882.0000	
313.	7390.0000	895.0000	
314	7410.0000	903.0000	
315	7530.0000	903.0000	
316	7510.0000	895.0000	
317	7490.0000	882.0000	
318	7450.0000	870.0000	
319	7420.0000	859.0000	
320	7400.0000	849.5000	
321	7390.0000	842.5000	
322	7510.0000	842.5000	
323	7520.0000	849.5000	
324	7550.0000	859.0000	
325	7580,0000	870.0000	
326	7600.0000	882.0000	
327	7620.0000	895.0000	
328	7650,0000	902.0000	
329	7770.0000	901.0000	
330	7740.0000	895.0000	
331	7710.0000	882.0000	
332	7690.0000	870.0000	
333	7660.0000	859.0000	
334	7630,0000	849,5000	
335	7610.0000	843.0000	
3 36	7710.0000	843.0000	
.337	7740.0000	849.5000	
338	7780.0000	859,0000	
339	7800.0000	870.0000	
340	7830.0000	882.0000	
341	7880.0000	895.0000	
342	7890.0000	900.5000	
343	7925.0000	900.5000	
344	7925.0000	895.0000	
345	7925.0000	882.0000	
346	7925,0000	870.0000	
347	7890.0000	859,0000	
348	7925.0000	860.0000	
349	7925.0000	850.0000	
350	7925.0000	843.0000	
351	7830.0000	843.0000	
352	7870,0000	849.5000	

ELEMENT INCIDENCE TABLE

ELEMENT	NODE	1 NODE	2 NODE	3 KX	KY	SLOPE
y-mark)	. 51	53	52	1.000	1.000	0.0
<mark>2</mark> .5	50		51	1.000	1.000	0.0
3	50	54	53	1.000	1.000	0.0
4	49	54	50	1.000	1.000	0.0
- 5	48	54	49	1.000	1.000	0.0
6	43	48	49	1.000	1.000	0.0
7	43	47	48	1.000	1.000	0.0
8	36	47	43	1.000	1.000	0.0
9	36	44	47	1.000	1.000	0.0
10	34	44	36	1.000	1.000	0.0
	41	50	51	1.000	1.000	0.0
12	41	42	50	1.000	1.000	0.0
13	40	42	41	1.000	1.000	0.0
14	39	42	40	1.000	1.000	0.0
15	42	49	50	1.000	1.000	0.0
16	38	49	42	1.000	1.000	0.0
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2000 - 1000 - 1000 - 1000 - 1000 - 1000 - 1000 - 1000 - 1000 - 1000 - 1000 - 1000 - 1000 - 1000 - 1000 - 1000 -		43	3.8	1.000	1.000	0.0	
18	3 7 36	43	37	1.000	1.000	0.0	
20		36	37	1.000	1.000	0.0	
21	33	34	36	1.000	1.000	0.0	
- 22	23	34	33	1.000	1.000	0.0	
23	38	42	39	1.000	1.000	0.0	
24	31	38	42	1.000	1.000	0.0	
25	3 4	37	38	1.000	1.000	0.0	
26	31	32	37	1.000	1.000	0.0	
27	18	19	27	1.000	1.000	0.0	
2.8	18	27	28	1.000	1.000	0.0	
29	5	de ferminale de fe	7	1.000	1.000	0.0	
30	7 7	Anny Anny Anny Anny Anny Anny Anny Anny	14	1.000	1.000 1.000	0.0	
31	14	25	15 15	1.000 1.000	1.000	0.0 0.0	
32 33	15	25 25	20	1.000	1.000	0.0	
34	20	25 25	26	1.000	1.000	0.0	
35	20	26	27	1.000	1.000	0.0	
36	2.6	29	27	1.000	1.000	0.0	
37	27	29	28	1,000	1.000	0.0	
38		29		1.000	1.000	0.0	
39	. 1	13	14	1.000	1.000	0.0	
40	13	22	14	1.000	1.000	0.0	
41		22	25	1.000	1.000	0.0	
42	22	32	25	1.000	1.000	0.0	
43	2.5	32	26	1.000	1.000	0.0	
44	26	32	31	1.000	1.000	0.0	
45 46	26 29	31 31	29 39	1.000 1.000	1.000 1.000	0.0 0.0	
47	29	39	30	1.000	1.000	0.0	
48	30	39	40	1.000	1.000	0.0	
49	13	23	22	1.000	1.000	0.0	
50	22	23	33	1.000	1.000	0.0	
51	22	33	32	1.000	1.000	0.0	
52	-32	33	37	1.000	1.000	0.0	
53	19	20	27	1.000	1.000	0.0	
54	1.6	20	19	1.000	1.000	0.0	
55	1.5	20	16	1.000	1.000	Q•Q	
56	- 6	15	16	1.000	1.000	0.0	
57	6	7	1.5	1.000	1.000	0.0	
58 59	3	5	6 . 7	1.000	1.000 1.000	0.0	
60	3	Ly.	5	1.000	1.000	0.0	
61	. 17	19	18	1.000	1.000	0.0	
62	16	19	17	1.000	1.000	0.0	
63	8	16	17	1,000	1.000	0.0	
64	6	16	8	1.000	1.000	0.0	
65	2	6	8	1.000	1.000	0.0	
-66	2	3	_{.,} 6	1.000	1.000	0.0	
67	- Journel	3	2	1.000	1.000	0.0	
6.8		4	3	1.000	1.000	0.0	
69	287	289	288	1.000	1.000	0.0 0.0	
70	275 275	287 276	288 287	1,000 1,000	1.000	0.0	
72	274	276	275	1.000	1.000	0.0	
73	273	276	274	1.000	1.000	0.0	
74	263	273	274	1.000	1.000	0.0	
75	263	264	273	1.000	1.000	0.0	
76	261	264	263	1.000	1.000	0.0	
77	261	262	264	1.000	1.000	0.0	
78	259	262	260	1.000	1.000	0.0	
79	243	259	260	1.000	1.000	0.0	
80	260	262	261	1.000	1.000	O•0	
81	232	244	243	1.000	1.000	0.0	
82			244	1.000	1.000	0.0	
	<u>الدي در آنځ از د از خام په ماه د د د د</u>	<u> </u>	<u>in a la constanta de la const</u>	<u>an an analysis see septimization of the art</u>	<u> Parlina ya kata ji Marini Baran Ba</u>		<u>i i kanala katapat Ma</u>

