

National Library of Canada

Canadian Theses Service

Service des thèses canadiennes

Bibliothèque nationale

du Canada

Ottawa, Canada K1A 0N4

The author has granted an irrevocable nonexclusive licence allowing the National Library of Canada to reproduce, loan, distribute or sell copies of his/her thesis by any means and in any form or format, making this thesis available to interested persons.

The author retains ownership of the copyright in his/her thesis. Neither the thesis nor substantial extracts from it may be printed or otherwise reproduced without his/her permission. L'auteur a accordé une licence irrévocable et non exclusive permettant à la Bibliothèque nationale du Canada de reproduire, prêter, distribuer ou vendre des copies de sa thèse de quelque manière et sous quelque forme que ce soit pour mettre des exemplaires de cette thèse à la disposition des personnes intéressées.

L'auteur conserve la propriété du droit d'auteur qui protège sa thèse. Ni la thèse ni des extraits substantiels de celle-ci ne doivent être imprimés ou autrement reproduits sans son autorisation.

ISBN 0-315-54927-0



MINERALOGICAL AND GEOCHEMICAL INVESTIGATION OF THE CORDIERITE-ANTHOPHYLLITE ROCKS AT STAR LAKE, NEAR SHERRIDON, MANITOBA

BY

MARC V. LEROUX, B.Sc.

A thesis Submitted to the Faculty of Graduate Studies in Partial Fulfillment of the Requirements for the degree of

MASTER OF SCIENCE

Department of Geological Sciences University of Manitoba Winnipeg, Manitoba.

(c) October, 1989

MINERALOGICAL AND GEOCHEMICAL INVESTIGATION OF THE CORDIERITE-ANTHOPHYLLITE ROCKS AT STAR LAKE, NEAR SHERRIDON, MANITOBA

ΒY

MARC V. LEROUX

A thesis submitted to the Faculty of Graduate Studies of the University of Manitoba in partial fulfillment of the requirements of the degree of

MASTER OF SCIENCE

© 1989

Permission has been granted to the LIBRARY OF THE UNIVER-SITY OF MANITOBA to lend or sell copies of this thesis, to the NATIONAL LIBRARY OF CANADA to microfilm this thesis and to lend or sell copies of the film, and UNIVERSITY MICROFILMS to publish an abstract of this thesis.

The author reserves other publication rights, and neither the thesis nor extensive extracts from it may be printed or otherwise reproduced without the author's written permission.

ABSTRACT

The Sherridon Group, a stratigraphic unit of the Kisseynew gneiss belt, near Sherridon, Manitoba is typically dominated by quartz-rich gneisses. Discontinuous lenses of cordierite-anthophyllite rocks, the focus of the present study, occur within these quartz-biotite gneiss units. These cordierite-anthophyllite gneisses are extremely coarse-grained in places, and are spatially associated with barren (Star Lake) and sulfide-mineralized zones (Sherridon; Elken Lake). The present chemical composition of the Star Lake-Elken Lake gneisses indicates hydrothermal alteration of a basaltic progenitor producing chlorite-rich alteration pipes. This alteration involves a depletion in alkali and alkaline earth elements with an apparent enrichment in Al and Mg as a function of constant summation. These chemical modifications are the result of reactions between heated seawater and basaltic rocks of the crust. A modern analogue of the Sherridon Group environment might be the shelf and slope deposits of modern island arcs.

i

ACKNOWLEDGEMENTS

Field work for this thesis was carried out under the auspices of the Manitoba Energy and Mines Department. Field work was provided by W.R.Gunter and P.H.Yamada during 1986, and assistance by W.R.Gunter, E.Froese, N.Halden and S.Lau during 1987.

Chemical analyses were performed by the University of Ottawa and Bondar-Clegg Laboratories. Mineral analyses were done at the National Museum of Canada - Mineral Sciences Division under the guidance of T.S.Ercit.

This thesis was supervised jointly by F.C.Hawthorne and N.Halden who gave constructive criticism during the working stages and during the preparation of the thesis. E.Froese was available for numerous discussions throughout the project and gave constructive criticism.

Personal communications with E.Froese, T.S.Ercit, W.R.Gunter, P.H.Yamada, S.Lau, E.Zaleski, L.Groat and M.Raudsepp are gratefully acknowledged.

ii

TABLE OF CONTENTS

		Page
ABSTI	RACT	i
ACKN	OWLEDGEMENTS	ii
LIST	OF TABLES	vi
LIST	OF FIGURES	vii
LIST	OF PHOTOGRAPHS	i×
LIST	OF APPENDICES	×iii
MINE	RAL ABBREVIATIONS	×i∨
Chap	ters	
I	INTRODUCTION	1
	(i) Statement of Problem	1
	(ii) Previous Work	4
II	GEOLOGICAL SETTING OF THE SHERRIDON AREA	6
	(i) Regional Geology and Stratigraphy	6
	(ii) Star Lake-Elken Lake Local Geology and Stratigraphy	12
III	MINERALOGY AND PETROLOGY	28
	Primary Minerals	28
	(i) Garnet (almandine)	28
	(ii) Orthoamphibole (anthophyllite/gedrite)	31
	(iii) Cordierite	36
	(iv) Sillimanite	36
	(v) Quartz	39
	(vi) Biotite	39
	Accessory Minerals	42

(i) Hercynite spinel
(ii) Staurolite42
(iii) Phosphates46
(iv) Rutile
(v) Magnetite46
(vi) Zircon48
(vii) Sericite and Chlorite
(viii) Sulfides48
IV WHOLE ROCK CHEMISTRY
(i) Chemical Parameters
(ii) Interpretation of Whole-Rock Geochemistry.71
V MINERAL CHEMISTRY
(i) Orthoamphibole
(ii) Garnet80
(iii) Cordierite80
(iv) Sillimanite
(v) Staurolite86
(vi) Biotite86
(vii) Hercynite
(viii) Phosphates
(ix) Chlorite91
VI MINERAL REACTIONS92
(i) Geothermometers
(a) Cordierite-garnet
(b) Biotite-garnet

1.1.1

A	CHEMICAL ANALYSES OF ROCKS FROM STAR LAKE112
в	REE ANALYSES OF ROCKS FROM STAR LAKE124
С	NIGGLI NORMS FOR THE ROCKS FROM STAR LAKE120
D	MINERAL CHEMISTRY OF ROCKS FROM STAR LAKE132

v

LIST OF TABLES

· ·

Tabl	e Page
1	Lithologic units of the Sherridon Group in the vicinity of Sherridon Manitoba
2	Star Lake saw-cut intersections, dimensions and locations25
3	Mineral assemblages from the Star Lake area29
4	Hand specimen modal analyses of representative samples
A1	Geochemical methods and accuracy of analysis114
D1	Amphibole-garnet element table
D2	Biotite-chlorite element table
D3	ZAF setup

.

LIST OF FIGURES

Fiqu	re	Page
1	Location and setting of the Sherridon area in the Kisseynew Complex	.2
2	Geological sketch map of the Sherridon area with the location of the study area	.3
3	Detailed outcrop map showing the position of the various saw cuts	.26
4	Detailed outcrop map showing the position of the various saw cuts	.27
5	Total rock triangle	.56
6	Silico-aluminate triangle	.57
7	Plot of Niggli mg verses c	. 59
8	Plot of Na ₂ O+K ₂ O vs. (K ₂ O/Na ₂ O+K ₂ O)100	.60
9	Plot of Niggli mg verses si	.62
10	Plot of Niggli mg verses k	.63
11	Plot of Niggli al-alk verses c	.65
12	Plot of Niggli al+alk verses c	.66
13	Plot of Niggli al+alk verses ti	. 68
14	The range of REE contents found in the Star Lake- Elken Lake gneisses	69
15	Silico-aluminate triangle	72
16	Plot of Niggli al+alk verses c	74
17	Plot of Niggli al+alk verses ti	.76
18	Plot of Na verses Al (orthoamphiboles)	79
19	Plot of Na verses Fe ²⁺ (orthoamphiboles)	81
20	Plot of Fe ²⁺ verses Mg (orthoamphiboles)	82
21	Plot of Mg/Mg+Fe orthoamphibole verses Mg0/Mg0+Fe0 rock	83

22	Almandine-spessartine-pyrope ternary plot84
23	Pyrope-grossular-spessartine ternary plot85
24	Spinel-gahnite-hercynite ternary plot
25	Plot of Fe ²⁺ verses Zn (hercynites)
26	Plot of Mg verses Zn (hercynites)
27	PT space after Grieve and Fawcett (1974)93
28	Log K ₁ for the cordierite-spinel-quartz equilibrium as a function of pressure for different temperatures99
29	Calculated compositions of anhydrous Mg-Fe cordierites in coexistence with quartz and spinels at three different temperatures

Page

LIST OF PHOTOGRAPHS

Phot

Photo	ograph	Page
1	Unit 1: Fine-grained needle-like anthophyllite, platy aggregates of brown-black biotite and garnets	.14
2	Unit 1: Pegmatite intruding unit 1	.14
3	Unit 1: Euhedral cordierite near pegmatite intrusion of unit 1	.15
4	Unit 1: Euhedral tourmaline found in the upper part of unit 1	.15
5	Unit 1: Abundant inclusions of quartz in large garnet porphyroblast of unit 1	.16
6	Unit 2: Circular to ellipsoidal `snowball' garnets, bladed to stellate groups of anthophyllite and locally abundant blue-white weathered quartz-cordierite aggregates	.16
7	Unit 2: `Snowball' garnets, dark brown groups of anthophyllite blades and quartz-cordierite aggregates	17
8a	Unit 2: Contact of coarser-grained unit 2 with the finer-grained quartz-rich unit 1	17
86	Unit 2: Contact of coarser-grained unit 2 with the finer-grained quartz-rich unit 1	19
9	Unit 2: Contact of unit 2 with pegmatite intrusion, showing no sign of biotite selvages typical of unit 1	19
10	Sillimanite zone: Sillimanite-rich zone located in stratiform tectonic contacts between units 1 and 2, and adjoining a finer-grained garnet-rich layer in unit 1	20
11	Sillimanite zone: Coarse-grained sillimanite from the sillimanite-rich zone	20
12	Unit 3: Elongated to slightly stellate bundles of anthophyllite with quartz-cordierite aggregates	

13	Unit 3: Cordierite occurring as anhedral quartz- cordierite aggregates which are interstitial to the stellate bundles of anthophyllite blades21
14	Unit 3: Gradational contact between units 2 and 3; elongate bundles of anthophyllite are oriented parallel to the bedding22
15a	Unit 3: Cordierite `pockets' with fine-grained myrmekitic hercynite spinel concentrations, surrounded by anthophyllite blades
156	Unit 3: Cordierite `pockets' with fine-grained myrmekitic hercynite spinel concentrations, surrounded by anthophyllite blades
16	Garnet porphyroblast extensively fractured, with chlorite filling the fractures
17	Almandine garnets with abundant inclusion trains of quartz, with minor magnetite, rutile, sillimanite and staurolite
18	Pre-tectonic garnet with sheaths of biotite along the foliation
19	Syntectonic garnet with `rotational' features preserved
20	Garnet with cordierite rims separating them from the surrounding sillimanite
21	Garnet replaced by cordierite and hercynite34
22a	Anthophyllite blades ranging in length from 0.1 -> 3 cm, and defining a schistosity together with intergrown biotite35
22ь	Anthophyllite blades ranging in length from 0.1 -> 3 cm, and defining a schistosity together with intergrown biotite
23	Anthophyllite in sheath-like groups with individual prisms ranging up to 9mm in length37
24	Chlorite reaction rims mantling anthophyllite blades
25	Quartz-cordierite-sillimanite pockets surrounded by amphibole blades and biotite grains

26	Cordierite-sillimanite pocket surrounded by biotite
27a	Fibrolitic sillimanite aggregations
27ь	Fibrolitic sillimanite aggregations
28	Fibrolitic sillimanite in elongate to elliptical aggregates of tiny needles oriented parallel to the foliation41
29	Coarse prismatic sillimanite crystals are diversely oriented in the rock and commonly cut other minerals and each other
30	Randomly oriented biotite grains studded with abundant zircon +/-sphene inclusions surrounded by pleochroic haloes43
31	Hercynite associated with cordierite and replacing garnet in the presence of sillimanite43
32	Fine-grained myrmekitic hercynite intergrown with minor magnetite and cordierite surrounded by anthophyllite; a `ring´ structure of anthophyllite is found isolated in a cordierite `pocket´44
33	Fine-grained myrmekitic hercynite replacing sillimanite and possibly pseudomorphing the sheaths
34	Staurolite as an anhedral prism rimmed by cordierite and replacing garnet
35	Euhedral staurolite grain replacing garnet and itself being altered to hercynite
36	Rounded to blocky, locally prismatic apatite grains are found mostly in the sillimanite zone; these tend to form a network associated with coarse-grained biotite and sillimanite
37	Wagnerite, a Mg-rich phosphate, is mostly fine- grained with minute inclusions of quartz and anthophyllite, both rimmed by apatite that separates them from the enveloping wagnerite47

x

38	Sericite as minute `shreds´ found rimming and altering cordierite, sillimanite and anthophyllite
39	Chlorite as an alteration product of anthophyllite49
40	Chlorite as an alteration product of cordierite50
41a	Chalcopyrite forming coarse-grained patches, mainly associated with pyrrhotite
41ь	Chalcopyrite forming coarse-grained patches, mainly associated with pyrrhotite
42	Pyrrhotite replaced by marcasite
43	Fine-grained network of sulfides infilling fractures in the cordierite-anthophyllite gneiss 53

,

LIST OF APPENDICES

an de service de la construcción d Construcción de la Construcción de l

-

Appendix		Page
A	CHEMICAL ANALYSES OF ROCKS FROM STAR LAKE NEAR SHERRIDON MANITOBA	.112
В	RARE EARTH ELEMENT ANALYSES OF ROCKS FROM STAR LAKE NEAR SHERRIDON MANITOBA	.124
С	NIGGLI NORMS CALCULATED FOR THE ROCKS FROM STAR LAKE NEAR SHERRIDON MANITOBA	.128
D	MINERAL CHEMISTRY OF ROCKS FROM STAR LAKE NEAR SHERRIDON MANITOBA	.132

•

MINERAL ABBREVIATIONS

 $\chi_{\rm c}$, the second second state of the descent second second state $\Delta_{\rm s}$

Alm	Almandine
Als	Aluminosilicate
And	Andalusite
Ath	Anthophyllite
Ар	Apatite
Bt	Biotite
Сср	Chalcopyrite
Chl	Chlorite
Cld	Chloritoid
Crd	Cordierite
Grt	Garnet
Ged	Gedrite
Hc	Hercynite
Ку	Kyanite
Mag	Magnetite
Mc	Marcasite
Oam	Orthoamphibole
Po	Pyrrhotite
Qtz	Quartz
Sil	Sillimanite
Spl	Spinel
St	Staurolite
Tur	Tourmaline
Wg	Wagnerite
Zrn	Zircon

xiv

•

(I) INTRODUCTION

(i) Statement of Problem

A prominent cordierite-anthophyllite assemblage occurs between Kississing Lake and Star Lake, Manitoba (known as the Sherridon structural basin, Goetz 1980). This structure is underlain by rocks of the Sherridon Group, a stratigraphic unit of the Kisseynew gneiss belt, near Sherridon, Manitoba.

The origin of cordierite-anthophyllite bearing rocks is obscure. They do not have any obvious chemical counterparts among common sedimentary or igneous rocks. Chemically, these rocks range from silica-poor to silicarich. They are all marked by deficiencies in Ca and alkalies, and relative enrichments in Al, Mg, and Fe. Cordierite- anthophyllite gneisses are almost always associated with Fe -sulfides and base metal mineralization. These rocks have been regarded as allochemical metamorphic products (Eskola 1914), and commonly form in high-grade regional metamorphic terranes, suggesting a common origin for all occurrences.

A detailed mineralogical and chemical study was done in the Star Lake - Elken Lake area of the Sherridon structure (Figures 1 and 2). The objectives of this work are (i) to identify the mineralogical and chemical character of the cordierite-anthophyllite unit; (ii) to compare it to other known occurrences; (iii) to define the processes

-1-



FIGURE 1: Location and setting of the Sherridon area in the Kisseynew Complex (Froese and Goetz, 1981).



FIGURE 2: Geological sketch map of the Sherridon area with the location of the study area (Froese and Goetz, 1981). leading to the observed lithologies. The origin and character of the original protolith can then be considered, with a view to establishing chemical and mineralogical parameters that discriminate between the barren assemblage at Star Lake and the sulfide-mineralized zones at Sherridon, Manitoba.

(ii) Previous Work:

Bateman and Harrison (1946) mapped the eastern portion of the Sherridon Structure. These authors primarily mapped west of the Batty Lake area (Figure 1), and discovered the first textural evidence of a volcanic origin for some of the amphibolites near Sherridon. They classified them as volcanic breccias and reported deformed pillow structures.

The anthophyllite-bearing rocks of the Sherridon Group were first recognized by Robertson (1953). In his study of the Batty Lake map-area, which adjoins the Sherridon map-area, he described two main rock types: quartz-rich and hornblende-rich gneisses. These units are associated with thin intercalations of limestone and anthophyllite-rich rocks. The band of anthophyllite-rich rock was traced from east to west in the map-area, and, although not mentioned by Bateman and Harrison (1946), was found at several localities in the Sherridon area (Figure 2). This band contained the characteristic mineral assemblages:

-4-

(i)anthophyllite-cummingtonite-biotite (ii)anthophyllite-biotite-almandine (iii)anthophyllite-cordierite-biotite (iv)anthophyllite-cordierite-biotite-almandine

varying in width from 15 to 30 meters and occupying a relatively constant stratigraphic horizon 150 to 250 meters above the base of the Sherridon Group. Robertson (1953) suggested that the mineral assemblages in the anthophyllite band represented the products of iron-magnesium metasomatism of an aluminium-rich sediment or tuff, with the latter option possibly favoured by the wide lateral development and apparent continuity of the band.

Wilkinson (1976) did detailed petrographic work on the anthophyllite-bearing rocks of the Sherridon Group, with emphasis on the Elken Lake area. He noted the similarity of enrichment in Mg, Fe and Al in the Sherridon cordieriteanthophyllite rocks with alteration zones related to massive sulfide deposits, suggesting that the rock type could have been deposited as a chemical sediment on the sea floor.

Other studies of the area include those by Goetz (1980), Froese and Goetz (1981, 1982) and Froese (1985). These authors suggest that the protolith of the cordieriteanthophyllite rocks could have been a chlorite-rich rock representing transported and altered material from an unknown source.

-5-

Alteration zones in volcanic rocks are commonly associated with sulfide mineralization, whereas many stratiform lenses of cordierite-anthophyllite rocks in the Sherridon Group occur structurally far removed from sulfide deposits. According to these authors, as most cordieriteanthophyllite rocks lie on the lower ore-bearing horizon (Figure 2), and due to their representative compositions, these rocks required a chlorite -rich protolith, associated with sulfide deposition (Froese and Goetz, 1981; 1982).

This study will, therefore, focus on the mineralogical and chemical character of the cordieriteanthophyllite unit in order to define possible processes leading to the observed lithogies.

(II) GEOLOGICAL SETTING OF THE SHERRIDON AREA

(i) Regional Geology and Stratigraphy

The Flin Flon region, including the southern edge of the Kisseynew Complex, was first mapped by Bruce (1918). The first subdivision of the Kisseynew gneisses, north and east of Flin Flon in northern Manitoba and Saškatchewan, was made by Bateman and Harrison (1946) for the rocks of the Sherridon area. They recognized three groups:

 (i) the Pre-Sherridon Group consisting of quartzofeldspathic gneisses in the centre of the Sherridon structure, pre-dating the Sherridon Group in which the orebodies occur;

-6-

(ii) the Sherridon Group consisting of a group of distinctive quartzites interbedded with hornblendeplagioclase gneisses that are metamorphosed volcanic flows; these lie in places upon Sherridon gneisses but elsewhere rest on still older sedimentary gneisses;

(iii) and the Post-Sherridon Group; these dark green hornblende-rich gneisses overlie the Sherridon Group quartzites.

Robertson (1953) later suggested that the Pre-Sherridon Group of Bateman and Harrison (1946) is older than the Sherridon Group, ("Sherridon rocks occupy a synclinal structure"), and renamed it the Nokomis Group. Later investigations in the Kisseynew Complex (Bailes, 1971) suggested that the Sherridon Group correlates with the Missi Group of the Flin Flon belt, and that the Nokomis Group correlates with the Amisk Group sediments. The main subdivisions of the Kisseynew Complex thus include Missi Group paragneisses (metasedimentary rocks with some associated volcanic rocks) surrounded by older Amisk Group metasedimentary rocks.

Grey, medium grained, quartz-plagioclase-biotite gneisses and migmatites predominate in the Nokomis Group. In addition to a gneissic fabric, the rocks locally have compositional layering 10 to 100 cm wide, defined by slight variations in biotite content.

The Sherridon Group has been divided into five

-7-

stratigraphic units (Froese and Goetz, 1981; Figure 4). The lower three units are:

(i)calc-silicate gneisses;

(ii)lower quartz-rich gneisses (with layers of amphibolite, pelitic gneiss, calc-silicate gneisses. Massive sulfide deposits occur along two stratigraphic horizons associated (in some localities) with cordierite-anthophyllite rocks);

(iii)impure marble-calc-silicate gneisses with thin layers of calc-silicate gneisses arranged in a progressively younger sequence from the edge to the centre of the `Sherridon structural basin', with no repetition of units by folding observed.

The most abundant rock in the Sherridon Group is a fine- to medium-grained quartz-rich gneiss consisting of quartz, oligoclase-andesine, K-feldspar, biotite and fine-grained almandine (Table 1). This unit is characterized by prominent quartz ridges on weathered surfaces, and layers of biotite (compositional layering on a scale of 10 to 30 cm) near the base of the Sherridon Group. Cordieriteanthophyllite rocks occur as discontinuous lenses within these quartz-rich gneisses. They are coarse-grained with garnet and anthophyllite crystals up to 10 cm in size. Most occurrences lie in the lower ore-bearing horizon, in a few

-8-

Table of Formations

Group	Lithology
Intrusive Rocks	(12) Felsic pegmatite.
	 Granodiorite, medium grained, composed of quartz, plagioclase, K feldspar, and biotite.
	(10) Pyroxenite, massive, composed of hornblende pseudomorphs after pyroxene.
	(9) Gabbro, massive to foliated, composed of plagioclase and hornblende.
Sherridon Group	 (8) Massive amphibolite composed of plagioclase, hornblende, and garnet.
	(7) Amphibolite with local layering and presence of felsic clasts; composed of quartz, plagioclase, hornblende, and garnet. Minor amount of felsic fragmental rock.
	(6) Impure marble and calc-silicate rocks; marble beds distinguished by a high (50%) content of calcite.
	(5) Calc-silicate rocks, composed of quartz, plagioclase, hornblende, diopside, and calcite.
	Cordierite-anthophyllite rocks, composed of cordierite, anthophyllite, and garnet, with minor amounts of quartz and biotite; some occurrences associated with sulphide mineralization.
	(4) Pelitic schists, composed of quartz, plagioclase, biotite, sillimanite, garnet, and cordierite; typically rusty weathering, associated with sulphide mineralization.
	(3) Biotite-garnet schist, composed of quartz, plagioclase, biotite and garnet; characterized by small euhedral garnet porphyroblasts.
	(2) Quartz-rich gneisses, composed of quartz, plagioclase, K feldspar, biotite, and garnet. Some interlayered calc-silicate and pelitic gneisses.
Nokomis Group	(1) Quartzofeldspathic gneisses, migmatites, and granitoid gneisses.

TABLE 1: Lithologic units of the Sherridon Group in the vicinity of Sherridon Manitoba (Froese and Goetz, 1981).

places below the sulfide layer, but more commonly along unmineralized sections of the cordierite-anthophyllite horizon - Star Lake, (Figure 2). Calc-silicate rocks are also interlayered with guartz-rich gneisses. These are fine- to medium-grained, greyish-green rocks composed mainly of quartz, andesine-labradorite and hornblende, with K-feldspar, biotite, scapolite, diopside and sporadic calcite. The interlayered impure marble and calc-silicate rocks have a similar mineralogical composition but with higher contents of calcite and the presence of Cr-bearing grossular. The amphibolites form several continuous layers within the quartz-rich gneisses; they are mainly dark grey, medium-grained rocks of hornblende, garnet, fine- to quartz, plagioclase and trace amounts of biotite.

Later investigations of the Kississing Lake map area (Schledewitz, 1988) have resulted in four mappable components for the Kisseynew gneisses. These four components are considered to be sedimentary and/or volcanic derived metamorphic rocks. Two major lithostratigraphic units are the Burntwood River (oldest metamorphic site (Nokomis Group; Robertson, 1953)); an aluminous - graphite - bearing biotite gneiss and the younger Missi metamorphic suite which is predominantly composed of quartzofeldspathic rocks that are variably magnetiferous and hornblende - bearing.

The two remaining lithostratigraphic components have been placed into a category of rocks of uncertain

-10-

affinity and/or age. These include a suite of amphibolite +/- garnet, calc - silicates and hornblende - biotite garnet feldspar - quartz gneiss lying discontinuously between the Burntwood and Missi metamorphic suites. It is uncertain whether or not they are part of the Burntwood or Missi suite of rocks. The second is the Sherridon metamorphic suite (Sherridon Group, Bateman and Harrison, 1946; Froese and Goetz, 1981) which is considered to be a highly recrystallized and tectonized varied suite of rocks

A new structural interpretation suggests that the Sherridon metamorphic suite may be older than the Missi suite (Zwanzig, 1988). Sherridon gneisses have been considered to be younger than the surrounding gneisses (Bateman and Harrison, 1946; Froese and Goetz, 1981) but the type section at Sherridon is surrounded by granitoid rocks and has an unknown stratigraphic position. Similar gneisses north and east of Walton Lake lie between two belts of Burntwood River metagreywacke (contact locally gradational). The metagreywacke is stratigraphically overlain by the Missi suite, which suggests that the Sherridon suite lies in the core of an anticline and is equivalent to or older than the Burntwood River suite.

Southwest of Meat Lake, Sherridon - type gneiss is traced laterally into less coarsely recrystallized rocks that appear to represent basaltic, dacitic and intermediate

-11-

metavolcanic rocks, locally interlayed with greywacke derived gneiss. This part of the Sherridon suite is tentatively correlated with the Amisk Group (Zwanzig, 1988).

Recent work in the Wildnest Lake area. Saskatchewan has also led to the recognition of a suite of Kisseynew gneisses which are strikingly similar to those of the Sherridon Group in the vicinity of Sherridon, but which have been assigned to the Amisk Group (Ashton et al, 1986; 1987). These geological observations, therefore, cast some doubt on the correlation between the Sherridon Group in the vicinity of Sherridon and the Missi Group. The possibility must be considered that the Sherridon Group, like the Wildnest Lake suite, is largely of volcanic origin and is equivalent to a part of the Amisk Group (Ashton and Froese, 1988).

(ii) Star Lake - Elken Lake: Local Geology and Stratigraphy

The garnet-anthophyllite gneiss zone at Star Lake is considered to be, by some authors, an extention of the cordierite-anthophyllite rocks near Sherridon, Manitoba, where massive sulfide deposits are found along the same stratigraphic horizon. In places, this prominent garnetanthophyllite gneiss zone is extremely coarse-grained and is underlain by a sillimanite-bearing unit. Within the garnetanthophyllite rocks, three mappable units have been recognized by Gunter and Yamada (1986). From the structural

à

-12-

base upwards, these units contain:

(1) quartz-garnet-anthophyllite-biotite

(2) garnet-anthophyllite +/- cordierite

(3) anthophyllite-cordierite

Unit 1 is characterized by 1 to 2 cm sprays of fine-grained needle-like anthophyllite and platy aggregates of brown to black biotite, both of which tend to concentrate in monominerallic bands parallel to the bedding planes. The garnets are generally 2 to 4 cm in diameter, and occur in a very fine-grained quartz-rich groundmass (photographs 1, 5). Quartz-cordierite aggregates are locally present, and pegmatites intruding the unit tend to develop biotite-rich selvages in contact with euhedral cordierite grains (photographs 2, 3). Euhedral crystals of K-feldspar, quartz, biotite, cordierite and garnet are present. In addition, tourmaline aggregates are also found in the upper part of unit 1 (photograph 4).

Unit 2 is composed mainly of circular to ellipsoidal snowball garnets 5 to 6 cm in diameter, stellate groups of dark brown anthophyllite, 2 to 5 cm in length and locally abundant 2 to 3 cm blue-white weathered quartzcordierite aggregates (photographs 6, 7, 8). This unit is conspicuously zoned, with a garnet-rich base grading upward into an approximately equal mixture of garnet, anthophyllite and cordierite. Pegmatites intruding this unit do not seem to have reacted with it, as there are no biotite selvages

PHOTOGRAPH 1

Unit 1: 1 - 2 cm sprays of fine-grained needle-like anthophyllite, platy aggregates of brown-black biotite and garnets ranging from 2 to 4cm in diameter in a fine-grained quartz-rich cordierite groundmass.

PHOTOGRAPH 2

Unit 1: pegmatite (pgm) intruding unit 1. Note the biotite (Bt) - rich shelvages in contact with euhedral cordierite.



PHOTOGRAPH 3

Unit 1: euhedral cordierite (Crd) in pegmatite intrusion of unit 1.

PHOTOGRAPH 4 Unit 1: euhedral tourmaline (Tur) found in the upper part of unit 1.



• •• •••

.

an in the second se

PHOTOGRAPH 5
Unit 1: abundant inclusions of quartz in large garnet (Grt)
porphyroblast of unit 1.

PHOTOGRAPH 6

Unit 2: circular to ellipsoidal `snowball' garnets, bladed to stellate groups of anthophyllite, and locally abundant blue-white weathered quartz-cordierite aggregates.


PHOTOGRAPH 7

Unit 2: 5-6 cm `snowball' garnets (Grt), 2-5 cm dark brown groups of anthophyllite (Ath) blades and 2-3 cm quartz-cordierite aggregates.

PHOTOGRAPH 8a

Unit 2: contact of coarser-grained unit 2 (circular to ellipsoidal `snowball' garnets, bladed to stellate groups of anthophyllite and blue-white weathered quartz-cordierite aggregates) with the finer-grained quartz-rich unit 1.



found near the contacts (photograph 9). Sillimanite-rich zones occur at contacts between units 1 and 2, and adjoin a finer-grained garnet-rich layer in unit 1 (photographs 10, 11). In the northwest area of Elken Lake (unit 2), there is disseminated pyrrhotite, chalcopyrite and minor pyrite along the same main ore-bearing horizon as the Sherritt Gordon orebodies. However, in this case, the sulfide is thin, stratiform, weakly disseminated and rarely attains a thickness greater than a few meters.

Unit 3 is composed of coarse-grained (5 to 10 cm) elongated to slightly stellate bundles of anthophyllite oriented parallel to bedding (photograph 12). Cordierite occurs as 3 to 5 cm anhedral quartz-cordierite aggregates interstitial to the anthophyllite aggregates (photograph 13). Both the lower and upper contacts of units 2 and 3 are transitional over 1 to 3 m; and are conformable with the flattened garnets and anthophyllite lineations (photograph 14, 15).

In general, these units contain large, poikiloblastic radiating sheaves of anthophyllite, which may be as much as 10 cm long, poikiloblastic dark-blue to purple cordierite (frequently altered to pinite, and containing quartz inclusions), large poikiloblastic pink garnets, with biotite and quartz making up much of the remainder of the rock. Intergrown in the gneisses are magnetite, green spinel, staurolite and chlorite, with minor

-18-

РНОТОGRAPH ВЬ

Unit 2: contact of coarser-grained unit 2 (circular to ellipsoidal `snowball' garnets, bladed to stellate groups of anthophyllite and blue-white weathered quartz-cordierite aggregates) with the finer-grained quartz-rich unit 1.

PHOTOGRAPH 9

Unit 2: contact of unit 2 with pegmatite (pgm) intrusion, showing no sign of biotite selvages that are typical of unit 1.



-19-

.,

PHOTOGRAPH 10

Sillimanite-rich (Sil) zone located in stratiform tectonic contacts between units 1 and 2, and adjoining a finer-grained garnet-rich layer in unit 1.

PHOTOGRAPH 11 Coarse-grained sillimanite (Sil) from the sillimanite-rich zone.



