

THE UNIVERSITY OF MANITOBA
SEASONAL CHANGES IN BEACH MORPHOLOGY,
GRAND BEACH, MANITOBA

by
A. BADERL

A THESIS
SUBMITTED TO THE FACULTY OF GRADUATE STUDIES IN
PARTIAL FULFILLMENT OF THE REQUIREMENTS FOR THE DEGREE
OF MASTER OF ARTS

DEPARTMENT OF GEOGRAPHY

WINNIPEG, MANITOBA
FALL, 1977

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GRAND BEACH, MANITOBA

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ALFONS BADERL

A dissertation submitted to the Faculty of Graduate Studies of
the University of Manitoba in partial fulfillment of the requirements
of the degree of

MASTER OF ARTS

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ACKNOWLEDGEMENTS

The writer is indebted to Professor L. P. Stene, who first proposed this topic, and under whose supervision this thesis was prepared. His guidance and assistance during the study is greatly appreciated.

Dr. J. A. Richtik, Head of the Department of Geography at the University of Winnipeg is thanked for the extensive use of laboratory facilities.

For additional guidance and many helpful suggestions the writer wishes to thank Dr. A. R. Lockery.

Sincere thanks are also expressed to Mr. T. Ball, lecturer, at the University of Winnipeg, who was always prepared to help and share his knowledge.

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Chapter IINTRODUCTION

A beach is one of the most dynamic and variable land-forms of a coastline. It is an accumulation of non-cohesive material which responds to various energy regimes (King, 1972).

Krumbein and Slack (1956) divide a beach into various zones; the backshore, foreshore, near shore bottom, and the dune belt zone. King's (1972) morphologic division is very similar, excluding only the dune belt.

The backshore is defined as a shore zone, or belt, which seldom experiences wave activity, and only during storms or periods of unusual wave reach does the backshore become inundated.

The foreshore, or swash limit, is defined as the zone of a beach which is alternately covered with water and exposed to the air.

The offshore zone (King, 1972), or near shore bottom (Krumbein 1956) is that section of the water-land interface which is completely submerged and extends to a point offshore where little or no sediment transport occurs.

A beach may be considered as a geometric element produced by the energy and material transfers between the offshore and backshore. A beach environment, therefore, represents an

open system involving energy, water and sediment. Associated with this system are the open system lateral transfers with longshore and littoral drifts (Figure 1). Limnic beaches exposed to these variables may be considered as having a seasonal dynamic equilibrium.

Researchers have postulated and examined processes occurring on oceanic coasts. Most of the results associated erosion with winter or high energy waves which are prevalent during the winter season. Dubois (1972) noted the seasonal variation in morphology on limnic beaches and directed his research to the Lake Michigan area. The results show that the beach retrograded from spring to summer concomitantly with the rise in lake level, and prograded from summer to winter with decreasing water levels.

The objective of this thesis is to investigate the seasonal changes of a limnic beach, including changes in profiles and grain size parameters.

Area of Research

The Grand Beach area was chosen for its accessibility. It is located on the east shore of Lake Winnipeg, approximately fifty miles northeast of Winnipeg (Figure 2). The East Beach of Grand Beach was selected as the sample site as it can be divided into two similar beach areas separated by vegetation and sand dunes. These two areas appear to

MODEL DEPICTING DYNAMIC EQUILIBRIUM OF A BEACH

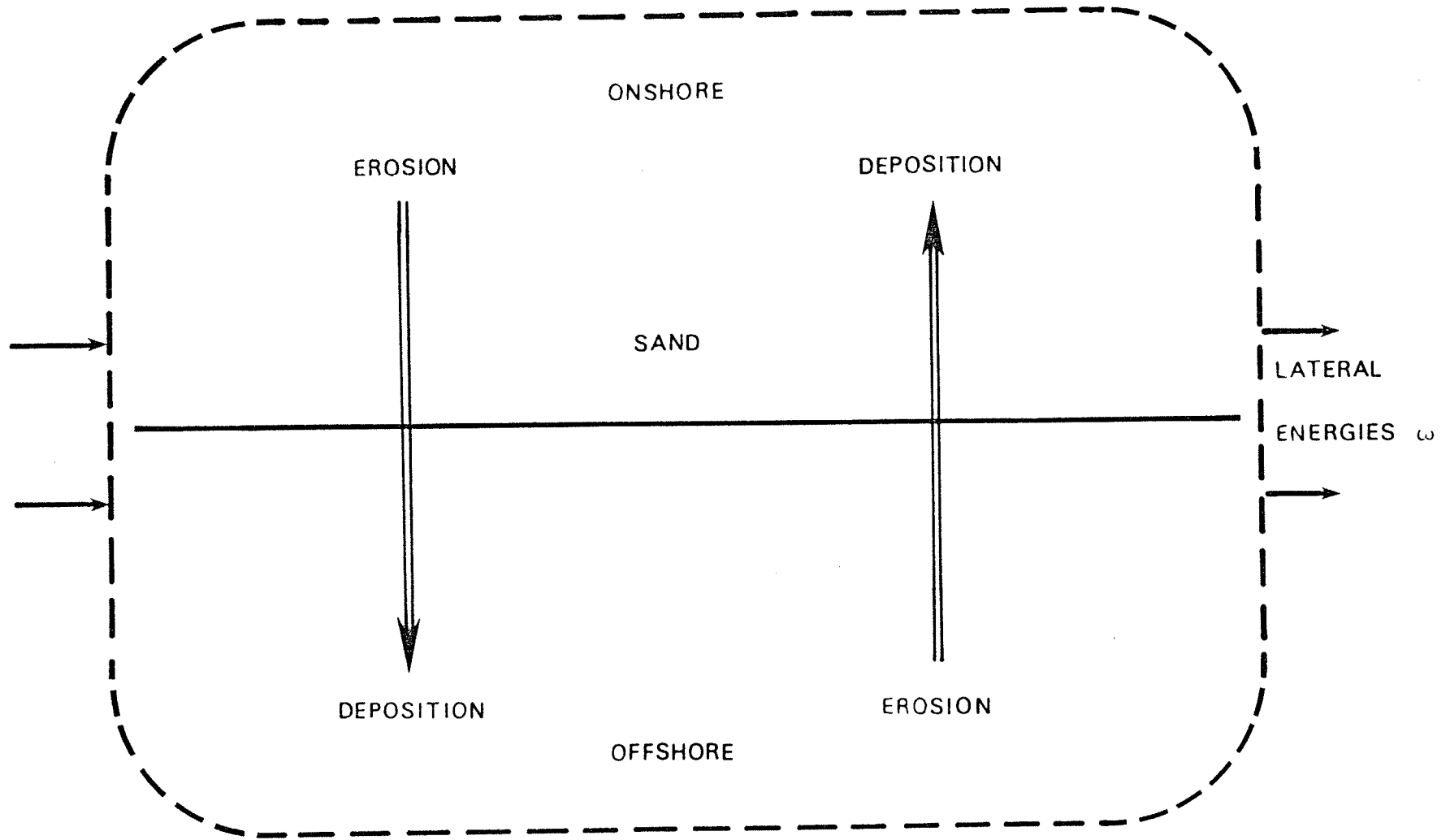


Figure 1

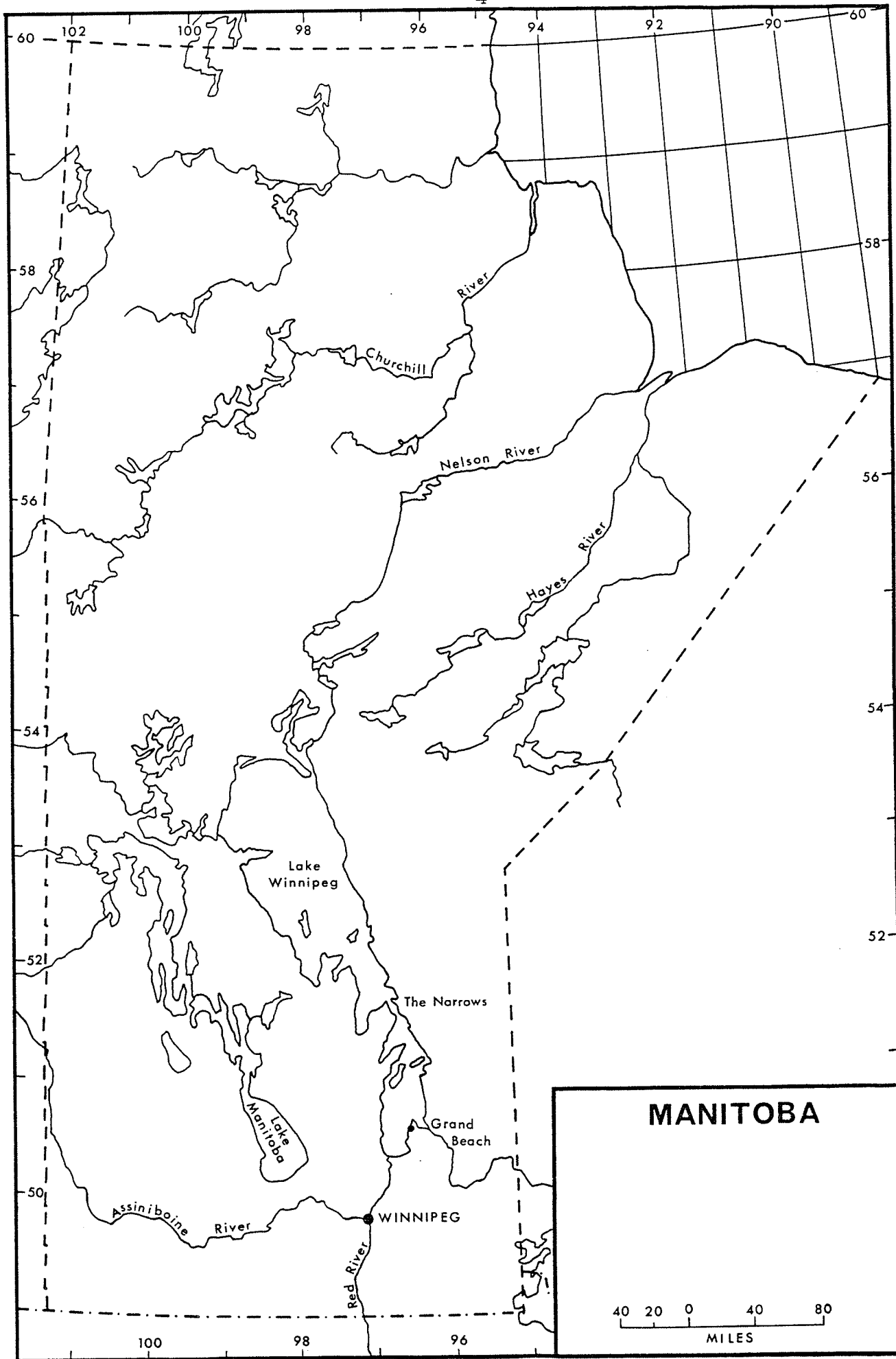


Figure 2

experience the same natural limnic processes. The East Beach also does not have the heavy influx of summer visitors which populate the other areas of Grand Beach.

Lake Winnipeg

Lake Winnipeg is a large fresh water body located in the Manitoba lowlands. Its boundaries are latitude $50^{\circ} 20'$ and $53^{\circ} 50'$ North, and longitude $96^{\circ} 20'$ and $99^{\circ} 15'$ West. It has a maximum length of 250 miles and a variable width of 25 to 70 miles creating a surface area of approximately 9,430 square miles.

The lake can actually be viewed as two bodies of water, divided into a northern and a southern section separated by narrows and several large islands. The northern part has the Saskatchewan River as the largest contributor. The southern part is approximately 55 miles long and from 20 to 30 miles in width (Figure 3). The Red River, Winnipeg River and Assiniboine River are the main sources of influx of water into this part of the lake.

The total drainage basin for Lake Winnipeg has an area of 380,000 square miles (Lake Winnipeg, Churchill and Nelson Rivers Study Board 1971-75) including portions of Alberta, Saskatchewan, Manitoba, Ontario, North Dakota and Minnesota. Lake Winnipeg is discharged by the Nelson River flowing northward into Hudson Bay (Figure 4).