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85	230	233	231	1.000	1.000	0.0	
86	227	233	gregory 23 0+	1.000	1.000	0.0	
87	227	230	229	1.000	1.000	0.0	
8.8	227	229.	2284	1.000	1.000	0.0	
89	225	227	228	1.000	1.000	0.0	
90	218	228	229	1.000 1.000	1.000 1.000	ି. ୦ ତ - ୦	
91 92	218 219	219 225	228 228	1.000	1.000	0.0	
93	219	220	225	1.000	1.000	0.0	
94		219	218	1.000	1.000	0.0	
95	216	219	217	1.000	1.000	0.0	
96	216	220	219	1.000	1.000	0.0	
97	214	220	216	1.000	1.000	0.0	
98		216	217	1.000	1.000	0.0	
99	200	201	216	1.000	1.000	0.0 0.0	
100	199 197	201 201	200 1 99	1.000 1.000	1.000 1.000	0.0	
101 102	197	199	198	1.000	1.000	0.0	
103	196	197	198	1.000	1.000	0.0	
104		196	198	1,000	1.000	0.0	
105	182	183	196	1.000	1.000	0.0	
106	181	183	182	1.000	1.000	0.0	
107	180	183	181	1.000	1.000	0.0	
108		180	181	1.000	1.000	0.0	
109		180	175	1.000	1.000	0.0	
110		215	216	1.000	1.000	0.0	
111	201	202	215 215	1.000 1.000	1.000 1.000	0.0 0.0	
112 113	202	214	214	1.000	1.000	0.0	
114		203	195	1.000	1.000	0.0	
115		203	202	1.000	1.000	0.0	
-116		202	197	1.000	1.000	0.0	
117		202	201	1.000	1.000	0.0	
118		197	196	1.000	1.000	0.0	
119		195	196	1.000	1.000	0.0	
120		195	194	1.000	1.000	0.0	
121		193	194	1.000	1.000 1.000	0.0	
122		186 192	192 194	1.000 1.000	1.000	0.0	
123 124		184	194	1.000	1.000		
125			196	1.000		0.0	
126		186	178	1.000	1.000	0.0	
127			185	1.000	1.000	0.0	
128	178	185	179	1.000		0.0	
129			184	1.000		0.0	
130			180			0.0	
131			183	1.000		0.0 0.0	
132		177 177	172 178	1.000 1.000		0.0	
133 134		178					
135			179				
136			174	1.000		0.0	
137			180	1.000		0.9	
138		174	175	1.000	1.000	0.0	
139	168		174			0.0	
140			168	1.000		0.0	
141				1.000		0.0	
142			164			0.0	
143			163	1.000 1.000		0.0	
144 145							
146		172	173	1.000		0.0	
147			167	1.000		0.0	
148		171	172	1.000	1.000	0.0	
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218		122	123	1.000	1.000	0.0	
219	110		122	1.000	1.000	0.0	
220		111	110	1.000	1.000	0.0	
221			109	1.000	1.000	0.0	
		114	122	1.000	1.000	0.0	
222 223		112	114	1.000	1.000	0.0	
225 224		112		1.000	1.000	0.0	
225		113	114	1.000	1.000	0.0	
		113	112	1.000	1.000	0.0	
226					1.000		
227		113	106	1.000		0.0	
228		112	108	1.000	1.000	0.0	
229		112	107	1.000	1.000	0.0	
230	99	106	107	1.000	1.000	0.0	
231	99	100	106	1.000	1.000	0.0	
232	100	105	106	1.000	1.000	0.0	
233	100	103	105	1.000	1.000	0.0	
234		101	103	1.000	1.000	0.0	
235	90	101	91	1.000	1.000	0.0	
236		101	100	1.000	1.000	0.0	
237	91	100	99	1.000	1.000	0.0	
238		108	108	1.000	1.000	0.0	
239		97	108	1.000	1.000	0.0	
240		107	108	1.000	1.000	0.0	
241	97	98	107	1.000	1.000	0.0	
242	98	99	107	1.000	1,000	0.0	
243	92	99	98	1.000	1.000	0.0	
244	91	99	92	1,000	1.000	0.0	
245	8.6	91	92	1.000	1.000	0.0	
246	86	90	91	1.000	1.000	0.0	
247		87	90	1.000	1.000	0.0	
248	95	97	96	1.000	1.000	0.0	
249	94	97	95	1.000	1.000	0.0	
:250	82	94	95	1.000	1.000	0.0	
251	82	83	94	1.000	1.000	0.0	
252	83	93	94	1.000	1.000	0.0	
253	93	98	94	1.000	1.000	0.0	
254	94	98	97	1.000	1.000	0.0	
255	92	98	93	1.000	1.000	0.0	
2,56	85	92	93	1.000	1.000	0.0	
257	85	86	92	1.000	1.000	0.0	
258	77	86	85	1.000	1.000	0.0	
259	77	87	86	1.000	1.000	0.0	
260	75	- 4.8.7	77	1.000	1.000	0.0	
261	83	84	. 93	1.000	1.000	0.0	
262	84	85	93	1,000	1.000	0.0	
263	78	85	84	1.000	1.000	0.0	
264	7.7	85	78	1.000	1.000	0.0	
265	72	77	78	1.000	1.000	0.0	
266	72	75 sa		1.000	1.000	O. • O -	
267		73	75	1.000	1.000	0.0	
268		83	82	1.000	1.000	0.0	
269		83	81	1.000	1.000	0.0	
270		80	81	1.000	1.000	0.0	
271		69	80	1.000	1.000	0.0	
272		84	83	1.000	1.000	0.0	
273		84	80	1.000	1.000	0.0	
274		84	79	1.000	1.000	0.0	
275		78	79	1.000		0.0	
276		72	78	1.000	1.000	0.0	
277		72	71	1,000	1.000	0.0	
278		73	72	1.000	1.000	0.0	
279		73	63	1.000	1.000	0.0	
280		79	80	1.000	1.000	0.0	
		v dje populav vojit italije e					
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232	70	71	79	1.000	1.000	0.0
283	64	71	7 ()	1.000	1,000	0.0
284	63	71	. 64	1.000	1.000	0.0