PHOTOGRAPH 12

Unit 3: 5-10 cm elongated to slightly stellate bundles of anthophyllite (Ath) with quartz-cordierite (Crd) aggregates interstitial to the anthophyllite blades.

PHOTOGRAPH 13

Unit 3: 3-5 cm anhedral quartz-cordierite (Crd) aggregates interstitial to the stellate bundles of anthophyllite blades.



••

.

PHOTOGRAPH 14 Unit 3: gradational contact between units 2 and 3; elongate bundles of anthophyllite are oriented parallel to the bedding.

PHOTOGRAPHS 15a

Unit 3: cordierite (Crd) `pockets' with fine-grained myrmekitic hercynite (Hc) spinel concentrations, surrounded by anthophyllite (Ath) blades.



PHOTOGRAPHS 15b

٠

Unit 3: cordierite `pockets' with fine-grained myrmekitic hercynite (Hc) spinel concentrations, surrounded by anthophyllite (Ath) blades.

.



pyrite, pyrrhotite, and chalcopyrite.

In the field a rock saw was used to sample the units, providing as continuous a sample as possible. Saw-cut intersections, dimensions and locations are given in Table 2 and Figures 3 and 4.

SAN CUT #	INTERSECTION OF ROCK UNITS	LENGTH (CM)	LOCATION
SC - 1 SC - 2a SC - 2b SC - 3a SC - 3b SC - 3c SC - 3d	UNIT 1 - pegmatite UNIT 1 - pegmatite - UNIT 1 UNIT 1 - pegmatite - UNIT 1 UNIT 1 UNIT 1 UNIT 1 UNIT 1 - UNIT 2 UNIT 2 - UNIT 1 (?)	$\begin{array}{rrrrrrrrrrrrrrrrrrrrrrrrrrrrrrrrrrrr$	NW - fig 3 NW - fig 3 NW - fig 3 N - fig 3 N - fig 3 N - fig 3 N - fig 3
SC - 3e	UNIT 2	773 - 870	NE - fig 3
SC - 3f	UNIT 2	870 - 1027	NE - fig 3
SC - 4a	UNIT 3	0 - 100	SF - fig 3
SC - 4b	UNIT 3	100 - 150	SE - fig 3
SC - 5a	UNIT 2	0 - 172	N - fig 4
SC - 5b	UNIT 2 (?)	470 - 685	NF - fig 4
SC - 5c SC - 5d SC - 5e	UNIT 3 UNIT 3 quartz - biotite gneiss	762 - 853 880 - 1022 1230 - 1317	NE - fig 4 NE - fig 4 NE - fig 4 NF - fig 4
SC - 5f	quartz - biotite gneiss	1556 - 1635	NE - fig 4
SC - 6	UNIT 1 (?) - UNIT 2	0 - 76	NW - fig 4
SC - 7	UNIT 1 - sillimanite/quartz	0 - 152	SW - fig 4
SC - 88	quartz vein - UNIT 3	170 - 476	centre - fig 4
SC - 86	UNIT 2 - UNIT 3-quartz vein	0 - 96	centre - fig 4

TABLE 2	•	STAR	IAKE	SAM-CHT	INTERSECTIONS	DIMENSIONS	AND	LOCATIONS
	••	• • • • • • •		01111 001	THICHOCOLLONG	DINCHOIVNO	nny	COMPTOND

UNIT 1:	Quartz - Garnet - Anthophyllite - Biotite
UNIT 2:	Garnet - Anthophyllite - Cordierite
UNIT 3:	Anthophyllite - Cordierite

-25-

· •

•



yanı, çokurana ar Öradild Ar Antori



(III) MINERALOGY AND PETROLOGY:

The detailed mineral assemblages from the Star Lake - Elken Lake area are summarized in Table 3.

The cordierite-garnet-anthophyllite bearing assemblages occur in a set of well-banded, coarse-grained porphyroblastic rocks which are generally modally dominated by amphibole (over 40-50 %). Schistosity is generally defined by the parallel orientation of long amphibole blades and biotite plates.

Due to the coarse grained nature of the specimens, modal analyses were done on hand specimens in order to give an estimate of bulk rock composition. In practice, a minimum area 100 times the area of the largest grain in the hand specimen (with a minimum of a 1000 point counts in such an area) was used (Jackson and Ross, 1956). Amphibole (where present), garnet, cordierite and quartz are the most abundant minerals, with anthophyllites > garnets > cordierites > quartz (units 2 & 3) (Table 4). Biotite is the next most abundant mineral, followed by sillimanite, hercynite, magnetite, rutile, apatite, staurolite and zircon +/- titanite.

Primary Minerals

(i) <u>Garnet (almandine)</u>

Porphyroblasts of almandine garnet ranging in size from 0.25 mm (broken pieces) to > 25 mm in diameter are very abundant in the gneiss zone. The garnets are usually

-28-

SAMFLE #	Qtz	Sil	St	Grt	Amp	Crd	Bt	Chl	Hc	Ар	Wg	Mag	Plag
SC 1-7 1-14a 2-1 2-9 3-10 3-25a 3-26a 3-26a 3-26a 3-26a 3-26a 3-26a 3-26a 3-26a 3-27 3-34 3-43 3-54a 4-9 4-12 4-12 5-12 5-126 5-216 5-216 5-25a 5-359 5-67 6-3 6-4 7-1 7-4 7-10 8-7 8-15 8-17 8-28	0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0	0×0××0× 0 ×0× 0×0 000×0×0×0×0×0×××	× × 0 ×	×××××× ×××××××××××××××××××××××××××××××	· · · · · · · · · · · · · · · · · · ·	××××××××××××××××××××××××××××××××××××××	×××××× o × o × × × × × × × × × × × × ×	o oo xooooo oo oooxoxox x		×	×××		0
ML 01 08 15 33 37 49 52 55 93		× ×	× × ×	0 0 × × ×	000 × × × ×	- × ×	0 0 0 × ×		× × ×		×	0	
x - analyzed mineral o - mineral present													

TABLE 3: MINERAL ASSEMBLAGES FROM THE STAR LAKE AREA

[· · · · · · · · · · · · · · · · · · ·	1	1	r	T
Specimen	Unit	GARNET	QUARTZ	CORD -	ANTHO -	BIOTITE	SILLI -
NO.	NO.			IERLIE	PHYLLITE		MANITE
SC 1-2	1	2.0	11.8			84.7	
SC 1-14A	1	3.0	27.9			59 6	95
SC 2-6	1	29.3	40.4			30.3	1,10
SC 2-11	1	26.8	24.0			49.2	
SC 2-21	1	37.0	22.6			37.0	3.4
SC 2-25	1	32.4	37.3			22.0	8.3
SC 3-8	1	13.2	26.0	5.9	50.	0	4.9
SC 3-10	1	9.1	7.7	20.3	52.	.8	10.1
SC 3-18	1		2.4	21.5	76.	1	
SC 3-26	2	19.4	10.4		60.1		10.1
SC 3-28	2	25.5	10.8	11.2	52.5		
SC 3-34	2	28.1	4.7	13.6	53.6		
SC 3-41	2	26.9	3.7	10.3	59.1		
SC 3-44	2	38.1	10.6	5.5	45.8		
SC 3-51	2	44.1	8.1	8.6	38.9		
SC 3-53	2	7.8	5.2	16.7	70.3		
SC 4-5	3	9.5	17.4	11.7	61.4		
SC 4-13	3		32.1	12.7	55.2		
SC 5-2	2	50.2	12.1			37.7	
SC 5-6	2	58.2	1.3		40.	5	
SC 5-11	2	56.2	6.4		37.	4	
SC 5-18	2	9.8	2.4	29.3	58.5		
SC 5-20	2	28.6	3.0	18.6	49.	8	
SC 5-25	2	12.4	13.1	15.1	59.	4	
SC 5-29	3		3.0	24.4	72.6		
SC 5-31	3		3.3	34.1	62.6		
SC 5-42	3		2.6	15.5	81.9		
SC 6-1	1		31.9			59.2	8.9
SC 6-4	2	37.0	8.8			41.2	13.0
SC 7-4	.1	28.6				13.1	58.3
SC 7-10	1					29.3	70.7
SC 8-4	3			31.0	69.0		
50 8-6	3		11.2	18.2	70.6		
SC 8-13	ے ح		21.9		78.1		
ar 8-1/	د		10.3		89.7		

TABLE 4: HAND SPECIMEN MODAL ANALYSES OF REPRESENTATIVE SAMPLES

.

anhedral to subhedral, and are extensively fractured with chlorite infilling and replacing them (photograph 16). This replacement was possibly caused by a later low-grade metamorphic event. These porphyroblasts contain abundant inclusion trains of quartz, with minor magnetite, staurolite, sillimanite, rutile and hercynite spinel (photograph 17).

Two distinct periods of garnet crystallization can be inferred in these units:

(i) pre-tectonic crystals (photograph 18), with sheaths of biotite along the foliation;

(ii) syntectonic crystals (photograph 19), with `rotational' features preserved. In general, the garnets show no sign of reaction except in the presence of sillimanite, where they seem to be replaced by cordierite and hercynite spinel (photographs 20 and 21).

(ii) Orthoamphibole (anthophyllite / gedrite)

Orthoamphiboles occur either radially, ranging in length from 0.1 to > 3.0 cm, or in randomly oriented sheathlike groups with individual prisms up to 9 mm long. The blades are dark-brown in hand-specimen and define a schistosity together with intergrown biotite (photograph 22). The sheath-like aggregates have pronounced fractures and prismatic cleavage (photograph 23), both of which are quite characteristic for orthoamphiboles. Cordierite and chlorite reaction rims are found mantling the orthoamphibole

-31-

PHOTOGRAPH 16 Extensively fractured garnet (Grt) porphyroblast, with chlorite filling the fractures (ppl - X15).

PHOTOGRAPH 17

Almandine garnets (Grt) with abundant inclusion trains of quartz, and minor magnetite (Mag), rutile, sillimanite (Sil) and staurolite (St) (ppl - X5).





PHOTOGRAPH 18 Pre-tectonic garnet (Grt) with sheaths of biotite (Bt) along the foliation (ppl - X4).

PHOTOGRAPH 19

Syntectonic garnet (Grt) with `rotational' features preserved (crystals contain a set of discontinuous internal S-surface (S1) which are composed of trails of round to elongate blebs of quartz (Qtz) and grains of magnetite) (ppl - X5).



PHOTOGRAPH 20 Garnet (Grt) with cordierite (Crd) rims separating garnet from the surrounding sillimanite (Sil) (ppl - X20).

PHOTOGRAPH 21

Garnet (Grt) replaced by cordierite (Crd) and hercynite (Hc) (ppl - X20).



.



PHOTOGRAPHS 22a & 22b Anthophyllite (Ath) blades ranging in length from 0.1->3 cm, and defining a schistosity together with intergrown biotite (Bt) (ppl - X5).



grains (photograph 24). The orthoamphiboles also show reaction rims of cordierite when in contact with hercynite and sillimanite (photograph 25) revealing their incompatibilities. Textural relationships between staurolites and orthoamphiboles were not found.

(iii) <u>Cordierite</u>

A considerable proportion of the matrix material between amphibole prisms is cordierite; the anhedral grains range from 0.5 to 4 mm in diameter and are purple in hand specimen. Many grains and larger crystals of cordierite are twinned, with both polysynthetic and penetration twins present, and are studded with minute hercynite inclusions. Chlorite is locally found rimming the cordierite, whereas pinite occurs along the fractures. The inclusions give the larger grains characteristic `cloudy' appearance, a distinguishing them from quartz (photograph 26). Cordierite-quartz-sillimanite aggregates are also commonly found surrounded by amphibole and biotite grains (photograph 25).

(iv) <u>Sillimanite</u>

In the gneiss zone, there are two varieties of sillimanite. Fibrolitic sillimanite occurs in elongate to elliptical aggregates of tiny needles oriented parallel to the foliation. They also tend to be associated with other elongate features such as pegmatite and quartz veins. Most aggregates are 0.5 - 10 mm long, and seem to suggest a

-36-

PHOTOGRAPH 23

Anthophyllite (Ath) in sheath-like groups with individual prisms ranging up to 9mm in length: the sheath-like aggregates are randomly oriented and have pronounced fractures and prismatic cleavages (ppl - X5).

PHOTOGRAPH 24

Chlorite (Chl) reaction rims mantling anthophyllite (Ath) blades (ppl - X20).



an saida 👘

PHOTOGRAPH 25

Quartz-cordierite(Crd)-sillimanite (Sil) pocket surrounded by amphibole (Ath) blades and biotite (Bt) grains (ppl - X5).

PHOTOGRAPH 26

Cordierite(Crd)-sillimanite (Sil) pocket surrounded by biotite (Bt); inclusions give the cordierite a 'cloudy dusting' appearance, distinguishing it from quartz grains (ppl - X4).



-28-

metamorphic reaction process adjacent to the various veins (photographs 27, 28). Coarse prismatic sillimanite crystals are also present. These are randomly oriented in the rock, commonly cross-cutting other minerals and each other (photograph 29). Single prismatic sillimanite crystals are generally 0.1 - 0.2 mm thick and 1.5 - 2 mm long. Aggregates of these crystals are locally as much as 12 cm long and 1 cm in diameter. Some of the sillimanite is replaced and possibly pseudomorphed by hercynite spinel. Sillimanite is found isolated from orthoamphibole by a cordierite mantle indicating their incompatibilities.

(v) <u>Quartz</u>

Quartz occurs as fine- to coarse-grained, anhedral and mostly strained aggregates scattered amongst the amphiboles and garnets in the gneiss zone. Quartz is prominent in unit 1; there is no evidence of mylonitic textures, and it is a major constituent of inclusions in the garnet porphyroblasts. Various inclusions of sillimanite, magnetite and rutile occur in the quartz; hercynite spinel inclusions are also found (SC 4-9; SC 4-12) indicating their compatibility.

(vi) Biotite

Most of the biotite in the cordieriteanthophyllite gneiss is light- to dark-brown in thin section. It usually occurs as subhedral to euhedral crystals studded with abundant zircon inclusions which are

-39-

PHOTOGRAPHS 27a & 27b

Fibrolitic sillimanite (Sil) aggregates are 0.5-10 mm long and seem to be due to leaching processes adjacent to pegmatite intrusions (ppl - X20).


Fibrolitic sillimanite (Sil) in elongate to elliptical aggregates of tiny needles oriented parallel to the foliation (ppl - X20).

PHOTOGRAPH 29

Coarse prismatic sillimanite (Sil) crystals are diversely oriented in the rock and commonly cut other minerals and each other (ppl - X20).



surrounded by pleochroic haloes (photograph 30). The biotites are 0.5 to 9 mm long and typically define a schistosity associated with anthophyllite. Biotite inclusions occur in amphiboles and garnets with no reaction rims present.

Accessory Minerals

(i) <u>Hercynite spinel</u>

Hercynite is present in all rock units in various amounts, probably dependent upon the extent of reaction between garnet, sillimanite, cordierite and amphibole. The spinel occurs as minute inclusions in cordierite, and as fine- grained myrmekitic intergrowths with cordierite which seem to be the result of pseudomorphing sillimanite sheaths (photographs 31, 32, 33). The mantling relationship between hercynite and cordierite may be explained by the reaction:

sillimanite + garnet = cordierite + hercynite
(ii) <u>Staurolite</u>

Staurolite occurs mainly as anhedral prisms up to 9 mm long, closely associated with amphibole and garnet. In its association with garnet, the grains are usually <1 mm, and occur at the contacts of both. Cordierite rims the staurolite, indicating some reaction (photograph 34), particularly at the contact with anthophyllite. Staurolite frequently replaces almandine garnet. Less commonly, it is itself replaced by hercynite spinel (photograph 35), indicating the compatibilities of the two.

Randomly oriented biotite (Bt) grains studded with abundant zircon (Zrn) inclusions surrounded by pleochroic haloes (ppl - X20).

PHOTOGRAPH 31

Hercynite (Hc) associated with cordierite (Crd) and replacing garnet in the presence of sillimanite and quartz (ppl - X20) (garnet + sillimanite = hercynite + cordierite + quartz) 0.25(Fe_{1.9}Mg_{1.1})Al₂Si₃O₁₂ + 0.5Al₂SiO₃ = 0.65(Fe_{0.7}Mg_{0.3})Al₂O₄ + 0.05 Fe_{0.4}Mg_{1.6}Al₄Si₃O₁₈ + SiO₂



.

.



Fine-grained myrmekitic hercynite (Hc) Cordierite (Crd) with minor magnetite (Mag), surrounded by anthophyllite (Ath); a 'ring' structure of anthophyllite is found isolated in a cordierite 'pocket' (ppl - X20).

PHOTOGRAPH 33

Fine-grained myrmekitic hercynite (Hc) replacing sillimanite (Sil) and possibly pseudomorphing the sheaths (ppl - X20).



Staurolite (St) as an anhedral prism rimmed by cordierite (Crd) and overgrown by garnet (Grt) (ppl - X40) (garnet = staurolite + cordierite).

PHOTOGRAPH 35

Euhedral staurolite (St) grain overgrown garnet (Grt); the Staurolite (St) is itself altered to hercynite (Hc) (ppl - X20).



-45-

(iii) Phosphates

Two phosphate minerals occur in the cordieriteanthophyllite gneiss zone, apatite and wagnerite. Apatite occurs as individual round to blocky grains, locally prismatic and clear to milky-white in thin-section. This minor phase is present in most sections and is 0.1 to 2 mm in size (photograph 36).

Wagnerite, a Mg-rich phosphate (Mg, Fe $^{2+})_2$ PO₄ F is orange to red in hand-specimen and yellow in thinsection. It is usually fine-grained, anhedral to subhedral with a vitreous to slightly resinous lustre. Some of the larger grains of wagnerite show cloudy areas at the centre due to minute inclusions of quartz and amphibole (photograph 37). Identification of these tiny inclusions proved impossible by ordinary petrographic methods but was possible by the S.E.M. (scanning electron microscopy) capabilities of the electron microprobe. Both the anthophyllite and quartz inclusions are rimmed by apatite, separating them from the enveloping wagnerite.

(iv) Rutile

Rutile is a characteristic minor constituent of the gneiss zone, ranging from 0.1 to 0.8 percent. It occurs as fine anhedral grains in various sections, with no reaction rims present.

(v) Magnetite

Magnetite shows anhedral outlines, is generally

Rounded to blocky, locally prismatic apatite (Ap) grains are mostly in the sillimanite (Sil) zone; these tend to form a network associated with coarse-grained biotite (Bt) and sillimanite (Sil) (ppl - X5).

PHOTOGRAPH 37

Wagnerite (Wg), a Mg-rich phosphate $(Mg,Fe^{2+})_{\pm}PO_{4}F)$, is generally fine-grained, with minute inclusions of quartz (Qtz) and anthophyllite (Ath), both rimmed by apatite (Ap) that separates them from the enveloping wagnerite (Wg) (reflected light).



<1 mm in size and present in all rock units. In its
association with cordierite, hercynite spinel is present. In
some areas, magnetite is altered to hematite, a secondary
oxidation process.</pre>

(vi) <u>Zircon</u>

Zircon occurs as minute grains with typically euhedral to subhedral habit and characteristic pleochroic haloes. These grains occur exclusively as inclusions in biotite (photograph 30).

(vii) <u>Sericite and Chlorite</u>

Minute shreds' of sericite rim and penetrate cordierite, sillimanite and occasionally anthophyllite (photograph 38).

Reaction rims of chlorite mantle garnet porphyroblasts and cordierite grains, separating them from the surrounding amphiboles. It also occurs as an alteration product after amphibole and/or biotite (photographs 39, 40). (viii) Sulfides

In the vicinity of Elken Lake, disseminated chalcopyrite, pyrrhotite and marcasite are found in association with the cordierite-anthophyllite gneisses. The small uneconomic sulfide showing occurs in thin stratiform layers that rarely attain a thickness greater than a few metres.

Chalcopyrite forms coarse-grained patches and is usually associated with pyrrhotite, which occurs as blebs in PHOTOGRAPH 38 Minute `shreds' of Sericite rimming and penetrating cordierite (Crd), Sillimanite (Sil) and anthophyllite (Ath) (ppl - X20).

PHOTOGRAPH 39

Chlorite (Chl) as an alteration product of anthophyllite (Ath) (ppl - X15).



Chlorite (Chl) as an alteration product of cordierite (Crd) (ppl - X20).

.



the latter (photograph 41). Pyrrhotite itself shows evidence of replacing marcasite (photograph 42). These sulfides form a fine-grained network which infills various fractures in the cordierite-anthophyllite gneiss (photograph 43).

Ś

PHOTOGRAPHS 41a & 41b

Coarse-grained patches of Chalcopyrite (Ccp) mainly associated with pyrrhotite (Po) (blebs in the chalcopyrite) (X20).





PHOTOGRAPH 42 Pyrrhotite (Po) replacing marcasite (Mc) (X20).

. .

PHOTOGRAPH 43

Fine-grained network of sulfides infilling fractures in the cordierite (Crd) - anthophyllite (Ath) gneiss (X20).



(IV) WHOLE ROCK CHEMISTRY

(i) <u>Chemical Parameters</u>

This section describes the use of chemical parameters to help indicate the origin of the rocks of the Star Lake - Elken Lake area, near Sherridon, Manitoba.

These rocks range from silica-poor (unit 3) to silica-rich (units 1 and 2), as seen in Appendix A. Units 1,2 and 3 typically exhibit low CaO (0.01-1.12 wt%) and alkalies (Na₂O: 0-0.58 wt%; K₂O: 0.02-2.2 wt%) and high amounts of Al₂O₃ (8.0-22.0 wt%), MgO (5.0-15.0 wt%) and FeO (5.0-17.0 wt%). Trace element patterns are fairly constant with some localized anomalies (SC-4-12). REE show characteristic basaltic trends with LREE depleted and HREE enriched values.

Chemical analyses are plotted on different types of diagrams to indicate the chemical fields occupied by igneous and sedimentary rocks, or to distinguish between rocks of igneous and sedimentary origins. The use of these diagrams to indicate the progenitors of the rocks involves the assumption that the metamorphism was essentially isochemical, except for the volatile constituents. The isochemical nature of the metamorphism is suggested by the consistency in chemical composition of the various units.

The rocks from the Star Lake - Elken Lake area have been divided into seven units for easy identification on most diagrams. They are:

-54-

- (1) " " unit 1 (quartz-garnet-anthophyllite),
- (2) " □ " unit 2 (garnet-anthophyllite),
- (3) "▲ " unit 3 (anthophyllite-cordierite),
- (4) " ¥ " quartz -biotite gneiss,
- (5) " ☆ " sillimanite zone,
- (6) " 举 " mineralized unit 2 zone, Elken Lake area
- (7) " Δ " unit 2 zone, Sherridon East orebody area

Figures 5 and 6 are after La Roche (1974) and are intended to discriminate between metamorphic, igneous and sedimentary rocks. Figure 5, the' total rock triangle' shows the greater scatter of the two diagrams. Most of the samples from units 1, 2 and 3 plot near the arkose field and felsic termination of the igneous trend, as do the sillimanite and quartz-biotite gneisses. Figure 6, the `silico-aluminate triangle', shows the most consistency for all the data. Samples representing all units plot in the left part of the diagram, the Mg-rich and common-clay regions. The cordierite- anthophyllite gneisses of units 2 and 3 indicate minor Mg enrichment relative to any `normal' igneous rock. Possible basalt/arkose mixture relationships exist which seem to indicate progenetic contamination (discussed in the 'Discussion of Geochemistry' section).

Many authors have used Niggli diagrams to show trends and positions of recent igneous and sedimentary rocks in comparison with high-grade metamorphic rocks to help classify them as sedimentary or igneous in character (eg.

-55-





•

Fe203+Ti02+Ca0

A1203

-56-

K20

1



Na20

-57-

Leake, 1964, 1969; Van de Kamp, 1968, 1969, 1970; Van de Kamp, Leak and Senior, 1976; Stamatelopoulou - Seymour and Maclean, 1977).

Figure 7, a plot of Niggli c vs. Niggli mg., should show a strong trend parallel to the mg ordinate for volcanic rocks, and a large scatter with respect to c for sedimentary rocks. A noticeable parallel trend to the mg ordinate is found for samples of units 2 and 3, suggesting possible development (alteration or weathering) from a volcanic parent. Samples from unit 1 seem to scatter somewhat with respect to the Niggli c axis.

Figure 8 is a plot of (Na_2O+K_2O) vs. $(K_2O/(Na_2O+K_2O))$ K₂O))x100 (Hughes, 1973) which discriminates relatively unaltered igneous rocks (igneous spectrum) from altered igneous rocks (spilites) and sediments. The rocks interpreted as having been formed from a volcanic parent (Figure 8) all plot outside the igneous spectrum, with the exception of one sample from unit 1 and one quartz-biotite gneiss sample. The occurrence of the data outside the igneous spectrum may be the result of contamination by terrigenous material, (rich in K_2O - as total alkalies increase K₂O becomes dominant) or alteration of a volcanic ash or tuff (derived from felsic to intermediate volcanics or related sedimentary rocks (greywakes or transported This contamination can be substantiated with tuffs)). similar mixtures (basaltic/arkose) in the total rock

-58-



FIGURE 7: NIGGLI MG VERBES NIGGLI C for Star Lake rocks. Legend as in Figure 5. From Leake (1964) and Van de Kamp (1968).



FIGURE 8: Na2O+K2O VERSES (K2D/Na2O+K2O)100 for Star Lake rocks. Legend as in Figure 5. Diagram in wt% oxide. From Hughes (1973). triangle and silicoaluminate triangle.

Van de Kamp et al. (1976) used a plot of Niggli mg vs. Niggli si to differentiate between sedimentary and igneous precursors for metamorphosed gneisses. Because of the concentration of fine mafic minerals in the high SiO_2 fraction of sands, positive slopes represent sediments, whereas negative slopes represent igneous rocks. Figure 9, a plot of Niggli mg vs. Niggli si, shows a positive distribution for units 1 and 2 and a negative distribution for unit 3. This diagram seems to suggest that initial precursor rocks were formed by mixed sedimentary and igneous processes resulting in variable mafic mineral concentrations.

Leake (1964) used a plot of Niggli k as a function of Niggli mg to distinguish amphibolites of sedimentary origin from those of igneous origin. Goetz (1980) and Gale et al. (1980) showed Sherridon host rocks and unmineralized gneisses to be distributed between the igneous and pelitic fields. Figure 10, a plot of Niggli k vs. Niggli mg, shows similar mixing effects and considerable scatter along the k axis with some slight mg variations. This plot again suggests mixed precursors that were partly sedimentary (greywacke-clay- metasediments) and partly igneous (arkose/basalt mix).

Van de Kamp (1970) and Leake (1969) used a plot of Niggli al-alk against Niggli c to separate metamorphic rocks

-61-



FIGURE 9: NIGGLI MG VERSES NIGGLI SI for Star Lake rocks. Legend as in Figure 5. From Van de Kamp et al. (1975).

÷.



FIGURE 10; NIGGLI K VERSES NIGGLI MG for Star Lake rocks. Legend as in Figure S. From Leake (1964).

· · · ·

of igneous origin from those of sedimentary origin. The value al-alk indicates available alumina present in excess of that used to produce all micaceous minerals and feldspar. Small variations of al-alk with composition produce a broad igneous field (of positive slope) which overlaps other fields. Figure 11, a plot of Niggli al-alk versus Niggli c, indicates how samples from unit 1 plot mainly in the 'normal clay' region, except for one which plots in the intermediate igneous rock field. Samples from units 2 and 3 lie in the normal clay-greywacke fields. This plot suggests similar precursor characteristics to what can be observed in Figure 10.

In order to distinguish between igneous and sedimentary precursors to metamorphic rocks, a discriminant plot of Niggli al+alk versus Niggli c was developed by Van de Kamp (1976). Niggli al+alk decreases and Niggli c increases with an increase in ferromagnesium minerals and calcic plagioclase in going from felsic to more mafic igneous rocks (narrow field with negative slope). A shallow negatively sloping sedimentary field is also present, corresponding to sedimentary rocks of shale and greywacke composition. Figure 12 shows that all Star Lake-Elken Lake samples are low in Ca and have low Niggli al+alk values. This may be due to a process that resulted in alkali depletion, but may be an artifact of dilution by fm related

-64-



FIGURE 11: NIGELI AL-ALK VERSES NIGELI C for Star Lake rocks. Legend as in Figure 5. From Van de Kamp (1970) and Leake (1969).

.



FIGURE 12: NIGGLI AL+ALK VERSES NIGGLI C for Star Lake rocks. Legend as in Figure 3. From Van de Kamp (1970).

.
to the summation of fm, c, al and alk to 100. Some of the samples also plot in the igneous and sedimentary fields, which again indicates mixture.

The use of a discriminant plot of Niggli ti vs. Niggli al+alk offers some aid in the determination of igneous and to some degree, sedimentary progenitors bу virtue of the supposed immobility of titanium (Finlow-Bates and Stumpfl, 1981; Roberts and Reardon, 1978). Nigali al+alk is used along the X-coordinate to produce a broad igneous spectrum with felsic rocks of high al+alk values and mafic rocks of low al +alk values (high fm + c). Figure 13 shows that most samples fall in the low al+alk region, straddling the `lower limit' of the igneous field. Some of these gneisses also show intermediate-high Niggli ti values, characteristic of chemically weathered basalts. Goetz (1980) has reported such an instance in the Flin Flon area, where high titanium levels occur in a regolith on an old basaltic land surface.

In general, the chemistry of these rocks is variable throughout the stratigraphic horizon at Star Lake – Elken Lake. Many of the discriminant diagrams presented seem to suggest sedimentary and/or volcanic mixed progenitors for these rocks. This will be discussed further in the following section.

Figure 14 shows the range of REE (rare earth elements) found in the Star Lake - Elken Lake gneisses. In

-67-



-89-

•



FIGURE 14: The range of REE content found in the Star Lake - Elkon Lake gneisses (shaded) containing small negative to positive Eu anomalies. One representative basaltic pattern (dashed line) is also included to illustrate the similarity of REE contents.

all three mappable units and associated quartz-rich gneisses, LREE depletions and HREE enrichments are characteristic. According to Menzies et al. (1979), diagenesis and metamorphism have little relative effect on REE patterns. The REE elements, therefore, seem to be truly immobile. These authors also state that although some fractionation of REE may take place within weathering profiles, there seems to be no selective loss of REE during weathering. If we assume that REE levels were essentially unaffected we can, therefore, associate these trends with an igneous basaltic source (the LREE depleted pattern is indicative of a basaltic source (Taylor and McLennan, 1985)). Menzies et al. (1979) have also found that the most significant feature for metasedimentary rocks from early Proterozoic high-grade terranes (as found in the Star Lake -Elken Lake and surrounding areas) is the wide range in REE patterns ranging from nearly flat trends to LREE depleted trends. These trace element trends, therefore, provide added information in support of a mixed source for these anthophyllite-cordierite rocks.

-70-

(ii) Interpretation of Whole-Rock Geochemistry

La Roche's (1974) silico-aluminate and total rock triangle, and some of the plots of Niggli values seem to produce distinct fields for common igneous and sedimentary rocks. Star Lake cordierite-anthophyllite bearing gneisses and sillimanite rocks often plot outside these well-defined fields, and also straddle boundaries between major igneous trends and sedimentary rocks.

All cordierite-anthophyllite rocks at Star Lake -Elken Lake plot near or at the MgO apex of the La Roche (1974) `silicoaluminate' diagram. This field is also occupied by intensely altered rocks from the cores of the alteration pipes at the Millenbach Mine, Noranda area, Quebec (MacGeehan and MacLean, 1980, Figure 15). Both unit 2 and 3 gneissic rocks plot closer to the MgO apex compared to the unit 1 rocks, possibly indicating increasing alteration from unit 1 to unit 3 respectively. La Roche's (1974) 'total rock triangle' shows similar features, with the quartz-garnet-anthophyllite rocks from unit 1 plotting in the more felsic termination grading into cordieriteanthophyllite rocks of units 2 and 3 towards the Mg clay fields.