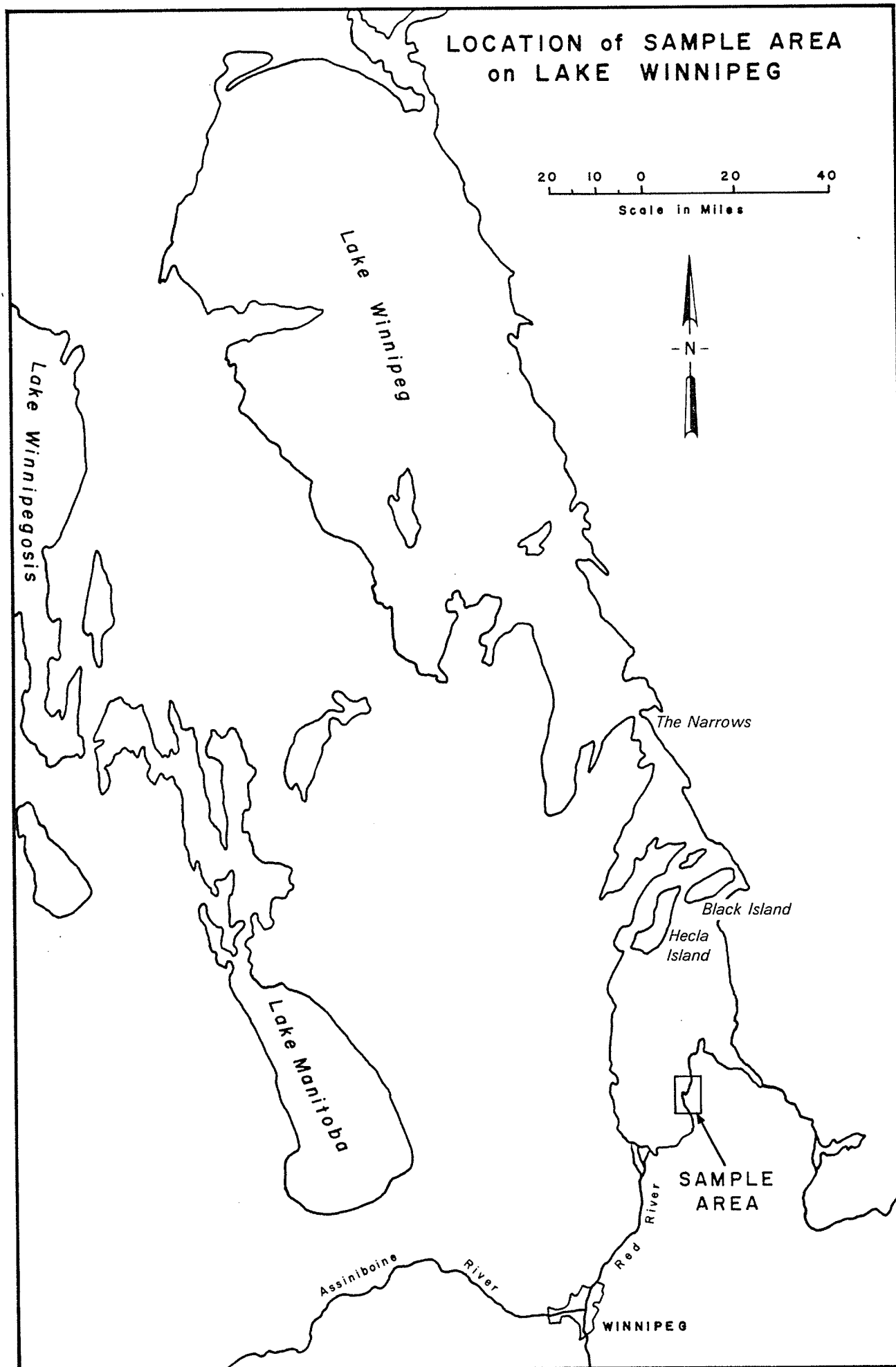


Figure 3

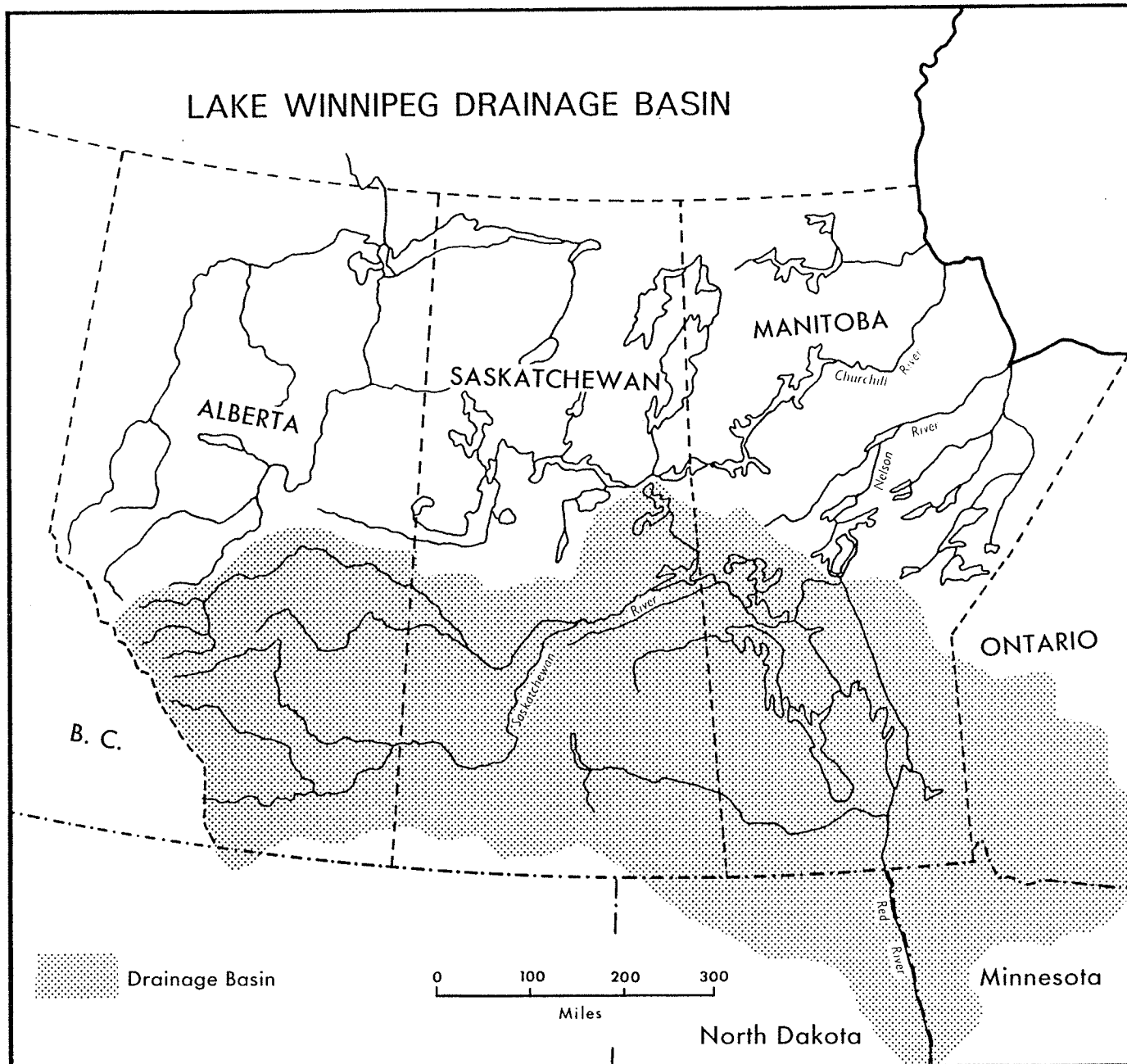


Figure 4

A brief examination of water levels of Lake Winnipeg (Figures 5a-5c) reveals a cyclic seasonal fluctuation. Water levels usually rise from spring breakup to a maximum in summer and then decrease to a minimum during the winter months.

Lunar or solar tidal forces do not significantly affect lake levels because of the Lake's relatively small size (Ball, 1972). In addition to the seasonal increase in water volume, wind set-up and waves are other factors influencing lake level.

Set-ups are created by either strong northerly or strong southerly winds. Daily maximum wind velocities and directions are listed on Tables 1a, 1b, 1c. A daily comparison of lake levels and peak wind velocities suggests a positive correlation although frequently a lag of several hours must be considered (Einarsson and Lowe, 1968). As the open water season progresses the frequency of set-ups increases. The probability of a set-up producing a one foot rise in lake level doubles from 5% in June to 10% in October (Lake Winnipeg, Churchill and Nelson River Study Board 1971-75).

During 1966, Lake Winnipeg experienced a high water level some four feet above the average which is 713.22 feet above mean sea level. No causes for this phenomenon were given (Province of Manitoba Water Bulletin May 1967).