285	62	63	64	1.000	1.000	0.0
286	61	63	62	1.000	1.000	0.0
287	57	61	62			
				1.000	1.000	0.0
288	67	6.9	68	1.000	1.000	0.0
289	66	69	67	1.000	1.000	0.0
290	66	70	69	1.000	1.000	0.0
291	65	70	66	1.000	1.000	0.0
292	64	70	65	1.000	1.000	0.0
293	55	64	65	1.000	1.000	0.0
294	55	6 2	64	1,000	1.000	0.0
295	5.5	56	62	1.000	1.000	0.0
296	56	57	62 = 7	1.000	1.000	0.0
297	46	57	56	1.000	1.000	0.0
298	52	66	67	1.000	1.000	0.0
299	52	53	66	1.000	1.000	0.0
300	5.3	65	66	1.000	1.000	0.0
301	53	54	65	1.000	1.000	0.0
302	54	55	65	1.000	1.000	0.0
303	48	55	54	1.000	1.000	0.0
	48					
304		56 57	55	1.000	1.000	0.0
305	47	56	48	1.000	1.000	0.0
306	46	56	47	1.000	1.000	0.0
307	44	46	47	1.000	1.000	0.0
308	Δį.	9	5	1.000	1.000	0.0
309	5	9	10	1.000	1.000	0.0
310	5	10	11	1.000	1.000	0.0
311	Power Company	12	11	1.000	1.000	0.0
312	12	21	**************************************			
				1.000	1.000	0.0
313	and I	12	13	1.000	1.000	្.0
314	13	21	23	1.000	1.000	0.0
315	21	24	23	1.000	1.000	0.0
316	2.3	24	34	1.000	1.000	0.0
317	24	35	34	1.000	1.000	0.0
318	34	35	44	1.000	1.000	0.0
319	35	45	L+ L2	1.000	1.000	0.0
320		45	46	1.000	1.000	0.0
	45	5 8				
321			46	1.000	1.000	0.0
322	46	5.8	57	1.000	1.000	0.0
323	57	5.8	59	1.000	1.000	0.0
324	57	59	61	1.000	1.000	0.0
325	59	60	61	1.000	1.000	0.0
. 326	60	73	61	1.000	1.000	0.0
327	60	74	73	1.000	1.000	0.0
- 328	73	74	75	1.000	1.000	0.0
329	74	76	75	1.000	1.000	0.0
330	75	76	87	1.000		0.0
331	76	38	87	1.000	1.000	0.0
332	8.7	88	90	1.000	1.000	0.0
333	88	89	90	1.000	1.000	0.0
334	89	101	90	1.000	1.000	0.0
335	89	102	101	1.000	1.000	0.0
336	101	102	103	1.000	1.000	0.0
337	102	104	103	1.000	1.000	0.0
338	103	104	116	1.000	1.000	0.0
339	104	117				
	the state of the s		116	1.000	1.000	0.0
340	116	117	119	1.000	1.000	0.0
341	- 117	118	119	1.000	1.000	0.0
342	118	130	119	1.000	1.000	0.0
343	118	1.31	130	1.000	1.000	0.0
344	130	131	133	1.000	1.000	0.0
345	131	132	133	1.000	1.000	0.0
346	132	143	133	1.000	1.000	0.0
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348	143	144	146	1.000	1.000	0.0	
349	144	145	146	1.000	1.000	0.0	
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350-	145	157	146	1.000			
351	145	158	157	1.000	1.000	0.0	
352	157	- 158	- 1 - 1160	1.000	1.000	0.0	
353	158	159	160	1.000	1.000	0.0	그 그 그 그 이 그 가 가 가 되는 그래요?
354	159	165	160	1.000	1.000	0.0	
				1,000	1.000	0.0	
355	159	170	165				
356	165	170	171	1.000	1.000	0.0	
357	170	176	171	1.000	1.000	0.0	
358	171	176	177	1.000	1.000	0.0	
				1.000	1.000	0.0	
359	176	177	187				
360	177	187	186	1.000	1.000	0.0	
361	186	187	189	1.000	1.000	0.0	E to the second of the second
362	187	188	189	1.000	1.000	0.0	
		190	189	1.000	1.000	0.0	
363	188						
364	189	190	191	1,000	1.000	0.0	
365	186	189	192	1.000	1.000	0.0	
366	189	191	192	1.000	1.000	0.0	
367	190	205	191	1.000	1.000	0.0	
							a de la companya de
,3,6.8	191	205	204	1.000	1.000	0.0	
369	191	204	193	1.000	1.000	0.0	
370	204	205	206	1.000	1.000	0.0	
371	204	206	207	1.000	1.000	0.0	
		204	203	1.000	1.000	0.0	
372	193						
373	203	204	267	1.000	1.000	0.0	
374	203	207	214	1.000	1.000	0.0	
375	206	208	207	1.000	1.000	0.0	
376	207	208	209	1.000	1.000	0.0	
					1.000	0.0	
- 377	207	209	214	1.000			
378	208	210	209	1.000	1.000	0.0	
379	209	210	211	1.000	1,000	0.0	
380	210	222	211	1.000	1.000	0.0	
381	211	222	223	1.000	1.000	0.0	
						0.0	
-382	222	237	223	1.000	1.000		
-383	223	237	236	1.000	1.000	0.0	
384	236	237	238	1.000	1.000	0.0	
385	236	238	239	1.000	1.000	0.0	
	238	249	239		1.000	0.0	
386							
387	239	249	248		1.000	0.0	
388	248	-249	250	1.000	1.000		
389	248	250	251	1.000	1.000	0.0	
390	209	212	213		1.000	0.0	
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391	212						
392		-220			1.000		그가 그 그는 그 이미를 모으면 바다왔다.
393	209	211	212	1.000	1.000	0.0	
394	211	221	212	1.000	1.000	0.0	
395	211	223			1.000	0.0	
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396	221	223					
397	223	236	224		1.000	0.0	i de la companya de