Some of the diagrams seem to suggest partly sedimentary and igneous (volcanic) precursors for these cordierite-anthophyllite rocks at Star Lake. Hugh's (1973) plot of (Na_20+K_20) vs. $(K_20/(Na_20+K_20))100$ shows that most

-71-

K20



MgO



of the samples plot outside the igneous spectrum, possibly resulting from the contamination of volcanic rocks by terrigenous material or alteration process. The plot of Niggli si vs. mg (Van de Kamp et al., 1976) differentiates between sedimentary and igneous precursors for metamorphosed gneisses, due to the concentration of fine mafic minerals in the high SiO₂ fraction of sands ((+) slope for sands, (-) slope for igneous rocks). This diagram also suggests that the process which formed the precursors of the Star Lake gneiss were partly sedimentary and partly igneous (positive and negative slopes present). Leake's (1964) Niggli k vs. mg plot shows that the data straddle both the igneous and pelitic fields, also suggesting precursors that were partly sedimentary and partly igneous (basaltic) in nature.

Due to greater MgO and FeO contents, cordieriteanthophyllite rocks have lower values of al-alk and al+alk than the associated sillimanite-bearing rocks. However, they do have similar low values of Niggli c and, therefore, plot in fields which approach the al-alk and al+alk axes. These fields lie in the subvolcanic to pelitic/greywacke sediment or subvolcanic felsic to intermediate volcanic rock regions, with unit 2 and 3 rocks plotting in the lower al+ alk fields. These patterns closely resemble those at the Millenbach Mine, Quebec, with increasing alteration intensity towards the core of the pipe corresponding (by analogy) to rocks from units 2 and 3 (Figure 16).

-73-





Niggli ti values in the cordierite-anthophyllite rocks at all locations have similar values to altered rocks from the cores of the alteration zones at the Millenbach and Mattagami Lake Mines (Figure 17). Most of these gneisses have Niggli ti values and TiO₂ levels (Appendix A & C) below those of the most felsic igneous rocks. If these low titanium levels reflect the original precursor levels and titanium is assumed to be immobile (Finlow-Bates and Stumpfl, 1981; Roberts and Reardon, 1978), then some dilution of original titanium with a low or zero titanium component has occurred. This component may have been added to the gneisses by co-deposition of a possibly altered felsic volcaniclastic or greywacke sedimentary rock or by some post-depositional hydrothermal alteration.

Another characteristic feature of these Ca and Na depleted gneisses is the lack of feldspars (plagioclase and K feldspar) associated with the sillimanite. The absence of plagioclase in these rocks suggests that it was probably not present prior to the metamorphism. The lack of K feldspars could also be explained by a dealkalization reaction followed by a silicate-water reaction; $3 \text{ KAlSi}_{3}O_{6} + 2H^{+} = \text{ KAl}_{3}\text{Si}_{3}O_{10}(OH)_{2} + 6 \text{ Si}O_{2} + 2K^{+}$ K feldspar K mica guartz

 $2 \text{ KAl}_3 \text{Si}_3 0_{10} (\text{OH})_2 + 3 \text{ H}_2 0 + 2\text{H}^+ = 3 \text{ Al}_2 \text{Si}_2 0_3 (\text{OH})_4 + 2\text{K}^+$

K mica

kaolinite

-75-



FIGURE 17: PLOT OF NIGGLI TI VERSES NIGGLI AL+ALK Snowing the general increase in hydrothermally altered rocks towards the core of the alteration zone at the Nillenbach Thine, Noranda area, Guebuc (circle) and samples representing the Mattagami Lake Mine, Guebuc (squares) related to Star Lake rocks (arrow indicates incruasing intensity of alteration towards the core of the zone). From MacGeehan and MacLean (1980).

.

These reactions could possibly explain the production of aluminosilicate layers beneath the more amphibole-rich sections of unit 2. But, as the sillimanite-bearing rocks plot as common felsic to intermediate volcanic rocks or greywackes, these may be the metamorphic equivalents of altered and/or chemically modified versions of these felsic to intermediate volcanics or greywackes.

(\vee) MINERAL CHEMISTRY

Analyses of coexisting minerals from the cordierite orthoamphibole rocks at Star Lake - Elken Lake are listed in Appendix D.

(i) Orthoamphiboles

Two orthoamphiboles occur in the Star Lake - Elken Lake units, anthophyllite and gedrite.

Variation in the amphibole compositions can be described in terms of end-member anthophyllite $[Mg_7Si_{B}O_{22}$ $(OH)_2]$ and three substitution mechanisms : Fe<=>Mg (Fe - Mg exchange), Al~1+Al1<<=>Mg+Si (tschermak exchange) and Na+Al1<<=>Si (edenite exchange); other substitutions such as those involving Mn, Ti, Fe³⁺ or Ca are of minor importance (Spear, 1980). Evaluation of the analyses indicates that the anthophyllite-gedrite series is a solid solution between two end compositions

 $R^{2+}_{2}R^{2+}_{3}Si_{0}O_{22}(OH)_{2}$ and

 $Na_{0.5}R_2^{2+}(R_{3.5}^{2+}R_{1.5}^{3+})Si_6Al_2O_{22}(OH)_2$

where R²⁺=Mg, Fe²⁺, Mn²⁺, Ca

 $R^{3+=A1}$, Fe^{3+} , $(Ti_{1/2}^{4+}+Fe_{1/2}^{2+})$

There is apparently complete solid solution in the anthophyllite-gedrite series at high temperature, as shown in specimens from the Star Lake - Elken Lake area. On cooling, members with intermediate Al and Na contents (Figure 18) exsolve to an anthophyllite-gedrite intergrowth. Most intergrowths are optically homogeneous but show strong

-78-





blue, green or yellow schiller effect believed to be due to submicroscopic exsolution.

As shown in Figure 19, amphiboles from unit 3 (associated with cordierite) show a slight Fe^{2+} enrichment compared to those in unit 2, which are mainly associated with garnets. Most of the amphiboles associated with sulfides from Elken Lake have slightly lower Fe^{2+} contents. Figure 20 shows similar Fe^{2+} enrichments and a restricted field of the Star Lake orthoamphiboles. The compositions of the orthoamphiboles generally show enrichment in Mg relative to the host rocks (Figure 21).

(ii) <u>Garnet</u>

In all samples, garnet is essentially almandine – pyrope (Figures 22 and 23). In most cases, garnets are zoned from magnesium cores (+/- calcium) to rims richer in iron (+/- manganese) (Appendix D). This zoning is rather erratic due to later fracturing, possible resorption later in the metamorphic history or due to variation of growth rate in different directions during formation. MnO zoning, although minor, has been related to the marked affinity of Mn for garnet in metamorphosed argillaceous rocks (Harte and Henley, 1966). No apparent correlation exists between the garnet and host rock compositions.

(iii) <u>Cordierite</u>

In most samples, cordierite is associated with -80-



- ÷.

FIGURE 19: FEP* VERSES NA, showing the amphiboles (anthophyllite and gedrite) from the Star Lake strata. Filled circle - unit 1, filled square - unit 2, filled triangle - unit 3, filled star - quartz-biotite gneiss, asterisk - sillimanite zone, star - Elken Lake unit 2, and triangle - Shorridon unit 3.



FIGURE 20: Fe=+ VERSES Mg, showing the amphiboles (anthophyllite and gedrite) from the Star Lake strata. Filled circle - unit 1, filled square - unit 2, filled triangle - unit 3, filled star - querts-biblite gness, asteriss - sillimanite zone, star - Elken Lake unit 2, and triangly - Sherridon unit 3.





Spessartine (Mn3Al2Si3012)

Pyrope (Mg3Al2Si3012)

FIGURE 22: ALMANDINE - SPESSARTINE - PYROPE TERNARY PLOT, showing the garnets from the Star Lake strata. Filled circle - unit 1, filled square - unit 2, filled triangle - unit 3, filled star - quartz-biotite gneiss, asterisk - sillimanite zone, star - Elken Lake unit 2, and triangle - Sherridon unit 3.

-84-



Pyrope (Mg3Al2Si3012)

Grossular (Ca3Al2Si3012)

Spessartine (Mn3Al2Si3012)

FIGURE 23: PYROPE - GROSSULAR - SPESSARTINE TERNARY PLOT, showing the garnets from the Star Lake strata. Filled circle - unit 1, filled square - unit 2, filled triangle - unit 3, filled star - quartz-biotite gneiss, asterisk - sillimanite zone, star - Elken Lake unit 2, and triangle - Sherridon unit 3. orthoamphiboles. In any individual specimen of a particular rock unit, the cordierites have fairly constant composition with no evidence of zoning. In general, the magnesium content is greater in unit 2, and the iron content is greater in units 1 and 3. The compositions of cordierite do not follow those of the host rocks but show a marked enrichment of Mg relative to the host rocks.

Cordierite is an important indicator mineral in progressive metamorphism. It has very low density for a ferromagnesian mineral and it is regarded as being characteristic of low-pressure terranes. It often occurs with sillimanite as the characteristic Al₂SiO₅ polymorph.

(iv) <u>Sillimanite</u>

There is no compositional zoning, except for small amounts of Fe³⁺ replacing Al.

(v) <u>Staurolite</u>

Staurolite is a very minor phase in the gneissic rocks from Star Lake. It is fairly uniform, including sample ML49 associated with the Elken Lake mineralization. Sample ML01 shows no indication of zoning of any kind.

(vi) <u>Biotite</u>

Biotite is a minor phase in most samples. Most biotites have moderate TiO_2 (0.5-2.5 wt.%), but some are enriched (TiO_2 reaching 3.5 wt.%, Appendix D). Magnesiumrich biotites occur in unit 3, decreasing in Mg content from unit 2 to unit 1 respectively, probably due to changes in

-86-

extensive conditions that alter Fe-Mg element partitioning preferences. The H_2O in the biotite analyses are determined by stoichiometry, and may be too high, causing the analyses totals to be somewhat erratic. The composition of biotites generally follow that of the host rocks, with Mg and Al enriched in the biotite relative to the host rock. The poor correlation is a result of the fact that, in addition to rock chemistry, biotite composition is dependent on the physical conditions of crystallization, the coexisting minerals, and the extent of secondary alteration.

(vii) <u>Hercynite</u>

Spinels from Star Lake and the Sherridon anthophyllite-cordierite gneisses are mainly hercynitic (Fe^{2+}) in composition (Figure 24). Zinc contents vary from 0 to 11.78 wt.% in rocks from unit 3 (anthophyllitecordierite). The increase in Zn is marked by decreases in Fe^{2+} (Figure 25) and Mg (Figure 26). Samples from Sherridon (ML93) show similar patterns for Mg-Zn contents but have evidence of increasing Fe^{2+} with increasing Zn (cordierites associated with these samples have lower Fe^{2+} contents). The Zn-Mg and Zn-Fe relationships are problably due to hercynite/ cordierite partitioning effects. In spinelbearing rocks the compositions of the spinel reflect bulk composition.

(viii) <u>Phosphates</u>

Two phosphates occur as very minor phases in these



Gahnite (ZnAl2O4)

Hercynite (FeAl204)

FIGURE 24: SPINEL - GAHNITE - HERCYNITE TERNARY PLOT, showing the hercynite spinels from the Star Lake strata. Filled circle - unit 1, filled square - unit 2, filled triangle - unit 3, filled star - quartz-biotite gneiss, asterisk - sillimanite zone, star - Elken Lake unit 2, and triangle - Sherridon unit 3.









.

gneisses. Apatite is closely associated with the sillimanite zone and is compositionally homogeneous throughout the unit. In areas of extreme calcium depletion and magnesium enrichment, wagnerite $((Mg,Fe^{2+})_2PO_4F)$ is found. This phosphate is fairly constant in composition throughout the samples.

(ix) <u>Chlorite</u>

Chlorite is found in unit 3 samples, showing evidence of later retrograde metamorphism. The composition is fairly consistent throughout all grains (Appendix D).

(VI) MINERAL REACTIONS

Nearly uniform conditions of high-grade metamorphism exist throughout the Kisseynew gneiss belt, decreasing with steep metamorphic gradient into the mediumand low-grade rocks of the adjacent volcanic belts (Bailes and McRitchie, 1978). The rocks of the Star Lake - Elken Lake area contain mineral assemblages characteristic of the amphibolite facies. Retrograde metamorphism is also common, with the presence of chlorite alteration of ferromagnesian minerals.

There is a considerable amount of experimental data pertinent to the system MgO-FeO-Al₂O₃-SiO₂-H₂O, some of which is relevant to the present study. A representative petrogenetic grid for quartz-bearing assemblages in this system from Star Lake is shown in Figure 27. Addition of Mg to this system will generally shift invariant points to higher pressures (Fe-Mg partitioning), whereas a decrease in X_{H20} will shift them to lower temperatures. Within the constraints of this system, cordierite is limited to low temperatures, whereas orthoamphibole and staurolite both become unstable with increasing temperature. The cordierite - amphibole rocks of unit 3 are stable in a field defined at high temperature by the reaction:

orthoamphibole = garnet + cordierite + quartz

-92-





and at high pressure by the reactions:

cordierite = orthoamphibole + staurolite + quartz cordierite = orthoamphibole + sillimanite + quartz

This petrogenetic grid can only indicate approximate conditions of metamorphism. An attempt to account for the shift in P-T conditions resulting from Fe-Mg partitioning or substitutions may be made with the analytical data on phase compositions.

(i) Geothermometers

(a)<u>Cordierite-garnet</u>

The Fe-Mg partitioning between cordierite and garnet observed both in natural assemblages (Thompson, 1976) and experimental runs (Hensen and Green, 1973; Holdaway and Lee, 1977) has been used for temperature calibration. The equilibrium of the exchange reaction

 $2Mg_3Al_2Si_3O_{12} + 3Fe_2Al_4Si_3O_{10}(OH) =$

 $2Fe_3Al_2Si_3O_{12} + 3Mg_2Al_4Si_3O_{10}(OH)$

Mg-Grt + Fe-Crd = Fe-Grt + Mg-Crd is represented by the distribution coefficient

$$K_{\mathbf{D}}^{\mathbf{Grt} - \mathbf{Crd}} = (\underline{X \ \mathbf{Grt}}_{\mathbf{Fe}}) (\underline{X \ \mathbf{crd}}_{\mathbf{Mq}})$$
$$(\underline{X \ \mathbf{grt}}_{\mathbf{Mq}}) (\underline{X \ \mathbf{crd}}_{\mathbf{Fe}})$$

This coefficient is an exponential function of T, which may

-94-

be combined with the P-T position of the reaction

```
5FesAl4Si6022(OH)2 + 3Fe4Al18Si8046(OH)2
```

orthoamphibole + staurolite

= 10.8Fe₂Al₄Si₅O₁₆O.5H₂O + 15.4FeAl₂O₄ + 2.6H₂O cordierite + spinel

This may account for a shift in P-T conditions resulting from Mg-Fe substitution.

T = 2725 + 0.0155P + - 50 k

 $\ln k_{p} + 0.896$ (Thompson, 1976)

The equilibrium temperatures for the representative garnet and cordierite rims, obtained for samples SC2-9 (unit 1) and SC4-12 (unit 3) are 658°C and 655°C respectively for 6 Kbar, 661°C and 658°C respectively for 6.5 Kbar and 663°C and 660°C respectively for 7 Kbar.

(b)<u>Biotite-garnet</u>

The Fe-Mg exchange reaction between biotite and garnet is represented by the distribution coefficient:

 $K_{D} \text{ Grt} - \text{Bt} = \frac{(X \text{ Grt}_{Fe}) (X \text{ Bt}_{Me})}{(X \text{ Grt}_{Me}) (X \text{ Bt}_{Fe})}$

for the exchange reaction

garnet + biotite

= $Fe_3Al_2Si_3O_{12}$ + $KMg_3AlSi_3O_{10}(OH)_2$

garnet + biotite

For this reaction, Thompson (1976) has shown that the following relationship holds:

$$T = 2740 + 0.0234P + - 50 K$$

$$\ln K + 1.56$$

Equilibrium temperatures for the representative garnet and biotite rims, obtained from sample SC2-1 (unit 1) and SC5-12 (unit 2) are 665°C at 6 Kbar, 669°C at 6.5 Kbar and 673°C at 7 Kbar for both samples. The application of this thermometer is made easier by the lack of zoning in the garnets from both these samples.

Both cordierite-garnet and biotite-garnet derived temperatures show that metamorphic temperatures were fairly constant throughout the Star Lake - Elken Lake area.

(ii) <u>Geobarometers</u>

Despite the variety of their mineral assemblages, cordierite-anthophyllite rocks are not particularly suited to derive pressure estimates from conventional geobarometers, due to the structural and chemical complexities of the orthoamphiboles. Seifert and Schumacher (1986) have shown that the assemblage cordierite-(Mg-Zn-Fe 2+) aluminate spinel-quartz can be successful as a geobarometer.

-96-

In the system $FeAl_2O_4-SiO_2$, the phases spinel (spl), cordierite (crd) and quartz (qtz) are related by the univariant reaction:

 $Fe_2Al_4Si_5O_{10}(Fe-crd) = 2 FeAl_2O_4(Hc) + 5 SiO_2(qtz)$

The equilibrium constant for this reaction is

$$K_2 = \frac{(\chi = p_{F_0})^2 (\chi = \chi = g_{102})}{(\chi = r_{e_0})}$$

Considering quartz as a pure phase and assuming ideality for the spinels

$$K_2 = \frac{(X \text{ spl}_{Fe})^2}{(X \text{ crd}_{Fe})^2}$$
 ie. $X^{qtz}_{siO2}=1$

Equilibrium pressures for the representative spinels and cordierites from samples SC5-35 (unit 3) and ML33 (unit 3) range from 6.6 Kb at 650°C to 6.8 Kb at 700°C (as shown in Figure 27).

In the system $MgAl_2O_4$ -SiO₂, the phases spinel (spl), cordierite (crd) and quartz (qtz) are related by the univariant reaction:

 $Mg_2Al_4Si_8O_{18} (Mg-crd) = 2 MgAl_2O_4 (spl) + 5 SiO_2 (qtz)$

and
$$K_1 = \frac{(\chi \approx p_1 m_0)_2}{(\chi \approx m_0)_2} (\chi q t_{E102})^3$$

where $X^{qtz}_{sto2} = 1$ where $X^{crd}_{Mq} = 1 - X^{crd}_{Fq}$.

Samples SC5-35 and ML33 with equilibrium constant values of $K_1 = 0.35$ (log $K_1 = -0.45$) and $K_1 = 0.34$ (log $K_1 = -0.47$) respectively, give estimates of equilibrium pressures of 7.1 Kb at 650°C to 7.3 Kb at 700°C (Figure 28). These later estimates may be lower than true values due to the stabilizing effect of H₂O in cordierite, which was not considered in the calculations used to derive K1. According to Mirwald (1982), hydrous cordierites would plot at +1.5 Kb compared to anhydrous cordierite (Seifert and Schumacher, 1986). The addition of Mg to the system (as occurs at Star Lake) generally shifts invariant points to similar higher pressure regimes (Figure 29). It is therefore obvious that added components (FeO) to the initial dry system MgO-Al_2O_3-SiO₂, reflect variations as a function of pressure, temperature and bulk rock chemistry.

-98-



FIGURE 28: Log K1 for the cordierite - spinel - quartz equilibrium as a function of pressure for different temperatures. From Seifert and Schumacher (1986).





-100-

(VII) DISCUSSION AND CONCLUSION

(i) Possible Cordierite-Anthophyllite Rock Origins

The anthophyllite-bearing rocks of the Sherridon Group form discontinuous layers that have both gradational and sharp contacts with the lithic arenites which enclose them. Hypotheses for the origin of this rock type have been numerous.

The classic study is that of the Orijarvi area (Eskola ,1914). He indicated that the composition of cordierite-anthophyllite bearing rocks does not correspond to that of any sedimentary or igneous rock. For this reason, their origin was attributed to contact metasomatism involving relative enrichment in iron and magnesium and depletion in calcium and alkalis. Later, Wegmann and Kranck (1931) also suggested a metasomatic origin related to regional metamorphism. Tuominen and Mikkola (1950) further suggested the that in Orijarvi area, cordieriteanthophyllite rocks represented isochemically metamorphosed quartz-chlorite rocks.

In the case of the Star Lake - Elken Lake area, Robertson (1953) described an anthophyllite band of wide distribution at a constant stratigraphic horizon. He stated that the character of the mineral assemblage suggests Fe-Mg metasomatism of an aluminum-rich sediment or tuff. Robertson favoured the tuff progenitor because of the wide lateral development and apparent continuity of the band

-101-

(normal sediments in these types of rocks would probably be less continuous and would not offer such a permeable medium for ascending solutions).

Wilkinson (1976) noted that this assemblage was very similar to the high-temperature assemblages found in metamorphosed volcanic alteration pipes. In alteration pipes, this is the result of the conversion of the original assemblage to one consisting almost exclusively of Fe-Mg chlorite, similar to the Shakami Mine occurrence in Japan (Shirozu, 1974).

Goetz and Froese (1982) proposed that the cordierite-anthophyllite rocks required chloritic rocks as precursors; these could be produced by hydrothermal "Chloritization is a common alteration. proccess in hydrothermal alteration, which involves the selective dissolution and removal of calcium and sodium, producing a rock depleted in these elements and enriched in aluminum. Thus , mafic rocks could acquire the composition of a quartz -chlorite mixture and felsic rocks could become chloritized by a combination of some aluminum with iron and magnesium from the hydrothermal solution." (Goetz and Froese, 1982). They also suggest that the protolith of the cordieriteanthophyllite rocks might have been detritus nf hydrothermally altered rocks, transported to their present site either as a fine suspension in a brine (Wilkinson, 1976; Goetz, 1980) or as a sediment gravity flow (Middleton

-102-
and Hampton, 1976). In the search for chloritic precursors of anthophyllite-bearing rocks, two main processes of chloritization have been considered; hydrothermal alteration related to mineralization (Froese, 1969; Goetz and Froese, 1982) and chloritization as part of the complex alteration leading to the formation of spilites (Vallance, 1967, 1969).

In the case of the cordierite-anthophyllite rocks at the Star Lake - Elken Lake area, the available evidence seems to point to an altered basaltic protolith with mixed injections of sedimentary material. The hydrothermal alteration of basalt by seawater thus forming alteration pipes is the most likely process leading up to the present lithologies. These alteration pipes probably formed due to rapid heating of large volumes of cold undepleted seawater in the vicinity of vent areas. The heated seawater then reacted with the basaltic rocks of the crust, recrystallizing them and modifying the chemistry of both rocks and solution and eventually exiting the rocks as warm or hot springs. In the process, with a very large water to rock volume ratio, chemical components were transferred from the seawater into the altered crust and vice versa. Wolery and Sleep (1976) have suggested that water/rock ratios in submarine hydrothermal systems may be as high as 100 (total mass of water which has passed through the system during its lifetime divided by the mass of rock within the system that has been altered).

-103-

In general, upon heating, the cold seawater lost its magnesium very rapidly to form chlorite and smectite mixtures in varying proportions. Seyfried and Mottl (1982) have shown that, with very large water to rock ratios and elevated temperatures, Mg was removed from solution and therefore supplied to the crust whereas Na and Ca had been leached from the basaltic rocks. As sodium is very depleted in these alteration zones, this would reflect the virtual total destruction of feldspars. As the Mg supply exceeds the amount of Mg that the available rock can absorb, the concentration of Mg increases in solution. Under these conditions, the solution becomes acid and rich in heavy metals promoting the formation of sulfides in isolated localities (Elken Lake and Sherridon deposits may be typical examples). The only silicates which could form from basalt and coexist stably with such a solution are quartz, smectite and chlorite. This results from the fact that nearly all the Ca, Na, K, Ba, Mn, Cu and Zn are leached from the altered silicates and reside in the acid solution (Seyfried and Mottl, 1982; Seyfried et al., 1988).

The size and vertical distribution of a chlorite pipe is probably a function of the permeability of the crust. This permeability may have been the result of highly fractured zones associated with major faults or a function of fragmental accumulations being altered. These compositionally variable alteration zones most likely

-104-

reflect the changing composition of the Star Lake - Elken Lake units, following individual periods of tectonism and metamorphism. Following this accumulation of altered material, high-grade regional metamorphism took place for a long period of time, promoting the growth of coarse crystals of some of the minerals.

It could possibly be inferred that a progressive alteration sequence from unit 1 to unit 3 rocks, from Star Lake, represent a progressive trend towards chlorite rich alteration pipe cores.

(ii) Tectonic and Petrologic History

The Kisseynew meta-sedimentary belt lies between two volcanic belts from which its sediments were derived (Bailes, 1971). Dickinson (1974) stated that a much larger amount of sediments can be derived from a volcanic belt than the amount of volcanic rocks preserved in the belt. This is true of the Kisseynew belt, which contains at least 10 times the amount of sediments as there are volcanics preserved in the Lynn Lake and Flin Flon belts. Due to the volcanic rock and unit size distribution, this setting resembles a modern back-arc basin, suggesting tectonic activity and volcanic alteration. From these relations, it can therefore be inferred that the volcanic and sedimentary belts of central Manitoba and Saskatchewan could have been created by the process of plate tectonics (Karig, 1972).

-105-

REFERENCES

- Ashton , K. E. and Froese, E. (1988). GS-18 Could the Sherridon Group at Sherridon be a high-grade equivalent of the Amisk Group? Report of field activities 1988, Minerals division, Manitoba Energy and Mines.
- Ashton, K. E., Wheatly, K. J., Moser, D., Paul, D. and Wilcox, K. H. (1986). The Kisseynew gneisses of Saskatchewan: Update; in Summary of Investigations 1986; Saskatchewan Geological Survey. Miscellaneous Report 86-4, p. 145-148.
- Ashton, K. E., Wilcox, K. H., Wheatly, K. J., Paul, D. and de Tombe, J. (1987). The boundary zone between the Flin Flon domain, Kisseynew gneisses and Hanson Lake block, northern Saskatchewan; in Summary of Investigations 1987; Saskatchewan Geological Survey, Miscellaneous Report 87-4, p. 131-134.
- Bailes, A.H. (1971). Preliminary compilation of the geology of the Snow Lake - Flin Flon - Sherridon area; Manitoba Mines Branch, Geological Paper 1/71.
- Bailes, A.H. and McRitchie, W.D.(1978). The transition from low to high grade metamorphism in the Kisseynew sedimentary gneiss belt, Manitoba; <u>in</u> Metamorphism in the Canadian Shield, Geol. Surv. Can. Paper 78-10, p.155-178.
- Bateman, J.D. and Harrison, J.M.(1946). Sherridon, Manitoba; Geological Survey of Canada, Map 862A.
- Bruce, E.L.(1918). Amisk Athapapuskow Lake district; Geological Survey of Canada, Memoir 105.
- Dickinson, W.R.(1974). Sedimentation within and beside ancient and modern magmatic arcs: <u>in</u> Modern and ancient geosynclinal sedimentation, <u>ed</u>. R.H.Dott and R.M.Shaver. Soc. Econ. Paleon. Miner., Spec. Pub. 19, p.230-239.
- Eskola, P.(1914). On the petrology of the Orijarvi region in southwestern Finland; Bulletin de la Commission Geol. de Finlande, No.40.
- Finlow Bates, T. and Stumpfl, E.F.(1981). The behaviour of so called immobile elements in hydrothermally altered rocks associated with submarine exhalative ore deposits; Mineral.Deposita, v.16,pp.319 - 328.

-106-

- Froese, E.(1969). Metamorphic rocks of the Coronation mine and surrounding area; Geological Survey of Canada, Paper 68-5, p.55 - 77.
- Froese, E. and Goetz, P.A.(1981). Geology of the Sherridon
 Group in the vicinity of Sherridon, Manitoba;
 Geological Survey of Canada, Paper 80 21.
- Froese, E.(1985). Anthophyllite bearing rocks in the Flin Flon - Sherridon area, Manitoba; <u>in</u> Current Research Part B, Geological Survey of Canada, Paper 85 - 1B, pp.541 - 544.
- Gale, G.H., Baldwin, D.A. and Koo, J.(1980). A geological evaluation of Precambrian massive sulfide potential in Manitoba; Manitoba Dept. of Energy and Mines, Minerals Resources Div., Economic Geology Report ER 79-1.
- Gittos, M.F., Lorimer, G.W. and Champness, P.E.(1976). The phase distribution in some exsolved amphiboles. In Wenk, H.R. et al.(eds) <u>Electron microscopy in</u> <u>Mineralogy</u>, 238 - 247. Berlin:Springer - Verlag.
- Goetz, P.A.(1980). Depositional environment of the Sherridon Group and related mineral deposits near Sherridon, Manitoba; unpublished Ph.D. thesis, Carleton University, Ottawa.
- Goetz, P.A. and Froese, E.(1982). The Sherritt Gordon massive sulfide deposit; Geological Association of Canada, Special Paper 25, pp.557 - 569.
- Grieve, R.A.F. and Fawcett, J.J.(1974). The stability of chloritoid below 10kb p_{H20}. J.Petrol.15,113-39.
- Gunter, W.R. and Yamada, P.H.(1986). Evaluation of industrial mineral occurrences in the Flin Flon Snow Lake area. Report of field activities 1986, Minerals division, Manitoba Energy & Mines.
- Harte, B. and Henley, K.J.(1966). Occurrence of compositionally zoned almanditic garnets in regionally metamorphosed rocks. Nature v.210, pp.689 - 692.
- Haskin, L.A.(1977). On rare earth element behaviour in igneous rocks; in Symposium on the origin and distribution of the elements, 2d, Paris.

- Hensen, B.J. and Green, D.H.(1973). Experimental study of the stability of cordierite and garnet in pelitic compositions at high pressures and temperatures. Synthesis of experimental data and geological applications. Contr.Mineral.Petrology 38,151-166.
- Holdaway, M.J. and Lee, Sang Man(1977). Fe Mg cordierite stability in high - grade pelitic rocks, based on experimental, theoretical and natural observations Contr.Mineral.Petrology 63, 175 - 198.
- Hughes, C.J.(1973). Spilites, keratophytes and the igneous spectrum; Geol.Mag., v.6, pp.513 527.
- Jackson, E.D. and Ross, D.C.(1956). A technique for modal analyses of some medium-and coarse-grained (3mm-10mm) rocks: Am. Min. 41, 648-651.
- Karig, D. E. (1972). Remnant arcs. Bull. Geol. Soc. Am., v.83, p. 1057-1068
- La Roche, Hubert de(1974). Geochemical characters of the metamorphic domains, survival and testimony of their premetamorphic history; Sci. de la terr., Nancy, France, t XIX, n.2, pp.103 - 117.
- Leake, B.E.(1964). Chemical distinction of ortho- and paraamphibolites; Jour. of Petrology, v.5,pp.238 - 254.
- Leake, B.E.(1969). Chemical distinction of ortho and para charnockitic rocks, anorthosites and amphibolites; Indian Mineralogist, v.10, pp.89 - 104.
- Leake, B.E.(1978). Nomenclature of amphiboles. Can. Mineral. 16, pp.501 - 520.
- MacGeehan, P.J. and MacLean, W.H.(1980). Tholeiitic basalt-rhyolite magmatism and massive sulfide deposits at Matagami, Quebec. Nature v.283 p.153-157.
- Menzies, M. et al.(1979). Experimental evidence of rare earth element immobility in greenstones. Nature, 282, 398.
- Middleton, G.V. and Hampton, M.A.(1976). Subaqueous sediment transport and deposition by sediment gravity flows; <u>in</u> Marine sediment transport and environmental management ed. D.J. Stanley and D.J.P. Swift, pp.197 - 218.

Mirwald, P.W.(1982). High - pressure phase transitions in cordierite. Phys. Earth Planetary Int. 29, 1 - 5.

- Muecke, G.K., Pride, C. and Sarkar, P.(19). Rare earth element geochemistry of regional metamorphic rocks; <u>in</u> Symposium on the origin and distribution of the elements, 2d, Paris.
- Niggli, P.(1954). <u>Rocks and Mineral Deposits</u>; San Francisco, W.H. Freeman and Co, 559 pages.
- Roberts, R.G. and Reardon, E.J.(1978). Alteration and ore forming processes at Mattagami Lake Mine, Quebec; Can. Jour. Earth Sci., v.15, pp.1 - 21.
- Robertson, D.S.(1953). Batty Lake map area, Manitoba; Geological Survey of Canada, Memoir 271.
- Schledewitz, D. C. P. (1988). GS-6 Kisseynew Project: Kississing Lake. Report of field activitites 1988, Minerals division, Manitoba Energy and Mines.
- Seifert, F. and Schumacher, J.C.(1986). Cordierite spinel quartz assemblages: A potential geobarometer
 Bull. Geol. Soc. Finland 58, Part 1, 95 108.
- Seyfried, W. E., Berndt, M. E. and Seewald, J. S. (1988). Hydrothermal alteration processes at mid-ocean ridges: Constraints from diabase alteration experiments, hotspring fluids and composition of the oceanic crust. Can. Min., v.26, p.787-804.
- Seyfried, W. E. and Mottl, M. J. (1982). Hydrothermal alteration of basalt by seawater under seawater dominated conditions. Geochemica et Cosmochimica Acta.,v.46, pp.985-1002.
- Shirozu, H.(1974). Clay minerals in altered wall rocks of the Kuroko-type deposits: Japan Min. Geol., Spec. Iss., No.6, p.303-310.
- Spear, F.S.(1980). The gedrite-anthophyllite solvus and the composition limits of orthoamphibole from the Post Pond Volcanics, Vermont. Am. Min., v.65, pp.1103-1118.
- Stamatelopoulou Seymour, K.S. and MacLean, W.H.(1977).
 The geochemistry of possible metavolcanic rocks
 and their relationships at Montoubon Les Mines
 Quebec: Can. Jour. Earth Sci., v.14, pp.2440-2452.