1973

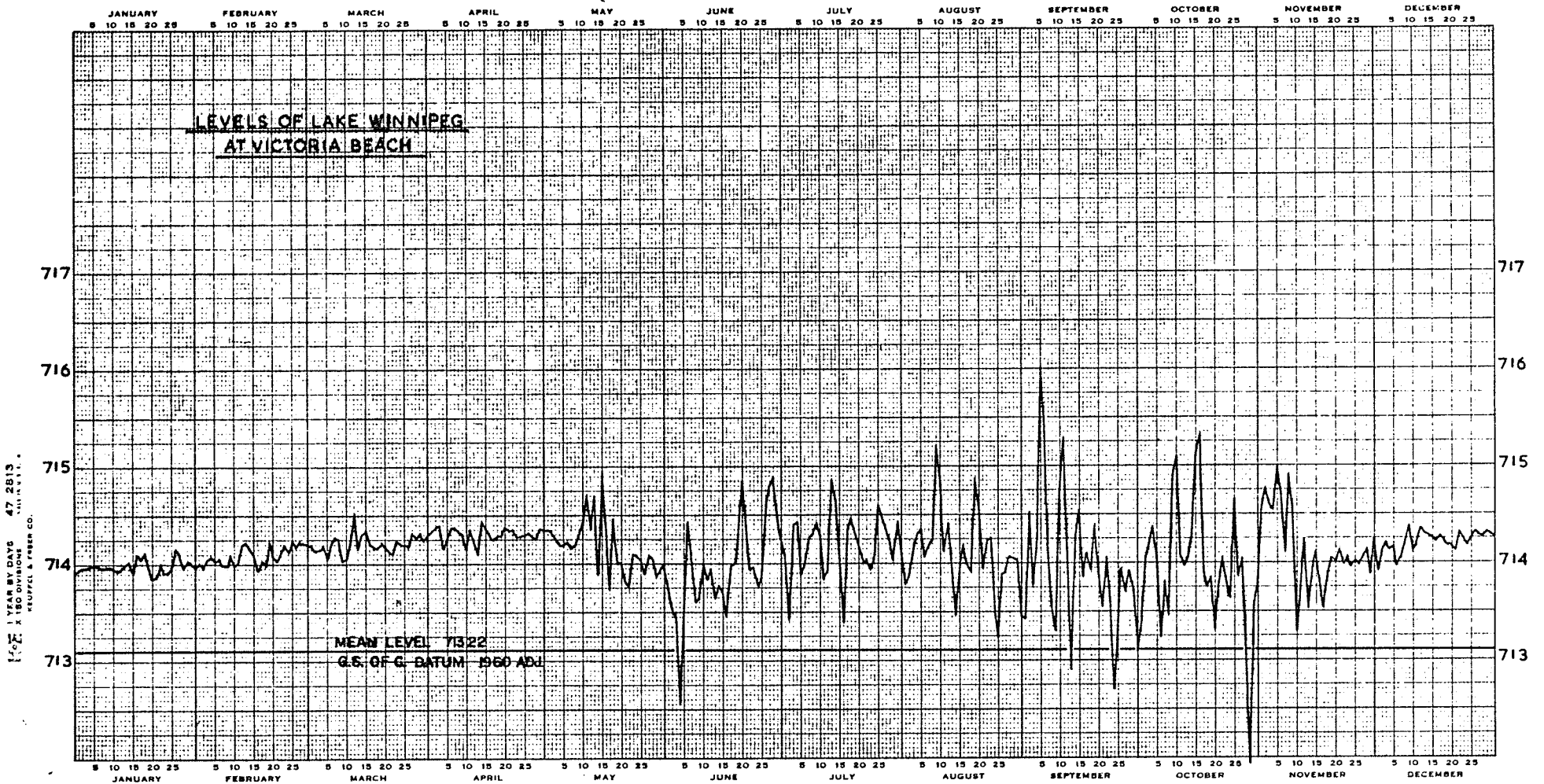


Figure 5a

1974

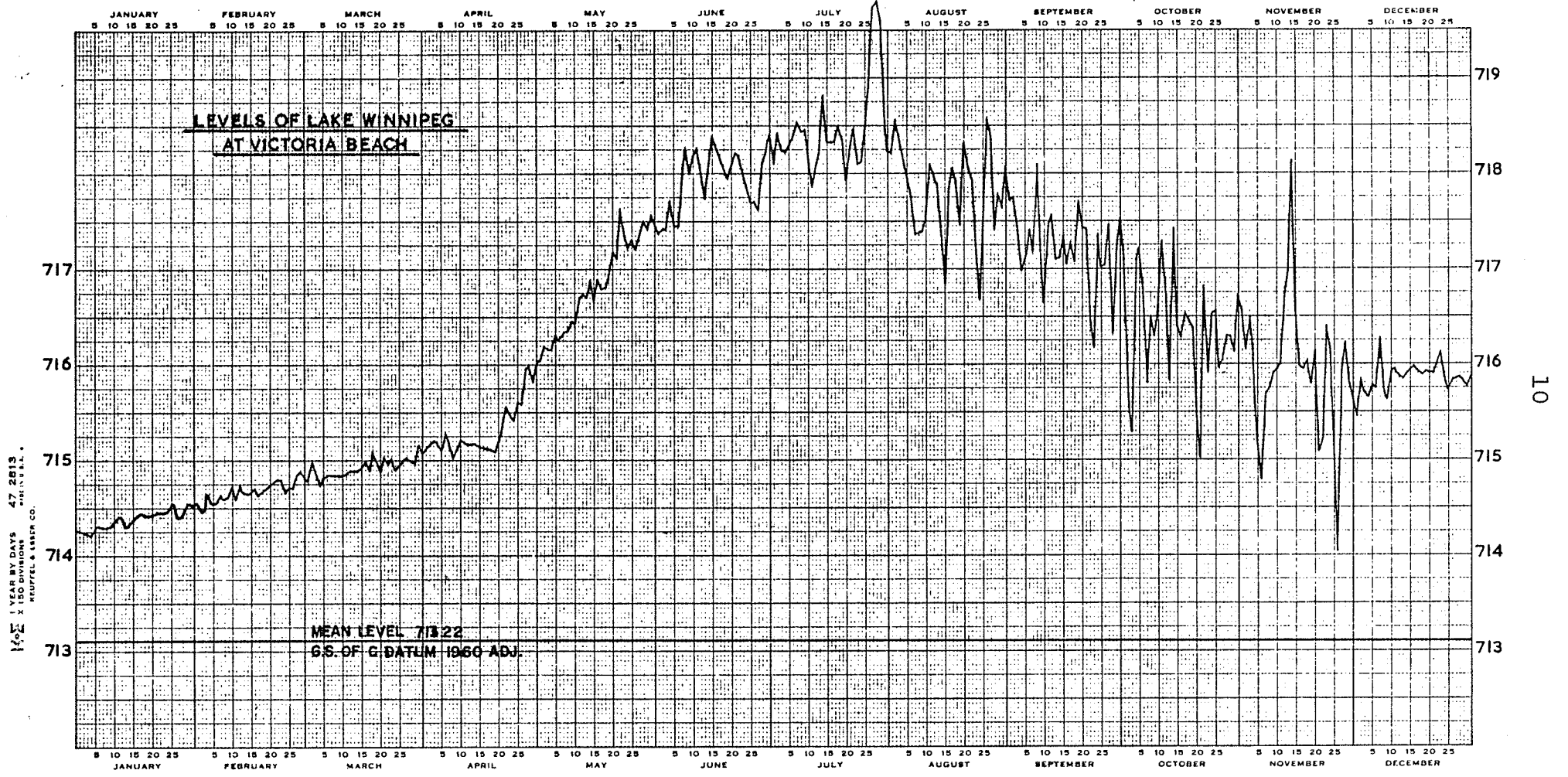


Figure 5b

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1975

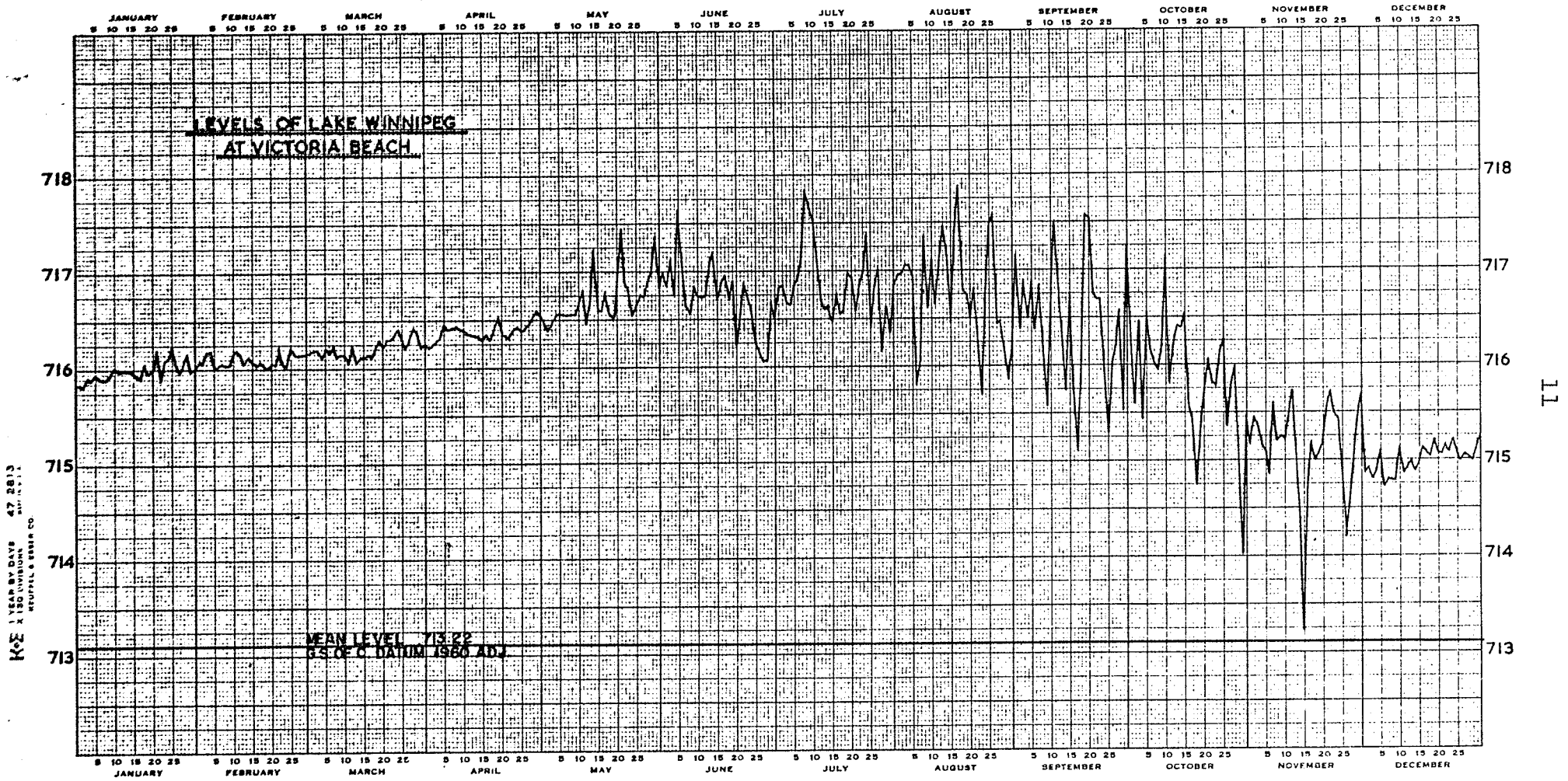


Figure 5c

TABLE 1a - WIND VELOCITIES

Day	<u>MAY 1973</u>			<u>AUGUST 1973</u>			<u>NOVEMBER 1973</u>		
	Prev. Dir.	Mean Speed	Max. Vel.	Prev. Dir.	Mean Speed	Max. Vel.	Prev. Dir.	Mean Speed	Max. Vel.
1	N	11.6	16	SE	6.2	11	NNW	12.3	18
2	SVL	10.0	15	SSE	5.0	11	NW	11.4	14
3	NNE	6.6	10	SSE	4.4	8	NW	11.9	15
4	NNE	5.6	11	WNW	7.8	13	NW	10.3	13
5	ESE	9.6	17	SVL	4.2	8	WNW	10.4	15
6	WSW	8.3	11	ENE	6.8	12	WNW	6.1	12
7	SW	7.4	11	SSE	4.5	10	SVL	10.0	16
8	ESE	5.3	9	SVL	6.9	13	W	13.6	17
9	NNE	9.0	15	SVL	11.9	18	W	8.4	13
10	NNE	11.6	17	SVL	7.1	12	S	10.8	17
11	NNW	11.9	15	W	4.9	7	SVL	7.0	12
12	SVL	10.7	16	SVL	4.5	10	NNW	9.8	18
13	NNE	12.0	18	SE	4.9	9	E	14.2	23
14	WSW	7.7	17	SVL	7.1	10	NE	9.6	16
15	NNW	14.4	25	SSW	5.0	8	ENE	8.7	14
16	NW	10.3	19	SVL	2.6	6	SE	8.3	15
17	S	8.3	14	ESE	4.5	7	SE	9.6	15
18	NW	8.5	14	SE	6.8	17	W	7.3	12
19	NE	6.2	9	NW	10.2	18	NNW	5.0	8
20	SVL	6.5	11	W	3.8	6	NW	7.1	11
21	E	9.4	13	SSE	4.3	9	N	12.5	16
22	ESE	8.7	17	SVL	4.2	11	WSW	11.0	16
23	NNE	5.9	13	ESE	4.3	8	WSW	8.0	13
24	NNE	12.3	16	SSE	10.2	15	ENE	4.6	9
25	NE	13.0	20	SVL	9.1	16	S	8.4	14
26	ENE	5.6	11	WSW	5.8	10	SE	8.8	13
27	N	7.1	13	SVL	7.5	15	WSW	10.2	16
28	NNE	10.2	13	W	10.3	28	WNW	8.3	12
29	SVL	6.0	14	WSW	7.2	13	WSW	10.7	16
30	NNE	6.3	14	ESE	7.5	12	NW	7.5	18
31	NNE	7.5	12	SVL	11.0	18			

TABLE 1b - WIND VELOCITIES

Day	<u>MAY 1974</u>			<u>AUGUST 1974</u>			<u>NOVEMBER 1974</u>		
	<u>Prev. Dir.</u>	<u>Mean Speed</u>	<u>Max. Vel.</u>	<u>Prev. Dir.</u>	<u>Mean Speed</u>	<u>Max. Vel.</u>	<u>Prev. Dir.</u>	<u>Mean Speed</u>	<u>Max. Vel.</u>
1	ESE	8.8	16	NNE	9.0	16	NE	14.0	21
2	NW	10.5	18	NNE	10.4	16	NW	7.6	10
3	SVL	4.9	13	SVL	5.3	8	NW	6.3	12
4	SSE	8.9	18	SVL	4.7	7	SSW	6.3	10
5	SVL	7.5	12	ESE	5.0	9	SSW	12.8	16
6	ESE	5.1	11	SVL	7.9	12	SSW	13.2	16
7	EVE	6.0	11	SSE	9.6	13	SSW	7.7	15
8	E	6.2	12	SSE	9.1	13	SSE	10.8	19
9	NE	5.0	9	SSE	11.1	15	WSW	9.6	17
10	ESE	7.3	11	SE	6.5	13	ESE	6.8	11
11	NE	12.8	20	NNE	8.3	14	NNW	8.5	14
12	SVL	6.5	10	SVL	5.2	9	NNW	11.5	17
13	NNE	9.7	17	ESE	9.0	15	NNW	13.9	21
14	NNE	10.5	14	SE	11.4	19	NW	14.5	21
15	SVL	7.4	11	SSW	15.7	27	WSW	6.1	9
16	NNE	9.7	13	WSW	11.3	18	SVL	7.1	15
17	NW	7.0	13	W	7.4	14	WSW	10.8	18
18	NNE	8.3	13	ESE	5.4	10	ENE	7.6	16
19	ESE	7.0	10	NNE	9.4	15	ENE	11.4	16
20	NNE	8.5	16	NNE	12.7	19	NW	6.5	11
21	NNE	8.7	16	NE	11.7	16	SSE	12.2	18
22	SVL	6.9	12	NW	10.5	14	NNW	9.3	15
23	NNW	12.5	16	ESE	7.3	15	WNW	4.9	9
24	SVL	6.4	9	SE	11.3	16	NW	5.9	13
25	NE	4.3	8	W	10.9	22	SSE	13.3	20
26	SVL	4.5	9	W	13.2	22	SSW	9.7	16
27	ESE	7.7	12	WNW	8.6	14	NW	9.9	14
28	NW	6.2	9	SVL	5.6	13	NNW	9.2	12
29	NNW	8.0	14	NNW	8.2	12	NW	3.5	6
30	SVL	8.2	16	NNW	15.8	25	SSW	5.5	11
31	WSW	7.5	15	NNW	8.2	12			

TABLE 1c - WIND VELOCITIES

Day	<u>MAY 1975</u>			<u>AUGUST 1975</u>			<u>NOVEMBER 1975</u>		
	Prev. Dir.	Mean Speed	Max. Vel.	Prev. Dir.	Mean Speed	Max. Vel.	Prev. Dir.	Mean Speed	Max. Vel.
1	WNW	6.8	12	SW	5.1	14	SSW	10.3	16
2	WSW	7.6	14	W	8.5	16	SVL	10.1	15
3	SW	2.8	7	SVL	7.6	14	W	12.3	22
4	NNE	6.5	10	NNW	5.3	9	WSW	6.1	12
5	NNE	7.8	12	E	5.2	9	SSW	8.1	13
6	NE	6.8	12	S	10.8	16	SSW	9.8	15
7	ENE	7.7	10	ENE	7.5	15	NW	6.3	14
8	NE	4.7	7	SVL	12.6	30	WNW	4.7	9
9	SVL	3.8	7	WSW	10.8	21	E	6.5	13
10	SVL	7.9	16	W	6.8	12	N	6.0	10
11	NE	6.3	13	SSE	6.7	21	NNE	11.8	17
12	SW	8.4	16	WNW	10.1	16	NNW	7.9	16
13	SVL	9.3	23	WNW	8.3	12	S	11.9	16
14	N	10.2	16	SVL	2.8	7	S	8.8	15
15	SVL	6.5	10	SW	7.4	14	W	5.3	12
16	SE	9.3	20	WNW	10.8	16	SW	8.3	14
17	NNW	8.1	13	WNW	7.5	13	WSW	4.3	10
18	SVL	8.9	20	WNW	2.9	6	SVL	7.8	13
19	SSE	6.7	13	ESE	6.6	13	SVL	7.5	13
20	NNW	10.0	16	ESE	9.5	16	NNW	7.2	12
21	NNW	9.2	15	NW	5.8	10	WNW	5.6	12
22	NE	5.6	9	SSE	9.5	15	SVL	6.3	10
23	NNE	6.6	10	SSE	10.5	15	WNW	7.8	13
24	SSW	9.0	15	ENE	8.5	17	NW	4.3	7
25	SSW	8.7	15	W	12.9	20	SSW	5.1	7
26	W	9.9	16	WNW	7.4	13	S	10.1	14
27	SSW	7.1	10	SVL	6.6	9	S	6.8	10
28	SVL	4.3	13	SVL	8.9	14	S	5.3	13
29	NW	7.8	14	SE	10.8	16	NNW	11.3	16
30	NNW	8.1	14	SVL	8.5	14	NW	9.2	15
31	NE	4.3	10	SSW	6.9	12			

Unfortunately, records of wave heights on Lake Winnipeg do not exist. Wave heights can be computed from wind speed, wind duration, wind direction and characteristics of the water body and shoreline configuration. However,

"since shoreline configuration is a significant factor in computing wave height and associated wave uprush, the values vary widely from one beach area to the next."

(Lake Winnipeg, Churchill and Nelson River Study Board, 1971-75, p.58).

During the period of December to April, the lake and the foreshore of the beach are frozen and therefore not subjected to major water fluctuations so that none of the normal limnic processes take place.

Climate

The area experiences a continental climate. It is characterized by long, cold winters and relatively short, cool summers. Monthly mean temperatures may vary from an average of 66^oF for July, the warmest month. Annual precipitation for the area is around 20 inches. The annual snowfall is approximately 60 inches and the summer rainfall for May, June, July approximately 8 inches (Table 2).

Regional Geology and Physiography

Pleistocene glaciations are mainly responsible for the surficial deposits covering the immediate Grand Beach

TABLE 2 - ANNUAL PRECIPITATION IN INCHES
AT GIMLI, 1975

<u>MONTHS</u>	<u>1973</u>	<u>1974</u>	<u>1975</u>
January	.07	1.61	1.94
February	.48	.42	.75
March	.50	1.04	1.13
April	1.53	2.48	1.43
May	1.16	3.27	1.47
June	6.67	1.19	3.45
July	1.60	1.01	2.49
August	5.86	3.40	4.69
September	4.65	3.36	1.98
October	3.02	.24	1.50
November	2.37	.38	.59
December	<u>1.02</u>	<u>.26</u>	<u>.85</u>
Total	28.93	18.66	22.27

Department of Transport Weather Records

area. The glacial drift occurs over a bedrock of Ordovician limestone, dolomite, red shale and sandstone, which is rarely exposed. The thickness of the glacial drift varies from 50 to a maximum of 200 feet.

The Grand Beach area is divided into the West and East Beach (Figure 6). Most of the sand deposit is in the form of a bay-mouth bar enclosing a lagoon. The East Beach, north-east of the bar, was chosen as the study area. The area was then divided into two sub-regions. Since the shore line is oriented in a northeast-southwest direction, these two sample areas are designated as the North and South Beach (Figure 7 and 8).

Both the North and South beaches are approximately 60 feet wide with well defined foreshores and backshores. The foreshore slope of the North beach was not as steep as that of the South beach. The average overall gradient was a gentle 1.2° for the North beach, and a significantly steeper 2.3° for the South beach. Both beaches had shrub vegetation including birch, alder and willow. Only the North beach, however, had growth within the sample area.

Solohub and Klovan (1970) state that the Grand Beach area experiences a strong southwest littoral drift which is responsible for the transportation of material in this area. Most of the sediment derived from glacial deposits

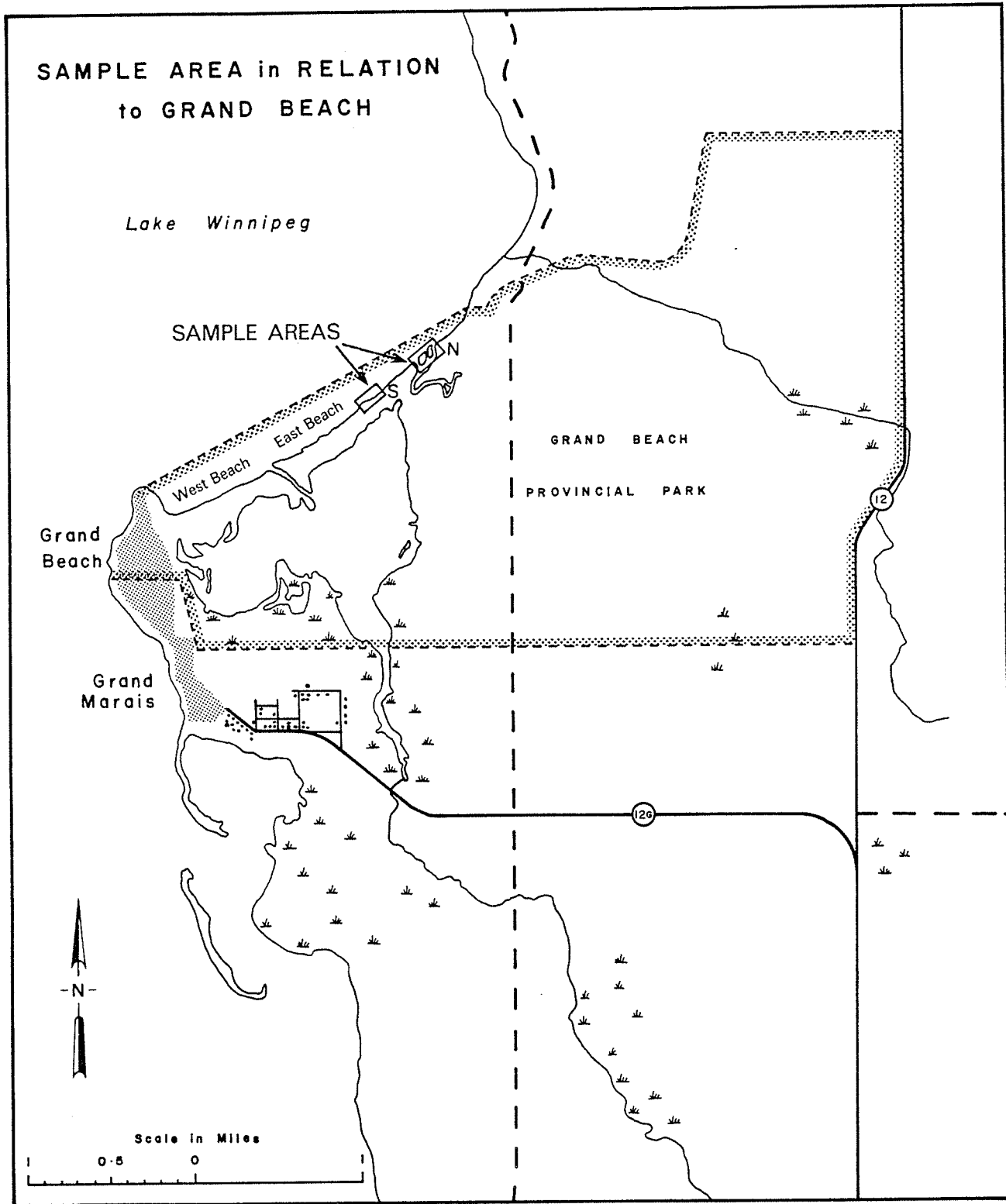


Figure 6



Figure 7

Looking Northeast Depicting North Beach



Figure 8

Looking Southwest Depicting South Beach

originally comes from local Paleozoic bedrock, particularly the Ordovician Winnipeg Sandstone formation.