39.8		236	235	1.000	1.000	0.0	하는 아이지도 아니아의 아이들의 사람들 없다.
399		236			1.000	0.0	그 그 그 아내는 아이를 내고 있다.
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400	235						
401	239				1.000	0.0	
402	240	2.48	247		1.000	0.0	
403	247	248	251	1.000	1.000	0.0	
404	247				1.000	0.0	
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405	250						
406					1.000	0.0	
407	252	268	254	1.000	1.000	0.0	- 1 - 1 - 1 - 1 - 1 - 1 - 1 - 1 - 1 - 1
408	254		267	1.000	1.000	0.0	
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411	269		270		1.000	0.0	
-412	270	281	2.80	1.000	1.000	0.0	
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4]			93	294		295		1.000		1.000	0.0
4			93	295		296		1.000		1.000	0.0
4			95	307		296		1.000		1.000	0.0
42			96	30 7		306		1.000		1.000	0.0
42			16	307		308		1.000		1.000	0.0
42			06	308		309		1.000		1.000	0.0
4.			09	321		320		1.000		1.000	0.0
42			0.8	321		309		1.000		1.000	0.0
47			20	321		322		1.000		1.000	0.0
42			20	322		323		1.000		1.000	0.0
42	27		22	335		323		1.000		1.000	0.0
42	2.8	3	23	335		334		1.000		1.000	0.0
42	29	31	34	335		336		1.000		1.000	0.0
43	3 (C	3.	34	336		337		1.000		1.000	0.0
43	3]	3	36	351		337		1.000		1.000	0.0
43		3:	3.7	351		352		1.000		1.000	0.0
43		3 !	50	352		351		1.000		1.000	0.0
43		3.	49	352		350		1.000		1.000	0.0
43		3	47	352		349		1,000		1.000	0.0
.43			47	349		348		1.000		1.000	0.0
4 =			3.8	352		347		1.000		1.000	0.0
43			3 7	352		338		1,000		1.000	0.0
43			3 3	337		338		1.000		1.000	0.0
44			3 3	334		337		1.000		1.000	0.0
44			24	334		333		1.000		1.000	0.0
44			23	334		324		1.000		1.000	0.0
44			19	323		324		1.000		1.000	0.0
44			19	320		323		1.000		1.000	0.0
- 44 - 44			10	320		319		1.000		1.000	0.0
44			9	320		310		1.000		1.000	0.0
			15	309		310		1.000		1.000	0.0
44)5 ?7	306		309		1.000		1.000	0.0
45			96 -	306 306		305 297		1.000		1.000	0.0
45			92	296		297		1.000		1.000	0.0 0.0
45			92	293		296		1.000		1.000	0.0
45			34	293		292		1.000		1.000	O.O
45			33	293		284		1.000		1.000	0.0
45			79	283		284		1.000		1.000	0.0
45			79	280		283		1.000		1.000	0.0
45		2		280		279		1.000		1.000	0.0
. 45	8		70.	280		271		1,000		1.000	0.0
45	9		56	270		271		1.000		1.000	0.0
46	0		6	267		270		1.000		1.000	0.0
46	1	23	56	267		266		1.000		1.000	0.0
46	2	2.5	54	267		256		1.000		1.000	0.0
46	3	25	54	256		255		1.000		1.000	0.0
46	4		33.,	254		255		1.000		1.000	
46		24	+6	253		255		1.000		1.000	0.0
46		24	1-6	247		253		1.000		1.000	0.0
46		24		247		246		1.000		1.000	0.0
46		2 1		221		220		1.000		1.000	0.0
46		2.2		221		225		1.000		1.000	0.0
47		22		224		225		1,000		1.000	0.0
47		22		227		225		1.000		1.000	0.0
47				235		226		1.000		1.000	0.0
47		- 22		234		227		1.000		1.000	0.0
47		2.2		235		234		1.000		1.000	0.0
47		2.2		234		233		1.000		1.000	0.0
47		23		235		240		1.000		1.000	0.0
47		23		240		241		1.000		1.000	0.0
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498	496	264	265	272	1.000	1.000	0.0	
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546	317	318	3.25	1.000	1.000	0.0
547	317	325	326	1.000	1.000	0.0
548	. 316	317	326	1.000	1.000	0.0
549	316	326	327	1.000	1.000	0.0
550	315	316	327	1.000	1.000	0.0
551	315	327	328	1.000	1.000	0.0
552	327	330	328	1.000	1.000	0.0
553	328	330	329	1.000	1.000	0.0
554	327	331	330	1.000	1.000	0.0
555	326	331	327	1.000	1.000	0.0
556	326	332	331	1.000	1.000	0.0
557	325	332	326	1.060	1.000	0.0
558	325	333	332	1.000	1.000	0.0
559	324	333	325	1.000	1.000	0.0
560	332	333	338	1.000	1.000	0.0
561	332	338	339	1.000	1.000	0.0
562	338	347	339	1.000	1.000	0.0
563	339	347	346	1.000	1.000	0.0
564	346	347	348	1.000	1.000	0.0
565	331	332	339	1.000	1.000	0.0
56.6	- 331	339	340	1.000	1.000	C.O
567	330	331	340	1.000	1.000	0.0
568	330	340	341	1.000	1.000	0.0
569	329	330	341	1.000	1.000	0.0
570	329	341	342	1.000	1.000	0.0
571	341	344	342	1.000	1.000	0.0
572	342	344	343	1.000	L.000 .	0.0
573	341	345	344	1.000	1.000	0.0
574	340	345	341	1.000	1.000	0.0
575	340	346	345	1.000	1.000	0.0
576	339	346	340	1.000	1.000	0.0