- Taylor, S.R. and McLennan, S.M.(1985). <u>The Continental</u> <u>Crust: its Composition and Evolution</u> Blackwell Scientific Publications, Oxford, London, Edinburgh, Boston, Palo Alto, Melbourne.
- Thompson, A.B.(1976). Mineral reactions in pelitic rocks: Calculation of some P-T-X (Fe-Mg) phase relations. Am.J.Sci. 276, 425 - 454.
- Tuominen, H.V. and Mikkola, T.(1950). Metamorphic Mg-Fe enrichment in the Orijarvi region as related to folding; Bull. Comm. Geol. Finlande, No.150, pp.67 - 92.
- Vallance, T.G.(1969). Spilites again: some consequences of the degradation of basalts; Linnean Society of New South Wales, Proceedings, v.94, pp.8 - 51.
- Van de Kamp, P.C.(1968). Geochemistry and origin of metasediments in the Haliburton - Madoc area, southeastern Ontario; Can. Jour. Earth Sci., v.5, pp.1337 - 1372.
- Van de Kamp, P.C.(1969). Origin of amphibolites in the Beartooth Mountains, Montana and Wyoming: new data and interpretation: Geol. Soc. Amer. Bull., v.80, pp.1127 - 1136.
- Van de Kamp, P.C.(1970). The Green Beds of the Scottish Dalbradian series: geochemistry, origin and metamorphism of mafic sediments: Jour. Geol., v.78, pp.218 - 303.
- Van de Kamp, P.C., Leake, B.E. and Senior, A.(1976). Petrography and geochemistry of some California arkoses with application of the data to gneisses of metasedimentary origin: Jour. Geol., v.84, pp.195 - 212.
- Wegmann, C.E. and Kranck, E.H.(1931). Beitrage zur Kenntnis der Svecofenniden in Finland; Bull. Comm. Geol. Finlande, No.163.
- Wilkinson, S.J.(1976). The petrography of the anthophyllite rocks at Sherridon Manitoba; unpublished B.Sc. thesis, University of Western Ontario, London.

-110-

Wolery, T. J. and Sleep, N. H. (1976). Hydrothermal circulation and geochemical flux at mid-ocean ridges. J. Geol. v.84, p. 249-275.

Zwanzig, H. V. (1988). GS-7 Kisseynew Project: Batty Lake region. Report of field activities 1988, Minerals division, Manitoba Energy and Mines. APPENDIX A CHEMICAL ANALYSES OF ROCKS FROM STAR LAKE, NEAR SHERRIDON MANITOBA

Ninety-nine chemical analyses of rocks from the Star Lake - Elken Lake area, near Sherridon, are compiled in this appendix. The analyses are listed by field sample identification number and unit number. Field sample numbers refer to figures 3 - 4 and Table 2; unit numbers refer to the field units:

> U-1: unit 1 (quartz-garnet-anthophyllite), U-2: unit 2 (garnet-anthophyllite), U-3: unit 3 (anthophyllite-cordierite), QBG: quartz-biotite gneisses, SIL: the sillimanite zone, U-2(M): the mineralized unit 2 (garnet-

anthophyllite) of Elken Lake and

U-3(M): unit 3 (anthophyllite-cordierite)

near the east orebody at Sherridon

Manitoba. (Sherridon unit 3)

A rock saw was used to cut sections within the three mappable Units and adjacent quartz-biotite gneisses in order to provide as continuous a sample as possible for chemical analysis and thin-sectioning. All weathered surfaces were removed, and the sample was sawn in half; one half was saved for making a thin-section, polished thinsection, chemical analyses and hand specimen, and the other half was catalogued at the Manitoba Energy and Mines Bureau.

-112-

The sample was then cut for thin-sectioning. Another representative part of the sample was crushed and sent to the University of Ottawa Analytical Geochemistry Section for major oxide analyses, and Bondar - Clegg Geochemical Laboratory for trace element analyses. Methods, lower detection limits and reliability of analyses for each element and reported standards are listed in Table A-1.

ELEMENT	LOWER DETEC.	HETHOD OF		ACCU (sta	RACY OF Ndard Di	ANALYS Eviatio	IS N)
	(+/- % CONC)	ANALISIS	SY - 2	BCC	SC2-12	SC5-14	SC8-17
Si02 Ti02 Al203 T Fe203 T Fe203 T Fe203 Mn0 Ca0 Na20 P205 Ba Cr Zr Sr V Sco Cs U Pb F S Lae Sm U D Y D Lu Y D Lu	+/- 1.0% +/- 1.0% +/- 1.5% +/- 5.0% 0.01% +/-10.0% +/-10.0% +/-5.0% +/-10.0	XRF XRF XRF XRF XRF XRF XRF XRF XRF XRF	0.168 0.005 0.103 0.036 ic 0.000 0.047 0.059 0.067 0.040 0.015 5.170 8.100 3.670 7.150 3.920 2.530 31.66 Act. Act. Act. Act. Act. Act. Act. Act.	0.000 1.410 0.420 0.495 0.580 2.650 4.950 8.490 0.707 0.640 0.707 0.640 0.071 0.350 0.042	0.400 0.021 0.064 0.622 1.263 0.014 0.000 0.000 0.014 0.035 26.80 2.830 2.830 2.120 1.410 0.000 2.620 2.9.70 0.000 2.620 2.9.70 0.000 0.000 0.5660 2.620 2.9.70 0.000 0.000 0.000 0.000 0.577 1.414 31.20 0.000	0.950 0.050 0.014 0.000 0.112 0.020 0.000 0.000 0.000 0.000 0.000 2.830 3.540 0.000 2.830 3.540 0.000 2.830 3.540 0.000 2.830 3.540 0.000 2.830 3.540 0.000 2.830 3.540 0.000 2.830 3.540 0.000 2.830 0.000 2.830 0.000 2.830 0.000 2.830 0.000 2.830 0.000 2.830 0.000 0.000 2.850 0.000 0	2.020 0.130 0.270 0.520 0.792 0.000 0.360 0.060 0.110 0.040 0.010 37.50 2.120 1.410 2.800 0.000 19.09 17.00 2.900 1.410 0.000 0.000 0.707 0.000 0.707 0.000 0.212 0.000 0.707 0.000 0.212 0.000 0.707 0.000 0.212 0.000 0.707 0.000 0.212 0.000 0.707 0.000 0.212 0.000 0.212 0.000 0.212 0.000 0.212 0.000 0.212 0.000 0.212 0.000 0.001 0.000 0.212 0.000 0.001 0.000 0.001 0.000 0.001 0.000 0.000 0.000 0.001 0.000 0

TABLE A1: GEOCHEMICAL METHODS AND ACCURACY OF ANALYSIS

Unit # I <th>Specimen No.</th> <th>SC 1-1</th> <th>SC 1-2</th> <th>SC 1-7</th> <th>SC 1-10</th> <th>SC 1-14a</th> <th>SC 1-14b</th> <th>SC 2-1</th> <th>SC 2-6</th> <th>SC 2-10</th> <th>SC 2-12c</th> <th>SC 2-12d</th>	Specimen No.	SC 1-1	SC 1-2	SC 1-7	SC 1-10	SC 1-14a	SC 1-14b	SC 2-1	SC 2-6	SC 2-10	SC 2-12c	SC 2-12d
Major Oxides (wt%)	Unit #	1	1	1	1	SIL	SIL	1	1	1	1	1
$ \begin{array}{cccccccccccccccccccccccccccccccccccc$	Major Oxides	(wtZ)										
Total 98.52 99.97 98.12 98.61 96.16 97.57 98.58 99.33 98.64 100.3 98.17 Minor Elements (ppm)	Si02 Ti02 Al203 Fe203 Fe0 Mn0 Mg0 Ca0 Na20 K20 P205	$\begin{array}{c} 75.55\\ 0.10\\ 13.10\\ 0.27\\ 1.52\\ 0.03\\ 1.16\\ 1.12\\ 3.30\\ 2.29\\ 0.08 \end{array}$	69.36 0.48 13.02 3.64 6.43 0.28 4.74 0.59 - 1.02 0.41	69.35 0.47 12.89 0.99 6.44 0.09 4.77 0.30 0.03 2.64 0.15	70.02 0.36 11.19 0.88 8.06 0.14 4.36 0.29 0.11 3.07 0.13	67.50 0.29 13.28 1.03 6.18 0.16 4.55 0.20 0.01 2.85 0.11	71.81 0.35 10.41 1.29 5.93 0.14 3.97 0.22 0.17 3.12 0.16	68.86 0.47 12.92 0.52 9.18 0.28 4.36 0.58 - 1.02 0.39	71.73 0.41 12.19 0.85 7.44 0.11 4.88 0.22 0.11 1.29 0.10	72.41 0.37 10.69 0.80 7.95 0.16 4.86 0.22 - 0.96 0.22	73.21 0.37 10.81 9.68 0.11 4.65 0.24 0.02 1.00 0.17	72.65 0.40 10.72 0.61 7.44 0.09 4.79 0.24 - 1.11 0.12
Minor Elements (ppm) Image: constraint of the system of t	Total	98.52	99.97	98.12	98.61	96.16	97.57	98.58	99.33	98.64	100.3	98.17
$ \begin{array}{c ccccccccccccccccccccccccccccccccccc$	Minor Elements	(ppm)										
	Ba Co Cr Cs Cu F Mo Ni Pb Rb S Sc Sr Y Y	186 80 28 2 3 250 3 - 7 66 - 5 52 - 33	442 86 33 9 1 1064 4 - 3 5 - 21 11 3 39	568 81 25 5 765 3 - 4 123 - 21 11 - 26	360 74 30 8 2 765 3 - 2 159 - 24 6 - 28	329 67 29 7 2 870 3 - 6 155 - 19 7 15 16	319 87 27 8 1 814 3 - 5 194 - 20 - 36	392 120 29 1 4 730 4 - 4 8 - 32 9 3 36	490 100 26 1 12 705 3 - 3 8 - 21 3 5 19	346 95 28 1 500 3 - 3 6 - 26 6 - 25	374 140 29 1 2 760 3 - 2 4 - 22 6 - 21	412 98 33 745 3 - 4 7 - 18 3 8 19

× *

			· · · · · · · · · · · · · · · · · · ·								
Specimen No.	SC 2-21a	SC 2-21b	SC 2-25a	SC 2-25b	SC 2-26	SC 3-2	SC 3-4	SC 3-8	SC 3-10	SC 3-13c	SC 3-13d
Unit #	1	1	1	1	1	1	1	1	1	1	1
Major Oxides	(wt%)										
Si02 Ti02 Al203 Fe203 Fe0 Mn0 Ca0 Na20 K20 P205	$\begin{array}{c} 72.59\\ 0.43\\ 11.55\\ 0.63\\ 7.50\\ 0.10\\ 4.18\\ 0.24\\ 0.09\\ 1.05\\ 0.13\\ \end{array}$	$\begin{array}{c} 72.68\\ 0.41\\ 12.03\\ 0.65\\ 7.82\\ 0.10\\ 4.24\\ 0.22\\ 0.06\\ 1.13\\ 0.09\\ \end{array}$	81.29 0.26 7.52 0.44 5.23 0.12 3.06 0.24 0.12 0.76 0.19	74.95 0.35 11.33 0.63 7.08 0.09 4.16 0.31 0.03 0.80 0.21	74.63 0.38 10.93 0.71 5.39 0.05 5.34 0.24 - 1.45 0.22	69.04 0.32 11.48 0.75 7.63 0.05 8.48 0.18 0.32 0.65 0.15	69.63 0.39 16.19 0.52 5.38 0.05 6.42 0.25 - 0.51 0.23	68.81 0.42 11.86 0.66 7.51 0.05 8.34 0.19 0.25 0.85 0.15	68.32 0.27 10.73 1.17 9.40 0.14 8.03 0.18 0.35 1.01 0.08	76.72 0.20 8.75 0.61 4.56 0.03 6.06 0.07 - 1.27 0.10	78.61 0.18 9.37 0.46 3.49 0.02 5.01 0.06 0.02 1.09 0.06
Total	98.49	99.43	99.23	99.94	99.34	99.05	99.57	99.09	99.68	98.37	98.37
Minor Elements	(ppm)										
Ba Co Cr Cs Cu F Mo Ni Pb Rb S Sc Sr Y Y Zn Zr	484 130 29 1 1 655 4 - 3 6 - 20 4 16 20 36 46	$517 \\ 120 \\ 28 \\ 1 \\ 1 \\ 725 \\ 3 \\ - \\ 2 \\ 10 \\ - \\ 23 \\ 6 \\ 9 \\ 22 \\ 33 \\ 50 \\ 0 \\ 0 \\ 0 \\ 0 \\ 0 \\ 0 \\ 0 \\ 0 \\ $	$\begin{array}{c} 320\\ 140\\ 34\\ 1\\ 1\\ 555\\ 4\\ -\\ 2\\ -\\ 24\\ 5\\ -\\ 41\\ 12\\ 30\\ \end{array}$	382 140 25 1 600 3 - 2 3 - 19 7 13 24 25 44	504 110 25 1 1125 3 - 2 8 - 17 9 - 24 21 42	233 110 29 1 1000 3 - 2 - 16 5 - 84 43 45	221 120 29 1 630 3 - 2 - 13 6 14 23 12 48	298 97 29 1 1130 2 - 2 5 - 15 5 - 21 29 47	348 120 31 1 1130 4 26 2 28 3 28 29 30 52	446 130 36 1 1195 4 - 2 8 - 17 9 25 22 25 56	406 110 40 1 935 3 - 5 6 - 11 7 26 29 32 59

Specime No.	n SC 3-18	SC 3-25a	SC 3-251	SC 3-29	SC 3-33	SC 3-39	SC 3-42	SC 3-44	SC 3-51a	SC 3-54b	SC 4-2c
Unit #	1	2	2	2	2	1	2	2	2	2	3
Major Oxides	(wt%)										
Si02 Ti02 A1203 Fe0 Mn0 Mg0 Ca0 Na20 K20 P205	75.09 0.17 9.76 0.45 5.01 0.05 5.57 0.09 0.99 0.08	64.71 0.51 11.15 0.81 11.30 0.10 10.22 0.25 0.48 0.05 0.16	61.42 0.64 13.95 1.37 11.01 0.16 10.91 0.24 0.22 0.14	56.71 0.65 14.35 1.00 14.00 0.27 11.97 0.39 0.24 0.03 0.62	61.05 0.67 14.90 1.04 10.43 0.17 11.14 0.23 0.27 0.03 0.15	70.97 0.27 13.96 0.69 6.16 0.07 5.58 0.16 0.05 1.05 0.09	58.67 0.66 13.21 1.52 12.86 0.25 11.98 0.30 0.27 0.18 0.20	59.48 0.61 15.32 0.56 13.04 0.32 10.23 0.39 0.24 0.02 0.27	57.12 0.64 15.52 0.84 14.13 0.28 10.52 0.42 0.04 0.05 0.19	59.35 0.65 12.78 1.13 12.32 0.14 12.69 0.26 0.37 0.11 0.15	62.95 0.67 14.36 1.01 10.39 0.09 9.22 0.28 0.19 0.18 0.20
Total	97.26	99.74	100.1	100.2	100.1	99.05	100.1	100.5	99.75	99.95	99.54
Minor Elements	(ppm)										
8a Co Cr Cs Cu F Mo Ni Pb Rb S Sc Sr Y Y Zn Zr	355 110 27 1 960 3 - 3 8 - 18 6 12 32 28 51	80 140 24 1 1125 3 - 2 - 40 4 313 11 76 24	93 120 20 1 970 2 - 2 - 44 1 349 13 91 34	$ \begin{array}{c} 114\\ 130\\ 23\\ 1\\ 1365\\ 2\\ -\\ 2\\ -\\ 66\\ -\\ 324\\ 22\\ 41\\ 32\\ \end{array} $	93 110 17 1 1205 2 - 2 - 39 1 369 22 64 31	411 999 33 1 900 3 - 2 6 - 17 6 42 27 44 47	$ \begin{array}{c} 115\\ 91\\ 23\\ 1\\ 5\\ 1060\\ 2\\ -\\ 2\\ -\\ 63\\ -\\ 339\\ 30\\ 41\\ 32\\ \end{array} $	102 100 28 1 2 810 1 - 2 - 76 - 316 47 69 32	92 130 21 1 944 835 2 - 2 - 48 4 346 15 80 33	75 110 25 1 1205 3 - 2 - 32 - 362 15 88 30	68 82 26 1 12 730 3 - 2 - 43 4 225 31 50 37

betoeteb ton -

	55 122 52 162 - 22 -	I\ I0I 56 I07 8⊄ -	50 528 155 45 45 758	16 69 12 522 1 22 22	16 54 15 188 188 48	24 508 45 162 56 56 56	44 528 52 1 58 1 7 2	25 1246 182 182 - 56	21 919 510 12 12 12	22 50 52 24 24	40 22 19 52 52 52 52 52	۲۲ ۲۳ ۲۳ ۲۶ ۲۶ ۲۶
	- 5 1 2 2 4 2 8 7 8 7 6 9 1 50 8 6 3 8	2 2 5 2 5 2 5 2 5 2 5 9 5 0 5 0 5 0 5 0 5 0 5	↓ ↓ 592 50 120 50 50 50 50 50 50 50 50 50 5	- 5 3 3 5 5 5 5 5 5 5 5 5 5 5 5 5 5 5 5	- 5 9 2 5 8 5 8 5 8 5 5 5 5 5 5 5 5 5 5 5 5 5	- - - 822 4 - 00 822 10 52 110 25	- 5 1 3 5 8 1 10 3 5 8 5 8 5 2 5 8 2 5 8 5 5 8 5 5 5 5 5 8 5 5 8 5 5 8 5 7 8 5 8 5	- 9 22 225 325 54 1 20 54 120	- - 3 - 3 - 5 3 - 5 3 - 5 3 5 3 5 8 8 1 2 0 8 8 8 8 8	- 5 - 52 1160 58 86	- 5 1095 100 120 120 120	B8 CC CC CC CC CC CC CC CC CC CC CC CC CC
······											(udd)	Minor Elements
	16.66	100.3	0.001	9°00T	8-00I	12°66	2.001	IZ.96	OS.99	82.99	0.001	Istol
	0.08 0.18 0.50 0.54 0.54 1.15 1.15 1.15 1.25 0.45 20.45	0'10 0'12 0'12 0'12 8'21 1'52 1'52 1'52 1'52 1'52 1'52 1'52 1	0.06 0.82 0.85 0.13 0.16 12.86 0.56 12.43 0.56 9.22 0.58	0°04 0°81 0°06 8°10 8°10 13°20 14°60 14°60 9°75	0°08 0°28 0°28 0°58 0°48 0°48 1°20 0°58 0°48 0°48 0°60 86°09	0°1¢ 0°2¢ 0°3¢ 0°3¢ 0°3¢ 1°20 1°20 1°20 1°20 1°20 0°3¢ 0°20 0°2¢ 0°20 0°2¢ 0°20 0°2¢	0'15 0'04 0'07 0'45 0'45 12'10 1'02 1'02 1'02 1'02 1'02 1'02 1'0	0°20 0°12 0°54 0°12 0°12 0°13 1°50 1°50 1°50 1°50 1°50 1°50 1°50 1°50	0°5¢ 0°5¢ 0°52 0°22 0°52 0°52 0°52 0°50 0°2¢	0°12 0°58 0°58 0°58 0°58 0°58 0°58 1°58 0°58 0°58 0°58 0°58 0°58 0°58 0°58	0'18 0'28 0'28 0'28 0'57 0'27 0'20 15'82 0'20 15'82	6202 620 620 620 620 820 8202 6502 6502 1105 81202 8120 8120
											(21M)	Major S9bix0
	5	5	5	5	5	2	2	2	2	2	2	# JinU
P	2-1¢ 3C	2-1¢C 2C	2-10 2C	9-S JS	I-S JS	¢-12P 2C	4-139 SC	¢-15 2C	6-† 35	\$-7 20	¢-59 2C	Specimen No.

Specime No.	n SC 5-18	SC 5-19	SC 5-251	SC 5-31	SC 5-39	SC 5-55	SC 5-60	SC 5-65	SC 5-67	SC 6-1c	SC 6-1d
Unit #	2	2	2	3	3	QBG	QBG	QBG	OBG	1	1
Major Oxides	(wt2)										
Si02 Ti02 Al203 Fe203 Fe0 Mn0 Mg0 Ca0 Na20 K20 P205	60.04 0.50 15.74 0.67 10.83 0.12 10.24 0.11 0.14 0.62 0.09	66.47 0.43 10.02 0.86 9.58 0.07 10.81 0.08 - 0.64 0.05	64.52 0.45 13.87 1.05 9.74 0.11 8.64 0.14 0.07 0.31 0.07	56.85 0.58 13.72 1.27 12.23 0.11 14.63 0.12 0.34 0.16 0.10	56.63 0.50 14.84 0.59 12.12 0.10 13.95 0.11 0.36 0.54 0.09	79.71 0.24 8.47 0.55 3.58 0.03 4.41 0.13 0.09 0.93 0.08	77.58 0.18 9.05 0.38 3.69 0.02 4.48 0.09 - 2.02 0.06	82.00 0.15 8.78 0.31 3.25 0.02 2.50 0.05 0.01 0.91 0.02	86.46 0.13 7.74 0.26 1.32 0.01 0.96 0.01 - 0.66	64.69 0.30 16.73 1.51 5.15 0.05 3.69 0.27 0.49 2.54 0.15	69.16 0.25 14.03 1.39 4.29 0.05 3.26 0.27 0.75 2.08 0.08
Total	99.10	99.01	98.97	100.1	99.83	98.22	97.55	98.00	97.55	95.57	95.61
Minor Elements	(ppm)										
Ba Co Cr Cs Cu F Mo Ni Pb Ni Sc Sr Sc Sr Y Y Zn Zr	188 110 57 1 1 1170 1 - 39 1 225 18 38 17	220 86 53 1 1425 1 4 2 - 22 - 195 10 38 20	160 160 60 1 985 2 - 2 - 47 - 47 - 201 17 46 17	82 90 68 1 1225 - 2 - 47 - 257 15 42 22	$ \begin{array}{c} 116\\ 76\\ 63\\ 1\\ 1200\\ 1\\ -\\ 3\\ -\\ 43\\ 4\\ 229\\ 14\\ 59\\ 17\\ \end{array} $	209 89 22 1 3 650 2 - 4 1 - 18 13 8 20 23 39	$269 \\ 110 \\ 24 \\ 1 \\ 3 \\ 1110 \\ 2 \\ - \\ 4 \\ 17 \\ - \\ 21 \\ 6 \\ 4 \\ 19 \\ 16 \\ 46 \\ 46 \\ $	229 100 25 1 85 555 3 - 2 6 - 21 6 2 36 15 57	178 110 24 1 36 415 2 - 4 1 - 12 - 16 9 47	541 67 27 8 5 730 1 - 5 57 - 8 9 34 25 63 82	442 68 29 7 3 640 1 - 8 45 - 6 9 19 25 54 72

Specimen SC 6-3 7-2 7-4 No. 6-4 7-10 7-11 8-4 8-7 8-11 8-17c 8-17d Unit # SIL SIL Major Oxides (wt%) Si02 Ti02 A1203 0.59 0.30 0.13 0.92 0.27 0.45 0.46 0.67 0.92 1.60 1.08 6.80 8.07 6.53 11.59 1.69 4.73 12.57 10.23 9.05 11.98 13.10 Fe203 Fe0
 0.01
 0.09
 0.08
 0.07
 0.12
 0.12

 1.61
 4.83
 13.50
 11.87
 10.62
 15.01
 15.52

 1.78
 0.84
 0.15
 0.19
 0.12
 0.37
 0.46
 0.11 0.21 3.87 4.57 Mn0 0.10 0.10 4.13 3.84 MgO CaO 0.20 0.19 0.73 0.35 Na20 0.26 0.23 0.41 0.44 0.58 0.44 0.29 0.06 0.55 1.01 0.24 0.29 0.02 0.05 --1.72 0.82 K20 0.44 1.52 1.04 2.75 1.67 0.79 0.10 0.13 0.07 0.09 0.11 P205 0.16 0.20 0.54 0.14 Total 98.84 99.99 98.07 99.61 99.02 97.34 100.0 100.4 97.24 98.50 97.33 Minor Elements (ppm) Ba Co Cr Cs Cu F Мо Ni _ --_ ---_ _ Pb Rb --S _ _ ------_ --Sc 7 Sr 52 98 17 -_ V 30 ---23 ~ Y Zn Zr

	1	· · · · · · · · · · · · · · · · · · ·	· · · · · · · · · · · · · · · · · · ·	,							
Specimen No.	SC 8-21	SC 8-28	ML 05	ML 06	ML 11	ML 12	ML 15	ML 18	ML 23	ML 31	ML 33
Unit #	3	3	SIL	1	SIL	2	3	QBG	2	2	3
Major Oxides	(wt%)										
Si02 Ti02 Al203 Fe203 Fe0 Mn0 Mg0 Ca0 Na20 K20 P205	$\begin{array}{c} 57.37\\ 0.55\\ 17.92\\ 1.22\\ 10.15\\ 0.09\\ 10.55\\ 0.19\\ 0.32\\ 0.16\\ 0.14\\ \end{array}$	56.63 0.66 13.85 1.53 11.32 0.08 12.86 0.22 0.28 0.29 0.11	81.31 0.13 14.50 0.24 0.98 0.01 0.82 0.27 	44.13 0.66 22.19 1.39 18.57 0.41 11.36 0.45 0.37 1.06 0.12	94.69 0.01 2.46 0.02 0.10 - 0.06 0.12 0.09 -	63.05 0.47 13.41 3.29 14.00 0.47 5.77 0.28 0.06 0.05 0.08	44.54 0.64 22.01 1.73 13.07 0.11 15.52 0.20 0.49 0.23 0.33	74.03 0.39 10.18 1.80 4.88 0.03 5.60 0.10 0.33 0.85 0.07	53.84 0.33 16.90 2.11 12.86 0.07 12.82 0.16 0.37 0.52 0.13	59.06 0.45 17.81 2.01 10.38 0.32 5.48 0.22 0.31 3.17 0.14	49.65 0.67 20.59 1.66 11.62 0.12 13.77 0.23 0.32 0.11 0.17
Total	98.66	97.83	99.13	100.7	97.55	100.9	98.87	98.26	100.1	99.35	98.91
Minor Elements	(ppm)										
Ba Co Cr Cs Cu F Mo Ni Pb Rb Sc Sr Y Y Zn Zr	136 63 67 1 2 1030 1 - 2 - 38 1 228 18 51 21	167 80 71 1 1315 2 - 2 - 47 3 235 12 70 26	228 62 23 1 480 2 - 2 10 - 2 8 35 7 29 9	310 84 66 1 10555 1 - 4 7 7 76 15 214 27 36 24	70 60 20 1 40 1 - 2 - - 5 - 1 -	64 100 56 1 97 115 1 - 2 - 59 1 73 18 - 16	82 66 61 1 28 1460 1 - 2 - 53 7 273 18 75 27	176 74 44 1121 1 - 6 5 - 29 8 146 12 101 13	199 63 37 1 20 1485 1 - 2 5 - 18 8 349 20 259 9	431 58 27 6 2 960 1 - 4 122 - 47 4 76 55 111 82	51 65 61 1 1320 1 - 2 - 46 8 247 16 75 24

Specimen No.	ML 34a	ML 34b	ML 46	ML 47	ML 53	ML 61	ML 65	ML 67	ML 68a	ML 68b	ML 69a
Unit #	QBG	QBG	QBG	2(M)	2(M)	2	QBG	2	QBG	QBG	2
Major Oxides	(wt%)										
Si02 Ti02 Al203 Fe203 Fe0 Mn0 Mg0 Ca0 Na20 K20 P205	83.70 0.17 9.00 0.63 1.92 0.01 1.78 0.07 0.12 0.99 0.02	83.48 0.17 8.94 0.56 2.11 1.83 0.10 0.27 1.03 0.02	$76.36 \\ 0.15 \\ 12.38 \\ 0.88 \\ 3.44 \\ 0.03 \\ 4.02 \\ 0.14 \\ 0.23 \\ 1.25 \\ 0.03 \\ 0.03 \\ 0.03 \\ 0.03 \\ 0.03 \\ 0.03 \\ 0.00 $	54.17 0.66 16.63 1.01 11.53 0.07 12.80 0.87 0.70 0.40 0.42	55.05 0.60 16.37 1.15 14.23 0.20 10.19 0.60 0.45 0.03 0.37	58.76 0.56 15.52 0.92 11.86 0.14 11.04 0.28 0.34 0.39 0.17	70.46 0.21 11.78 0.45 2.74 0.06 1.07 1.13 3.25 1.61 0.09	62.58 0.44 13.41 0.48 9.62 0.06 11.72 0.15 0.32 0.51 0.12	75.72 0.35 11.52 0.53 3.32 0.02 4.49 0.14 0.11 1.50 0.14	77.66 0.34 10.51 0.47 3.22 0.02 4.32 0.14 0.19 1.37 0.10	59.93 0.49 16.34 1.23 10.46 0.12 9.28 0.20 0.14 0.60 0.19
Total	98.41	98.52	98.91	99.26	99.24	99.98	92.85	99.41	97.84	98.34	98.98
Minor Elements	(ppm)										
Ba Co Cr Cs Cu F Mo Ni Pb Rb S Sc	193 74 18 1 21 745 1 - 2 7 - 14	181 70 18 1 23 700 1 - 2 11 - 13	203 63 22 1 320 1 - 2 11 - 2 5	65 85 21 2 560 1 - 2 - 48	54 110 63 1 2780 665 1 - 2 - 2 - 68	186 70 52 1 6 1760 1 - 2 - - 29	333 57 17 15 250 1 - 2 18 - 14	$ \begin{array}{r} 172 \\ 78 \\ 48 \\ 1 \\ 1905 \\ 1 \\ - \\ 2 \\ 1 \\ - \\ 2 \\ 1 \\ - \\ 21 \end{array} $	439 57 22 1 1 1130 3 - 2 15 - 9	423 74 22 1 1165 - 2 10 - 9	298 63 57 1 1395 1 - 2 1 - 2 1 - 55
Sr Y Y Zn Zr	29 4 22 17 62	43 - 23 19 60	9 64 47 95 99	13 288 32 140 14	6 295 17 75 12	5 142 28 82 27	85 7 29 55 77	9 192 10 31 20	15 - 19 20 57	10 - 15 10 57	55 8 186 17 63 19

~ •

	· · · · · · · · · · · · · · · · · · ·	1	7								
Specimen No.	ML 69b	ML 70	ML 72	ML 74	ML 78	ML 79	ML 80	ML 81	ML 82	ML 91	ML 104
Unit #	2	QBG	QBG	2	2	1	3	3	3	3(M)	QBG
Major Oxides	(wtZ)										
Si02 Ti02 Al203 Fe203 Fe0 Mn0 Mg0 Ca0 Na20 K20 P205	59.90 0.49 15.85 1.50 10.59 0.11 9.59 0.19 0.11 0.53 0.13	74.47 0.21 11.44 0.96 4.62 0.06 3.73 0.23 0.22 1.78 0.18	58.95 0.19 10.31 1.11 3.48 0.12 8.66 11.65 0.23 3.80	68.43 0.40 11.04 9.28 0.08 8.37 0.12 0.37 0.10 0.08	59.08 0.50 15.89 0.94 11.88 0.09 10.33 0.18 0.31 0.12 0.11	61.61 0.43 14.63 0.46 14.96 0.25 5.34 0.43 0.03 1.11 0.19	64.15 0.64 12.78 0.96 9.68 0.03 10.25 0.34 0.36 0.02 0.21	63.98 0.62 14.15 0.95 9.05 0.04 9.46 0.34 0.03 0.22	62.80 0.66 14.15 1.19 8.38 0.07 10.04 0.27 0.22 0.28 0.12	70.98 0.23 10.82 1.17 7.17 0.21 6.73 0.24 0.37 0.36 0.07	76.71 0.18 11.20 0.14 0.65 0.07 0.66 4.17 0.55 3.22
Total	98.99	97.90	98.50	98.91	99.43	99.44	99.42	99.18	98.18	98.35	97.55
Minor Elements	(ppm)										
Ba Bo Dr Dr Dr Dr Dr Dr Dr Dr Dr Dr Dr Dr Dr	249 64 57 1 1320 1 - 2 1 - 44 6 193 15 50 16	265 54 30 5 11 870 1 - 2 57 - 20 10 42 21 49 45	647 33 20 1 74 520 1 - 54 31 - 18 72 - 39 228 59	98 81 47 1 1090 1 - 29 4 151 12 36 13	80 84 58 1 710 1 - 2 - 35 5 178 18 26 20	614 84 57 1 2 770 1 - 4 10 - 53 137 20 47 22	65 65 16 1 985 1 - 2 22 4 223 8 465 35	49 72 14 15 1020 1 - 2 208 40 734 36	119 75 20 1 64 730 1 - 2 - 36 3 198 25 53 36	98 69 21 280 730 3 - 2 7 - 15 6 - 26 98 32	337 52 20 1 280 1 280 1 - 2 32 - 11 113 22 23 32 67

- not detected

.