Literature Review

Researchers have recognized that the dynamic processes associated with the land-water interface are complex and constitute an area of continuing research. The majority of this research has involved the marine environment. Coates (1972), however, suggests that many of the characteristic marine geomorphic processes may also occur in lacustrine environments. Studies of limnic environments are not as prolific as marine studies, pointing out the possibility for further intensive studies in this area.

Darling (1964) recorded the seasonal changes in morphology of eight beaches along the eastern seaboard of the U.S.A. between southern New Jersey and Rhode Island. The positional displacement of index contours from one survey period to the next was recorded. In this study, recordings of beach profile and tide and wave data were made, and only that portion which is above sea level was included.

Rohrbough, Koehr and Thompson (1964) recorded profile changes on a quasi-weekly and daily basis, from July 1963 to January 1964, along two profiles in Monterey Bay, California. Each profile consisted of a series of 5 cm. pipes driven into the sands perpendicular to the beach. Measurements of sand

elevation were then taken relative to the tops of the pipes, the heights of which were then referred to a common datum. During the study period they experienced an erosion-deposition range in sand height of 100 cm. at the low-tide and mid-tide levels and diminishing to zero at the back area of the beach where waves did not reach.

Strahler (1964) employing a profile-sampling interval of one half-hour on a beach in New Jersey, found a semi-diurnal cycle of cut and fill and associated this with the semi-diurnal tide.

Ingle (1966) studies the lateral movement of beach sand along the coast of California. To trace sand movements under a wide range of foreshore-inshore conditions he made use of fluorescent dyes. A significant percentage of the dyed grains were transported obliquely offshore under all wave conditions.

Inman, Komar and Bowen (1968) also investigated the longshore transport of sand along California beaches. The data indicated that the longshore transport of sand is directly proportional to the longshore component of wave power.

Thompson and Harlett (1968) investigated the relationship between the daily beach profile and wave frequency of Del Monte Beach, California. The data demonstrated that

the daily tidal range was several times larger than the wave height and associated runup during most of the study. They concluded that the general shapes of the profiles significantly reflected the local tidal characteristics.

Suryaprakaso Rao, and Kassim (1970) observed the seasonal changes of a beach at Surathkal, along the west coast of India about 20 km. north of Mangalore, during the six month period from February to August, 1969. Profile observations and sediment sampling were carried out at frequent intervals. Profile measurements were taken at 3 to 9 day intervals depending on the magnitude of the changes which had taken place. Sediment samples were also collected. The results indicated that while the beach was subjected to low-wave steepness up to the middle of May, buildup occurred. From May through August there were periods of erosion and deposition with the overall effect being one of erosion. Sediment characteristics were recorded and the relationship between grain size (median diameter in mm.) and the foreshore slope was shown; the curve being inserted for comparison with other beaches.

McCann (1972) investigated the special characteristics of Arctic beaches. Due to the long, near total ice cover each year, Arctic beaches are termed as low energy beaches, exhibiting low rates of longshore sediment transport and

little change in beach material. Most areas show yearly variations in ice cover conditions revealing similar variations in wave action. During certain years Arctic beaches may show very little change due to the inhibiting effect of the ice, whereas infrequent catastrophic storms will show decisive changes in beach characteristics. In the particular area of Radstock Bay, S.W. Devon Island, strong winds from the southeast quadrant generated those waves which had the greatest effect on the beach. Dominant longshore movement of material took place during occasional storms such as that of August 11-12, 1969.

Hume and Schalk (1967) in the Point Barrow area of Alaska, recorded the effects of a catastrophic storm which produced a movement of twenty years of normal transport of beach sediment.

Saylor and Hands (1970) investigated the movement of longshore bars of Lake Michigan. Results showed that a significant migration of the offshore bars occurred due to a change in lake level. During 1967-1969 Lake Michigan recorded a rise of one-half meter, constituting a shoreward movement of bar crests and troughs over a distance averaging 30 meters. Furthermore they noted extensive shore erosion because longshore bars were now not as effective in dissipating wave energy.

Dubois (1972) recorded seasonal variations in beach and near shore environments along a profile of Lake Michigan. Dubois concluded that changes occur on most marine beaches due to seasonal variations of wave regimes, whereas limnic beaches under study respond to seasonal fluctuations of lake level.

Engstrom (1974) examined foreshore sediments and slopes of 39 beaches in the Apostle Islands of northern Wisconsin. He attempted to link beach foreshore parameters with selected coastal processes characterizing the individual beaches. Engstrom concluded that in time, relationships based on statistical analyses are possible, stipulating however, that repeated testing is still required.

In summary, the literature review reveals that beaches are indeed constantly undergoing changes. The studies range from observations taken on a daily to a seasonal basis. Under normal environmental processes a cycle of deposition and erosion is observable.

There is, however, some evidence that most movement of material may take place during catastrophic events. In these studies, only slight consideration has been given to disastrous storms.

Chapter IIMETHODOLOGYFieldwork

Topographic surveys were carried out three times during the 1974 summer season. The first survey corresponded with lake ice breakup and was taken May 29. The second, or midsummer, survey took place August 29 and the third, pre-freezeup, was taken November 29. These times were assumed to coincide with the open water season, thus covering the low, high, and low energy cycle.

The two study sections are 400 yards apart and are separated by a dense growth of willow bushes and fairly high sand dunes. Grid systems (6 columns x 5 rows) were established over each sample beach. The grid includes the following morphologic zones as defined by King (1959); the back shore, the foreshore, and a small part of the near-shore bottom zone. This system was chosen as it appeared to be the most accurate and easily identifiable method of obtaining samples and profiles for comparison over three surveys.

Grid lines orthogonal to the shore line are referred to as columns (C) and lines parallel to the shore as rows (R) (Figure 9). Columns and rows were uniformly spaced; the intervals between the columns being sixty feet, and

LAY-OUT OF GRID SYSTEM

position approx.- not to scale

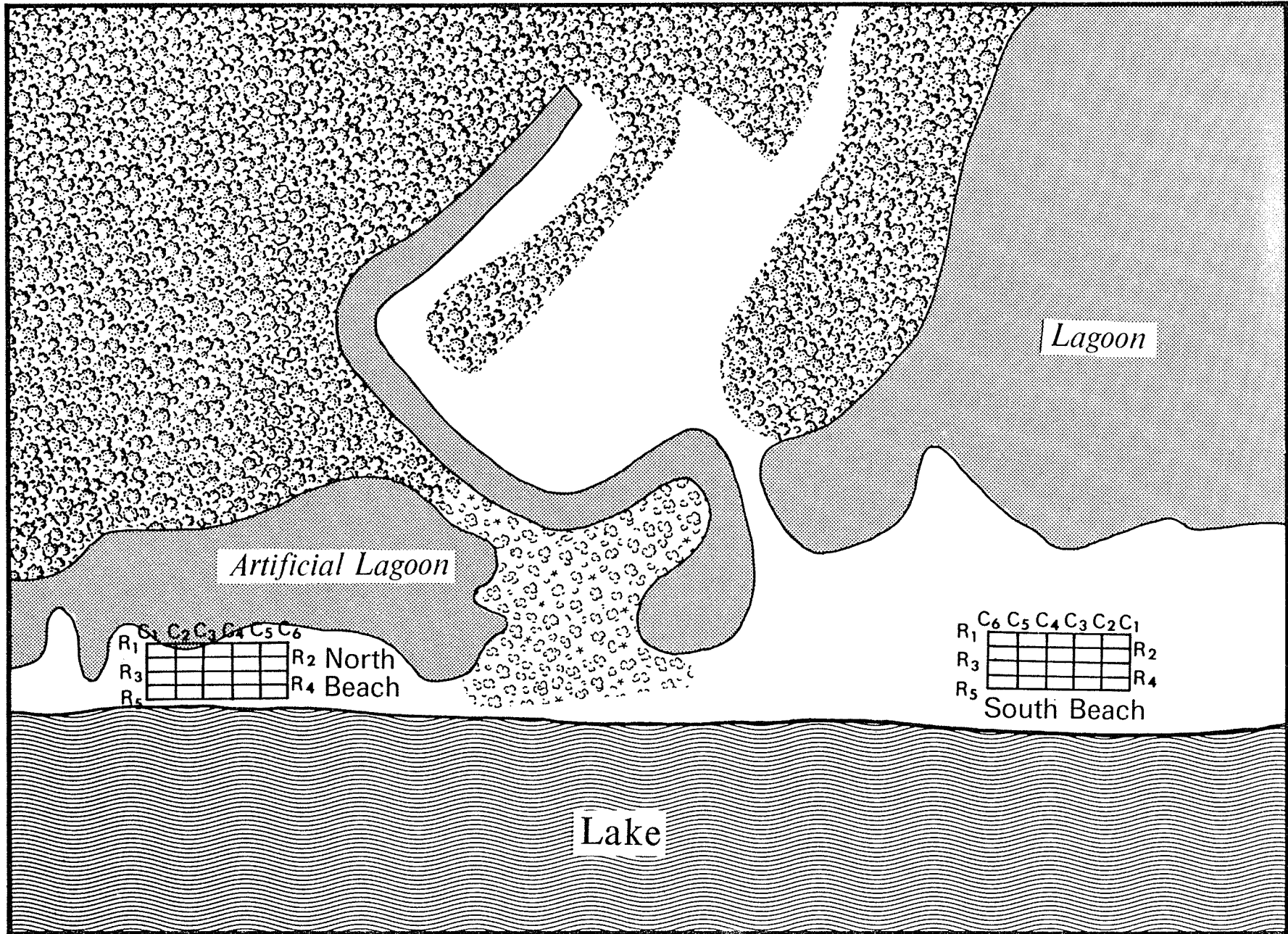


Figure 9

between the rows, fifteen feet. Stakes of 2" x 4" wood, capped by a 1/8" thick sheet metal top, were driven 5 feet into the ground in each area. Permanent reference stakes (C_1R_1 and C_6R_1) defined a base line with origin at C_1R_1 . This established base line facilitated repetitive surveys.

Knowledge of the lake level on the given dates allowed actual measurements to be given to the obtained levelling data. Accurate measurements of distance from shore line to Row One were also taken at each column.

A total of 30 samples from each Beach, each weighing 500 grams, was collected during each survey. All samples consisted of the upper 16 mm. of sediment. Subaqueous samples were taken from a depth of water of about 12 to 16 inches. In this way, each sample is believed to represent the same sedimentation unit.

Laboratory Work

Profiles were drawn to scale along each column from the data obtained from the survey. All 180 samples were split using an Endicott sample splitter. The sample size selected was 100 grams; it was weighed out very accurately and sieved for 15 minutes on a RO-Tap shaking machine. The grade sizes are based on the phi (ϕ) scale (Krumbein, 1938). Eight inch diameter screens were used at one quarter phi (ϕ) intervals ranging from -1.00ϕ to 4.00ϕ inclusive. The resulting

sediment fractions were weighed to 0.01 grams on a Sartorius electric balance.

Chapter III

Results of the Field and Laboratory Study

Data Presentation

Locations of the samples, thirty from the North Beach and thirty from the South Beach, are shown in Figure 9.

Profiles and elevations are shown in Figures 10 to 13. Each profile and elevation is identified as to which area, column and time period of the season it belongs. The seasonal changes of volume of sand are recorded in Table 3.

The figures shown in Table 4 (Appendix B) represent the grain size data from the sieve analyses as used in the computer program for statistical analyses. Table 5 (Appendix C) shows the graphical parameters according to columns and rows. Figures 14 to 15 are the cumulative curves of selected samples from both areas at different time (seasonal) intervals.

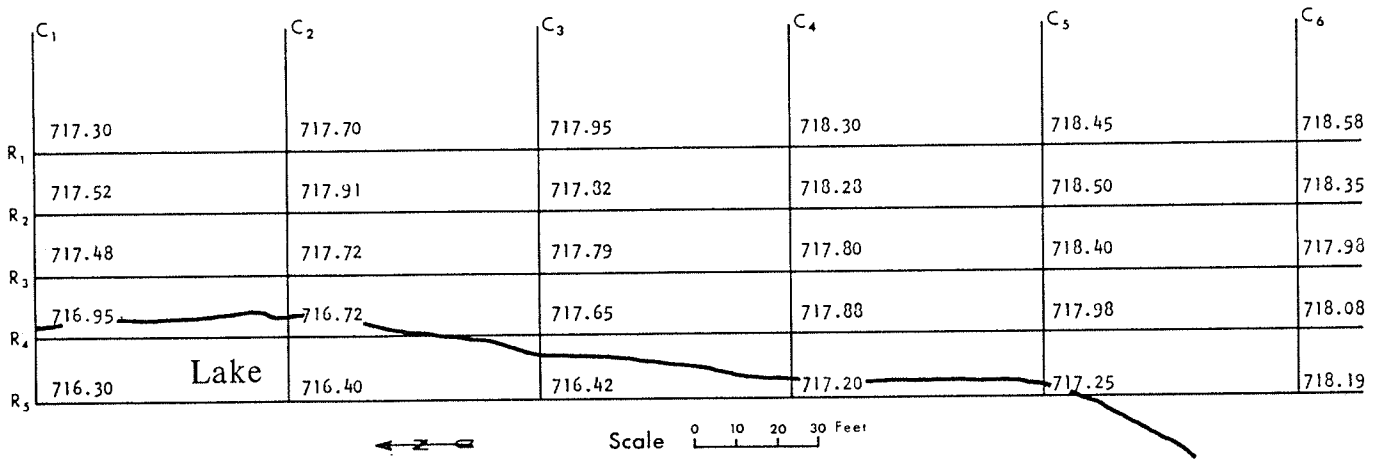
Composition of Sediments

The mineral contents of the sand samples were found to be almost identical to the results of Solohub and Klován (1970) and Ball (1972).

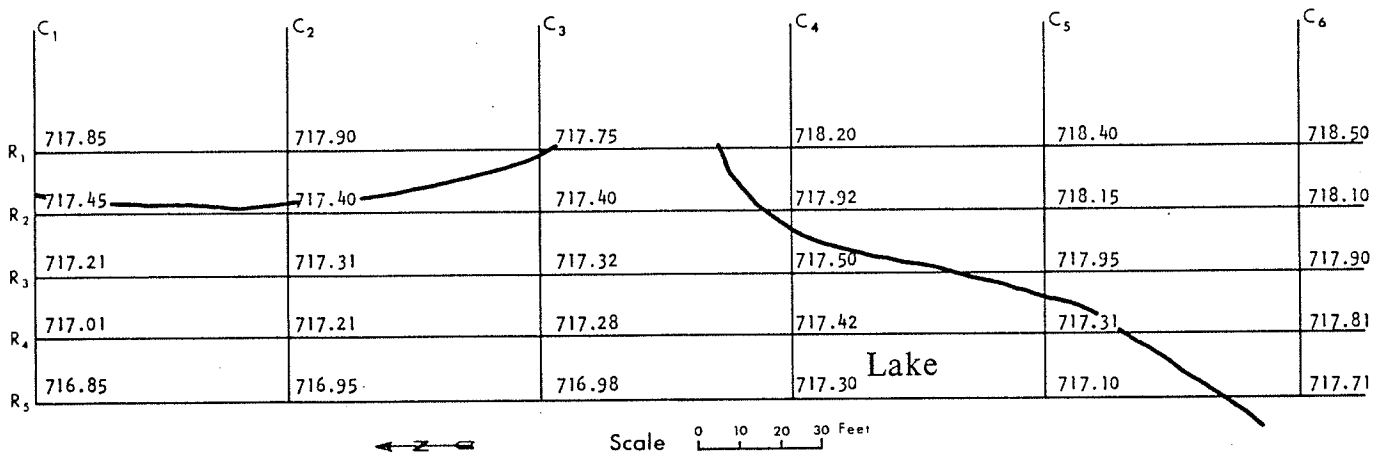
The sample material can be broken down into three groups, with quartz being the largest, accounting for 96 percent. Feldspar amounts to approximately 2 percent and heavy minerals constitute the balance of the sample material.