PRESCRIBED NODAL HEADS

NODE PRESCRIBED HEAD

. 1	839.0000
2	849.0000
- 8	860,0000
17	865,0000
18	668.0000
28 .	869.0000
30	869,0000
40	869.0000
41	869.0000
51	869,5000
52	270.0000
67	870.0000
68	870.5000
81	870.5000
82	870.5000
95	871.0000
96	871.0000
L0.9	871,0000
10	871.0000
123	871,0000
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314	900									
315	900									
328	900									
329	900									
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FINITE ELEMENT SEEPAGE PROGRAM BESWICK PINAWA

RESULTS

NODAL HEADS

NODE HEAD (FT)

- 7 838,9998
- 2 848,9995
- 3 851.0352
- 4 840.5310
- 5 852.9375
- 6 859.7358
- 7 860,3801
- 8 859,9998 9
- 841.3079 10 858.8625
- 863.7202
- 11 12
- 866.5659
- 13 868.3914
- 14 867.1941
- 15 866.0562
- 16 864.4456
- 17 864,9998
- 18 867,9995
- 19 866.8091 20 868.3201
- 21
- 869.0432 22 869.3037
- 23 869.4353
- 24 869,4233 25 869,0679
- 26 869.2236
- 27 868.9602
- 28 868,9995
- 29 869.0750
- 30 868.9995
- 31 869,0200
- 32 869.2710
- 33 869,2639
- 34 869,1909
- 35 869:0786
- 36 869.0361
- 37 869.0376
- 38 868.9929 39 868.9902
- 40 868,9995
- 41 868,9995
- 42 868.9946
- 43 869.0417
- 44 869.0420
- 45 869.0916
- 46 869.2551
- 47 869.1563
- 48 869.3413
- 49 869.1831 50 869.3196
- 51 869,4998
- 52 869.9998