APPENDIX B RARE EARTH ELEMENT ANALYSES OF ROCKS FROM STAR LAKE NEAR SHERRIDON MANITOBA

Ninety-nine rare earth element analyses of rocks from the Star Lake - Elken Lake area, near Sherridon, are compiled in this appendix. The analyses are listed by field sample identification number and unit number. Field sample numbers refer to figures 3 - 4 and Table 2; unit numbers refer to the field units:

U-1: unit 1 (quartz-garnet-anthophyllite),

U-2: unit 2 (garnet-anthophyllite),

U-3: unit 3 (anthophyllite-cordierite),

QBG: quartz-biotite gneisses,

SIL: the sillimanite zone,

U-2(M): the mineralized unit 2 (garnet-

anthophyllite) of Elken Lake and

U-3(M): unit 3 (anthophyllite-cordierite)

near the east orebody at Sherridon

Manitoba. (Sherridon unit 3)

Methods, lower detection limits and reliability of analyses for each element and reported standards are listed in Table A-1.

Specimer No.	SC 1-1	SC 1-2	SC 1-7	SC 1-10	SC 1-14a	SC 1-14b	SC 2-1	SC 2-6	SC 2-10	SC 2-12c	SC 2-12d
Unit #	1	1	1	1	SIL	SIL	1	1	1	1	1
La Ce Sm Eu Tb Yb Lu	21 40 4 1 2 3 1	13 31 4 1 2 7 1	6 9 2 1 1 3 -	9 16 3 1 2 4 1	5 10 2 1 1 2 -	14 22 4 1 2 4 1	7 10 3 1 2 4 1	6 16 2 1 2 3 1	8 12 3 1 1 2 1	7 7 2 1 2 2 1	7 14 3 1 1 2 -
Specimen No.	SC 2-21a	SC 2-21b	SC 2-25a	SC 2-25b	SC 2-26	SC 3-2	SC 3-4	SC 3-8	SC 3-10	SC 3-13c	SC 3-13d
Unit #	1	1	1	1	1	1	1	1	1	1	1
La Ce Sm Eu Tb Yb Lu	7 8 3 1 2 2 -	8 10 3 1 2 2 -	4 8 2 1 2 5 1	11 19 4 1 2 3 1	3 5 2 1 2 2 -	3 9 3 1 3 8 1	4 9 1 2 2 -	4 9 2 1 1 2 -	3 8 2 1 2 3 1	2 6 1 1 3 -	3 5 2 1 2 2 1
Specimen No.	SC 3-18	SC 3-25a	SC 3-25b	SC 3-29	SC 3-33	SC 3-39	SC 3-42	SC 3-44	SC 3-51a	SC 3-54b	SC 4-2c
Unit #	1	2	2	2	2	1	2	2	2	2	3
La Ce Sm Eu Tb Yb Lu	6 12 3 1 2 4 1	4 6 2 1 1 2 -	5 6 2 1 2 -	5 7 2 1 2 2 -	5 7 2 1 1 4 1	6 9 2 1 1 2 1	5 7 2 1 1 4 1	5 5 2 1 2 5 1	6 5 3 1 2 -	5 8 1 1 2 -	6 9 1 1 3 1

.....

a tala a distriction di transferito de l'estre este di 👘 e

,

Specimen No.	SC 4-2d	SC 4-5	SC 4-9	SC 4-12	SC 4-13a	SC 4-13b	SC 5-1	SC 5-6	SC 5-10	SC 5-14c	SC 5-14d
Unit #	3	3	3	3	3	3	2	2	2	2	2
La Ce Sm Eu Tb Yb Lu	6 7 1 1 2 -	5 7 2 1 2 -	6 5 3 1 2 -	10 65 48 5 49 150 21	7 13 1 1 2 -	3 12 2 1 2 8 1	3 6 1 1 2 -	351122-	2 5 1 1 2 -	3 5 1 1 5 1	3 5 1 1 3 1
Specimen No.	SC 5-18	SC 5-19	SC 5-25b	SC 5-31	SC 5-39	SC 5-55	SC 5-60	SC 5-65	SC 5-67	SC 6-1c	SC 6-1d
Unit ∦	2	2	2	3	3	QBG	QBG	QBG	QBG	1	1
La Ce Sm Eu Tb Yb Lu	3 5 2 1 2 -	2 5 1 1 2 -	4 5 2 1 2 -	3 5 1 1 2 -	4 5 1 1 2 -	6 5 2 1 2 -	6 11 2 1 1 3 1	13 26 4 1 2 5 1	5 9 2 1 1 2 -	7 11 2 1 1 3 1	7 12 2 1 1 2 1
Specimen No.	SC 6-3	SC 6-4	SC 7-2	SC 7-4	SC 7-10	SC 7-11	SC 8-4	SC 8-7	SC 8-11	SC 8-17c	SC 8-17d
Unit #	2	2	1	1	SIL	SIL	3	3	3	3	3
La Ce Sm Eu Tb Yb Lu	8 15 3 1 1 3 1	6 8 2 1 2 3 1	6 8 3 1 2 -	2 5 1 1 2 -	18 45 8 3 2 2 -	9 18 4 1 1 2 -	4 10 2 1 1 2 -	5 10 2 1 1 2 -	3 5 1 1 2 -	3 7 1 1 2 -	4 7 2 1 1 2 -

			•								
Specimen No.	SC 8-21	SC 8-28	ML 05	ML 06	ML 11	ML 12	ML 15	ML 18	ML 23	ML 31	ML 33
Unit #	3	3	SIL	1	SIL	2	3	QBG	2	2	3
La Ce Sm Eu Tb Yb Lu	3 7 1 1 1 2 -	4 10 2 1 1 2 -	3 5 2 1 1 2 1	5 5 2 1 3 1 3	2 5 - 1 1 2 1	3 5 2 1 1 2 1	2 5 1 1 1 2 1	3 6 2 1 1 2 1	2 5 1 1 2 1	21 63 7 1 2 7 1	2 5 1 1 2 1
Specimen No.	ML 34a	ML 34b	ML 46	ML 47	ML 53	ML 61	ML 65	ML 67	ML 68a	ML 68b	ML 69a
Unit #	QBG	QBG	ØBG	2(M)	2(M)	2	QBG	2	QBG	QBG	2
La Ce Sm Eu Tb Yb Lu	8 25 3 1 1 3 1	8 14 3 1 1 2 1	8 24 5 1 2 7 1	4 12 3 1 1 4 1	3 5 2 1 1 2 1	2 5 2 1 1 3 1	8 29 4 1 1 4 1	2 5 2 1 2 1 2 1	7 16 3 1 2 1	7 15 3 1 1 2 1	3 5 2 1 2 1
Specimen No.	ML 69D	ML 70	ML 72	ML 74	ML 78	ML 79	ML 80	ML 81	ML 82	ML 91	ML 104
Unit #	2	QBG	QBG	2	2	1	3	3	3	3(M)	<u></u> Øbg
La Ce Sm Eu Tb Yb Lu	3 5 2 1 1 2 1	5 6 2 1 1 3 1	6 12 4 1 2 5 1	2 5 1 1 2 1	3 5 2 1 1 2 1	3 5 2 1 1 2 1	6 9 3 1 1 2 1	3 6 4 1 2 4 1	8 14 3 1 1 4 1	7 16 3 1 1 3 1	8 22 3 1 1 3 1

APPENDIX C NIGGLI NORMS CALCULATED FOR THE ROCKS FROM STAR LAKE NEAR SHERRIDON MANITOBA

Niggli norms were calculated for 99 samples from Appendix A, and are compiled in this appendix. These samples represent samples from the Star Lake - Elken Lake area, near Sherridon. The analyses are listed by field sample identification number and unit number. Field sample numbers refer to figures 3 - 4 and Table 2; unit numbers refer to the field units:

> U-1: unit 1 (quartz-garnet-anthophyllite), U-2: unit 2 (garnet-anthophyllite), U-3: unit 3 (anthophyllite-cordierite), QBG: quartz-biotite gneisses, SIL: the sillimanite zone, U-2(M): the mineralized unit 2 (garnetanthophyllite) of Elken Lake and U-3(M): unit 3 (anthophyllite-cordierite) near the east orebody at Sherridon Manitoba. (Sherridon unit 3)

> > -128-

Specimer	n SC	SC	SC	SC	SC	SC	SC	SC	SC	SC	SC
No.	1-1	1-2	1-7	1-10	1-14a	1-14t	2-1	2-6	2-10	2-12c	2-12d
Unit #	1	1	1	1	SIL	SIL	1	1	1	1	1
fm	19.3	63.2	58.1	60.9	56.7	58.5	62.6	62.9	67.3	67.4	65.7
alk	27.6	2.7	7.5	9.1	8.0	10.5	2.8	4.2	2.8	3.0	3.4
al	45.9	31.4	33.1	28.6	34.3	29.9	32.1	31.9	28.8	28.4	29.7
c	7.2	2.7	1.3	1.4	1.0	1.1	2.5	1.0	1.1	1.2	1.2
si	451.6	283.3	303.6	303.9	296.9	351.2	290.3	317.4	332.3	327.0	342.9
ti	0.5	1.5	1.6	1.2	1.0	1.3	1.5	1.4	1.3	1.2	1.4
p	0.2	0.7	0.3	0.2	0.2	0.3	0.7	0.2	0.4	0.3	0.2
mg	0.5	0.5	0.5	0.5	0.5	0.5	0.4	0.5	0.5	0.5	0.5
k	0.3	1.0	1.0	1.0	1.0	0.9	1.0	0.9	1.0	1.0	1.0
Specimen	SC										
No.	2-21a	2-21b	2-25a	2-25b	2-26	3-2	3-4	3-8	3-10	3-13c	3-13d
Unit #	1	1	1	1	1	1	1	1	1	1	1
fm	62.7	62.5	64.0	62.7	63.4	71.8	59.0	70.9	73.6	68.8	63.0
alk	3.6	3.6	4.1	2.7	4.4	2.7	1.3	2.9	3.5	4.3	4.3
al	32.5	32.8	30.1	33.0	31.0	24.8	38.6	25.5	22.2	26.5	32.3
c	1.2	1.1	1.8	1.6	1.2	0.7	1.1	0.7	0.7	0.4	0.4
si	347.7	336.9	551.7	371.1	360.8	252.2	281.7	252.5	240.6	394.4	460.5
ti	1.6	1.4	1.3	1.3	1.4	0.9	1.2	1.2	0.7	0.8	0.8
P	0.3	0.2	0.5	0.5	0.4	0.2	0.4	0.2	0.1	0.2	0.1
mq	0.5	0.5	0.5	0.5	0.6	0.7	0.7	0.7	0.6	0.7	0.7
k	0.9	0.9	0.8	0.9	1.0	0.6	1.0	0.7	0.7	1.0	1.0
Specimen	SC										
No.	3-18	3-25a	3-25b	3-29	3-33	3-39	3-42	3-44	3-51a	3-54b	4-2c
Unit #	1	2	2	2	2	1	2	2	2	2	3
fm	66.5	77.7	75.4	77.0	73.9	60.8	78.0	73.6	74.6	78.7	72.1
alk	3.4	1.5	0.7	0.6	0.8	3.0	1.0	0.7	0.2	1.1	0.9
al	29.6	20.0	23.2	21.3	24.6	35.4	20.2	24.6	24.0	19.5	26.1
c	0.5	0.8	0.7	1.1	0.7	0.8	0.8	1.1	1.2	0.7	0.9
si	386.5	197.5	173.2	142.8	171.3	305.4	152.1	162.6	150.1	154.3	194.3
ti	0.7	1.2	1.4	1.2	1.4	0.9	1.3	1.2	1.3	1.3	1.6
p	0.2	0.2	0.2	0.7	0.2	0.2	0.2	0.3	0.2	0.2	0.3
mg	0.7	0.6	0.6	0.6	0.6	0.6	0.6	0.6	0.6	0.6	0.6
k	1.0	0.1	0.1	0.1	0.1	0.9	0.3	0.1	0.5	0.2	0.4

Specimen	SC										
No.	4-2d	4-5	4-9	4-12	4-13a	4-13b	5-1	5-6	5-10	5-14c	5-14d
Unit #	3	3	3	3	3	3	2	2	2	2	2
fm	75.6	76.7	74.4	70.8	78.8	74.9	71.5	72.1	72.6	74.6	75.5
alk	1.7	1.0	0.8	1.0	1.3	1.2	1.1	1.8	1.7	0.3	0.3
al	22.0	21.4	23.6	27.6	18.8	23.0	26.5	25.4	25.1	24.5	23.6
c	0.7	0.9	1.2	0.6	1.1	0.9	0.9	0.7	0.6	0.6	0.6
si	182.4	179.7	204.9	262.0	142.5	222.6	181.9	175.3	200.3	157.3	148.4
ti	1.5	1.3	1.3	1.4	1.4	1.5	1.1	1.1	1.1	0.9	1.0
p	0.2	0.2	0.3	0.5	0.1	0.2	0.1	0.1	0.1	0.1	0.1
mg	0.6	0.6	0.7	0.6	0.6	0.6	0.4	0.5	0.5	0.5	0.5
k	0.4	0.2	0.1	0.1	0.1	-	1.0	0.9	1.0	0.9	1.0
Specimen	SC										
No.	5-18	5-19	5-25b	5-31	5-39	5-55	5-60	5-65	5-67	6-1c	6-1d
Unit #	2	2	2	3	3	QBG	QBG	QBG	QBG	1	1
fm	71.6	79.6	71.9	79.3	76.8	63.8	60.1	53.7	35.3	47.4	47.5
alk	1.5	1.3	0.9	1.0	1.7	4.3	7.5	4.7	5.4	9.0	10.1
al	26.6	18.8	26.7	19.4	21.2	31.0	31.8	41.2	59.1	42.4	41.0
c	0.3	0.3	0.5	0.3	0.3	0.9	0.6	0.4	0.2	1.2	1.4
si	172.3	212.7	211.4	135.9	138.0	496.1	462.3	654.4	1121.	278.4	342.2
ti	1.1	1.0	1.1	1.0	0.9	1.1	0.8	0.9	1.2	1.0	0.9
p	0.1	0.1	0.1	0.1	0.1	0.2	0.1	0.1	-	0.3	0.2
mg	0.6	0.7	0.6	0.7	0.7	0.6	0.7	0.6	0.5	0.5	0.5
k	0.7	1.0	0.8	0.2	0.5	0.9	1.0	1.0	1.0	0.8	0.7
Specimen	SC										
No.	6-3	6-4	7-2	7-4	7-10	7-11	8-4	8-7	8-11	8-17c	8-17d
Unit #	2	2	1	1	SIL	SIL	3	3	3	3	3
fm	60.7	65.2	45.3	55.2	20.3	45.8	77.6	70.3	71.3	76.3	76.7
alk	6.6	3.8	1.2	3.2	3.4	6.9	1.1	2.0	3.6	1.3	1.0
al	31.6	30.0	50.4	40.4	66.3	43.7	20.9	27.2	24.7	21.5	21.2
c	1.1	1.0	3.1	1.2	9.8	3.6	0.4	0.5	0.4	0.9	1.1
si	362.1	387.5	252.1	185.1	353.5	251.5	143.7	151.7	176.9	117.2	107.2
ti	1.4	1.1	0.3	0.8	0.9	2.1	1.2	1.1	0.9	1.0	1.2
p	0.3	0.4	0.9	0.2	3.7	1.3	0.1	0.1	0.1	0.1	0.1
mg	0.5	0.4	0.5	0.4	0.6	0.6	0.7	0.7	0.7	0.7	0.7
k	0.8	0.7	0.9	1.0	1.0	1.0	0.1	0.5	0.5	0.3	0.4

.

	1		· · · · · ·	····		- <u>,</u>					
Specimer	n SC	SC	ML	ML	ML	ML	ML	ML	ML	ML	ML
No.	8-21	8-28	05	06	11	12	15	18	23	31	33
Unit ≇	3	3	SIL	1	SIL	2	3	SIL	2	2	3
fm alk al c si ti p mg k	69.3 1.1 29.0 0.6 157.3 1.1 0.2 0.6 0.3	77.2 1.2 21.0 0.6 146.0 1.3 0.1 0.6 0.4	19.9 3.6 73.9 2.6 705.4 1.0 0.8 0.6 1.0	69.9 2.1 27.0 1.0 91.0 0.1 0.5 0.7	5.7 10.1 80.5 3.7 5295. 0.3 - 0.3	73.8 0.3 25.0 200.0 1.1 0.1 0.4 0.3	72.1 1.2 26.2 0.5 90.1 1.0 0.2 0.7 0.2	66.5 4.0 28.9 0.6 356.0 1.4 0.1 0.6 0.6	74.5 1.7 23.4 0.4 127.0 0.6 0.1 0.6 0.5	58.9 7.4 33.0 0.7 186.0 1.1 0.2 0.4 0.9	71.3 0.9 27.3 0.6 112.0 1.1 0.1 0.7 0.2
Specimen	ML	ML	ML	ML	ML	ML	ML	ML	ML	ML	ML
No.	34a	34b	46	47	53	61	65	67	68a	68b	69a
Unit ≇	QBG	QBG	QBG	2(M)	2(M)	2	QBG	2	QBG	QBG	2
fm	43.9	44.2	53.4	71.8	72.5	73.3	26.1	75.1	55.1	56.2	69.7
alk	7.1	8.0	5.6	2.2	1.1	1.4	25.0	1.7	6.0	6.4	1.4
al	48.3	46.8	40.0	23.7	24.7	24.5	41.7	22.7	37.9	36.4	28.2
c	0.7	1.0	1.0	2.3	1.7	0.8	7.2	0.5	1.0	1.0	0.7
si	765.0	740.0	421.0	131.0	142.0	158.0	426.0	180.0	423.0	457.0	176.0
ti	1.1	1.1	0.7	1.2	1.2	1.1	1.1	1.0	1.3	1.4	1.1
p	0.1	0.1	0.1	0.4	0.5	0.2	0.2	0.2	0.3	0.3	0.2
mg	0.6	0.6	0.6	0.7	0.5	0.6	0.4	0.7	0.7	0.7	0.6
k	0.9	0.7	0.8	0.3	-	0.4	0.3	0.5	0.9	0.8	0.8
Specimen	ML	ML	ML	ML	ML	ML	ML	ML	ML	ML	ML
No.	695	70	72	74	78	79	80	81	82	91	104
Unit #	2	QBG	OBG	2	2	1	3	3	3	3(M)	QBG
fm	71.0	55.0	44.4	74.8	72.6	68.3	74.6	71.5	71.7	70.3	11.3
alk	1.4	7.4	6.9	1.5	1.0	2.4	1.1	1.0	1.3	2.5	16.8
al	27.1	36.3	15.9	23.3	25.9	27.7	23.2	26.4	26.1	26.2	43.0
c	0.5	1.3	32.8	0.4	0.5	1.6	1.1	1.1	0.9	1.0	28.9
si	174.0	402.0	155.0	246.0	164.0	199.0	198.0	203.0	196.0	292.0	500.0
ti	1.0	1.0	0.3	1.1	1.0	1.0	1.5	1.5	1.5	0.7	0.8
p	0.2	0.3	-	0.1	0.1	0.2	0.3	0.3	0.2	0.1	-
mg	0.6	0.6	0.8	0.6	0.6	0.4	0.6	0.6	0.7	0.6	0.6
k	0.8	0.8	0.9	0.2	0.2	1.0	-	0.1	0.4	0.4	0.8

APPENDIX D MINERAL CHEMISTRY OF ROCKS FROM STAR LAKE NEAR SHERRIDON MANITOBA

Four hundred, seventy-two mineral analyses from the Star Lake - Elken Lake area, near Sherridon, are compiled in this appendix. The mineral analyses are listed by field sample identification number and unit number. Field sample numbers refer to figures 3 - 4 and Table 2; unit numbers refer to the field units:

> U-1: unit 1 (quartz-garnet-anthophyllite), U-2: unit 2 (garnet-anthophyllites), U-3: unit 3 (anthophyllite-cordierite), QBG: quartz-biotite gneisses, U-2(M): the mineralized unit 2 (garnet-

anthophyllite) of Elken Lake and

U-3(M): unit 3 (anthophyllite-cordierite)

near the east orebody at Sherridon

Manitoba. (Sherridon unit 3)

Analyses of minerals was done with a JEOL 733 electron microprobe with full Tracor Northern automation at the National Museum of Natural Sciences in Ottawa. Standards, spectrometer positions and crystals used are listed in Tables D1 and D2. The ZAF setup table is also listed in Table D3.

-132-

			·····							
EL	SPEC	XSTAL	POS	BKG1	BKG2	CF	BASLN	WIND	BIAS	STD
MG SI CA FE AL K CR ZN FL NA TI CL	1 4 3 2 2 2 3 4 4	TAP TAP PET LIF TAP PET TAP TAP LIF PET PET	107.735 77.700 107.580 134.545 90.895 119.815 159.145 133.450 199.555 129.675 146.085 88.065 151.370	-2.5 -2.2 -3.5 -2.5 -2.5 -2.5 -3.5 -6.0 -4.5 -3.0 -3.5 -2.0 -2.0	$\begin{array}{c} 2.5\\ 2.9\\ 2.0\\ 2.5\\ 2.0\\ 2.5\\ 3.5\\ 2.0\\ 3.5\\ 2.0\\ 2.0\\ 2.0\\ 2.0\\ 2.0\end{array}$	0.078 0.206 0.168 0.149 0.150 0.087 0.439 0.103 0.008 0.008 0.050 0.494 0.176 0.061	0.8 1.0 2.0 4.5 1.0 1.5 3.0 0.8 0.5 0.8 3.0 2.0 1.0	$\begin{array}{c} 6.0\\ 5.0\\ 5.0\\ 5.0\\ 4.0\\ 6.0\\ 6.0\\ 5.0\\ 6.0\\ 6.0\\ 6.0\\ 6.0\\ 6.0\\ 6.0\\ 6.0\\ \end{array}$	1750 1750 1750 1700 1750 1750 1750 1750	DIOPSI DIOPSI ALMAN2 GEHLEN SANIDI NICHRC BAZNGE FLRIEB ALBITE TEPHRO TITANI TUGTUP

TABLE D1: AMPHIBOLE - GARNET ELEMENT TABLE

1

TABLE D2: BIOTITE - CHLORITE ELEMENT TABLE

EL	SPEC	XSTAL	POS	BKG1	BKG2	CF	BASLN	WIND	BIAS	STD
AL SI FE MG CA CR NA TI MN K ZN FL	1 1 3 1 4 3 2 4 3 4 2 4 2 4 2	TAP TAP LIF TAP PET LIF TAP PET TAP PET TAP	90.895 77.690 134.545 107.735 107.575 159.150 129.675 88.065 146.070 119.810 133.450 151.355 199.555	-2.5 -2.5 -2.5 -2.5 -2.0 -3.5 -3.0 -3.5 -3.0 -3.5 -3.5 -3.5 -2.0 -3.5 -2.0 -3.5 -4.5	2.5 2.5 2.0 2.5 2.0 2.0 2.0 2.0 2.0 2.0 2.0 3.5 2.0 3.5	0.059 0.096 0.071 0.144 0.028 0.348 0.050 0.009 0.035 0.073 0.102 0.061 0.008	1.0 1.0 4.5 0.8 2.0 3.0 0.8 2.0 3.0 2.0 0.8 1.0 0.5	$\begin{array}{c} 5.0\\ 5.0\\ 5.0\\ 5.0\\ 6.0\\ 6.0\\ 6.0\\ 6.0\\ 6.0\\ 6.0\\ 6.0\\ 6$	1750 1750 1700 1750 1750 1750 1750 1750	CHLORI CHLORI BIOTIT CHLORI ALMAN2 CHROMI ALBITE BIOTIT BIOTIT BIOTIT BIOTIT BIOTIT FIRIFR

TABLE D3: ZAF SETUP

Acc. Vol	tage= 15KV	Take Off Angle= 40.0 Deg
ELEMENT	LINE	
NA20 K20 CA0 MG0 MN0 FE0 AL203 CR203 SI02 TI02 0	K K K K K K K K	WDS REF: NA WDS REF: K WDS REF: CA WDS REF: MG WDS REF: MN NDS REF: FE WDS REF: FE WDS REF: CR WDS REF: SI WDS REF: SI WDS REF: TI BY DIFFERENCE

-

•

	<u>-~</u>	1	,	·····							
Specimen No.	SC 3-10	SC 3-10	SC 3-10	SC 3-10	SC 3-10	SC 3-10	SC 3-25a	SC 3-25a	SC 3-25a	SC 3-25a	SC 3-25a
Unit #	1	1	1	1	1	1	2	2	2	2	2
Anth/Ged	ged	ged	ged	ged	ged	ged	anth	anth	anth	ged	anth
Si02 Ti02 Al203 Mg0 Ca0 Mn0 Fe0 Na20 F 0-E	47.37 0.19 12.59 16.37 0.08 0.12 20.12 1.34	47.95 0.20 10.72 17.30 0.11 0.09 19.94 1.12	46.00 0.12 14.26 16.92 0.09 	45.41 0.11 14.05 17.03 0.10 0.13 18.46 1.56	46.03 0.13 13.83 17.41 0.09 0.12 17.50 1.38	46.77 0.13 12.91 17.58 0.11 	49.41 0.26 9.57 19.12 0.17 	49.23 0.29 9.65 19.13 0.18 	49.13 0.33 9.52 19.19 0.17 - 17.73 0.99	48.12 0.30 11.44 19.06 0.18 	49.76 0.31 9.90 19.08 0.19 - 17.93 1.00
Total	98.18	97.43	- 97.14	- 96.85	- 96.49	- 96.93	- 97.79	- 97.11	- 97.06	97.40	- 98 17
Si * Al Z Al Ti Mg Ca Mn Fe Na	6.81 1.19 8.00 0.94 0.02 3.51 0.01 0.02 2.42 0.08	6.94 1.06 8.00 0.77 0.02 3.73 0.02 0.01 2.41 0.04	6.64 1.36 8.00 1.07 0.01 3.64 0.01 - 2.20 0.07	6.59 1.41 8.00 1.00 0.01 3.69 0.02 0.02 2.24 0.02	6.67 1.33 8.00 1.03 0.01 3.76 0.01 0.02 2.12 0.05	6.75 1.25 8.00 0.95 0.01 3.78 0.02 - 2.19 0.05	7.06 0.94 8.00 0.67 0.03 4.07 0.03 2.17 0.03	7.06 0.94 8.00 0.69 0.03 4.09 0.03 - 2.11 0.05	7.06 0.94 8.00 0.67 0.04 4.11 0.03 - 2.13 0.02	6.88 1.22 8.00 0.71 0.03 4.06 0.03 - 2.04 0.13	7.06 0.94 8.00 0.72 0.03 4.04 0.03 - 2.13 0.05
(x+y)	7.00	7.00	7.00	7.00	7.00	7.00	7.00	7.00	7.00	7.00	7.00
Na	0.30	0.27	0.36	0.42	0.34	0.31	0.27	0.24	0.26	0.21	0.23
W	0.30	0.27	0.36	0.42	0.34	0.31	0.27	0.24	0.26	0.21	0.23

* number of ions on the basis of 23 oxygens - not detected

.

.