Figure 10

(a) NORTH BEACH ELEVATIONS May 29, 1974



(b) NORTH BEACH ELEVATIONS Aug 29, 1974



(c) NORTH BEACH ELEVATIONS Nov 29, 1974

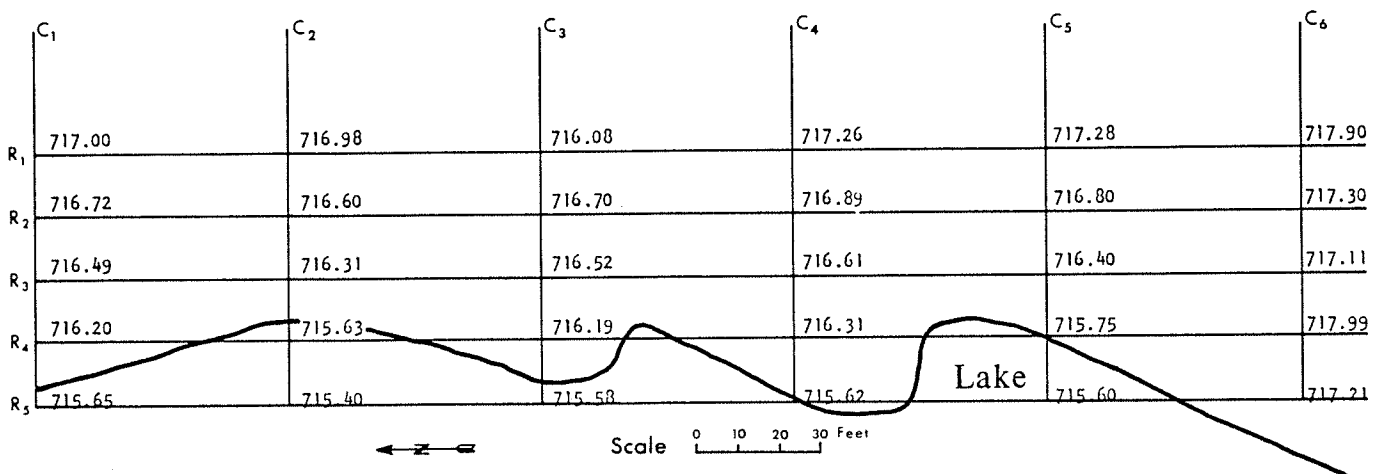
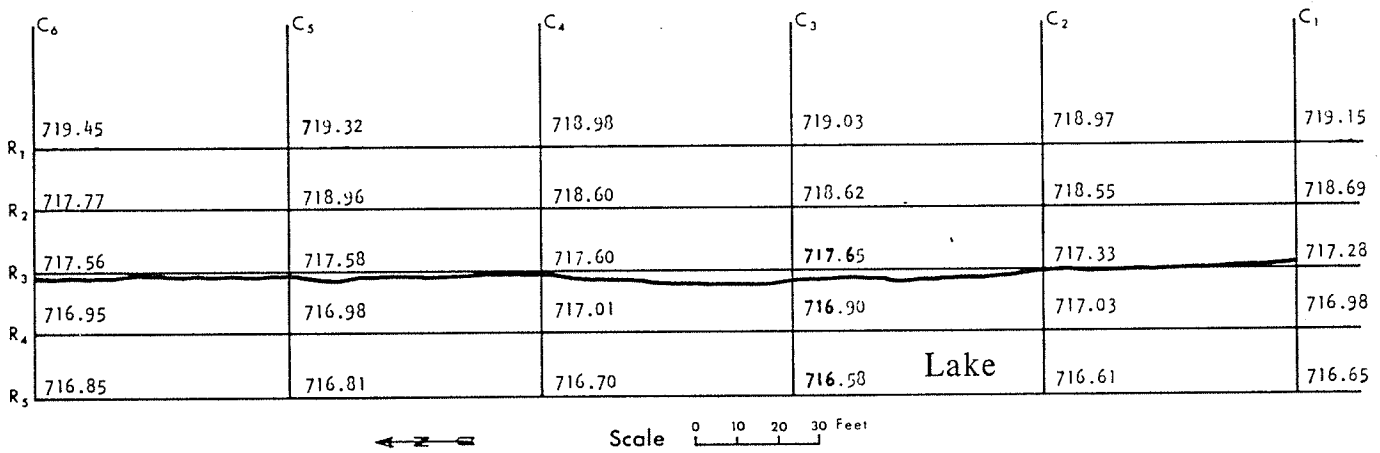
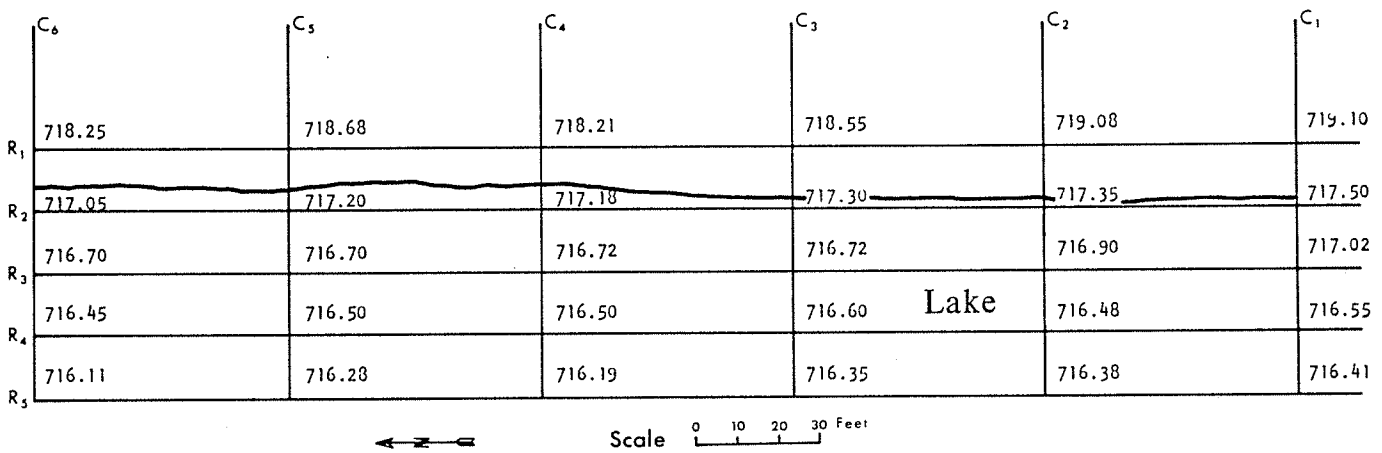


Figure 11

(a) SOUTH BEACH ELEVATIONS May 29, 1974



(b) SOUTH BEACH ELEVATIONS Aug 29, 1974



(c) SOUTH BEACH ELEVATIONS Nov 29, 1974

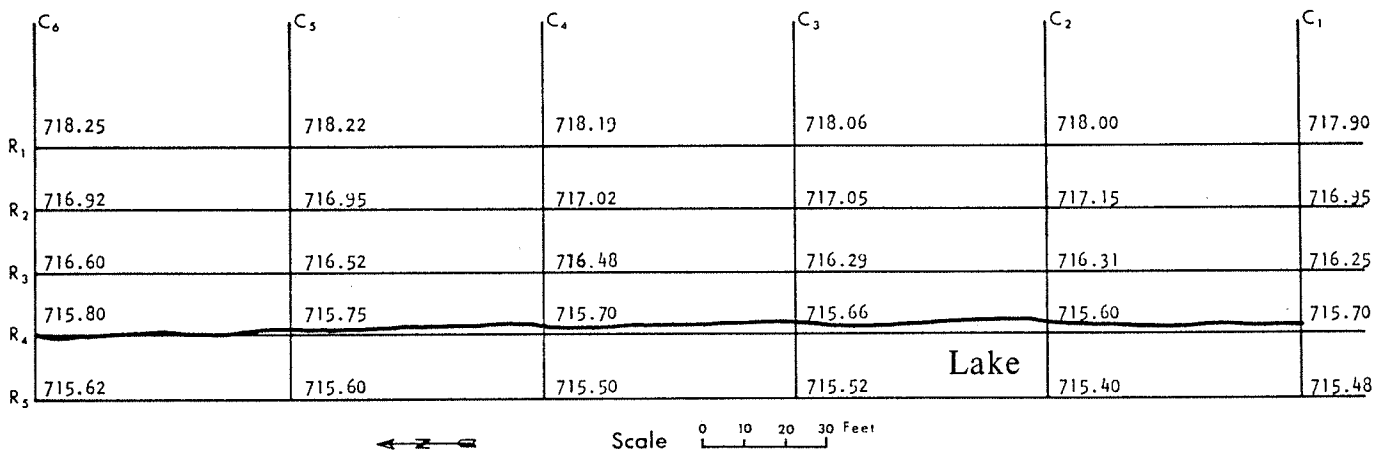
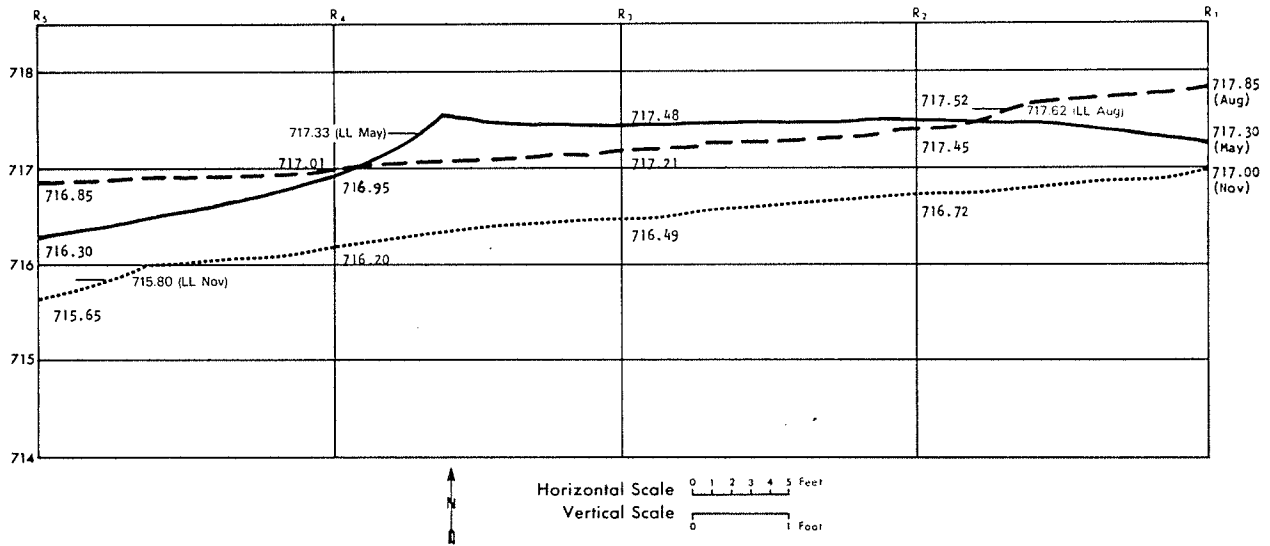


Figure 12

(a) PROFILE C₁ NORTH BEACH 1974



(b) PROFILE C₂ NORTH BEACH 1974

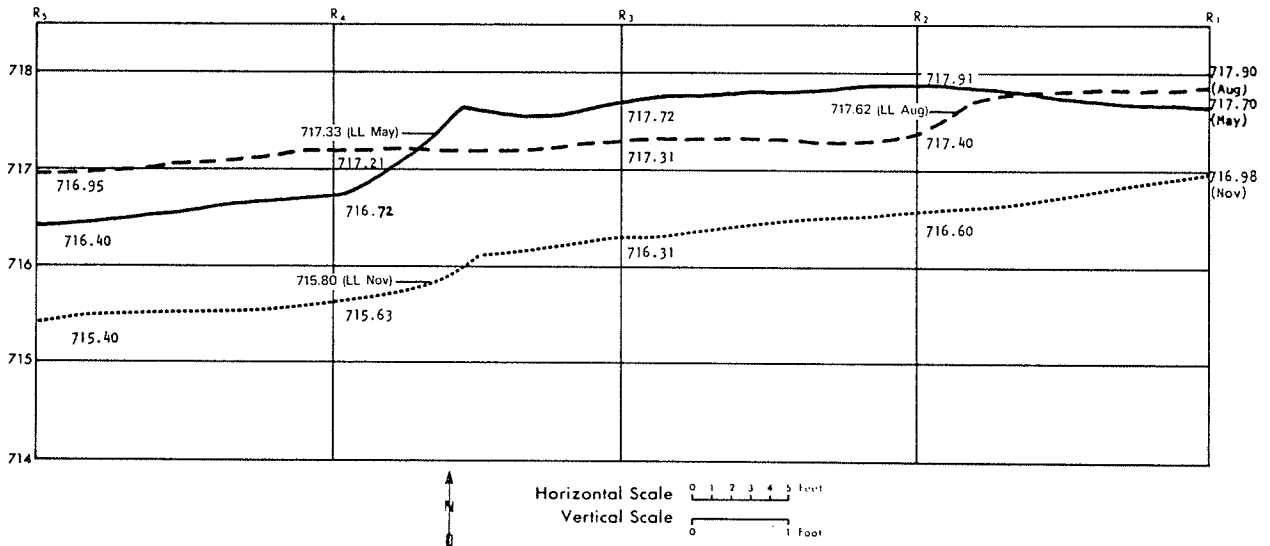
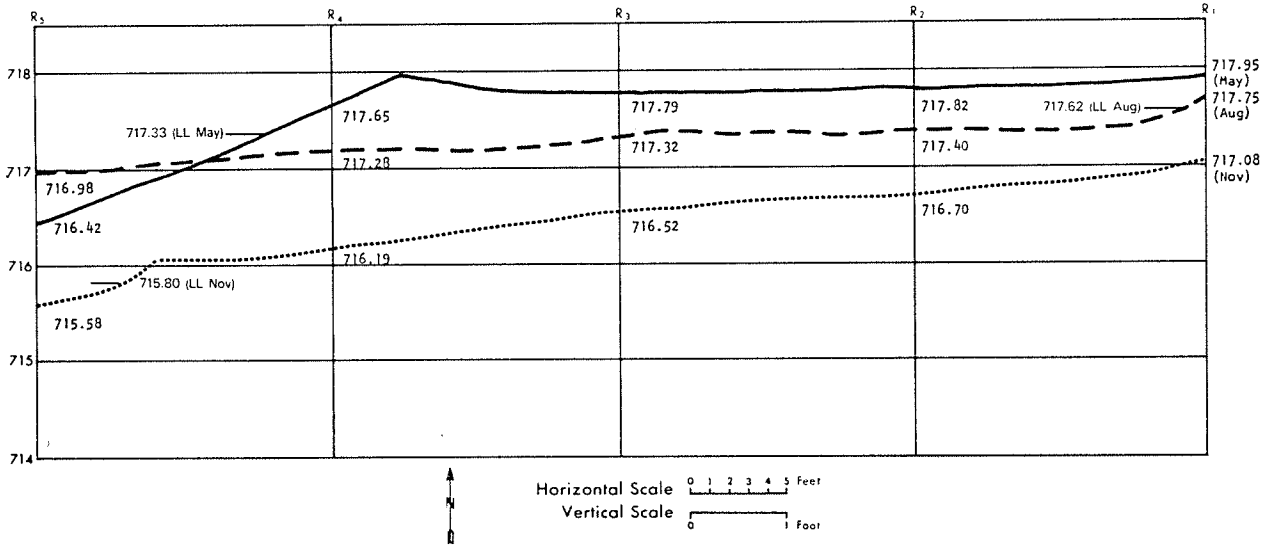