869.8462

54 869,7170

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869.5308
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       869.7058
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       869.7656
       870.0750
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       870.0295
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        869,9063
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       870.1614
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347	900.0232
348	900.0151
349	900.0181
350	900.0205
351	900,0374
352	900.0332

```
THIS PROGRAM CALCULATES TOTAL HEAD AT ALL NODES IN A REGION
    WHEN GIVEN HEADS AT SPECIFIED BOUNDARY NODES
    REAL NL(360), KX(576), KY(576)
    INTEGER ELEM(576)
    COMMON M, CJ (360, 2), IE (576, 3), L1, L2, L3, T, ST (3), KL, KX, KY, S (576)
    COMMON JRD.JWT
    DIMENSION LIST(360), A(360), ISR(576), HDG(20)
    DIMENSION STORE(6000), NODE(360), NEIN(360), NI(360,8)
    DIMENSION G(3,6),DISPL(6),STRESS(576)
    EQUIVALENCE (NE, NN)
    JR D=1
    JWT=3
    IPAGE=1
    LINE=0
    READ JOB TITLE
510 FORMAT([4,19A4)
                         \{HDG(I), I=1, 19\}
  I READ(JRD,510) IDENT,
    IF(IDENT.EQ.O) GO TO 9999
564 FORMAT("1",100X, "PAGE ", [4)
    WRITE(JWT, 564) IPAGE
    IPAGE=IPAGE+1
                 *NO.*,15,1X,19A4)
512 FORMAT(101,
    WRITE(JWT, 512) IDENT, (HDGII), I=1, 19)
    WRITE(JWT, 600)
600 FORMAT(/* TWO DIMENSIONAL SEEPAGE PROBLEM*//* INPUT DATA*//)
    READ NUMBER OF ELEMENTS, NUMBER NODES, NUMBER SPECIFIED NODAL
    HEADS, THICKNESS
    ALL INPUT PRINTED OUT FOR CHECKING PURPOSES
    READ(JRD, 500) NEL, NE, NBD, T
500 FORMAT(315,F10.3)
    IF(T_{\bullet}EQ_{\bullet}O_{\bullet}) T=1.
    WRITE(JWT.601) NEL, NN, NBD, T
601 FORMAT(14, * ELEMENTS 14, * NODES 14, * SPECIFIED NODAL HEADS 1// EL
   1EMENT THICKNESS = 'F8.3//)
    READ COORDINATES OF NODES
501 FORMAT(3(15,2F10.3))
    READ(JRD, 501) (NODE(J), (CJ(J,I),I=1,2),J=1,NN)
    LINE=5
    WRITE (JWT , 564) IPAGE
    IPAGE=IPAGE*1
    WRITE(JWT:602)
6020FCRMAT(1HO, NODAL COORDINATES (FT) 1/1 NODE X COORD
                                                                    Y COD
    DO 99 J=1.NN
    WRITE(JWT, 694)NODE(J), CJ(J.1), CJ(J.2)
    LINE=LINE+1
    IF(LINE.LT.60)GD TO 99
    WRITE(JWT, 564) IPAGE
    IPAGE=IPAGE+1
    LINE=5
 99 CONTINUE
    LINE=4
    ZERO CONSTANT VECTOR
```