<u> </u>											
Specimen No.	SC 3-25a	SC 3-25a	SC 3-25a	SC 3-25a	SC 3-25a	SC 3-25a	SC 3-25a	SC 3-26a	SC 3-26a	SC 3-26a	SC 3-26a
Unit #	2	2	2	2	2	2	2	2	2	2	2
Anth/Ged	ged	ged	ged	anth	ged	ged	anth	anth	ged	anth	anth
Si02 Ti02 Al203 Mg0 Ca0 Mn0 Fe0 Na20 F 0=F	48.49 0.26 11.16 19.26 0.20 - 16.97 1.17 -	49.35 0.28 10.42 19.57 0.16 0.11 16.79 1.08 0.39 -0.16	49.12 0.29 10.76 18.88 0.17 - 17.77 1.13 -	49.31 0.32 10.37 19.19 0.18 - 17.27 1.08 -	48.47 0.31 11.57 18.94 0.20 - 16.97 1.24 -	48.09 0.32 12.03 18.78 0.20 0.09 17.01 1.27	49.66 0.35 8.95 18.79 0.17 0.11 18.51 1.01 -	49.87 0.32 9.48 19.48 0.18 - 17.88 0.97 0.40 -0.17	48.99 0.33 10.73 19.44 0.14 - 17.40 1.07 -	49.14 0.28 9.43 19.89 0.18 0.10 17.37 0.93	49.76 0.24 9.45 19.81 0.16 0.12 17.31 0.90
Total	97.51	97.99	98.12	97.72	97.70	97.79	97.55	98.41	98.10	97.32	97.75
Si * Al Z	6.91 1.09 8.00	6.98 1.02 8.00	6.98 1.02 8.00	7.01 0.99 8.00	6.89 1.11 8.00	6.84 1.16 8.00	7.12 0.88 8.00	7.05 0.95 8.00	6.95 1.05 8.00	7.03 0.97 8.00	7.08 0.92 8.00
Al Ti Mg Ca Mn Fe Na	0.79 0.03 4.09 0.03 2.02 0.04	0.72 0.03 4.13 0.02 0.01 1.99 0.10	0.78 0.03 4.00 0.03 2.11 0.05	0.75 0.03 4.07 0.03 2.05 0.07	0.83 0.03 4.02 0.03 2.02 0.07	0.86 0.03 3.98 0.03 0.01 2.02 0.07	0.63 0.04 4.02 0.03 0.01 2.22 0.05	0.63 0.03 4.10 0.03 	0.74 0.04 4.11 0.02 2.06 0.03	0.62 0.03 4.24 0.03 0.01 2.08 0.00	0.66 0.03 4.20 0.02 0.01 2.06 0.02
(x+y)	7.00	7.00	7.00	7.00	7.00	7.00	7.00	7.00	7.00	7.00	7.00
Na	0.28	0.20	0.26	0.23	0.27	0.28	0.23	0.17	0.26	0.26	0.23
W	0.28	0.20	0.26	0.23	0.27	0.28	0.23	0.17	0.26	0.26	0.23
			1	1		1		1			1

* number of ions on the basis of 23 oxygens
- not detected

					_						
Specimen No.	SC 3-26a	SC 3-29	SC 3-29	SC 3-29	SC 3-29	SC 3-34	SC 3-34	SC 3-34	SC 3-34	SC 3-34	SC 3-34
Unit #	2	2	2	2	2	2	2	2	2	2	2
Anth/Ged	anth	ged	ged	anth	ged	ged	anth	anth	anth	ged	ged
Si02 Ti02 Al203 Mg0 Ca0 Mn0 Fe0 Na20 F 0=F	50.58 0.28 9.02 19.73 0.15 17.45 0.89	44.32 0.32 15.40 15.23 0.19 0.26 20.23 1.87 -	48.10 0.40 10.29 16.90 0.20 0.32 20.64 1.22	48.64 0.32 9.35 18.37 0.18 0.23 19.12 1.12 -	47.26 0.40 11.07 17.50 0.21 0.28 19.17 1.29	47.97 0.31 10.32 17.25 0.18 0.18 20.03 1.16	49.49 0.26 8.58 19.10 0.16 0.12 18.50 0.93 -	49.42 0.29 8.84 18.76 0.19 0.14 18.78 1.04	49.67 0.28 8.97 19.14 0.17 0.15 18.07 1.05 -	48.34 0.25 9.92 18.01 0.23 0.16 19.10 1.22	48.17 0.38 10.65 17.98 0.20 0.20 18.95 1.30 -
Total	98.10	97.82	98.07	97.33	97.18	97.40	97.14	97.46	97.50	97.23	97.83
Si * Al Z	7.16 0.84 8.00	6.44 1.66 8.00	6.95 1.05 8.00	7.03 0.97 8.00	6.86 1.14 8.00	6.96 1.04 8.00	7.13 0.87 8.00	7.11 0.89 8.00	7.11 0.89 8.00	6.99 1.01 8.00	6.92 1.08 8.00
Al Ti Ca Mn Fe Na	0.66 0.03 4.16 0.02 2.07 0.06	0.98 0.04 3.30 0.03 0.03 2.46 0.16	0.70 0.04 3.64 0.03 0.04 2.50 0.05	0.62 0.04 3.96 0.03 0.03 2.31 0.01	0.75 0.04 3.79 0.03 0.03 2.33 0.03	0.72 0.03 3.73 0.03 0.02 2.43 0.04	0.59 0.03 4.10 0.03 0.02 2.23 0.00	0.61 0.03 4.02 0.03 0.02 2.26 0.03	0.62 0.03 4.09 0.03 0.02 2.16 0.05	0.68 0.03 3.88 0.04 0.02 2.31 0.04	0.72 0.04 3.85 0.03 0.02 2.28 0.06
(x+y)	7.00	7.00	7.00	7.00	7.00	7.00	7.00	7.00	7.00	7.00	7.00
Na	0.18	0.37	0.29	0.30	0.33	0.29	0.26	0.26	0.24	0.30	0.30
W	0.18	0.37	0.29	0.30	0.33	0.29	0.26	0.26	0.24	0.30	0.30
i						1		1	1	1	

* number of ions on the basis of 23 oxygens
- not detected

•

•

AMPHIBOL	ES										
Specimen No.	SC 3-43	SC 3-43	SC 3-43	SC 3-43	SC 3-54a	SC 3-54a	SC 3-54a	SC 3-54a	SC 3-54a	SC 3-54a	SC 3-54a
Unit #	2	2	2	2	2	2	2	2	2	2	2
Anth/Ged	anth	anth	ged	ged	anth	ged	ged	anth	anth	ged	anth
Si02 Ti02 Al203 Mg0 Ca0 Mn0 Fe0 Na20 F 0=F	49.77 0.29 8.87 18.70 0.15 0.17 19.09 0.89	49.56 0.23 9.79 19.44 0.15 0.17 17.20 0.95 0.39 -0.16	48.99 0.29 10.78 18.98 0.16 0.14 17.45 1.08	47.91 0.32 11.64 18.01 0.20 0.13 18.31 1.20	51.71 0.22 7.25 20.42 0.15 	49.08 0.22 10.95 19.18 0.18 	47.24 0.29 12.70 18.67 0.18 0.09 17.25 1.31	51.47 0.17 7.54 20.51 0.13 - 17.50 0.75 0.47 -0.20	51.33 0.18 7.64 20.50 0.14 0.09 17.53 0.75 -	49.10 0.24 10.92 19.19 0.18 - 16.84 1.15 -	50.30 0.32 9.11 19.95 0.14 0.10 17.50 0.89
Total	97.93	97.72	97.87	97.72	97.55	97.36	97.73	98.34	98.16	97.62	98.31
Si * Al Z	7.13 0.87 8.00	7.04 0.96 8.00	6.97 1.03 8.00	6.86 1.14 8.00	7.34 0.66 8.00	6.99 1.01 8.00	6.74 1.26 8.00	7.25 0.75 8.00	7.26 0.74 8.00	6.98 1.02 8.00	7.11 0.89 8.00
Al Ti Mg Ca Mn Fe Na	0.63 0.03 3.99 0.02 0.02 2.29 0.02	0.68 0.03 4.12 0.02 0.02 2.04 0.09	0.78 0.03 4.03 0.02 0.02 2.08 0.04	0.83 0.03 3.85 0.03 0.02 2.19 0.05	0.55 0.02 4.32 0.02 2.03 0.06	0.83 0.02 4.07 0.03 1.99 0.06	0.88 0.03 3.97 0.03 0.01 2.06 0.02	0.50 0.02 4.31 0.02 	0.53 0.02 4.32 0.02 0.01 2.07 0.03	0.81 0.03 4.07 0.03 - 2.00 0.06	0.63 0.03 4.21 0.02 0.01 2.07 0.03
(x+y)	7.00	7.00	7.00	7.00	7.00	7.00	7.00	7.00	7.00	7.00	7.00
Na	0.23	0.17	0.26	0.28	0.12	0.22	0.34	0.12	0.18	0.26	0.21
W	0.23	0.17	0.26	0.28	0.12	0.22	0.34	0.12	0.18	0.26	0.21

* number of ions on the basis of 23 oxygens - not detected

~
	<u></u>		· · · · · · · · · · · · · · · · · · ·	· · · · · · · · · · · · · · · · · · ·							
Specimen No.	SC 3-54a	SC 3-54a	SC 3-54a	SC 3-54a	SC 3-54a	SC 4-2b	SC 4-2b	SC 4-2b	SC 4-2b	SC 4-2b	SC 4-9
Unit #	2	2	2	2	2	3	3	3	3	3	3
Anth/Ged	ged	ged	ged	ged	ged	ged	ged	ged	ged	ged	ged
Si02 Ti02 Al203 Mg0 Ca0 Mn0 Fe0 Na20 F 0=F	47.20 0.31 13.35 18.63 0.18 16.97 1.33 	48.69 0.34 10.97 19.10 0.19 	49.54 0.35 10.61 19.30 0.17 - 17.68 1.12 -	46.67 0.25 13.82 18.37 0.17 0.09 17.13 1.45	47.79 0.30 12.38 18.97 0.18 17.49 1.29	47.31 0.30 11.73 17.52 0.16 	48.65 0.31 11.65 18.16 0.18 - 18.23 1.14 -	48.50 0.30 11.30 18.06 0.20 	48.19 0.25 11.59 18.08 0.20 0.10 17.78 1.09	47.68 0.31 12.17 17.99 0.19 0.10 17.83 1.13	48.08 0.31 11.32 17.75 0.30 18.98 1.07
Total	97.97	97.55	98.77	97.95	98.40	96.64	98.32	97.32	97.28	97.40	97.81
Si * Al 7	6.71 1.29 8.00	6.94 1.06	6.98 1.02	6.65 1.35	6.77 1.23	6.86 1.14	6.91 1.09	6.95 1.05	6.91 1.09	6.84 1.16	6.90 1.10
Al Ti Mg Ca Mn	0.95 0.03 3.95 0.03	0.78 0.04 4.06 0.03	0.74 0.04 4.06 0.03	0.00 0.97 0.03 3.90 0.03 0.01	0.84 0.03 4.01 0.03	0.86 0.03 3.79 0.03	0.86 0.03 3.84 0.03	0.86 0.03 3.86 0.03	8.00 0.87 0.03 3.86 0.03 0.01	8.00 0.90 0.03 3.84 0.03	8.00 0.81 0.03 3.80 0.05
Fe Na	2.02 0.02	2.04 0.05	2.08 0.05	2.04 0.02	2.07 0.02	2.23 0.06	2.17 0.07	2.15 0.07	2.13	2.14	2.28 0.03
(x+y)	7.00	7.00	7.00	7.00	7.00	7.00	7.00	7.00	7.00	7.00	7.00
Na	0.35	0.27	0.26	0.38	0.34	0.28	0.24	0.22	0.23	0.26	0.27
W	0.35	0.27	0.26	0.38	0.34	0.28	0.24	0.22	0.23	0.26	0.27

AMPHIBOLES

* number of ions on the basis of 23 oxygens
- not detected

a P Santa Santa

AMPHIBOL	ES										
Specimen No.	SC 4-9	SC 4-9	SC 4-9	SC 4-9	SC 4-12	SC 4-12	SC 4-12	SC 4-12	SC 4-13b	SC 4-13b	SC 4-13b
Unit ≇	3	3	3	3	3	3	3	3	3	3	3
Anth/Ged	anth	anth	anth	ged	ged	ged	ged	ged	ged	ged	ged
SiO2 TiO2 Al2O3 MgO CaO MnO FeO Na2O F	49.35 0.21 10.26 18.43 0.29 	51.76 0.17 5.27 19.59 0.24 	48.67 0.36 10.20 17.79 0.30 	48.65 0.30 11.02 18.05 0.29 	47.98 0.28 10.13 16.14 0.31 	46.15 0.32 11.93 15.18 0.23 	47.09 0.30 12.44 15.40 0.24 21.73 1.25	45.45 0.32 13.14 14.39 0.28 0.12 22.30 1.42	47.91 0.27 10.16 17.45 0.32 	47.43 0.35 10.66 17.49 0.35 	47.62 0.29 11.62 16.88 0.34
Total	98.20	97.37	- 97.59	- 98.16	- 97.67	- 97.32	- 98.45	- 97.42	- 96.29	- 96.50	- 98.02
Si * Al Z Al Ti	7.03 0.97 8.00 0.75 0.02	7.46 0.54 8.00 0.36 0.02	7.00 1.00 8.00 0.73 0.04	6.94 1.06 8.00 0.79 0.03	6.99 1.01 8.00 0.73 0.03	6.78 1.22 8.00 0.85 0.04	6.80 1.20 8.00 0.92 0.03	6.68 1.32 8.00 0.96 0.04	6.99 1.01 8.00 0.74 0.03	6.92 1.08 8.00 0.75 0.04	6.86 1.14 8.00 0.83 0.03
Ca Mn Fe Na (x+y)	0.04 2.22 0.06 7.00	0.04 2.39 0.00 7.00	0.02 0.05 2.32 0.04 7.00	0.04 0.04 2.24 0.06 7.00	0.05 0.05 2.65 0.04 7.00	0.04 2.73 0.02 7.00	0.04 - 2.63 0.06 7.00	5.13 0.04 0.02 2.74 0.05 7.00	0.05 2.33 0.05 7.00	0.06 2.34 0.01 7.00	0.05 2.42 0.05 7.00
Na ₩	0.22 0.22	0.13 0.13	0.24 0.24	0.24 0.24	0.26 0.26	0.35 0.35	0.29 0.29	0.36 0.36	0.26	0.29	0.29

* number of ions on the basis of 23 oxygens
- not detected

		·	·····								
Specimen No.	SC 4-13b	SC 4-13b	SC 5-1	SC 5-1	SC 5-9	SC 5-16	SC 5-16	SC 5-21a	SC 5-21a	SC 5-21a	SC 5-21a
Unit #	3	3	2	2	2	2	2	2	2	2	2
Anth/Ged	ged	anth	anth	ged	anth	anth	ged	ged	anth	ged	ged
Si02 Ti02 Al203 Mg0 Ca0 Mn0 Fe0 Na20 F 0=F	48.31 0.33 10.29 17.16 0.29 	49.07 0.30 9.55 17.22 0.33 - 20.43 1.09	48.82 0.34 9.83 18.49 0.11 18.80 1.16	48.83 0.29 10.14 18.37 0.13 	50.53 0.20 9.42 19.64 0.12 17.57 0.69	49.68 0.26 10.65 19.21 0.12 18.04 0.75	48.53 0.31 11.38 17.90 0.11 19.80 0.93	48.17 0.14 13.06 18.36 0.15 17.23 0.90	51.12 0.12 7.55 19.46 0.10 18.75 0.43	48.95 0.16 11.68 18.87 0.12 0.11 17.65 0.72	48.50 0.26 12.46 18.30 0.15
Total	97.30	97.99	97.55	97.79	98.17	98.71	98.96	98.07	97.53	98.26	98.24
Si * Al Z	7.00 1.00 8.00	7.07 0.93 8.00	7.02 0.98 8.00	7.00 1.00 8.00	7.14 0.86 8.00	7.00 1.00 8.00	6.89 1.11 8.00	6.82 1.18 8.00	7.30 0.70 8.00	6.92 1.08 8.00	6.85 1.15 8.00
Al Ti Mg Ca Mn Fe	0.76 0.04 3.70 0.05 - 2.40	0.69 0.03 3.70 0.05 2.46	0.69 0.04 3.96 0.02 - 2.26	0.71 0.03 3.93 0.02 2.25	0.71 0.02 4.14 0.02 - 2.08	0.77 0.03 4.04 0.02 2.13	0.80 0.03 3.79 0.02 2.35	1.00 0.02 3.87 0.02 - 2.04	0.57 0.01 4.15 0.02 2.24	0.87 0.02 3.98 0.02 0.01 2.09	0.92 0.03 3.85 0.02 - 2.07
Na (x+y)	0.05	0.07	0.03	0.06	0.03	0.01	0.01	0.05	0.01	0.01	0.11
Na	0.27	0.23	0.29	0.28	0.16	0.20	0.25	0.20	0.11	0.19	0.12
W	0.27	0.23	0.29	0.28	0.16	0.20	0.25	0.20	0.11	0.19	0.12

AMPHIBOLES

AMPHIBOLES

	i										
Specimen No.	SC 5-25a	SC 5-25a	SC 5-25a	SC 5-39	SC 5-39	SC 5-39	SC 5-39	SC 5-35	SC 5-35	SC 5-35	SC 5-35
Unit #	2	2	2	3	3	3	3	3	3	3	3
Anth/Ged	anth	anth	ged	ged	ged	ged	ged	ged	ged	ged	anth
Si02 Ti02 Al203 Mg0 Ca0 Mn0 Fe0 Na20 F 0=F	49.90 0.26 9.67 18.60 0.11 0.09 18.96 0.66	49.69 0.26 9.76 18.31 0.15 0.10 19.08 0.75 -	47.30 0.24 12.53 17.02 0.12 - 19.30 0.91 -	47.46 0.28 12.92 17.36 0.14 0.16 19.01 1.06 -	48.32 0.23 10.23 18.56 0.14 0.13 18.74 0.92	48.06 0.23 11.91 18.45 0.14 0.20 18.17 1.04	48.67 0.19 10.89 18.27 0.19 0.12 17.87 0.93	46.66 0.40 12.82 16.77 0.17 0.17 20.18 1.31	46.59 0.32 13.03 16.53 0.18 0.28 20.85 1.33	46.61 0.63 12.28 16.02 0.15 0.19 21.16 1.30	49.38 0.31 8.32 18.43 0.13 0.18 20.10 0.80 -
Total	98.25	98.10	97.42	98.39	97.27	98.20	97.13	98.48	99.11	98.34	97.65
Si * Al Z	7.10 0.90 8.00	7.09 0.91 8.00	6.82 1.18 8.00	6.77 1.23 8.00	6.97 1.03 8.00	6.84 1.16 8.00	6.99 1.01 8.00	6.71 1.29 8.00	6.68 1.32 8.00	6.74 1.26 8.00	7.13 0.87 8.00
Al Ti Mg Ca Mn Fe Na	0.72 0.03 3.94 0.02 0.01 2.26 0.02	0.73 0.03 3.89 0.02 0.01 2.28 0.04	0.95 0.03 3.66 0.02 2.33 0.01	0.94 0.03 3.69 0.02 0.02 2.27 0.03	0.71 0.03 3.99 0.02 0.02 2.26 0.00	0.84 0.03 3.92 0.02 0.02 2.16 0.01	0.83 0.02 3.91 0.03 0.02 2.15 0.04	0.88 0.04 3.59 0.03 0.02 2.43 0.01	0.88 0.03 3.53 0.03 0.03 2.50 0.00	0.83 0.07 3.45 0.02 0.02 2.56 0.05	0.55 0.03 3.97 0.02 0.02 2.43 0.00
(x+y)	7.00	7.00	7.00	7.00	7.00	7.00	7.00	7.00	7.00	7.00	7.00
Na	0.16	0.17	0.24	0.26	0.26	0.28	0.22	0.36	0.37	0.32	0.22
W	0.16	0.17	0.24	0.26	0.26	0.28	0.22	0.36	0.37	0.32	0.22
			1	1		1			1		

Specimen No.	SC 5-35	SC 8-4	SC 8-4	SC 8-7	SC 8-7	SC 8-7	SC 8-7	SC 8-15	SC 8-15	SC 8-15	SC 8-17
Unit #	3	3	3	3	3	3	3	3	3	3	3
Anth/Ged	ged	anth	anth	anth	anth	anth	ged	anth	anth	anth	anth
SiO2 TiO2 Al2O3 MgO CaO MnO FeO Na2O F O=F	47.90 0.35 11.37 17.67 0.18 0.22 19.71 1.14	50.15 0.23 8.83 19.15 0.16 0.16 19.00 0.98 -	50.83 0.19 6.96 19.79 0.12 0.20 18.64 0.79	50.85 0.19 6.98 20.32 0.15 0.12 17.71 0.78	49.90 0.14 9.95 20.27 0.14 0.12 16.56 1.11	50.91 0.12 8.28 20.69 0.13 0.12 16.65 0.91	46.66 0.33 12.94 17.46 0.18 0.15 18.88 1.56	50.56 0.18 6.39 18.44 0.31 0.16 20.38 0.70	50.92 0.15 6.50 18.63 0.15 0.20 20.17 0.69	49.36 0.37 8.88 17.45 0.15 0.17 20.77 0.95	49.91 0.22 8.89 19.15 0.16 0.22 18.12 1.01
Total	98.54	98.66	97.52	97.10	98.19	97.81	98.16	97.12	97.41	98.10	97.68
Si * Al Z	6.85 1.15 8.00	7.12 0.88 8.00	7.29 0.71 8.00	7.29 0.71 8.00	7.04 0.96 8.00	7.20 0.80 8.00	6.69 1.31 8.00	7.34 0.66 8.00	7.36 0.64 8.00	7.11 0.89 8.00	7.13 0.87 8.00
Al Ti Mg Ca Mn Fe Na	$\begin{array}{c} 0.77\\ 0.04\\ 3.77\\ 0.03\\ 0.03\\ 2.36\\ 0.00\\ \end{array}$	0.60 0.03 4.05 0.02 0.02 2.26 0.02	$\begin{array}{c} 0.47\\ 0.02\\ 4.23\\ 0.02\\ 0.02\\ 2.24\\ 0.00\\ \end{array}$	$\begin{array}{c} 0.47\\ 0.02\\ 4.34\\ 0.02\\ 0.02\\ 2.12\\ 0.01\\ \end{array}$	0.70 0.02 4.26 0.02 0.01 1.95 0.04	0.58 0.01 4.37 0.02 0.01 1.97 0.04	0.88 0.04 3.73 0.03 0.02 2.27 0.03	0.43 0.02 3.99 0.05 0.02 2.48 0.01	0.47 0.02 4.01 0.02 0.02 2.44 0.02	0.62 0.04 3.75 0.02 0.02 2.50 0.05	0.63 0.02 4.08 0.03 0.03 2.17 0.04
(x+y)	7.00	7.00	7.00	7.00	7.00	7.00	7.00	7.00	7.00	7.00	7.00
Na	0.32	0.25	0.22	0.21	0.26	0.21	0.40	0.19	0.17	0.22	0.24
Ŵ	0.32	0.25	0.22	0.21	0.26	0.21	0.40	0.19	0.17	0.22	0.24
<u> </u>				i					1		

AMPHIBOLES

AMPHIBOL	ES										
Specimen No.	SC 8-17	SC 8-17	SC 8-17	SC 8-17	ML 55	ML 49	ML 49	ML 49	ML 49	ML 49	ML 49
Unit ∦	3	3	3	3	2(M)	2(M)	2(M)	2(M)	2(M)	2(M)	2(M)
Anth/Ged	anth	anth	anth	anth	ged	ged	ged	ged	ged	ged	ged
Si02 Ti02 Al203 Mg0 Ca0 Mn0 Fe0 Na20 F 0=F	49.09 0.19 10.29 18.85 0.17 0.20 17.74 1.13	51.96 0.15 6.00 20.12 0.16 0.18 18.45 0.63	52.42 0.16 5.35 20.09 0.13 0.22 18.83 0.56	49.48 0.22 8.96 19.03 0.19 0.20 18.76 0.99	40.83 0.43 17.64 12.17 0.12 0.19 22.65 1.84	47.03 0.40 11.99 18.06 0.19 0.11 17.82 1.15	44.83 0.32 15.12 17.39 0.19 0.08 17.16 1.49	45.14 0.41 15.09 17.25 0.20 	44.67 0.40 15.40 16.65 0.20 0.09 18.34 1.59	45.01 0.31 15.61 17.67 0.22 0.11 16.80 1.48	46.24 0.33 14.31 17.40 0.21 0.07 17.73 1.29
Total	97.66	97.65	97.76	97.83	95.87	96.75	96.58	96.85	97.34	97.21	97.58
Si * Al Z	7.01 0.99 8.00	7.42 0.58 8.00	7.48 0.52 8.00	7.09 0.91 8.00	6.15 1.85 8.00	6.80 1.20 8.00	6.49 1.51 8.00	6.51 1.49 8.00	6.46 1.54 8.00	6.46 1.54 8.00	6.62 1.38 8.00
Al Ti Mg Ca Mn Fe Na (x+y)	0.74 0.02 4.01 0.03 0.02 2.12 0.06 7.00	0.43 0.02 4.28 0.02 0.02 2.20 0.03 7.00	0.38 0.02 4.28 0.02 0.03 2.25 0.02 7.00	0.60 0.02 4.07 0.03 0.02 2.25 0.01 7.00	1.28 0.05 2.73 0.02 0.02 2.85 0.05 7.00	0.84 0.04 3.89 0.03 0.01 2.15 0.04 7.00	1.07 0.04 3.75 0.03 0.01 2.08 0.02 7.00	1.08 0.04 3.71 0.03 2.09 0.05 7.00	1.08 0.04 3.59 0.03 0.01 2.22 0.03 7.00	1.10 0.03 3.78 0.03 0.01 2.02 0.03 7.00	1.04 0.04 3.72 0.03 0.01 2.12 0.04
Na	0.25	0.14	0.14	0.27	0.49	0.28	0.40	0.35	0.42	0.38	0.32
W	0.25	0.14	0.14	0.27	0.49	0.28	0.40	0.35	0.42	0.38	0.32

ANFRIDUL	23									
Specimen No.	ML 49	ML 49	ML 52	ML 52	ML 52	ML 52	ML 52	ML 93	ML 93	
Unit #	2(M)	3(M)	3(M)							
Anth/Ged	ged	anth	anth							
Si02 Ti02 Al203 Mg0 Ca0 Mn0 Fe0 Na20 F 0=F	44.96 0.49 15.15 17.17 0.19 0.06 17.58 1.55	44.54 0.32 16.54 17.07 0.20 0.09 17.11 1.55 -	44.67 0.35 15.69 16.72 0.19 0.16 18.26 1.48	43.81 0.42 16.10 16.42 0.20 0.10 18.47 1.52	45.59 0.24 14.82 17.73 0.19 0.10 16.98 1.32	45.44 0.34 15.51 17.36 0.17 0.09 16.71 1.44 -	46.61 0.24 14.16 18.10 0.18 0.08 17.27 1.26	50.40 0.15 7.61 18.94 0.25 0.18 18.79 0.64	48.97 0.08 7.56 18.96 0.08 0.12 16.97 0.32	
Total	97.15	97.42	97.52	97.04	96.97	97.06	97.90	96.99	93.06	
Si * Al Z Al Ti Mg Ca Mn Fe Na	6.48 1.52 8.00 1.06 0.05 3.69 0.03 0.01 2.12 0.04	6.39 1.61 8.00 1.19 0.04 3.65 0.03 0.01 2.05 0.03	6.44 1.56 8.00 1.11 0.04 3.59 0.03 0.02 2.20 0.01	6.36 1.64 8.00 1.12 0.05 3.55 0.03 0.01 2.24 0.00	6.56 1.44 8.00 1.07 0.03 3.80 0.03 0.01 2.04 0.02	6.52 1.48 8.00 1.14 0.04 3.71 0.03 0.01 2.01 0.06	6.64 1.36 8.00 1.02 0.03 3.84 0.03 0.01 2.06 0.01	7.27 0.73 8.00 0.56 0.02 4.07 0.04 0.02 2.27 0.02	7.30 0.70 8.00 0.63 0.01 4.21 0.01 0.02 2.11 0.01	
(x+y)	7.00	7.00	7.00	7.00	7.00	7.00	7.00	7.00	7.00	
Na	0.39	0.40	0.40	0.43	0.35	0.34	0.34	0.16	0.08	
	0.39	0.40	0.40	0.43	0.35	0.34	0.34	0.16	0.08	

AMPHIRALES

GARNETS

GARNETS											
Specimen No.	SC 1-7	SC 1-7	SC 1-7	SC 1-7	SC 1-7	SC 1-7	SC 1-14a	SC 2-1	SC 2-1	SC 2-1	SC 2-1
Unit #	1	1	1	1	1	1	SIL	1	1	1	1
Si02 Ti02	38.09	38.54	38.42	38.55	38.72	38.20	38.20	38.82	39.33	39.24	38.85
Mg0 Ca0 Mn0 Fe0 Na20	4.18 0.44 1.54 35.59	4.68 0.47 1.32 34.77	4.94 0.42 1.34 34.76	4.05 0.50 1.58 35.46	21.47 5.70 0.50 0.82 34.26	21.03 4.14 0.47 1.78 35.09	4.23 0.58 5.11 32.00	21.73 8.22 0.90 1.00 29.51	21.79 8.76 0.65 1.05 29.29 -	21.83 8.04 0.64 1.14 30.46	21.71 7.61 0.69 1.14 30.55
Total	101.0	101.1	101.2	101.3	101.5	100.7	101.2	100.2	100.9	101.4	100.6
Si * Z Al Ti	3.02 3.02 1.98	3.03 3.03 1.98	3.02 3.02 1.97	3.04 3.04 1.97	3.02 3.02 1.97	3.03 3.03 1.97	3.02 3.02 1.97	3.01 3.01 1.99	3.02 3.02 1.97	3.02 3.02 1.98	3.01 3.01 1.96
Fe Fe Mg Mn Ca X	0.02 2.00 2.34 0.49 0.10 0.04 2.97	0.02 2.00 2.27 0.55 0.09 0.04 2.95	0.03 2.00 2.26 0.58 0.09 0.04 2.97	0.03 2.00 2.31 0.48 0.11 0.04 2.94	0.03 2.00 2.21 0.66 0.05 0.04 2.96	0.03 2.00 2.30 0.49 0.12 0.04 2.95	0.03 2.00 2.09 0.50 0.34 0.05 2.98	0.01 2.00 1.90 0.95 0.07 0.08 3.00	0.03 2.00 1.85 1.00 0.07 0.05 2.97	0.02 2.00 1.94 0.92 0.07 0.05 2.98	0.04 2.00 1.94 0.88 0.08 0.06 2.96
Specimen No.	SC 2-1	SC 2-1	SC 2-1	SC 2-9	SC 2-9	SC 2-9	SC 2-9	SC 2-9	SC 2-9	SC 2-9	SC 2-9
Unit #	1	1	1	1	1	1	1	1	1	1	1
Si02 Ti02 Al203 Mg0 Ca0 Mn0 Fe0 Na20	38.93 21.53 7.53 0.69 1.17 30.63 -	38.66 21.60 7.04 0.88 1.29 31.03	39.01 21.61 7.62 0.70 1.19 30.66	38.62 21.38 8.34 0.40 0.57 30.46	38.92 21.51 8.99 0.32 0.53 30.01	39.38 21.38 8.96 0.38 0.59 30.25 -	38.91 21.47 7.62 0.37 0.63 31.95	39.13 21.84 9.20 0.41 0.48 29.88	39.26 21.84 8.98 0.42 0.49 29.85 -	38.96 22.02 8.68 0.39 0.54 30.11	38.98 21.47 9.09 0.44 0.49 29.48
Total	100.5	100.5	100.8	99.77	100.3	100.9	101.0	100.9	100.8	100.7	99.95
Si * Z Al Ti	3.02 3.02 1.97	3.01 3.01 1.98	3.02 3.02 1.97	3.01 3.01 1.97	3.01 3.01 1.96	3.03 3.03 1.94	3.02 3.02 1.96	3.00 3.00 1.98	3.01 3.01 1.98	3.00 3.00 2.00	3.02 3.02 1.96
Fe Y Fe Mg Mn Ca X	0.03 2.00 1.96 0.87 0.08 0.06 2.97	0.02 2.00 2.00 0.82 0.09 0.07 2.98	0.03 2.00 1.96 0.88 0.08 0.08 0.06 2.98	0.03 2.00 1.96 0.97 0.04 0.03 3.00	0.04 2.00 1.90 1.04 0.04 0.03 3.01	0.06 2.00 1.88 1.03 0.04 0.03 2.98	0.04 2.00 2.03 0.88 0.04 0.03 2.98	$\begin{array}{c} 0.02\\ 2.00\\ 1.90\\ 1.05\\ 0.03\\ 0.03\\ 3.01 \end{array}$	0.02 2.00 1.90 1.03 0.03 0.04 3.00	- 2.00 1.94 1.00 0.04 0.03 3.01	0.04 2.00 1.87 1.05 0.03 0.04 2.99

GARNETS											
Specimen No.	SC 2-9	SC 2-9	SC 2-9	SC 2-9	SC 3-10	SC 3-10	SC 3-10	SC 3-25a	SC 3-25a	SC 3-29	SC 3-29
Unit #	1	1	1	1	1	1	1	2	2	2	2
Si02 Ti02 Al203 Mg0 Ca0 Mn0 Fe0 Na20	39.62 21.72 9.19 0.42 0.52 29.71	38.32 21.11 7.01 0.44 0.57 23.63	39.32 21.81 8.59 0.46 0.74 29.96	38.18 21.35 5.66 0.41 0.94 34.40	39.20 21.81 9.51 0.46 0.42 28.90	39.17 21.42 9.91 0.44 0.40 28.32	39.06 21.47 8.40 0.47 0.47 29.88	39.48 21.83 8.58 0.59 0.41 30.34	38.79 21.87 8.70 0.37 29.72	38.92 21.45 8.14 0.55 0.92 30.03	39.51 21.66 9.63 0.63 1.50 27.02
Total	101.2	91.08	100.9	100.9	100.3	99.66	99.75	101.2	99.45	100.0	99.95
Si * Z Al Ti	3.03 3.03 1.96	3.17 3.17 2.06	3.02 3.02 1.98	3.00 3.00 1.98	3.01 3.01 1.98	3.02 3.02 1.95	3.03 3.03 1.97	3.02 3.02 1.97	3.02 3.02 2.00	3.03 3.03 1.97	3.04 3.04 1.96
Fe Y Fe Mg Mn Ca X	0.04 2.00 1.87 1.05 0.03 0.03 2.98	2.00 1.63 0.86 0.04 0.04 2.57	0.02 2.00 1.91 0.98 0.05 0.04 2.98	0.02 2.00 2.24 0.66 0.06 0.04 3.00	0.02 2.00 1.84 1.09 0.03 0.04 3.00	0.05 2.00 1.78 1.14 0.03 0.04 2.99	0.03 2.00 1.91 0.97 0.03 0.04 2.95	0.03 2.00 1.91 0.98 0.03 0.05 2.97	- 2.00 1.93 1.01 0.02 - 2.96	0.03 2.00 1.92 0.94 0.06 0.05 2.97	0.04 2.00 1.70 1.10 0.10 0.05 2.95
Specimen No.	SC 3-29	SC 3-29	SC 3-29	SC 3-29	SC 3-34	SC 3-34	SC 3-34	SC 3-34	SC 3-34	SC 3-43	SC 3-43
Unit #	2	2	2	2	2	2	2	2	2	2	2
Si02 Ti02 Al203 Mg0 Ca0 Mn0 Fe0 Na20	39.66 21.72 9.69 0.55 1.67 27.07	39.85 21.58 10.02 0.55 1.03 27.62	39.34 21.80 9.82 0.58 1.22 27.64 -	38.97 21.48 9.46 0.55 0.87 28.32	39.96 22.09 10.25 0.57 0.72 27.89 -	39.64 21.82 9.62 0.64 1.39 27.22	39.63 21.82 9.82 0.59 1.00 28.03	39.31 21.89 9.29 0.54 0.69 29.06	39.17 21.31 8.62 0.59 0.92 29.35	39.28 0.08 21.17 7.43 2.06 2.43 27.78	39.30 21.45 7.71 1.81 2.10 28.48
Total	100.4	100.7	100.4	99.65	101.5	100.3	100.9	100.8	99.96	100.2	100.9
Si * Z Al Ti Fe Y Fe Mg Mn	3.04 3.04 1.96 0.04 2.00 1.69 1.11 0.11	3.04 3.04 1.94 0.06 2.00 1.70 1.14 0.07	3.01 3.01 1.97 0.03 2.00 1.74 1.12 0.08	3.01 3.01 1.96 0.04 2.00 1.79 1.09 0.04	3.02 3.02 1.97 0.03 2.00 1.73 1.16 0.05	3.03 3.03 1.97 0.03 2.00 1.71 1.10	3.02 3.02 1.96 0.04 2.00 1.75 1.12 0.07	3.01 3.01 1.98 0.02 2.00 1.84 1.06 0.05	3.04 3.04 1.95 0.05 2.00 1.85 1.00	3.05 3.05 1.94 0.01 0.05 2.00 1.75 0.86	3.03 3.03 1.95 0.05 2.00 1.79 0.89
Ca X	0.05	0.05	0.05	0.05	0.05	0.05	0.05	0.04 2.99	0.05	0.16	0.14 0.15 2.97