Figure 12

(c) PROFILE C₃ NORTH BEACH 1974



(d) PROFILE C₄ NORTH BEACH 1974

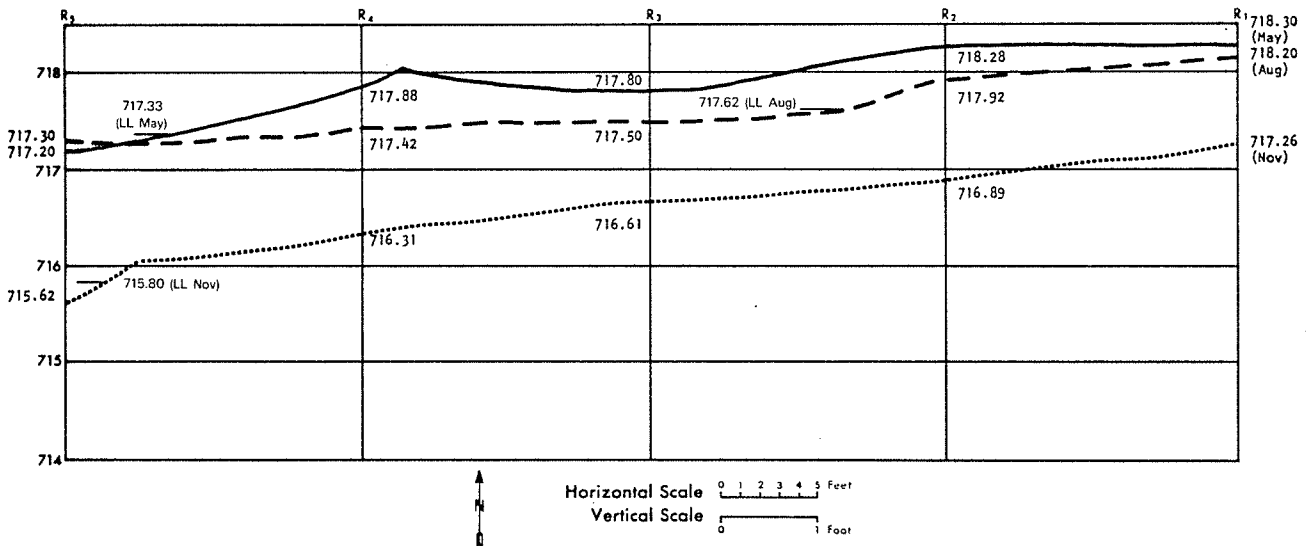
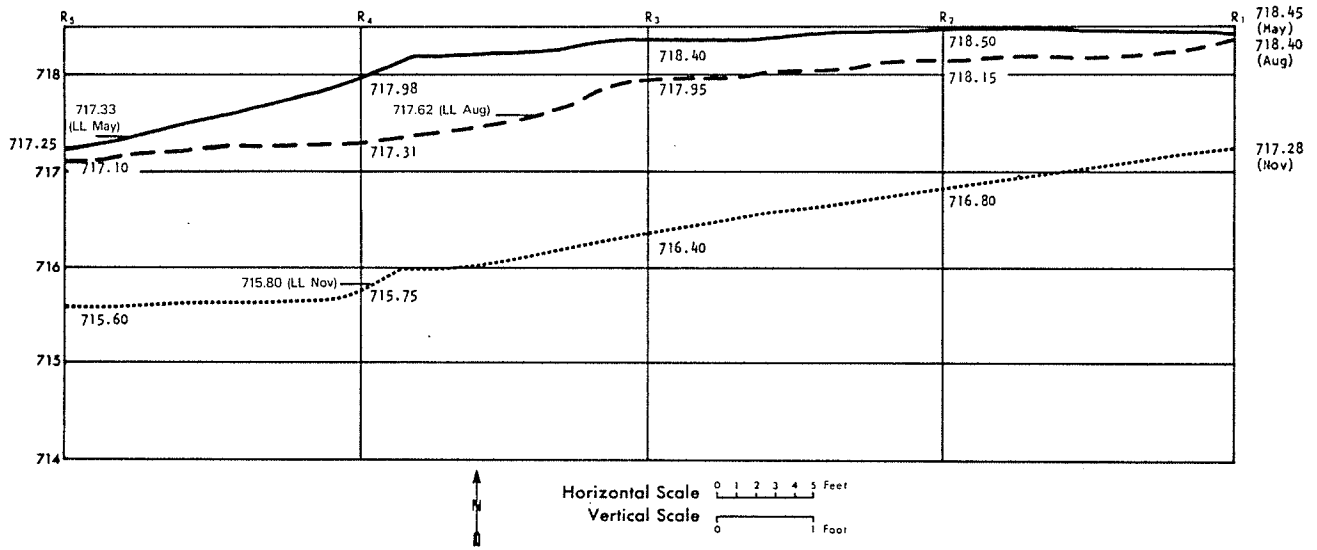


Figure 12

(e) PROFILE C₅ NORTH BEACH 1974



(f) PROFILE C₆ NORTH BEACH 1974

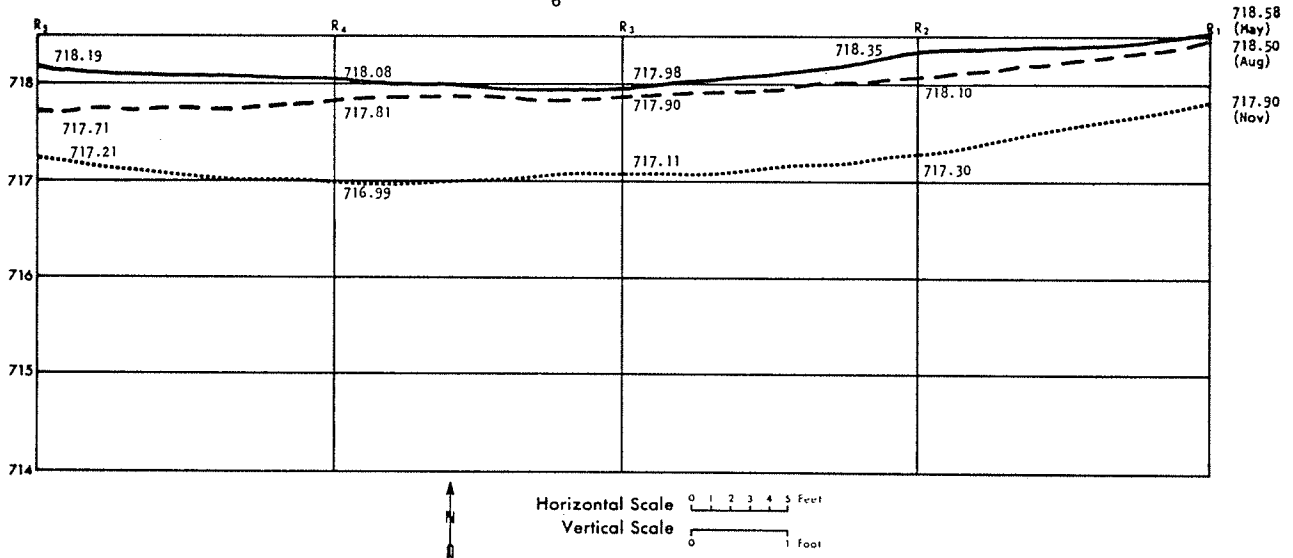
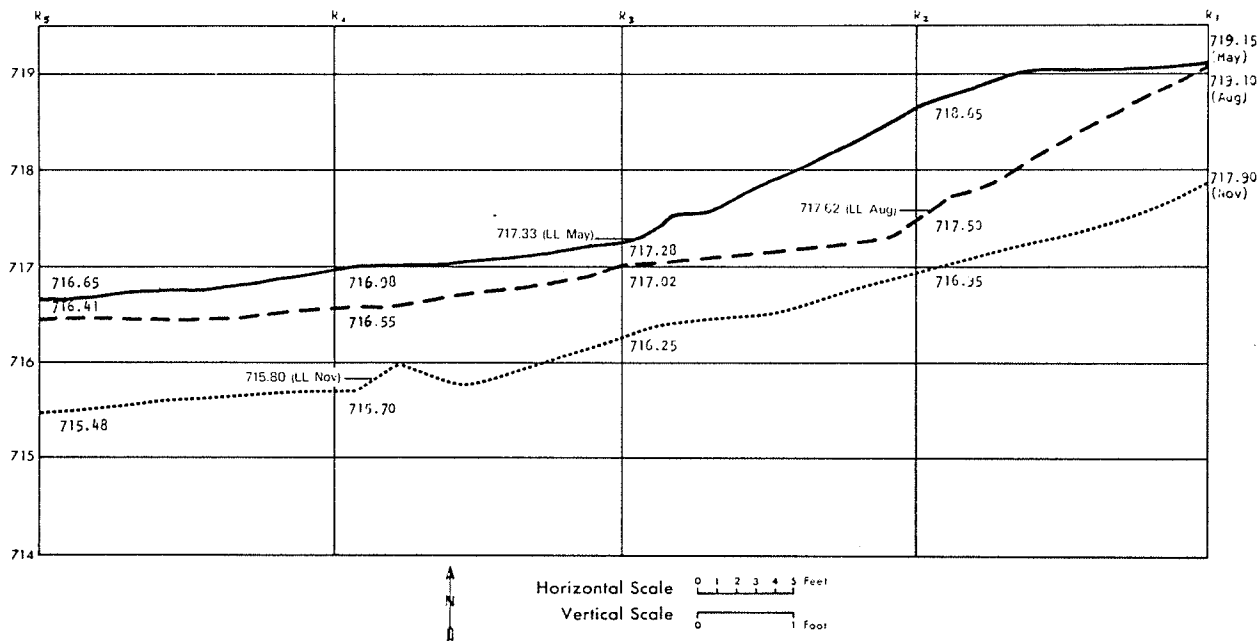


Figure 13

(a) PROFILE C₁ SOUTH BEACH 1974



(b) PROFILE C₂ SOUTH BEACH 1974

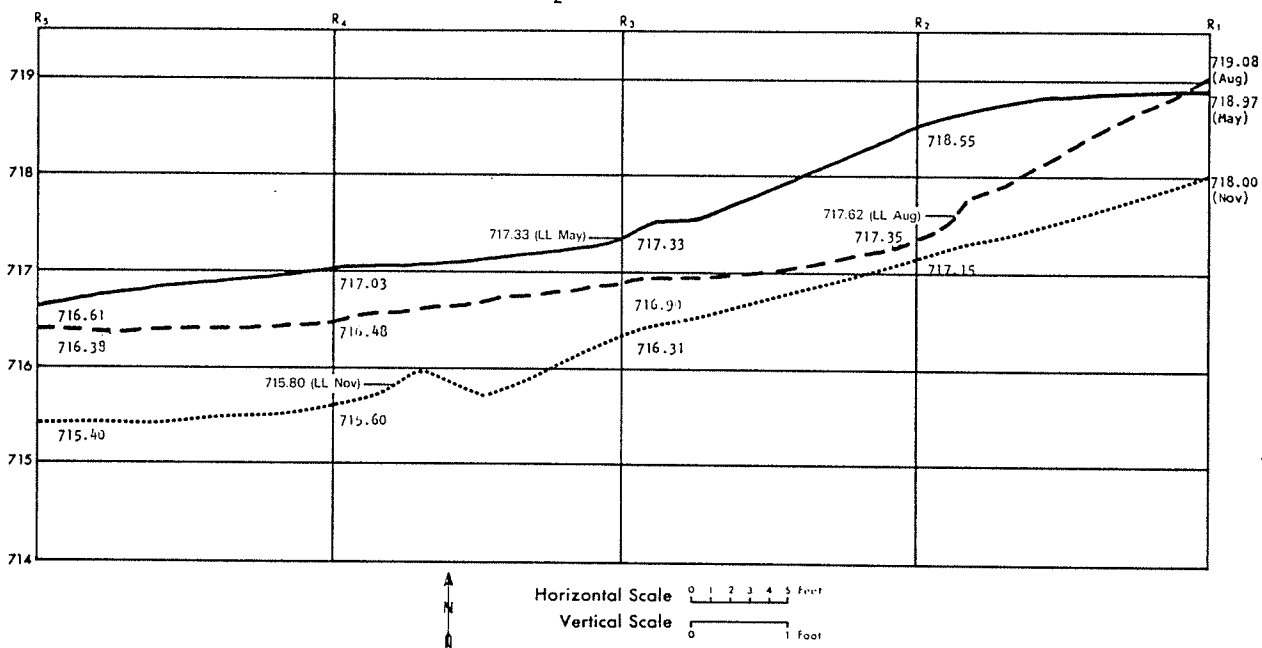
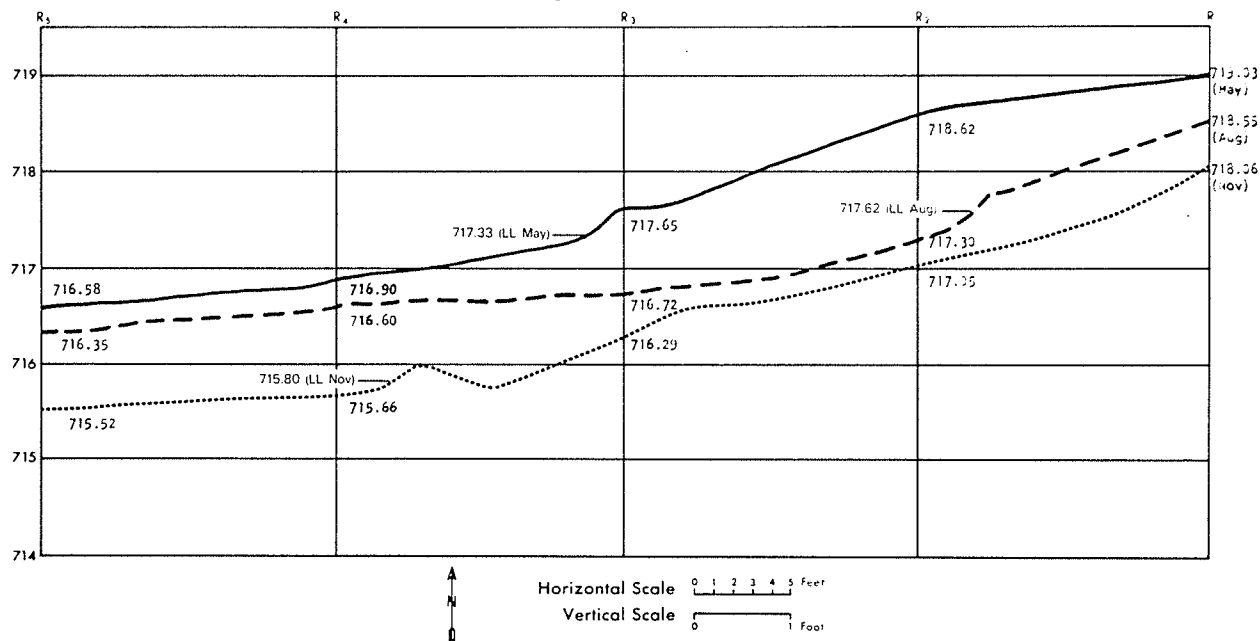