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25/08/70
                                                                TIME
360N-F0-479 3-1
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191 DO 39 J=1,360
 39 NL(J)=0.
    GENERATE NODAL INCIDENCE TABLE
    WRITE(JWT, 564) IPAGE
    IPAGE=IPAGE+1
    WRITE(JWT,686)
6860FORMAT(1HO, ELEMENT INCIDENCE TABLE 1// ELEMENT NODE 1 NODE
                                  ΚY
         NUDE 3
                        ΚX
                                      SLOPE'/)
   1 2
    SS=0.
    DO 17 M=1, NEL
505 FORMAT(15,315,3F10.3)
    READ(JRD, 505) ELEM(M), (IE(M, I), I=1,3), XK, YK, SL
    IF(XK.NE.C.)XX=XK
    IF(YK,NE,C.)YY=YK
    IF(SL.NE.O.) SS=SL
    S(M)=SS
    KX(M) = XX
    KY\{M\}=YY
    WRITE(JWT,687)ELEM(M),(IE(M,I),I=1,3),XX,YY,S(M)
687 FORMAT(419,3F10.3)
    LINE=LINE+1
    IF(LINE.LT.60) GO TO 2002
    WRITE(JWT,564) IPAGE
    IPAGE=IPAGE+1
    LINE=4
2002-1F(JE1H,1)-EQ-IE(M,2)-GO-TO-2000-
    1F(1E(M,1),E0,1E(M,3))-G0-T0-2000
    IF(IE(%,2).EQ.IE(M,3)) GO TO 2000
                                                      elements checked
                                                      duplication of values (debug)
    GO TO 17
2001 FORMAT(*O', CHECK-ELEMENT(*, 13, 1) INCIDENCES*)
2000-WRITE(JUIT, 2001)-ELEM(M)
    STOP 2000
 17 CONTINUE
    LINE = 5
    DO 18 J=1,360
    A(J)=0.0
    NEIN(J)=0
    DO 18 M=1.8
 18 NI(J,M)=0
    DO 19 M=1,NEL
    DO 19 K=1,3
    J=IE(M,K)
    NEIN(J)=NEIN(J)+1
    L=NEIN(J)
    NI(J,L)=H
 19 CONTINUE
694 FORMAT(15,6F12.4)
    READ ANY PRESCRIBED NODAL HEADS'S
    DO 376 J=1.576
376 ISR(J)=0
    WRITE (JWT . 564) IPAGE
    IPAGE = IPAGE + 1
```

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25/08/70
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360N-F0-479 3-1
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596 FORMAT(*O', *PRESCRIBED NODAL HEADS*// NODE PRES. HEAD*//)
    WRITE(JWT, 596)
    DO 574 I=1,NBD
    READ(JRD, 694)J, NL(J)
    ISR(J)=1
    WRITE(JWT,694)J; NL(J)
    LINE = LINE + 1
    IF(LINE.LT.60)GO TO 574
    WRITE (JWT, 564) IPAGE
    WRITE (JWT,596)
    IPAGE = IPAGE + 1
574 CONTINUE
   LINE=11
    GENERATION AND ELIMINATION OF NODAL EQUILIBRIUM EQUATIONS
    GENERATE I THE ROW OF STIFFNESS MATRIX AND STORE IN A TEMPORARILY
172 LIST(1)=1
   LINE=0
    DO 100 I=1,NE
    NON ZERO BAND OF ROW I IN STIFFNESS IS FROM KL TO KH. KL = LOWEST
    NODE NO FOR NODES INCIDENT ON ELEMENTS INCIDENT ON NODE I, KH = HIGHEST
    KH = I
    IM=NEIN(I)-
    DO 5 J=1, IM
    M=NI(I,J)
    DO 5 K=1,3
    JF=IF(M,K)
    IF(JF-NE)10,10,5
 10 IF (JF-KH) 14, 14, 12
 12 KH=JF
    GO TO 5
 14 IF(JF-KL)16,5,5
 16 KL=JF
  5 CONTINUE
    ZERO ALL A MATRICES IN NON - ZERO BAND
    K = KH - KL + 1
    DO 29 J=1,K
 29 A(J)=0.
    INSERT STIFFNESS MATRICES INTO NON ZERO BAND
    DO 200 J=1, IM
    M=NI(I,J)
    DO 20 L=1,3
    IF(I-IE(M,L))20,22,20
 22 KI=IE(M,L)
    L1=L
    L2=L+1-3*(L/3)
    KJ=IE(M,L2)
    L3=L+2-3*(L/2)
    KK = IE(M, L3)
 20 CONTINUE
    CALL STIFF
    NNI, NJ AND NK ARE NODES 1, J AND K FOR ELEMENT M - (POSITION IN ROW I
    RELATIVE TO KL = 1
```