GARNETS											
Specimen No.	SC 3-43	SC 4-2b	SC 4-2b	SC 4-2b	SI 4						
Unit #	2	2	2	2	2	2	2	3	3	3	
Si02 Ti02 Al203	38.81 21.34	39.58 21.71	39.66 21.77	39.63 21.61	39.40 - 21.89	39.90 21.53	39.31 - 21.83	39.78 - 21.88	39.23	39.16 	38 21
Mg0 Ca0 Mn0 Fe0 Na20	7.53 2.06 2.27 27.32 -	9.23 0.60 0.84 28.86 -	9.37 0.55 0.49 28.80	9.32 0.51 1.15 28.45	9.18 0.60 1.17 27.98 -	9.74 0.55 0.90 28.22 -	9.47 0.53 0.48 28.70 -	9.82 0.54 0.30 28.39 -	9.85 0.58 0.34 28.23	9.11 0.56 0.34 29.08	7 0 0 32
Total	99.33	100.8	100.6	100.7	100.2	100.8	100.3	100.7	100.1	99.75	10
Si * Z Al Ti	3.03 3.03 1.97	3.03 3.03 1.96	3.03 3.03 1.96	3.04 3.04 1.95	3.03 3.03 1.98	3.04 3.04 1.94	3.02 3.02 1.98	3.03 3.03 1.97	3.01 3.01 1.98	3.03 3.03 1.96	33
Fe Y Fe Mg Mn Ca X	0.03 2.00 1.76 0.88 0.15 0.17 2.96	0.04 2.00 1.81 1.05 0.05 0.05 2.96	0.04 2.00 1.80 1.07 0.03 0.05 2.95	0.05 2.00 1.77 1.06 0.08 0.04 2.95	0.02 2.00 1.78 1.05 0.08 0.05 2.96	0.06 2.00 1.74 1.11 0.06 0.05 2.96	0.02 2.00 1.82 1.08 0.03 0.04 2.97	0.03 2.00 1.78 1.12 0.02 0.04 2.96	0.02 2.00 1.79 1.13 0.02 0.05 2.99	0.04 2.00 1.84 1.05 0.02 0.05 2.96	0 2 0 0 0
Specimen No.	SC 4-9	SC 4-12	SC 4-12	SC 4-12	SC 5-1	SC 5-1	SC 5-1	SC 5-1	SC 5-1	SC 5-1	S(5-
Unit #	3	3	3	3	2	2	2	2	2	2	2
Si02 Ti02 Al203 Mg0 Ca0 Mn0 Fe0 Ha20	38.60 21.34 8.04 0.84 0.34 31.35	38.87 21.40 7.08 0.88 0.41 31.89	38.68 - 21.44 7.26 0.95 0.35 32.08	38.79 21.25 6.77 0.98 0.33 32.70	39.55 22.18 10.34 0.31 0.21 28.25	39.45 21.83 9.87 0.34 0.31 28.52	39.74 22.09 9.67 0.35 0.43 28.19	39.71 21.93 9.31 0.37 0.36 29.57	39.51 22.09 9.09 0.28 0.20 30.15	39.24 21.59 7.94 0.24 0.15 31.72	39 21 9 0 0 29
Total	100.5	100.5	100.8	100.8	100.8	100.3	100.5	101.3	-	100.9	-
Si * Z Al	3.00 3.00 1.96	3.03 3.03 1.96	3.01 3.01 1.97	3.02 3.02 1.95	3.01 3.01 1.99	3.02 3.02 1.97	3.03 3.03 1.99	3.03 3.03 1.97	3.01 3.01 1.99	3.03 3.03 1.97	3. 3. 1.
Fe Y Fe Mg Mn Ca X	0.04 2.00 2.00 0.93 0.02 0.07 3.02	0.04 2.00 2.04 0.82 0.03 0.07 2.96	0.03 2.00 2.06 0.84 0.02 0.08 3.00	0.05 2.00 2.08 0.79 0.02 0.08 2.97	0.01 2.00 1.79 1.17 0.01 0.03 3.00	0.03 2.00 1.80 1.13 0.02 0.03 2.98	- 0.01 2.00 1.79 1.10 0.03 0.03 2.95	0.03 2.00 1.85 1.06 0.02 0.03 2.96	0.01 2.00 1.91 1.03 0.01 0.02 2.97	0.03 2.00 2.02 0.91 0.01 0.02 2.96	0. 2. 1. 1. 0. 2.

GARNETS											
Specimen No.	SC 5-1	SC 5-9	SC 5-9	SC 5-9	SC 5-9	SC 5-12	SC 5-12	SC 5-12	SC 5-12	SC 5-12	SC 5-12
Unit ∦	2	2	2	2	2	2	2	2	2	2	2
Si02 Ti02 Al203 Mg0 Ca0 Mn0 Fe0 Na20	39.36 21.74 9.79 0.31 0.19 29.10	40.25 23.02 10.82 0.33 0.22 27.90	39.98 	40.18 - 22.45 11.10 0.33 0.18 27.46 -	39.48 22.10 8.64 0.31 0.36 30.83	40.32 	39.98 22.42 10.05 0.33 0.20 29.14	39.80 22.46 10.59 0.36 0.36 28.37	39.56 22.07 9.10 0.37 0.70 29.40	40.26 22.35 10.64 0.24 0.24 28.17	40.10 22.45 11.31 0.34 0.45 27.24
Total	100.5	102.5	101.2	101.7	101.7	102.3	102.1	101.9	101.2	101.9	101.9
Si * Z Al Ti	3.02 3.02 1.96	3.00 3.00 2.02	3.02 3.02 1.97	3.01 3.01 1.98	3.01 3.01 1.99	3.00 3.00 2.00	3.01 3.01 1.99	2.99 2.99 1.99	3.02 3.02 1.98	3.02 3.02 1.98	3.00 3.00 1.98
Fe Y Fe Mg Mn Ca X	0.04 2.00 1.83 1.12 0.01 0.03 2.99	0.00 2.00 1.74 1.20 0.01 0.03 2.98	0.03 2.00 1.71 1.23 0.02 0.03 2.99	0.02 2.00 1.70 1.24 0.01 0.03 2.98	0.01 2.00 1.96 0.98 0.02 0.03 2.99	0.00 2.00 1.71 1.26 0.01 0.02 3.00	0.01 2.00 1.82 1.13 0.01 0.03 2.99	0.01 2.00 1.78 1.19 0.02 0.03 3.02	0.02 2.00 1.86 1.04 0.05 0.03 2.98	0.02 2.00 1.75 1.19 0.02 0.02 2.98	0.02 2.00 1.69 1.26 0.03 0.03 3.01
Specimen No.	SC 5-12	SC 5-16	SC 5-16	SC 5-16	SC 5-16	SC 5-25a	SC 5-25a	SC 5-25a	SC 6-3	SC 6-3	SC 6-3
Unit #	2	2	2	2	2	2	2	2	2	2	2
Si02 Ti02 Al203 Mg0 Ca0 Mn0 Fe0 Na20	40.32 22.71 11.35 0.33 0.34 26.87	39.40 21.92 10.01 0.32 0.14 29.12	40.33 22.72 10.33 0.34 0.10 28.73	40.01 23.00 10.25 0.39 0.12 29.03	39.47 22.34 8.27 0.42 31.53	39.78 22.21 9.52 0.48 0.33 29.01	39.83 22.11 9.64 0.49 0.72 28.31	38.84 22.17 6.78 0.40 0.43 32.74	37.85 21.12 3.70 0.44 1.03 35.31	37.88 20.88 3.72 0.43 1.07 35.79	37.56 21.08 3.42 0.40 1.10 36.02
Total	101.9	100.9	102.6	102.8	102.0	101.3	101.1	101.4	99.45	99.77	99.58
Si * Z Al Ti	3.01 3.01 2.00	3.01 3.01 1.97	3.01 3.01 2.00	2.99 2.99 2.02	3.01 3.01 2.01	3.02 3.02 1.99	3.03 3.03 1.98	3.00 3.00 2.02	3.04 3.04 2.00	3.04 3.04 1.98	3.03 3.03 2.00
Fe Y Fe Mg Mn Ca X	0.00 2.00 1.68 1.26 0.02 0.03 2.99	0.03 2.00 1.83 1.14 0.01 0.03 3.01	0.00 2.00 1.79 1.15 0.01 0.03 2.98	0.00 2.02 1.81 1.14 0.01 0.03 2.99	0.00 2.01 2.01 0.94 0.00 0.03 2.98	0.01 2.00 1.83 1.08 0.02 0.04 2.97	0.02 2.00 1.78 1.09 0.05 0.04 2.96	0.00 2.02 2.12 0.78 0.03 0.03 2.96	0.00 2.00 2.37 0.44 0.07 0.04 2.92	0.02 2.00 2.38 0.45 0.07 0.04 2.94	0.00 2.00 2.43 0.41 0.08 0.04 2.96

GARNETS											
Specimen No.	SC 6-3	SC 6-3	SC 6-3	SC 6-3	SC 6-4	SC 6-4	SC 6-4	SC 6-4	SC 6-4	SC 6-4	SC 6-4
Unit #	2	2	2	2	2	2	2	2	2	2	2
Si02 Ti02 Al203 Mg0 Ca0 Mn0 Fe0 Na20	37.57 0.41 20.91 3.89 0.42 0.80 34.95	37.47 21.20 3.61 0.42 0.98 35.53	37.38 - 20.68 3.60 0.44 1.05 35.25 -	37.17 21.05 5.22 0.43 0.54 34.09	39.72 22.20 9.86 0.44 0.35 27.99	39.55 22.07 10.07 0.48 0.38 28.23	39.70 22.16 9.84 0.45 0.39 28.08 -	38.97 21.56 7.42 0.45 0.39 31.43	38.74 21.79 6.44 0.44 0.50 32.75	39.38 22.04 9.83 0.47 0.31 28.19	39.55
Total	98.95	99.21	98.40	98.50	100.6	100.8	100.6	100.2	100.7	100.2	100.4
Si * Z Al Ti	3.03 3.03 1.99 0.03	3.02 3.02 2.02	3.04 3.04 1.98	3.00 3.00 2.00	3.03 3.03 1.99	3.01 3.01 1.98 -	3.02 3.02 1.99 -	3.03 3.03 1.98	3.02 3.02 2.00	3.02 3.02 1.99	3.02 3.02 2.00
Fe Fe Mg Mn Ca X	0.00 2.02 2.36 0.47 0.06 0.04 2.93	0.00 2.02 2.40 0.43 0.07 0.04 2.94	0.02 2.00 2.38 0.44 0.07 0.04 2.93	0.00 2.00 2.30 0.63 0.04 0.04 3.01	0.01 2.00 1.77 1.12 0.02 0.04 2.95	0.02 2.00 1.78 1.14 0.03 0.04 2.99	0.01 2.00 1.78 1.12 0.03 0.04 2.97	0.02 2.00 2.03 0.86 0.03 0.04 2.96	0.00 2.00 2.14 0.75 0.03 0.04 2.96	0.01 2.00 1.80 1.12 0.02 0.04 2.98	0.00 2.00 1.79 1.11 0.02 0.04 2.96
Specimen No.	SC 6-4	SC 6-4	SC 7-1	SC 7-1	SC 7-1	SC 7-4	SC 7-4	SC 7-4	SC 7-4	SC 7-4	SC 7-4
Unit #	2	2	1	1	1	1	1	1	1	1	1
Si02 Ti02 Al203 Mg0 Ca0 Mn0 Fe0 Na20	37.74 21.36 6.40 0.43 0.50 32.45	39.73 22.09 9.71 0.48 0.32 28.43	38.54 21.55 7.02 0.87 1.10 31.91	38.73 21.72 7.85 0.58 1.30 30.65	38.96 - 21.77 7.60 0.61 1.40 30.93	38.71 - 21.69 6.78 0.58 0.62 32.50 -	38.68 - 21.97 7.73 0.56 0.44 31.25 -	38.68 21.51 6.31 0.77 0.59 32.60	39.18 22.04 8.20 0.70 0.57 30.23	38.62 21.94 6.75 0.65 0.61 32.47	38.68 21.67 7.11 0.60 0.64 32.01
Total	98.88	100.8	101.0	100.8	101.3	100.9	100.6	100.5	100.9	101.0	100.7
Si * Z Al Ti	3.00 3.00 2.00	3.03 3.03 1.98	3.00 3.00 1.98	3.00 3.00 1.98	3.01 3.01 1.98	3.01 3.01 1.99	3.00 3.00 2.01	3.03 3.03 1.98	3.01 3.01 2.00	3.00 3.00 2.01	3.01 3.01 1.99
Fe Y Fe Mg Mn Ca X	0.00 2.00 2.06 0.76 0.03 0.04 2.89	0.02 2.00 1.79 1.10 0.02 0.04 2.95	0.02 2.00 2.06 0.82 0.07 0.07 3.02	0.02 2.00 1.97 0.91 0.09 0.05 3.02	0.02 2.00 1.98 0.88 0.09 0.05 3.00	0.01 2.00 2.11 0.79 0.04 0.05 2.99	0.00 2.01 2.03 0.89 0.03 0.05 3.00	0.02 2.00 2.11 0.74 0.04 0.07 2.96	0.00 2.00 1.94 0.94 0.04 0.06 2.98	0.00 2.01 2.11 0.78 0.04 0.05 2.98	0.01 2.00 2.07 0.83 0.04 0.05 2.99

· · ·

.

GARNETS											
Specimen No.	ML 55	ML 55	ML 55	HL 55	ML 55	ML 49	ML 49	ML 52	ML 52	ML 52	ML 52
Unit ≇	2(M)										
SiO2 TiO2	39.18	38.68	39.13	38.88	38.60	38.51	38.87	39.07	39.59	39.21	39.62
A1203 Hg0 Ca0 Hn0 Fe0 Na20	21.90 9.51 1.01 0.68 27.34	21.54 9.57 0.93 0.68 27.62	21.92 9.40 0.93 0.64 27.65 -	21.56 9.37 1.09 0.57 27.70 -	21.93 9.00 0.81 0.57 28.80 -	21.44 7.38 0.63 0.51 32.14	21.79 9.05 0.61 0.39 29.39 -	21.98 9.49 0.55 0.38 28.95	22.23 9.20 0.62 0.36 29.22 -	22.27 9.90 0.57 0.36 28.48	22.25 10.44 0.65 0.45 27.62
Total	99.62	99.02	99.67	99.17	99.71	100.6	100.1	100.4	101.2	100.8	101.0
Si * Z Al Ti	3.02 3.02 1.99	3.01 3.01 1.97	3.02 3.02 1.99	3.02 3.02 1.97	2.99 2.99 2.00	3.00 3.00 1.97	3.00 3.00 1.98	3.00 3.00 1.99	3.02 3.02 2.00	2.99 2.99 2.00	3.00 3.00 1.99
Fe Fe Mg Mn Ca X	0.01 2.00 1.75 1.09 0.04 0.08 2.96	0.03 2.00 1.77 1.11 0.05 0.08 3.01	0.01 2.00 1.77 1.08 0.04 0.08 2.97	0.03 2.00 1.77 1.08 0.04 0.09 2.98	0.00 2.00 1.87 1.04 0.04 0.07 3.02	0.03 2.00 2.07 0.86 0.03 0.05 3.01	0.02 2.00 1.88 1.04 0.03 0.05 3.00	0.01 2.00 1.85 1.09 0.03 0.05 3.02	0.00 2.00 1.86 1.04 0.02 0.05 2.97	0.00 2.00 1.82 1.13 0.02 0.05 3.02	0.01 2.00 1.75 1.18 0.03 0.05 3.01
Specimen No.	ML 52										
Unit #	2(M)										
Si02 Ti02	38.88	39.53	39.34	39.55	39.44	39.93	39.40	39.02	39.48		
Al203 Mg0 Ca0 Mn0 Fe0 Na20	21.82 8.57 0.63 1.09 29.56	21.94 9.05 0.59 0.56 29.47	21.76 9.46 0.62 0.42 28.88 -	22.09 9.87 0.58 0.40 28.43	22.06 10.11 0.54 0.37 28.43	22.39 11.01 0.64 0.47 26.56 -	22.24 10.64 0.65 0.47 27.01	21.97 8.21 0.67 0.42 30.66 -	22.07 10.40 0.53 0.84 27.59 0.05		
Total	100.6	101.1	100.5	100.9	101.0	101.0	100.4	101.0	101.0		
Si * Z Al Ti	3.00 3.00 1.99	3.02 3.02 1.98	3.02 3.02 1.97	3.01 3.01 1.98	3.00 3.00 1.98	3.01 3.01 1.99	3.00 3.00 2.00	3.01 3.01 1.99	3.00 3.00 1.98		
Fe Y Fe Mg Mn Ca X	0.01 2.00 1.90 0.99 0.07 0.05 3.01	0.02 2.00 1.86 1.03 0.04 0.05 2.98	0.03 2.00 1.82 1.08 0.03 0.05 2.98	0.02 2.00 1.79 1.12 0.03 0.05 2.99	0.02 2.00 1.79 1.15 0.02 0.04 3.00	0.01 2.00 1.67 1.24 0.03 0.05 2.99	0.00 2.00 1.72 1.21 0.03 0.05 3.01	0.01 2.00 1.96 0.94 0.03 0.06 2.99	0.02 2.00 1.73 1.18 0.05 0.04 3.00		

-152-

CORDIERITES

-153-

CORDIERI	TES										
Specimen No.	SC 2-1	SC 2-1	SC 2-1	SC 2-9	SC 2-9	SC 3-10	SC 3-10	SC 3-10	SC 3-10	SC 3-10	SC 3-25a
Unit #	1	1	1	1	1	1	1	1	1	1	2
SiO2 Al2O3 MgO CaO MnO FeO	49.27 33.10 10.17 - 5.33	49.28 33.24 10.22 0.11 5.31	49.42 32.98 10.17 - 5.46	49.14 33.09 10.77 - 4.69	49.23 33.12 10.49 - 5.22	49.42 32.93 10.76 - - 4.68	49.23 33.06 10.29 - - 5.26	48.67 32.81 10.25 - 5.41	49.22 32.99 10.30 	49.05 32.60 10.35 - 5.31	49.32 33.05 10.95 - 4.33
Na20	0.16	0.20	0.16	0.22	0.32	0.26	0.25	0.19	0.19	0.27	0.19
Total	98.03	98.36	98.19	97.91	98.38	98.05	98.09	97.33	98.21	97.58	97.84
* Si Al B Mg Ca	5.02 3.97 8.99 1.54 -	5.00 3.98 8.98 1.55	5.03 3.95 8.98 1.54 -	5.00 3.97 8.97 1.63	5.00 3.96 8.96 1.59 -	5.02 3.94 8.96 1.63	5.01 3.97 8.98 1.56	5.00 3.97 8.97 1.57	5.01 3.96 8.97 1.56 -	5.02 3.93 8.95 1.58 -	5.01 3.96 8.97 1.66
Fe A Na C	0.45 1.99 0.03 0.03	0.01 0.45 2.01 0.04 0.04	0.46 2.00 0.03 0.03	0.40 2.03 0.04 0.04	0.44 2.03 0.06 0.06	- 0.40 2.03 0.05 0.05	0.45 2.01 0.05 0.05	0.46 2.03 0.04 0.04	- 0.47 2.03 0.04 0.04	- 0.45 2.03 0.05 0.05	- 0.37 2.03 0.04 0.04
Specimen No.	SC 3-25a	SC 3-25a	SC 3-25a	SC 3-25a	SC 3-26a	SC 3-26a	SC 3-29	SC 3-29	SC 3-29	SC 3-29	SC 3-34
Unit #	2	2	2	2	2	2	2	2	2	2	2
SiO2 Al2O3 MgO CaO MnO FeO	50.31 33.00 11.09 -	49.49 33.33 10.73	49.58 32.75 10.85	49.82 33.18 10.61 - - 4.76	49.63 33.76 10.85 0.04	50.34 33.67 11.00	48.07 32.12 9.46 0.19	49.15 32.97 10.58	48.86 32.89 10.58	48.84 32.66 10.67	49.46 32.85 10.72
Na20	0.22	0.19	0.27	0.20	0.14	0.19	0.20	0.14	0.21	0.19	0.17
Total	98.81	98.53	97.92	98.57	98.78	99.54	94.97	97.52	97.14	96.97	97.77
* Si Al B Mg Ca Mn	5.05 3.91 8.96 1.66	5.00 3.97 8.97 1.62	5.04 3.92 8.96 1.64	5.03 3.95 8.98 1.60 -	4.99 4.00 8.99 1.63 -	5.02 3.96 8.98 1.64	5.04 3.97 9.01 1.48 0.02	5.02 3.97 8.99 1.61	5.01 3.97 8.98 1.62	5.01 3.95 8.96 1.63	5.03 3.94 8.97 1.63
Fe A Na C	0.35 2.01 0.04 0.04	0.41 2.03 0.04 0.04	0.38 2.02 0.05 0.05	0.40 2.00 0.04 0.04	0.37 2.00 0.03 0.03	0.36 2.00 0.04 0.04	0.43 1.93 0.04 0.04	0.40 2.01 0.03 0.03	0.39 2.01 0.04 0.04	0.40 2.03 0.04 0.04	0.39 2.02 0.03 0.03

. . .

		T		T	1	1	1	1	1	1	-
Specimen No.	SC 3-34	SC 3-54a	SC 1 3-54a	SC 1 3-54a	SC 1 3-54a	SC 1 4-21	SC 4-2b	SC 4-2b	SC 4-9	SC 4-9	
Unit ∦	2	2	2	.2	2	3	3	3	3	3	
Si02 Al203 Mg0 Ca0 Mn0 Fe0	48.64 32.76 10.55 - 4.42	49.66 33.24 11.07 - - 4 22	49.71 32.89 11.06	49.57 32.96 10.91	49.90 33.17 10.94	48.84	50.05 33.34 10.68	49.11 33.08 10.39 - -	49.26 32.77 10.56	49.31 32.88 10.36	49
Na20	0.22	0.21	0.28	0.19	0.21	0.20	0.22	0.24	0.25	0.21	
lotal	96.59	98.40	198.09	97.75	98.23	97.09	99.14	97.93	98.02	97.76	97
* Si Al B Mg Ca Mn	5.01 3.98 8.99 1.62	5.01 3.95 8.96 1.67	5.03 3.92 8.95 1.67	5.03 3.94 8.97 1.65 -	5.04 3.95 8.99 1.65	5.01 3.99 9.00 1.56 0.01	5.03 3.95 8.98 1.60	5.00 3.97 8.97 1.58	5.02 3.93 8.95 1.60	5.03 3.95 8.98 1.58 -	577201
Fe A Na C	0.38 2.00 0.04 0.04	0.36 2.03 0.04 0.04	0.35 2.02 0.06 0.06	0.35 2.00 0.04 0.04	0.34 1.99 0.04 0.04	0.41 1.98 0.04 0.04	0.41 2.01 0.04 0.04	0.44 2.02 0.05 0.05	- 0.44 2.04 0.05 0.05	- 0.43 2.01 0.04 0.04	0 2 0 0
Specimen No.	SC 4-9	SC 4-12	SC 4-12	SC 4-12	SC 4-12	SC 4-12	SC 4-13b	SC 4-13b	SC 5-1	SC 5-16	S 5-
Unit #	3	3	3	3	3	3	3	3	2	2	
SiO2 Al2O3 MgO CaO MnO	48.50 32.26 10.45 0.06	48.90 33.16 9.83 - -	48.95 32.78 10.09 -	49.17 32.69 9.99 -	48.89 32.69 9.94 -	47.90 31.82 9.98 -	49.85 33.52 10.43 -	49.92 33.31 10.65	49.49 33.22 10.64	49.89 34.00 10.70	50 33 10
Fe0 Na20	4.73 0.29	5.79 0.26	5.23 0.23	5.94 0.26	5.78 0.39	5.66 0.21	4.92 0.26	4.83 0.24	4.73 0.10	4.89 0.12	4. 0.
Total	96.29	97.94	97.28	98.05	97.69	95.57	98.98	98.95	98.18	99.60	99.
Si Al B Mg Ca	5.02 3.94 8.96 1.61 0.01	5.00 3.99 8.99 1.50	5.02 3.96 8.98 1.54	5.02 3.94 8.96 1.52	5.01 3.95 8.96 1.52	5.02 3.93 8.95 1.56	5.02 3.98 9.00 1.57	5.02 3.95 8.97 1.60	5.02 3.97 8.99 1.61	4.99 4.01 9.00 1.60	5. 3. 8. 1.
Fe A Na C	0.41 2.03 0.06 0.06	0.50 2.00 0.05	0.45 1.99 0.05 0.05	0.51 2.03 0.05	0.50 2.02 0.08	0.50 2.06 0.04	0.41 1.98 0.05	0.41 2.01 0.05	0.40 2.01 0.02	0.41 2.01 0.02	0. 2. 0.

CORDIERI	TES										
Specimen No.	SC 5-21a	SC 5-25a	SC 5-35	SC 5-35	SC 5-35	SC 5-35	SC 5-35	SC 5-35	SC 5-39	SC 5-39	SC 6-4
Unit ≇	2	2	3	3	3	3	3	3	3	3	2
Si02 Al203 Mg0 Ca0 Mn0 Fe0 Na20	50.01 33.95 11.18 - 3.98 0.08	49.90 33.95 10.71 - 4.88 0.09	49.48 33.23 10.58 - 4.92 0.15	49.40 33.58 10.74 - 4.76 0.16	49.51 33.64 10.82 - 4.78 0.16	49.69 33.21 10.45 - 5.23 0.07	49.75 33.30 10.59 - 4.91 0.13	48.65 33.38 10.63 - 4.84 0.18	49.72 33.90 10.60 - 4.90 0.19	49.36 33.23 9.87 - 6.03 0.12	49.19 33.14 9.54 - 6.50 0.27
Total	99.20	99.53	98.36	98.64	98.91	98.65	98.68	97.68	99.31	98.61	98.64
* Si Al B Mg Ca Mn	5.00 4.00 9.00 1.67 -	4.99 4.00 8.99 1.60	5.01 3.97 8.98 1.60	4.99 4.00 8.99 1.62	4.99 3.99 8.98 1.62	5.02 3.96 8.98 1.58	5.02 3.96 8.98 1.59	4.97 4.02 8.99 1.62	4.99 4.01 9.00 1.59 -	5.01 3.98 8.99 1.49	5.01 3.98 8.99 1.45 -
Fe A Na C	0.33 2.00 0.02 0.02	0.41 2.01 0.02 0.02	0.42 2.02 0.03 0.03	0.40 2.02 0.03 0.03	0.40 2.02 0.03 0.03	0.44 2.02 0.01 0.01	0.41 2.00 0.03 0.03	0.41 2.03 0.04 0.04	0.41 2.00 0.04 0.04	0.51 2.00 0.02 0.02	0.55 2.00 0.05 0.05
Specimen No.	SC 8-4	SC 8-7	SC 8-15	SC 8-15	SC 8-15	SC 8-17	SC 8-17	SC 8-28	ML 49	ML 93	ML 93
Unit #	3	3	3	3	3	3	3	3	2(M)	3(M)	3(M)
Si02 Al203 Mg0 Ca0 Mn0 Fe0 Fe0	49.46 33.41 10.21 5.38	50.06 34.09 11.03 - 4.47	49.47 32.82 9.88 5.94	48.84 32.49 9.89 - 6.05	49.94 32.98 10.55 - 4.97	48.71 32.79 10.70 - 4.38	49.25 33.09 10.99 - 4.22	49.17 33.30 9.63 - 6.51	49.41 33.25 10.93 - 4.34	48.74 32.62 10.59 - 4.26	48.79 33.04 10.65 - 4.21
NdZU	0.20	0.2/	0.15	0.20	0.19	0.21	0.18	0.20	0.18	0.27	0.27
Si Al B Mg Ca Mn Fe	5.01 3.98 8.99 1.54 - 0.46	4.98 4.00 8.98 1.64 - 0.37	5.04 3.94 8.98 1.50 	5.02 3.94 8.96 1.52 	5.04 3.93 8.97 1.59 0.42	5.00 3.97 8.97 1.64	5.01 3.96 8.97 1.67 0.36	5.00 3.99 8.99 1.46 	28.11 5.01 3.97 8.98 1.65 - 0.37	96.48 5.02 3.96 8.98 1.63 - 0.37	96.96 5.00 3.99 8.99 1.63 - 0.36
Na C	0.04	0.05	2.01 0.03 0.03	2.04 0.04 0.04	2.01 0.04 0.04	2.02 0.04 0.04	2.03 0.04 0.04	2.01 0.04 0.04	2.02 0.04 0.04	2.00 0.05 0.05	1.99 0.05 0.05

C -1	SC 2-1	SC 2-1	SC
1		21	2-1
•	1	1	1
40 76 50 48 20 59 40 56 20 24	37.40 2.76 18.50 15.48 - 15.20 0.59 8.40 0.56 4.20 -0.24	37.60 3.18 18.54 14.99 	37.01 3.53 18.22 14.41 16.28 0.53 8.68 0.45 4.18 -0.19
.91	102.9	104.4	103.1
33 67 00	5.33 2.67 8.00	5.32 2.68 8.00	5.31 2.69 8.00
44 30 29 81	0.44 0.30 3.29 1.81	0.41 0.34 3.16 1.88	0.39 0.38 3.08 1.95
84	5.84	5.79	5.80
16 53 69	- 0.16 1.53 1.69	- 0.17 1.63 1.80	- 0.15 1.59 1.74
	12 18 15 0 4 -0 10: 5 2 8 0. 3 1. 5. 0. 1. 1. 1.	.76 .50 .48 - .20 .59 .40 .56 .20 .24 2.9 .33 .67 .00 .44 .30 .29 .81 .84 - .53 .67	.76 3.16 .50 18.54 .48 14.99 - - .20 15.94 .59 0.61 .40 9.01 .56 0.54 .20 4.24 .21 104.4 .33 5.32 .67 2.68 .00 8.00 .44 0.41 .30 0.34 .29 3.16 .81 1.88 .84 5.79 .16 0.17 53 1.63 .69 1.80

* number of ions on the basis of 24 (0,0H,F)
determined by stoichiometry
- not detected

.