Figure 13
(c) PROFILE C₃ SOUTH BEACH 1974



(d) PROFILE C₄ SOUTH BEACH 1974

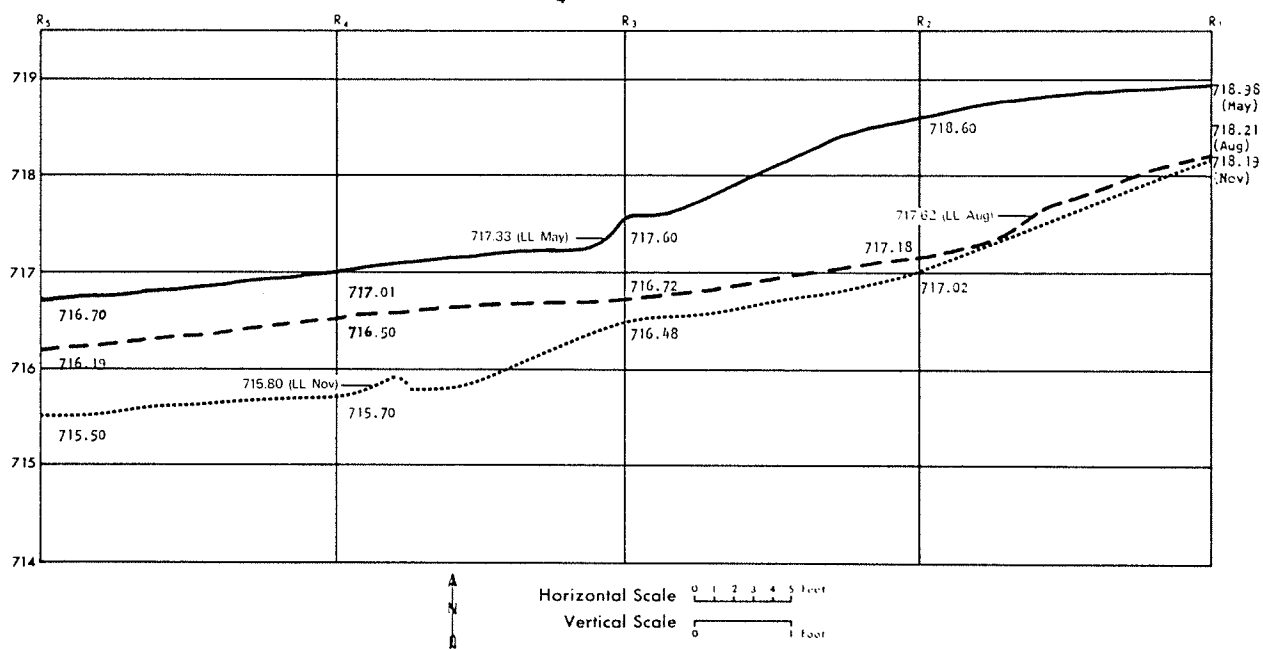
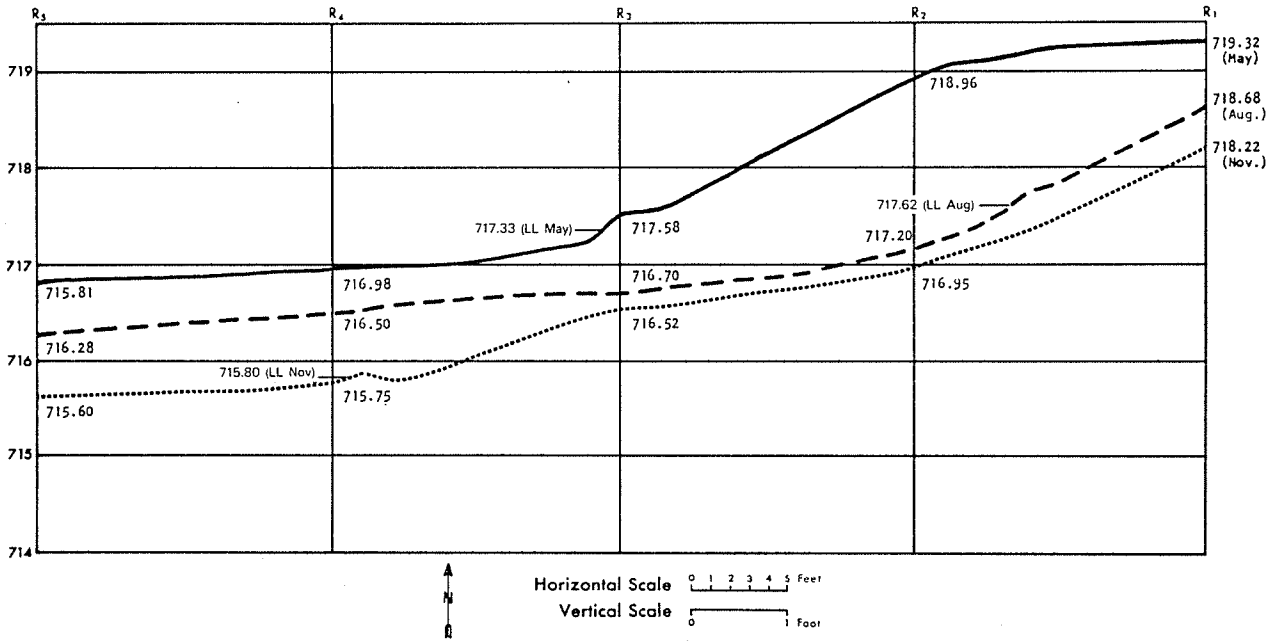


Figure 13

(e) PROFILE C₅ SOUTH BEACH 1974



(f) PROFILE C₆ SOUTH BEACH 1974

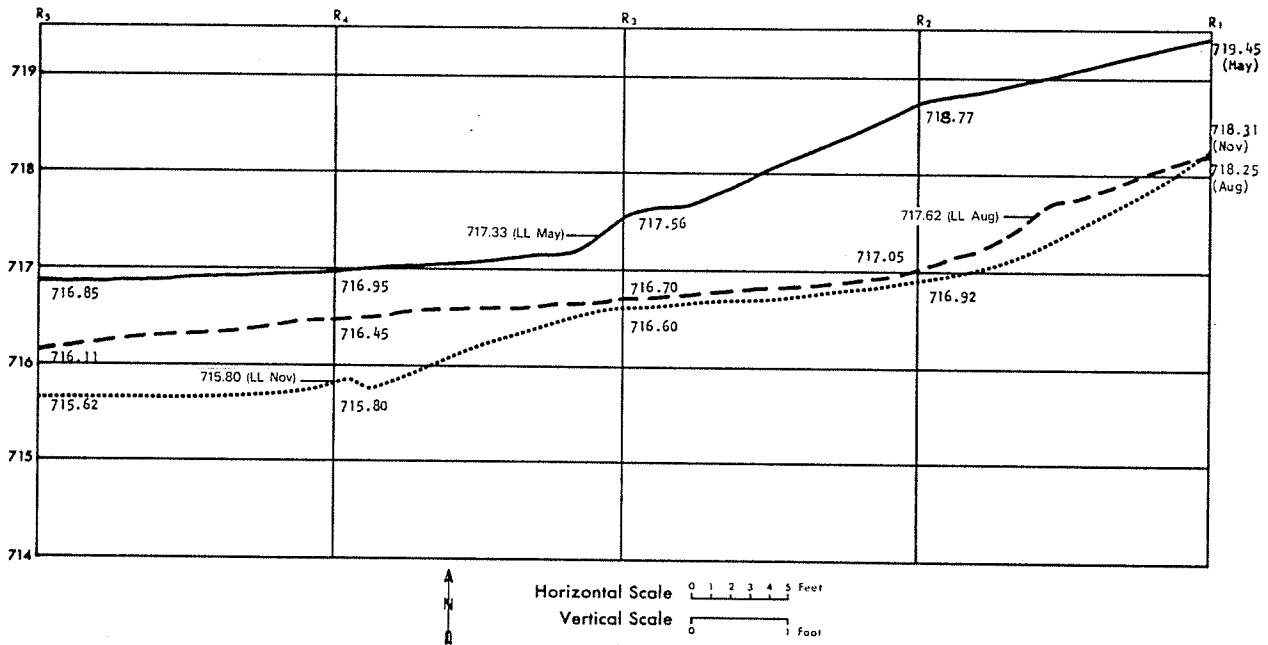


Table 3Volumetric Changes of Sand

North Beach

May - August	Erosion - -	168.38 cu. yds.
Aug. - Nov.	Erosion - -	<u>809.18</u> cu. yds.
	Total Erosion	977.56 cu. yds.

South Beach

May - August	Erosion - -	569.38 cu. yds.
Aug. - Nov.	Erosion - -	<u>345.83</u> cu. yds.
	Total Erosion	915.21 cu. yds.