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 360N-F0-479 3-1
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     NNI = I - KL + 1
     INSERT STIFFNESS MATRIX I
     A(NNI)=A(NNI)+ST(1)
     TEST WHETHER MODE NJ IS A FIXED SUPPORT
     IF(KJ-NE)47,47,6
     INSERT STIFFNESS MATRIX J
  47 NJ=KJ-KL+1
     A(NJ) = A(NJ) + ST(2)
     TEST WHETHER NODE NK IS FIXED SUPPORT
   6 IF(KK-NE)8,8,200
   8 \text{ NK} = \text{KK} - \text{KL} + 1
     INSERT STIFFNESS MATRIX K
     A(NK) = A(NK) + ST(3)
 200 CONTINUE
     MODIFY EQUATION IF SPECIFIED HEAD AT NODE I
     IF(ISR(I).EQ.0) GO TO 497
     A(NNI)=10.**25
     NL(I)=NL(I)*10.**25
 497 LINC=KH-I
     FOR FIRST EQUATION, BYPASS ELIMINATION
     IF(I-KL)60,60,50
     PERFORM ELIMINATION FOR ROW. I TO ZERU BELOW MAIN DIAGONAL
  50 KU=I-1
     DO 256 K=KL, KU
     IK = PIVOTAL COLUMN RELATIVE TO KL = 1
     IK=K+1-KL
     IM=LIST(K+1)-LIST(K)
     IJ=K+IM-I-LINC
     IF NON ZERO BAND FOR PIVOTAL EQ ENDS TO RIGHT OF THAT FOR EQ I: EXTEND
     FOR EQ I
      IF(IJ.LE.O.) GO TO 82
     KK=LINC+I-KL+2
     LINC=LINC+IJ
     LL=IJ+KK-1
     DO 83 L=KK,LL
  83 A(L)=0.
  82 IF(IM.LE.O.) GO TO 256
3000 FORMAT(101, LMAIN-LOOP-I=1, 14, 1-IN-1, 14, 1-****** , 1-LINC=1, 14/)
     WRITEIJHI-3000)-I. IM-LINC . Print check for main loop (debug)
     DO 55 J=1, IM
     IJ=IK+J
     KJ = LIST(K) + J - 1
     IF(ABS(STORE(KJ)).LE.10.**(-10)) STORE(KJ)=0.
     A(IJ)=A(IJ)-A(IK)*STORE(KJ)
    IF(ABS(A(IJ)).EQ.-0.) GO TO 55
    -1F(ABS(A(IJ)),GT.10.04(-50))-G0-10-55
30010FORMAT(1 13 IN 100P J=1,14,1 IJ=1,14,14,1 A(IJ)=1,E12,5
   1,1-4(1K)=1,E12,5,1-STORE(KJ)=1E12,5)
                                                   Print check for location of
                                                     under flow errors (debug)
    -IF(LINE . EQ. L000) GO TO 55
    WRITE(JWT, 3001) J, IJ, KJ, A(IJ), A(IK), STORE(KJ)...
    TIME=FINE+1
  55 CONTINUE
 256 NL(I)=NL(I)-A(IK)*NL(K)
     NORMALIZE ROW I. JJ = PIVCTAL ELEMENT RELATIVE TO KL = 1
```

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 60 JJ=I+I-KL
     IJ=JJ
    LIST(I+1)=LINC+LISī(I)
     IF(LINC)75,75,65
  65 DO 70 J=1,LINC
     IJ=IJ+1
     IK=LIST(I)+J-1
  70 STORE(IK)=A(IJ)/A(JJ)
     NORMALIZE HEAD AT NODE I
 75 NL(I)=NL(I)/A(JJ)
 100 CONTINUE
     START BACK SUBSTITUTION
     N2=NE-1
     IF(N2)210,210,220
 220 DO 250 K=1,N2
     I=NE-K
     KU=LIST(I+1)-LIST(I)
     SUM=0.
     DO 240 J=1,KU
     IK=LIST(I)+J-1
     IJ=I+J
240 SUM=SUM+STORE(IK)*NL(IJ)
 250 NL(I)=NL(I)-SUM
     WRITE NODAL HEADS
 210 WRITE (JWT, 564) IPAGE
     IPAGE=IPAGE+1
     WRITE(JWT, 560)(HDG(I), I=1,20)
 560 FORMAT(*O'1X2CA4,//' RESULTS'// NODAL HEADS*// NODE HEAD (F
    1T)'/)
     DO 397 L=1,NN
 562 FORMAT(16,F12.4)
     WRITE(JWT, 562)L, NL(L)
     LINE =LINE + 1
     IF(LINE.LT.60) GO TO 397
     WRITE(JWT, 564) IPAGE
     IPAGE=IPAGE+1
 561 FORMAT('0', RESULTS CONT. '//' NODE HEAD (FT)'/)
     WRITE (JWT, 561)
SONTINUE - CALL PLOTER (NL) - subroutine PLOTER may be introduced at this noist
                                        introduced at this point
563 FORMAT (//1x, *END OF RUN NO. *, 14)
     WRITE(JWT, 563) IDENT
     LINE=0
     IPAGE=1
     GO TO 1
9999 STOP 1000
     END
```

```
25/08/70
                                                                 TIME
                                                                          13.16.21
                                            DATE
360N-F0-479 3-1
                         STIFF
    SUBROUTINE STIFF
    THIS SUBROUTINE DEVELOPS STIFFNESS MATRICES II, IJ AND IK FOR
    CONSTANT STRAIN TRIANGLE I-J-K
    REAL NL[360], KX[576], KY[576]
    CCMMON M, CJ(360, 2), IE(576, 3), L1, L2, L3, T, ST(3), KL, KX, KY, S(576)
    COMMON JRD.JWT
    DIMENSION A(3), B(3)
 15-FORMAT(401,1CHECK-NOBAL-INCIDENCES$$$-I=1,14,15=1,14,1K=1,14)
    SL=ABS(S(M)/57.2958)
    COSA=COS(SL)
    SINA=SIN(SL)
    IF(S(M).LE.O.) SINA=-SINA
    I=IE(M,1)
    J=IE(M,2)
    K=IE(M,3)
    A(1) = (CJ(K,1) - CJ(J,1)) *COSA + (CJ(K,2) - CJ(J,2)) *SINA
    A(2) = (CJ(I,1) + CJ(K,1)) * COSA * (CJ(I,2) + CJ(K,2)) * SINA
    A(3) = (CJ(J,1) - CJ(I,1)) * COSA + (CJ(J,2) - CJ(I,2)) * SINA
    B(1) = (CJ(J,2) - CJ(K,2)) * COSA - (CJ(J,1) - CJ(K,1)) * SINA
    B(2)=[CJ(K,2)-CJ(I,2))*COSA-!CJ(K,1)-CJ(I,1))*SINA
    B(3)=(CJ(1,2)-CJ(J,2))*COSA-(CJ(1,1)-CJ(J,1))*SINA_
    DET=CJ(J,1)*CJ(K,2)-CJ(K,1)*CJ(J,2)-CJ(I,1)*CJ(K,2)+CJ(K,1)*CJ(I,2
   1)+CJ(I,1)*CJ(J,2)-CJ(J,1)*CJ(I,2)
    1F (DET NE 0.0) - GC-TO-16
   WRITE (JWT . 15) I J J K
    -0006-90-F2
-16-ST(1)=(KY(M)*A(L1)*A(L1)*KX(M)*B(L1)*B(L1)}/2./DET
    ST(2)=(KY(M)*A(L1)*A(L2)+KX(M)*B(L1)*B(L2))/2./DET
    ST(3)=(KY(M)*A(L1)*A(L3)+KX(M)*B(L1)*B(L3))/2,/DET
    RETURN
    END
                              print check for nodal values (debug)
```