•

DIVITES											
Specimen No.	SC 2-9	SC 2-9	SC 2-9	SC 2-9	SC 2-9	SC 2-9	SC 3-10	SC 3-10	SC 3-10	SC 3-25a	SC 3-29
Unit #	1	1	1	1	1	1	1	1	1	2	2
Si02 Ti02 Al203 Mg0 Ca0 Mn0	38.13 0.70 17.37 17.92 0.13	37.68 2.70 17.78 15.39 -	37.92 2.03 18.16 16.39	37.14 2.56 17.75 15.47	37.44 2.57 17.66 15.23	37.12 3.38 17.69 14.50	38.10 2.06 17.39 17.48	37.95 2.32 17.40 17.03	37.92 2.98 17.10 16.14	39.27 1.91 16.94 19.21	38.76 1.47 16.81 18.06 -
Fe0 Na20 K20 F H20 # 0=F	12.25 0.47 7.39 0.54 4.12 -0.23	15.27 0.68 8.79 0.57 4.19 -0.24	13.94 0.78 8.34 0.46 4.19 -0.19	15.05 0.67 8.67 0.43 4.14 -0.18	12.15 0.59 8.54 0.58 4.08 -0.24	15.41 0.64 8.89 - 4.13 -	12.56 0.67 8.33 0.61 4.18 -0.26	12.94 0.71 7.67 0.66 4.17 -0.28	13.64 0.67 8.39 0.61 4.17 -0.26	10.14 0.38 8.12 0.98 4.24 -0.41	13.32 0.37 7.04 0.76 4.19 -0.32
 Total	98.79	102.8	102.0	101.7	98.60	101.8	101.1	100.6	101.4	100.8	100.5
 Si * Al	5.55 2.45	5.39 2.61	5.42 2.58	5.38 2.62	5.50 2.50	5.39 2.61	5.46 2.54	5.46 2.54	5.45 2.55	5.55 2.45	5.55 2.45
 C	8.00	8.00	8.00	8.00	8.00	8.00	8.00	8.00	8.00	8.00	8.00
 Al Ti Mg Mn	0.53 0.08 3.89	0.39 0.29 3.28	0.48 0.22 3.49	0.41 0.28 3.34	0.56 0.28 3.34	0.42 0.37 3.14	0.40 0.22 3.73	0.41 0.25 3.65	0.35 0.32 3.46	0.37 0.20 4.05	0.39 0.16 3.86
 Fe	1.49	1.83	1.67	1.82	1.49	1.87	1.51	1.56	1.64	1.20	1.60
 В	5.99	5.79	5.86	5.85	5.67	5.80	5.86	5.87	5.77	5.82	6.01
 Ca Na K	0.02 0.13 1.37	- 0.19 1.61	- 0.22 1.52	- 0.19 1.60	0.17 1.60	0.18 1.65	0.19 1.52	- 0.20 1.41	- 0.19 1.54	- 0.10 1.47	- 0.10 1.29
 A	1.52	1.80	1.74	1.79	1.77	1.83	1.71	1.61	1.73	1.57	1.39
		•	•		:		•	•			

14000

* number of ions on the basis of 24 (0,0H,F)
determined by stoichiometry
- not detected

۰.					

Specimen No.	SC 3-54a	SC 3-54a	SC 4-2b	SC 4-2b	SC 4-2b	SC 5-1	SC 5-1	SC 5-1	SC 5-9	SC 5-9	SC 5-12
Unit #	2	2	3	3	3	2	2	2	2	2	2
SiO2 TiO2 Al2O3 MgO CaO MnO	38.97 2.71 17.01 18.77	39.00 2.59 17.20 18.88 -	39.09 1.73 17.97 19.32	38.62 1.22 17.90 20.26 0.18	38.82 2.32 17.40 18.45	38.71 2.76 17.20 17.34	38.34 2.67 17.60 17.40 0.06	38.56 2.23 17.75 18.03	38.89 1.90 17.30 18.21 0.08	38.89 2.32 17.37 18.63 -	38.44 2.67 18.00 17.50
Fe0 Na20 K20 F H20 # 0=F	10.67 0.54 8.28 1.13 4.27 -0.48	10.66 0.51 8.64 1.13 4.28 -0.48	10.98 0.37 8.35 0.58 4.28 -0.24	10.65 0.33 7.71 0.89 4.27 -0.37	12.11 0.54 8.02 0.52 4.25 -0.22	13.60 0.65 8.38 0.96 4.27 -0.40	13.40 0.61 7.92 0.73 4.25 -0.31	13.35 0.58 8.44 0.73 4.28 -0.31	12.32 0.54 8.29 1.01 4.26 -0.43	11.33 0.56 8.58 0.88 4.27 -0.37	11.97 0.59 8.17 0.69 4.24 -0.29
Total	101.9	102.4	102.4	101.7	102.2	103.5	102.7	103.6	102.4	102.5	102.0
Si * Al	5.48 2.52	5.46 2.54	5.47 2.53	5.42 2.58	5.48 2.52	5.44 2.56	5.41 2.59	5.40 2.60	5.48 2.52	5.46 2.54	5.43 2.57
C	8.00	8.00	8.00	8.00	8.00	8.00	8.00	8.00	8.00	8.00	8.00
Al Ti Mg Mn	0.30 0.29 3.93	0.30 0.27 3.94	0.44 0.18 4.03	0.38 0.13 4.24	0.37 0.25 3.88	0.29 0.29 3.63	0.34 0.28 3.66	0.33 0.24 3.77	0.35 0.20 3.83	0.34 0.25 3.90	0.43 0.28 3.69
Fe	1.25	1.25	1.29	1.25	1.43	1.60	1.58	1.56	1.45	1.33	1.41
B	5.77	5.76	5.94	6.00	5.93	5.81	5.86	5.90	5.83	5.82	5.81
Ca Na K	- 0.15 1.48	- 0.14 1.54	- 0.10 1.49	0.03 0.09 1.38	- 0.15 1.44	- 0.18 1.50	0.01 0.17 1.43	- 0.16 1.51	0.01 0.15 1.49	- 0.15 1.54	- 0.16 1.47
A	1.63	1.68	1.59	1.47	1.59	1.68	1.60	1.67	1.64	1.69	1.63

* number of ions on the basis of 24 (0,0H,F)
determined by stoichiometry
- not detected

	<u> </u>										
Specimen No.	SC 5-12	SC 5-16	SC 5-21a	SC 5-21a	SC 5-21a	SC 5-21a	SC 5-25a	SC 5-39	SC 5-67	SC 5-67	SC 6-3
Unit #	2	2	2	2	2	2	2	3	QBG	QBG	2
Si02 Ti02 Al203 Mg0 Ca0 Mn0 Fa0	38.74 1.63 17.81 18.30 - -	39.37 3.18 16.58 18.24 0.05	38.82 1.81 17.21 18.56	38.94 2.43 16.61 17.93	38.85 2.29 16.57 18.36 0.08	41.24 2.42 17.76 17.92	38.14 3.11 17.08 16.26	37.81 2.38 17.82 17.56	36.76 2.00 20.12 13.61	35.70 2.11 19.55 13.63 0.06	35.63 1.86 19.54 11.29
Na20 K20 F H20 # 0=F	0.46 9.03 0.50 4.24 -0.21	0.43 8.92 0.90 4.30 -0.38	0.37 8.81 0.89 4.24 -0.37	0.47 8.43 0.70 4.21 -0.29	11.44 0.40 8.37 1.01 4.22 -0.43	12.30 - 2.43 0.86 4.30 -0.36	13.78 0.37 8.24 0.50 4.18 -0.21	13.59 0.34 8.82 0.70 4.23 -0.29	15.43 0.49 8.77 0.78 4.16 -0.33	15.68 0.65 8.66 0.47 4.08 -0.20	20.63 0.58 8.59 - 4.05
Total	102.2	103.7	102.1	101.1	101.2	99.10	101.5	103.0	101.8	100.4	102.2
Si * Al	5.48 2.52	5.49 2.51	5.49 2.51	5.55 2.45	5.52 2.48	5.76 2.24	5.47 2.53	5.36 2.64	5.30 2.70	5.25 2.75	5.27 2.73
C	8.00	8.00	8.00	8.00	8.00	8.00	8.00	8.00	8.00	8.00	8.00
Al Ti Mg Mn	0.45 0.17 3.86	0.22 0.33 3.79	0.36 0.19 3.91	0.34 0.26 3.81	0.30 0.25 3.89	0.68 0.25 3.73	0.36 0.34 3.48	0.34 0.25 3.71	0.72 0.22 2.93	0.64 0.23 2.99	0.68 0.21 2.49
Fe	1.39	1.41	1.39	1.39	1.36	1.46	1.65	1.61	1.86	1.93	2.55
B	5.87	5.75	5.85	5.80	5.80	6.12	5.83	5.91	5.73	5.79	5.93
Ca Na K	0.13 1.63	0.01 0.12 1.59	- 0.10 1.59	- 0.13 1.53	0.01 0.11 1.52	- 0.43	- 0.10 1.51	- 0.09 1.59	0.14 1.61	0.01 0.19 1.63	- 0.17 1.62
A	1.73	1.72	1.69	1.66	1.64	0.43	1.61	1.68	1.75	1.83	1.79

* number of ions on the basis of 24 (0,0H,F)
determined by stoichiometry
- not detected

.....

:									ŝ.
i	Ð	г	ሳ	т	7	т	٣	c	ŧ
ŧ	D	Ł	υ	1	£	ł	E	Э	1
1	-	_	-		_		-	-	٤.

·----;

~

Specimen No.	SC 6-3	SC 6-3	SC 6-3	SC 6-4	SC 6-4	SC 7-1	SC 7-4	SC 7-4	SC 7-10	SC 7-10	SC 7-10
Unit #	2	2	2	2	2	1	1	1	SIL	SIL	SIL
Si02 Ti02 Al203 Mg0 Ca0 Mn0 Fe0 Na20 K20 F H20 # 0=F	36.10 2.00 19.13 11.30 0.13 - 19.90 0.54 7.70 0.47 4.06 -0.20	36.00 0.48 20.75 11.89 - 18.82 0.40 8.73 - 4.06	36.26 0.45 20.81 11.81 - 19.30 0.48 8.10 - 4.07	37.79 1.07 18.96 16.56 	37.55 1.21 19.36 16.58 - - 13.93 0.77 7.49 0.51 4.19 -0.21	35.86 2.35 19.70 12.82 - - 17.87 0.25 9.21 0.44 4.11 -0.19	35.93 1.66 19.83 14.29 - 16.56 0.25 9.34 0.47 4.13 -0.20	36.37 1.17 19.85 14.77 15.99 0.30 9.00 0.62 4.14 -0.26	36.52 2.49 19.74 14.43 	36.57 2.65 20.06 14.40 - 15.06 0.45 8.97 0.46 4.18 -0.19	36.60 2.51 20.03 14.47 - 14.65 0.32 9.42 0.55 4.18 -0.23
Total	101.1	101.1	101.3	102.5	101.4	102.4	102.3	102.0	102.0	102.6	102.5
Si * Al C	5.34 2.66 8.00	5.32 2.68 8.00	5.34 2.66 8.00	5.38 2.62 8.00	5.37 2.63 8.00	5.23 2.77 8.00	5.22 2.78 8.00	5.26 2.74 8.00	5.29 2.71 8.00	5.24 2.76 8.00	5.25 2.75 8.00
Al Ti Mg Mn Fe	0.67 0.22 2.49	0.94 0.05 2.62	0.95 0.05 2.59	0.56 0.12 3.51	0.63 0.13 3.54	0.61 0.26 2.79	0.61 0.18 3.09	0.65 0.13 3.19	0.66 0.27 3.11	0.63 0.29 3.08	0.64 0.27 3.09
B	5.84	5.94	5.97	5.94	5.97	5.84	5.89	5.91	5.85	5.81	5.76
Ca Na K A	0.02 0.16 1.45 1.63	- 0.12 1.65 1.77	- 0.14 1.52 1.66	- 0.23 1.46 1.69	- 0.21 1.37 1.58	- 0.07 1.71 1.78	- 0.07 1.73 1.80	- 0.08 1.66 1.74	- 0.08 1.75 1.83	- 0.13 1.64 1.77	- 0.09 1.72 1.81
					1					1	

1

* number of ions on the basis of 24 (0,0H,F)
determined by stoichiometry
- not detected

Unit #	3	3	2(M)	2(M)	2(M)	3(M)	3(M)	3(M)	3(M)	3(M)
Si02 Ti02 Al203 Mg0 Ca0 Mn0	38.02 3.37 17.26 17.85 -	37.73 1.84 17.79 18.11 -	33.94 3.46 18.80 12.02 1.25	36.50 3.23 18.17 13.45 0.05	36.53 3.45 18.19 13.47 0.04 -	37.14 1.02 16.63 16.97	38.16 1.13 17.85 17.00 -	38.45 1.18 17.80 17.15 -	37.43 1.31 18.23 17.65 -	38.59 1.18 17.73 17.15 0.15
Fe0 Na20 K20 F H20 # 0=F	12.09 0.91 8.42 0.53 4.24 -0.22	12.48 0.58 8.09 0.80 4.20 -0.34	17.81 0.25 6.53 0.48 3.97 -0.20	17.45 0.52 8.65 0.54 4.13 -0.23	16.50 0.39 8.73 0.59 4.12 -0.25	12.05 0.17 9.19 0.52 4.02 -0.22	13.42 0.21 9.08 0.58 4.17 -0.24	13.56 0.30 8.96 0.35 4.19 -0.15	13.90 0.19 7.71 0.47 4.17 -0.20	13.38 0.41 8.50 0.47 4.19 -0.20
Total	102.5	101.3	98.31	102.5	101.8	97.49	101.4	101.8	100.9	101.6
Si * Al	5.38 2.62	5.39 2.61	5.13 2.87	5.30 2.70	5.32 2.68	5.54 2.46	5.49 2.51	5.51 2.49	5.39 2.61	5.52 2.48
C	8.00	8.00	8.00	8.00	8.00	8.00	8.00	8.00	8.00	8.00
Al Ti Mg Mn	0.26 0.36 3.77	0.38 0.20 3.86	0.48 0.39 2.71	0.41 0.35 2.91	0.44 0.38 2.92	0.46 0.11 3.78	0.51 0.12 3.64	0.51 0.13 3.66	0.48 0.14 3.79	0.51 0.13 3.66
Fe	1.43	1.49	2.25	2.12	2.01	1.50	1.61	1.62	1.67	1.60
В	5.82	5.93	5.83	5.79	5.75	5.85	5.88	5.92	6.08	5.90
Ca Na K	- 0.25 1.52	- 0.16 1.47	0.20 0.07 1.26	0.01 0.15 1.60	0.01 0.11 1.62	- 0.05 1.75	- 0.06 1.67	- 0.08 1.64	- 0.05 1.42	0.02 0.11 1.55

* number of ions on the basis of 24 (0,0H,F)
determined by stoichiometry

1.77

- not detected

A

BIOTITES

Specimen

No.

SC

8-7

SC

8-7

ML 55

ML 55

ML

55

ML 93

ML 93

ML 93

ML 93

ML.

93

1.63 1.53 1.76 1.74 1.80 1.73 1.72 1.47 1.68

CHLORITES

Specimen No.	SC 4-2b	SC 4-2b	SC 8-4	SC 8-11	SC 8-11	SC 8-17	SC 8-17	SC 8-17	SC 8-28	SC 8-28	SC 8-28
Unit #	3	3	3	3	3	3	3	3	3	3	3
Si02 Ti02 Al203 Mg0 Ca0 Mn0 Fe0 Na20 K20 H20 # Total	26.18 0.07 23.15 21.91 - 18.71 - 12.21 102.2	28.19 0.09 21.91 21.30 - 19.22 0.10 - 12.25 103.1	26.22 0.13 23.28 23.44 	25.69 0.08 23.23 20.73 0.10 20.33 - 12.13 102.3	25.23 0.09 23.76 21.29 - 19.20 - 12.12 101.7	26.42 22.64 21.97 0.09 17.60 	25.70 22.46 20.99 0.06 0.09 18.37 - 0.05 11.88 99.60	31.51 21.41 20.25 	24.87 0.07 23.22 21.07 - 19.28 - 11.97 100.5	25.13 23.25 20.48 0.07 19.86 11.96 100.8	25.45 0.11 23.38 21.56 - 19.07 - 12.13 101.7
Si * Al B Al Ti Mg Ca Mn Fe Na K	5.14 2.86 8.00 2.50 0.01 6.42 - - 3.07 -	5.52 2.48 8.00 2.58 0.01 6.22 - 3.15 0.04 -	5.12 2.88 8.00 2.47 0.02 6.82 - - 2.70 -	5.08 2.92 8.00 2.50 0.01 6.11 - 0.02 3.36 -	4.99 3.01 8.00 2.53 0.01 6.28 - - 3.18 -	5.25 2.75 8.00 2.55 6.51 0.02 2.92 -	5.19 2.81 8.00 2.54 	6.26 1.74 8.00 3.28 6.00 - - 2.72 -	4.98 3.02 8.00 2.46 0.01 6.29 - 3.23 -	5.04 2.96 8.00 2.53 6.12 0.02 3.33 -	5.03 2.97 8.00 2.47 0.02 6.35 - 3.15 -
A	12.00	12.00	12.00	12.00	12.00	12.00	12.00	12.00	12.00	12.00	12.00

* number of ions on the basis of 20 cations
determined by stoichiometry
- not detected

SILLIMANITES

SILLIMAN	ITES										
Specimen No.	SC 1-14a	SC 1-14a	SC 1 2-9	SC 2-9	SC 3-10	SC 3-10	SC 3-10	SC 3-10	SC 3-26a	SC 3-26a	SC 1 3-26a
Unit ≇	SIL	SIL	1	1	1	1	1	1	2	2	2
Si02 Al203 Mg0 Fe203 Zn0 K20	37.33 59.36 0.16 0.46 - 0.10	37.04 57.02 0.16 0.49 	36.97 60.40 0.29 -	37.15 59.77 0.13 0.38 -	37.19 60.25 0.34	36.52 58.91 0.17 0.46 0.23	37.03 59.41 0.32 -	37.02 59.98 0.25 -	37.48 59.92 0.25 -	36.87 59.94 0.15 0.36 -	37.61 60.86 0.05 0.29 -
Total	97.41	95.15	97.66	97.43	97.78	96.29	96.76	97.30	97.65	97.32	98.81
Si B	1.04 1.04	1.05 1.05	1.02 1.02	1.03 1.03	1.03 1.03	1.03 1.03	1.03 1.03	1.03 1.03	1.04 1.04	1.02 1.02	1.03 1.03
Al Mg Fe3+ Zn	1.94 0.01 0.01	1.91 0.01 0.01	1.97 0.01	1.95 0.01 0.01 -	1.96 0.01	1.95 0.01 0.01 0.01	1.95 0.01	1.96 0.01	1.95 0.01	1.96 0.01 0.01	1.96 0.01
K A	_ 1.96	0.01 1.94	- 1.98	- 1.97	- 1.97	1.98	- 1.96	- 1.97	- 1.96	- 1.98	- 1.97
Specimen No.	SC 3-54a	SC 3-54a	SC 3-54a	SC 4-9	SC 4-9	SC 5-1	SC 5-1	SC 5-1	SC 5-35	SC 5-35	SC 5-35
Unit #	2	2	2	3	3	2	2	2	3	3	3
Si02 Al203 Mg0 Fe203 Zn0 K20	37.12 59.49 0.23 0.41 -	37.36 59.76 0.08 0.27 -	37.12 59.47 0.20 -	36.56 58.73 0.32 -	37.03 60.04 0.29 -	37.16 61.03 0.24 -	37.38 60.63 0.26 -	37.41 61.18 0.07 0.25 -	37.25 60.80 0.32 -	37.23 60.68 0.31 -	37.06 60.60 0.31
Total	97.25	97.47	96.79	95.61	97.36	98.43	98.27	98.91	98.37	98.22	97.97
* Si B	1.03 1.03	1.03 1.03	1.03 1.03	1.03 1.03	1.03 1.03	1.02 1.02	1.03 1.03	1.02 1.02	1.02 1.02	1.02 1.02	1.02
Al Mg Fe3+ Zn	1.95 0.01 0.01	1.95 0.01	1.95 0.01	1.95 0.01	1.96 0.01	1.97 0.01	1.96 0.01	1.97 0.01	1.97 0.01	1.97 0.01	1.97 0.01
K A	1.97	- 1.96	1.96	- 1.96	- 1.97	1.98	- 1.97	- 1.98	1.98	1.98	- 1.98

. .

SILLIMAN	ITES										
Specimen No.	SC 5-67	SC 7-1	SC 7-10	SC 7-10	SC 8-7	SC 8-15	SC 8-15	SC 8-17	SC 8-17	SC 8-28	ML 49
Unit #	QBG	1	SIL	SIL	3	3	3	3	3	3	2(M)
Si02 Al203 Mg0 Fe203 Zn0 K20	38.55 55.53 0.43 0.48 - 0.08	36.91 60.47 0.25 -	37.23 61.02 0.19 -	36.72 60.36 0.21 -	37.63 61.22 - 0.31 -	37.10 59.59 0.27 -	36.92 59.40 0.28 -	37.13 59.38 0.06 0.28 -	36.93 60.01 0.24 -	36.91 60.59 0.34 -	36.62 61.42 0.24 -
Total	95.26	97.63	98.44	97.29	99.16	96.96	96.60	96.85	97.18	97.84	98.28
× Si B	1.09 1.09	1.02 1.02	1.02 1.02	1.02 1.02	1.02 1.02	1.03 1.03	1.03 1.03	1.03 1.03	1.03 1.03	1.02 1.02	1.01 1.01
Al Mg Fe3+ 75	1.85 0.02 0.01	1.97 0.01	1.97 _ 0.00	1.97 _ 0.01	1.96 0.01	1.95 	1.96 	1.95 0.00 0.01	1.96 	1.97 0.01	1.99
K K	0.00 1.88	- - 1.97	- - 1.97	_ _ 1.98	_ 1.97	- - 1.96	- - 1.97	 1.96	- - 1.97	- 1.98	- 2.00
Specimen No.	ML 93										
Unit #	3(M)										
Si02 Al203 Mg0 Fe203 Zn0 K20	38.23 60.25 1.21 1.21 -										
Total	100.9										
si [*] B	1.03 1.03			-							
Al Mg Fe3+ Zn	1.91 0.05 0.03 -				· · · · · · · · · · · · · · · · · · ·						
K	1.99										

STAUROLITES

STAUROLI	TES										
Specimen No.	SC 5-9	SC 5-12	SC 5-25	ML 01	ML 01	ML 01	ML 01	ML 01	ML 01	ML 01	ML 01
Unit #	2	2	2	1	1	1	1	1	1	1	1
Si02 Ti02 Al203 Mg0 Mn0 Fe0 Zn0 Na20	27.10 1.11 54.42 3.74 - 9.91 1.51 -	27.39 0.84 54.46 3.95 - 10.25 0.30 -	27.45 0.99 52.92 3.25 12.06 -	27.60 0.71 53.20 3.31 0.11 13.32 -	27.98 0.72 53.64 3.36 0.07 13.35 -	27.37 0.64 53.32 3.51 0.09 13.44 -	27.55 0.70 53.58 3.52 0.12 13.44	27.37 0.62 53.71 3.58 0.12 13.39	27.31 0.70 53.59 3.53 0.09 13.23 - -	27.41 0.76 53.25 3.54 0.09 13.34 0.14	27.45 0.78 53.53 3.47 0.10 13.29 -
Total	97.79	97.19	96.67	98.25	99.12	98.42	98.91	98.79	98.45	98.53	98.62
Si * Al C Al Ti B Mg Mn Fe Zn A	3.73 0.27 4.00 8.56 0.12 8.68 0.77 - 1.14 0.15 2.06	3.77 0.23 4.00 8.61 0.09 8.70 0.81 - 1.18 0.03 2.02	3.83 0.17 4.00 8.53 0.10 8.63 0.68 - 1.41 - 2.09	3.81 0.19 4.00 8.47 0.07 8.54 0.68 0.01 1.54 - 2.23	3.83 0.17 4.00 8.48 0.07 8.55 0.69 0.01 1.53 - 2.23	3.78 0.22 4.00 8.46 0.07 8.53 0.72 0.01 1.55 -	3.78 0.22 4.00 8.45 0.07 8.52 0.72 0.01 1.54 - 2.27	3.76 0.24 4.00 8.47 0.06 8.53 0.73 0.01 1.54 - 2.28	3.77 0.23 4.00 8.48 0.07 8.55 0.73 0.01 1.53 - 2.27	3.78 0.22 4.00 8.44 0.08 8.52 0.73 0.01 1.54 0.01 2.29	3.78 0.22 4.00 8.47 0.08 8.55 0.71 0.01 1.53 - 2.25
Specimen No.	ML 01	ML 01	ML 01	ML 08	ML 37	ML 37	ML 49				
Unit #	1	1	1	1	2	2	2(M)				
Si02 Ti02 Al203 Mg0 Mn0 Fe0 Zn0 Na20	27.19 0.65 54.49 3.38 0.09 13.40	27.67 0.63 53.08 3.60 0.11 13.19	27.40 0.66 53.50 3.41 0.13 13.33	26.55 1.03 53.53 1.89 0.10 11.74 1.36	27.09 0.89 54.77 3.46 11.54	27.36 1.03 53.32 3.88 0.06 11.52 -	26.78 0.79 53.72 4.13 0.07 11.09 0.14				
lotal	99.20	98.28	98.43	96.20	97.75	97.17	96.72				.
Si * Al C Al Ti B Mg Mn Fe Zn A	3.72 0.28 4.00 8.51 0.07 8.58 0.69 0.01 1.53 - 2.23	3.82 0.18 4.00 8.46 0.07 8.53 0.74 0.01 1.52 - 2.27	3.78 0.22 4.00 8.48 0.07 8.55 0.70 0.02 1.54 - 2.26	3.75 0.25 4.00 8.65 0.11 8.76 0.40 0.01 1.39 0.14 1.94	3.73 0.27 4.00 8.61 0.09 8.70 0.71 - 1.33 - 2.04	3.79 0.21 4.00 8.50 0.11 8.61 0.80 0.01 1.34 2.15	$\begin{array}{c} 3.73\\ 0.27\\ 4.00\\ 8.54\\ 0.08\\ 8.62\\ 0.86\\ 0.01\\ 1.29\\ 0.01\\ 2.17\\ \end{array}$				

SPINELS

nervinite spinels											
Specimen No.	SC 4-9	SC 4-9	SC 4-12	SC 4-12	SC 4-12	SC 5-35	SC 5-35	ML 15	ML 15	ML 15	ML 01
Unit ≇	3	3	3	3	3	3	3	3	3	3	1
Si02 Ti02 Al203 Cr203 V203 Mg0 Mn0 Fe0 Zn0 Total	0.16 58.48 0.67 6.22 25.72 7.45 98.70	0.19 59.11 0.76 6.26 26.55 7.82 100.7	0.15 59.18 0.26 5.26 23.98 11.19 100.0	0.20 59.05 0.24 5.27 24.18 10.90 99.84	0.23 59.99 0.23 5.16 23.10 11.78 100.5	0.26 	0.20 59.96 0.17 7.14 0.10 32.95 0.45 99.97	0.05 59.35 0.15 6.06 33.07 0.17 98.85	0.18 0.17 55.70 0.17 5.15 0.11 34.27 0.51 96.26	- 59.15 0.14 5.37 33.82 -	- 59.54 - 8.07 0.10 30.63 0.71 999.05
Si * Al Cr V Ti B Mg Mn Fe Zn A	0.01 1.96 0.02 1.99 0.26 0.61 0.16 1.03	0.01 1.95 0.02 1.98 0.26 0.62 0.16 1.04	0.00 1.97 0.01 - 1.98 0.22 0.57 0.23 1.02	0.01 1.97 0.01 - 1.99 0.22 0.57 0.23 1.02	0.01 1.99 0.01 - 2.00 0.22 0.54 0.24 1.00	0.01 1.97 0.00 0.00 - 1.98 0.26 0.75 0.01 1.02	0.01 1.96 - 0.00 - 1.97 0.30 0.00 0.74 0.01 1.05	- 1.97 0.00 0.00 1.97 0.26 0.78 0.00 1.04	0.01 1.93 0.00 0.00 1.94 0.23 0.00 0.84 0.01 1.08	1.98 0.00 - 1.98 0.23 0.80 - 1.03	- 1.96 - - 1.96 0.34 0.00 0.71 0.02 1.07

and a second second

UEDOVNITE SDINELS

* number of ions on the basis of 4 oxygens
- not detected

•

NERGINITE SLINETS											
Specimen No.	ML 33	ML 33	ML 93	ML 93	ML 93	ML 33	ML 33				
Unit #	3	3	3(M)	3(M)	3(M)	3	3				
Si02 Ti02 Al203 Gr203 Y203 Mg0 Mn0 Fe0 Zn0	- 61.18 0.11 7.97 0.11 30.98 0.16	- 59.58 - 0.12 7.82 0.09 31.03 0.26	- 58.43 - 7.38 0.09 29.98 2.83	- 58.74 - 7.69 0.08 29.99 2.17	- 58.72 - 8.53 0.08 28.99 2.15	- 58.92 0.20 6.92 0.08 31.97 0.37	0.08 58.64 0.15 5.57 0.11 34.49 0.45				
Total	100.5	98.90	98.71	98.67	98.47	98.46	99.49				
Si * Al Cr V Ti B Mg Mn Fe Zn	1.97 - 1.97 0.33 0.71	1.96 - 1.96 0.33 - 0.72 0.01	1.95 - 1.95 0.31 0.71 0.06	1.95 - 1.95 0.32 0.71 0.05	1.94 - 1.94 0.36 0.68 0.05	- 1.96 0.01 - 1.97 0.29 0.75 0.01	1.95 - 1.95 0.24 0.82 0.01				
A	1.04	1.06	1.08	1.08	1.09	1.05	1.07				

HERCYNTTE SPINELS

APATITES

n haran waxan în sant a sant an eren. Referin

,
APATITES								
Specimen No.	SC 7-10	SC 7-10	SC 7-10	SC 7-10				
Unit #	SIL	SIL	SIL	SIL				
Si02 Mg0 Ca0 Mn0 Fe0 Na20 P205 F 0=F	0.33 52.33 0.13 0.81 0.10 43.61 5.10 -2.15	0.07 0.42 51.84 0.11 0.98 0.09 43.19 4.95 -2.08	0.33 55.18 1.02 0.08 43.72 4.72 -1.99	0.06 0.24 55.55 0.18 0.86 0.09 43.33 5.17 -2.18				
Total	100.3	99.57	103.1	103.3				
*								
Р	3.04	3.04	3.00	2.97		1		
В	3.04	3.04	3.00	2.97				
Si Mg Ca Mn Fe Na	- 0.04 4.62 0.01 0.06 0.02	0.01 0.05 4.61 0.01 0.07 0.01	0.04 4.79 0.07 0.01	$\begin{array}{c} 0.01 \\ 0.03 \\ 4.81 \\ 0.01 \\ 0.06 \\ 0.01 \end{array}$				
A	4.73	4.76	4.91	4.93				

* number of ions on the basis of 13 oxygens
- not detected

r.

WAGNERITES

.

ener e Berner en en 🔶

-176-

			,	·, · · · · · · · · · · · · · · · · · ·					
Specimen No.	SC 5-9	SC 5-9	SC 5-21	SC 5-21	ML 02	ML 02	ML 02		
Unit #	2	2	2	2	2	2	2		
Si02 Ti02 Mg0 Ca0 Mn0 Fe0 Na20 P205 F 0=F Total	2.00 42.82 0.17 7.11 44.09 9.11 -3.84 101.5	2.06 42.83 0.12 6.90 43.56 9.30 -3.92 100.9	- 1.34 42.60 0.09 - 7.96 43.38 9.59 -4.04 100.9	0.09 1.25 42.28 0.15 - 7.22 0.24 42.41 8.72 -3.67 98.69	- 0.91 43.76 0.06 0.10 7.14 42.60 8.05 -3.39 99.23	- 0.76 43.70 0.05 0.09 7.11 43.06 8.22 -3.46 99.53	- 0.61 44.46 0.05 0.09 7.05 - 42.68 8.28 -3.49 99.73		
P B Si Ti Mg Ca Mn Fe Na A	1.03 1.03 - 0.04 1.77 0.01 - 0.17 - 1.99	1.03 1.03 - 0.04 1.78 - 0.16 - 1.98	1.02 1.02 - 0.03 1.77 - 0.19 - 1.99	1.03 1.03 - 0.03 1.80 0.01 - 0.17 0.01 2.02	1.03 1.03 - 0.02 1.86 - 0.17 2.05	1.03 1.03 - 0.02 1.85 - 0.17 - 2.04	1.02 1.02 - 0.01 1.88 - 0.17 - 2.06		

WAGNERITES

` ;

* number of ions on the basis of 5 oxygens - not detected