FIGURE 14
CUMULATIVE CURVES OF BEACH SANDS
NORTH BEACH (a) C_1R_3

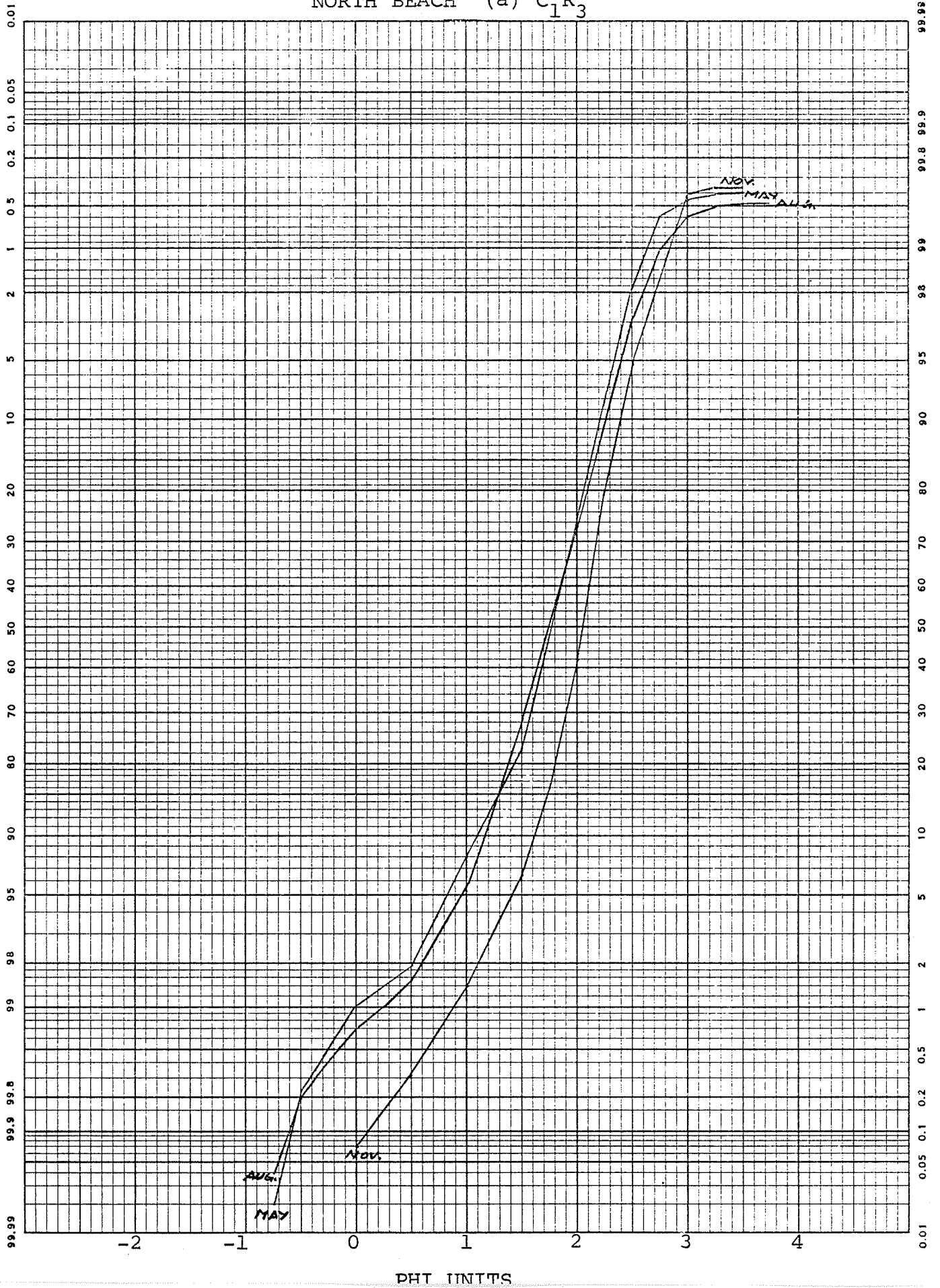


FIGURE 14 CONTINUED
(b) C_2R_5

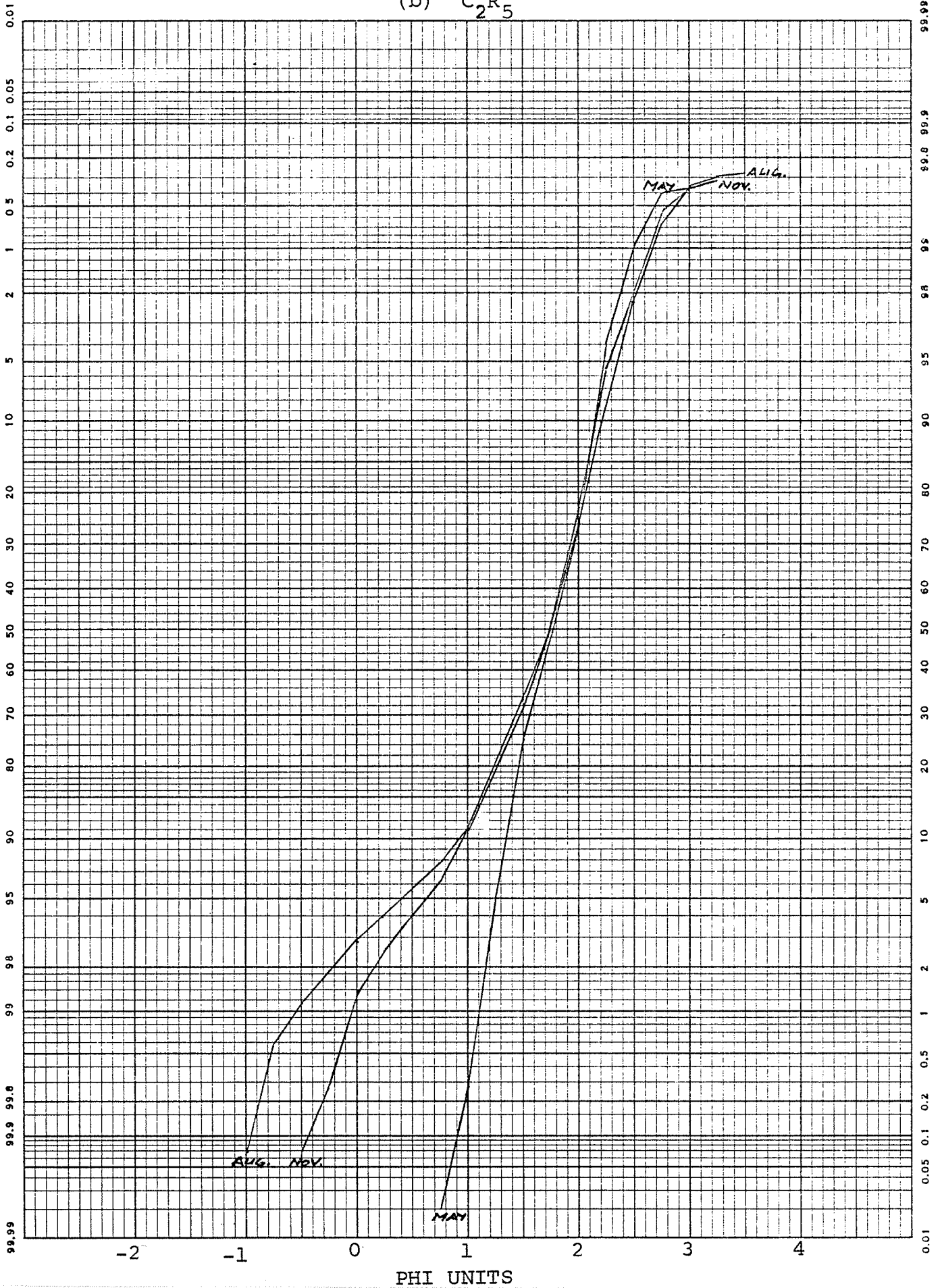


FIGURE 14 CONTINUED
(c) C_3R_1

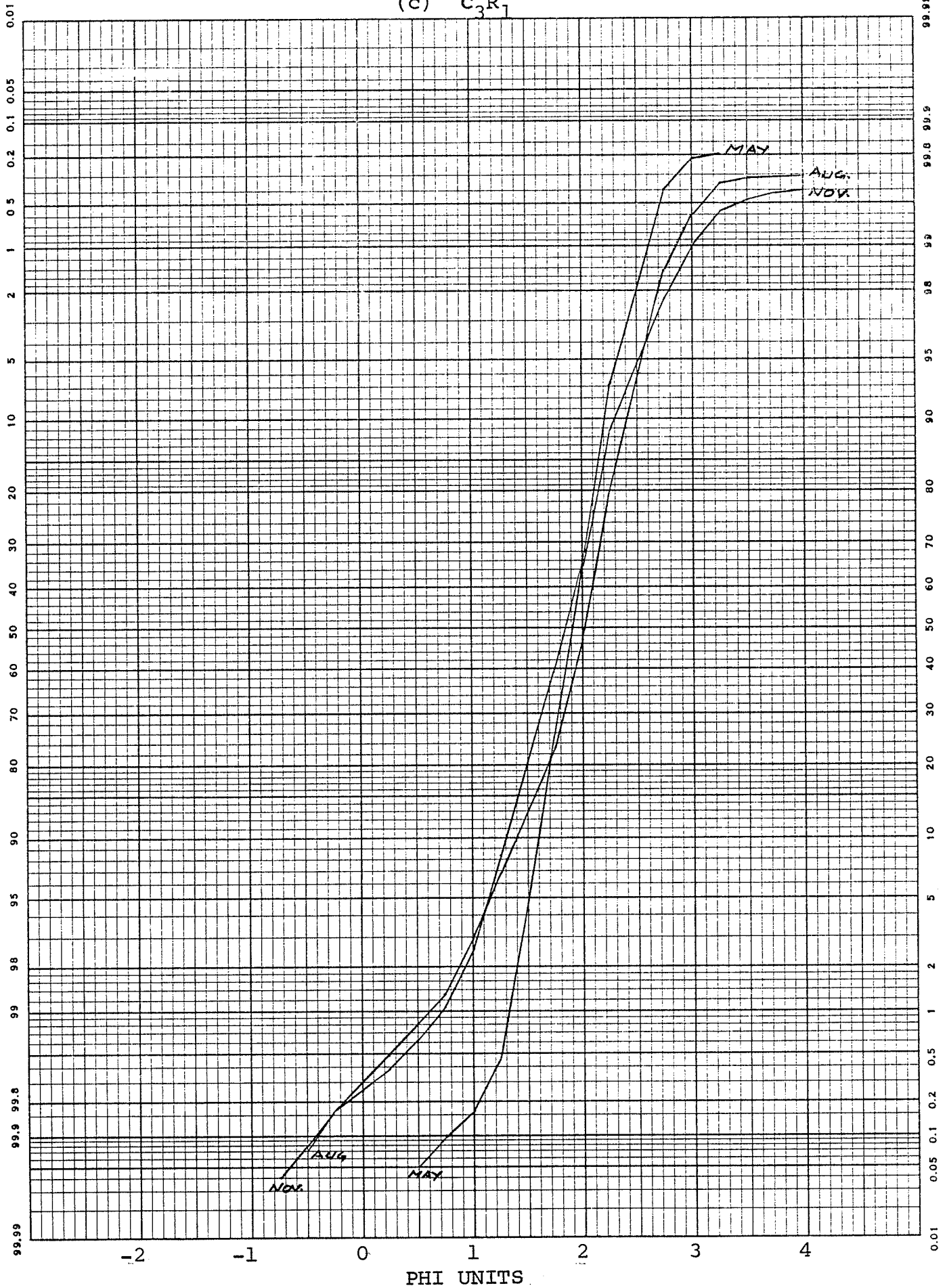


FIGURE 14 CONTINUED

(d) C_5R_2

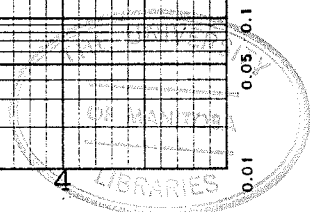
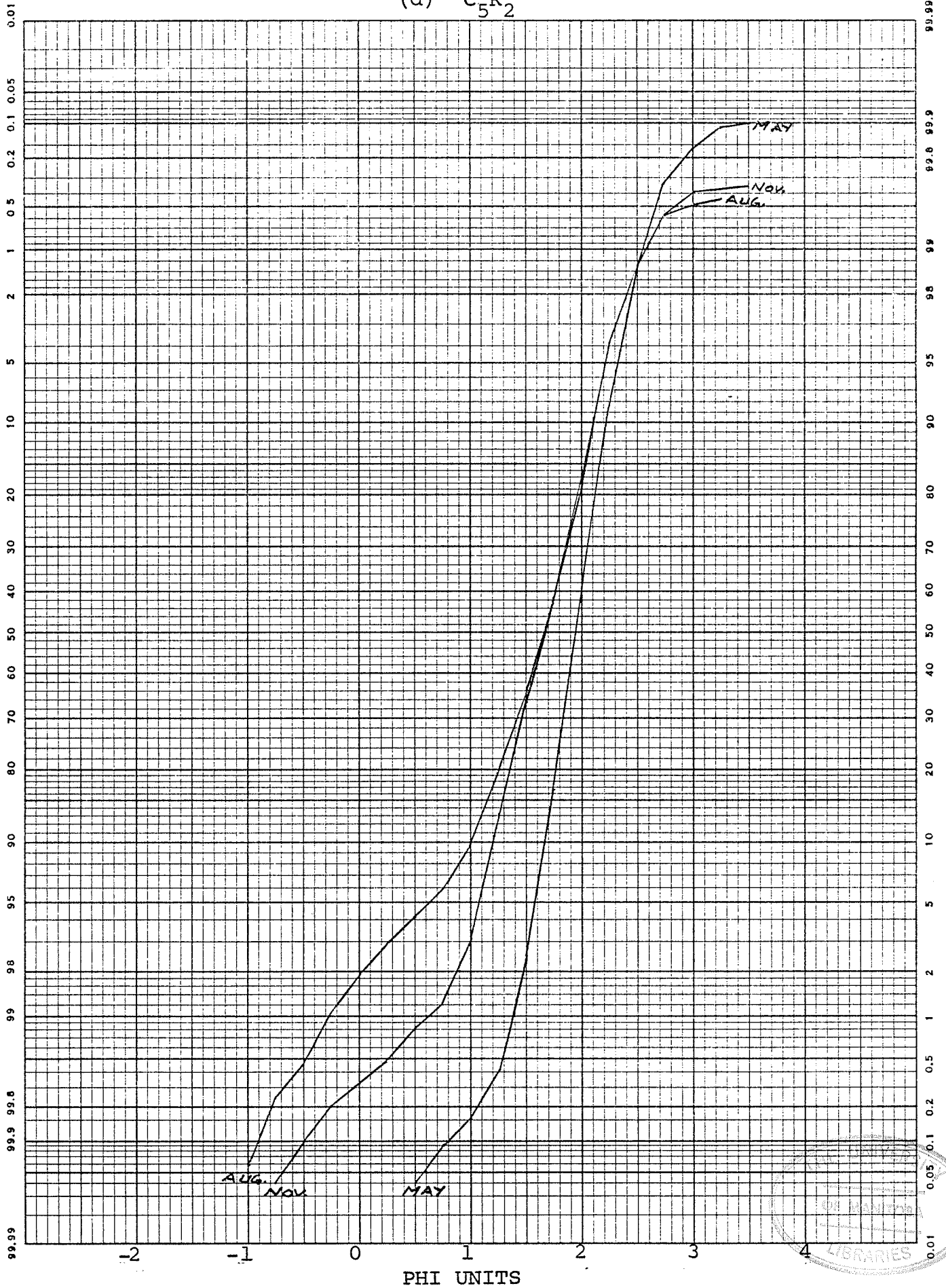


FIGURE 14 CONTINUED
(e) C_6R_4

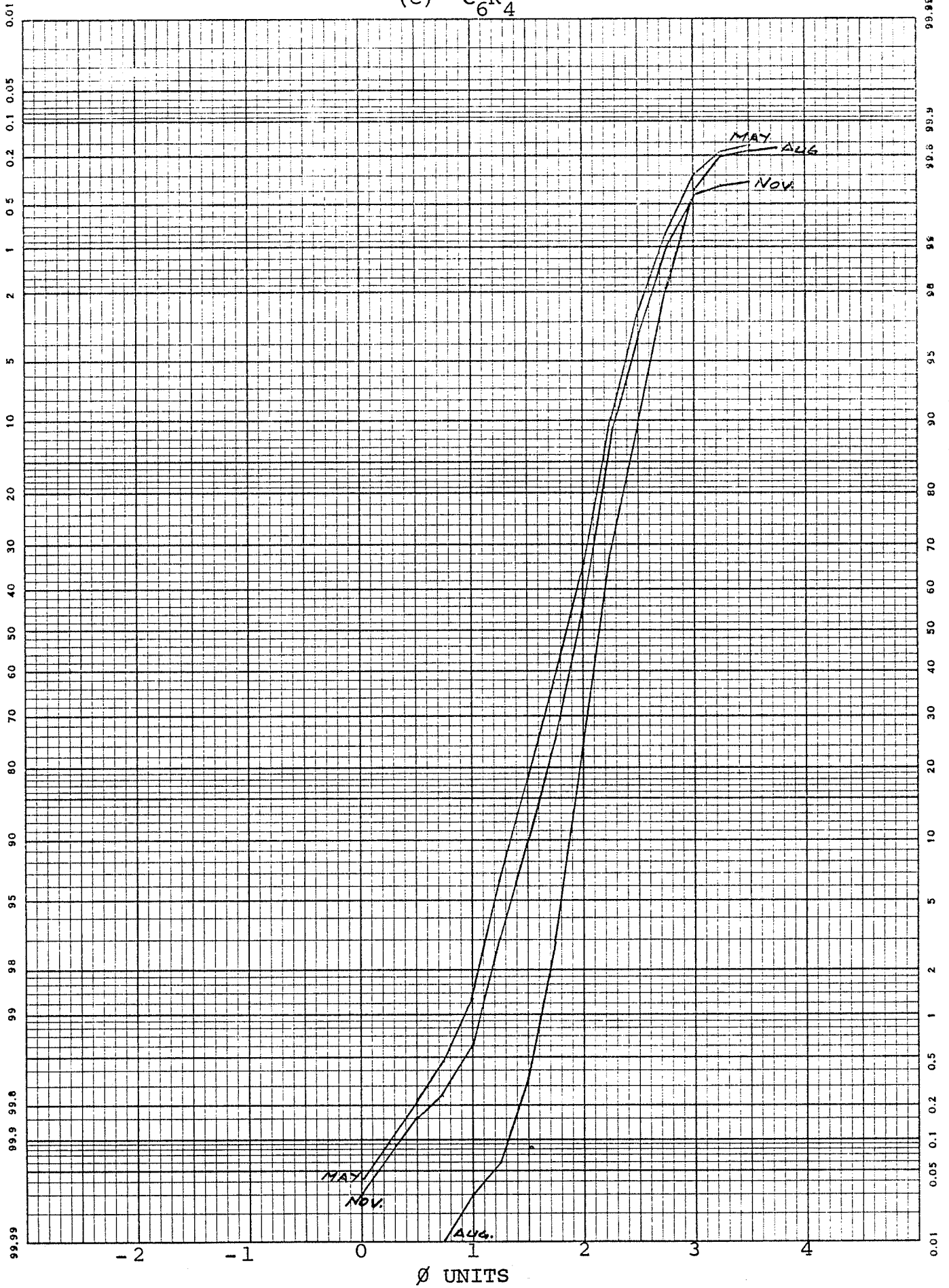


FIGURE 15
CUMULATIVE CURVES OF BEACH SANDS
SOUTH BEACH (a) C₁R₁

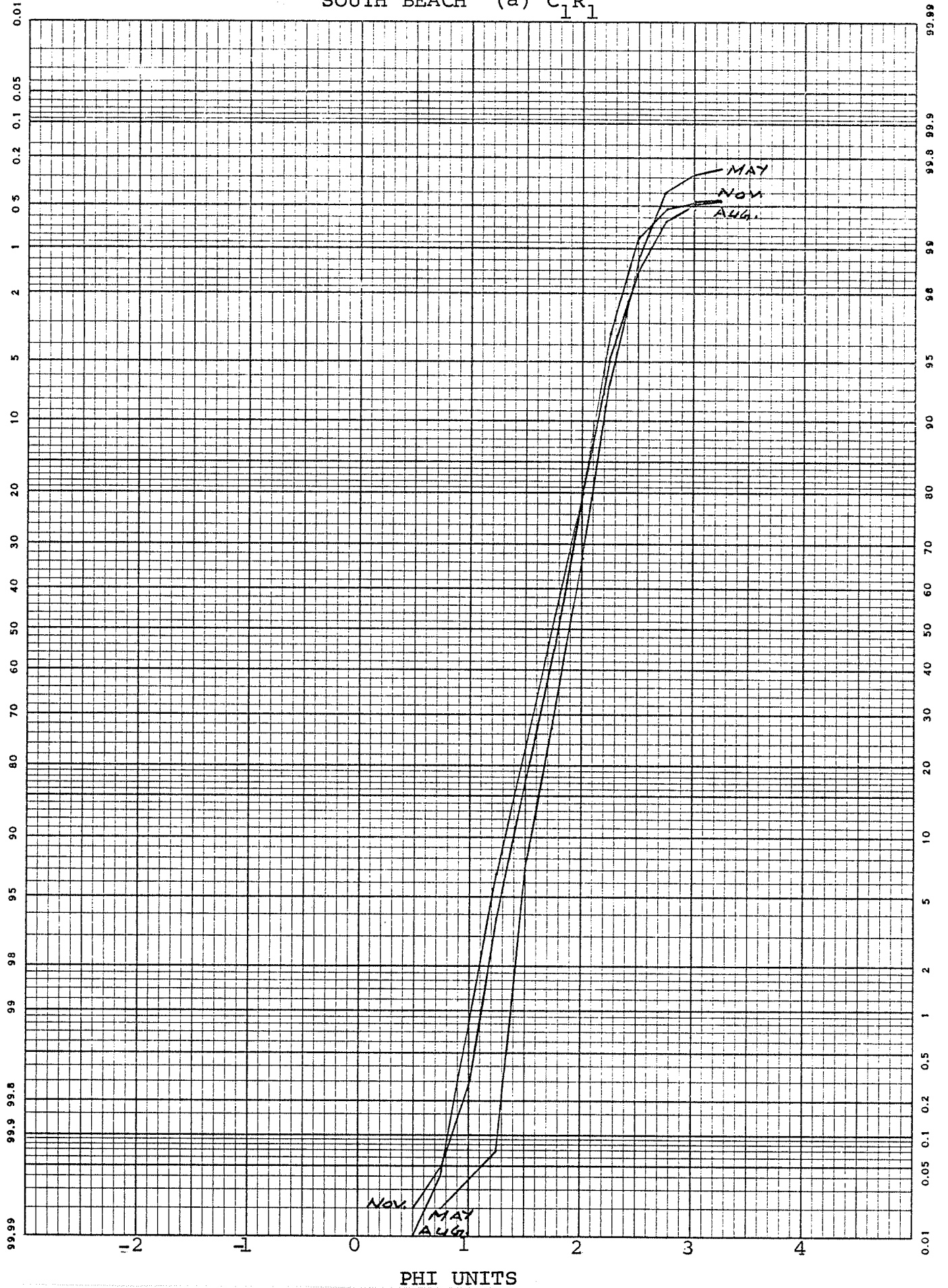


FIGURE 15 CONTINUED

(b) C_2R_5

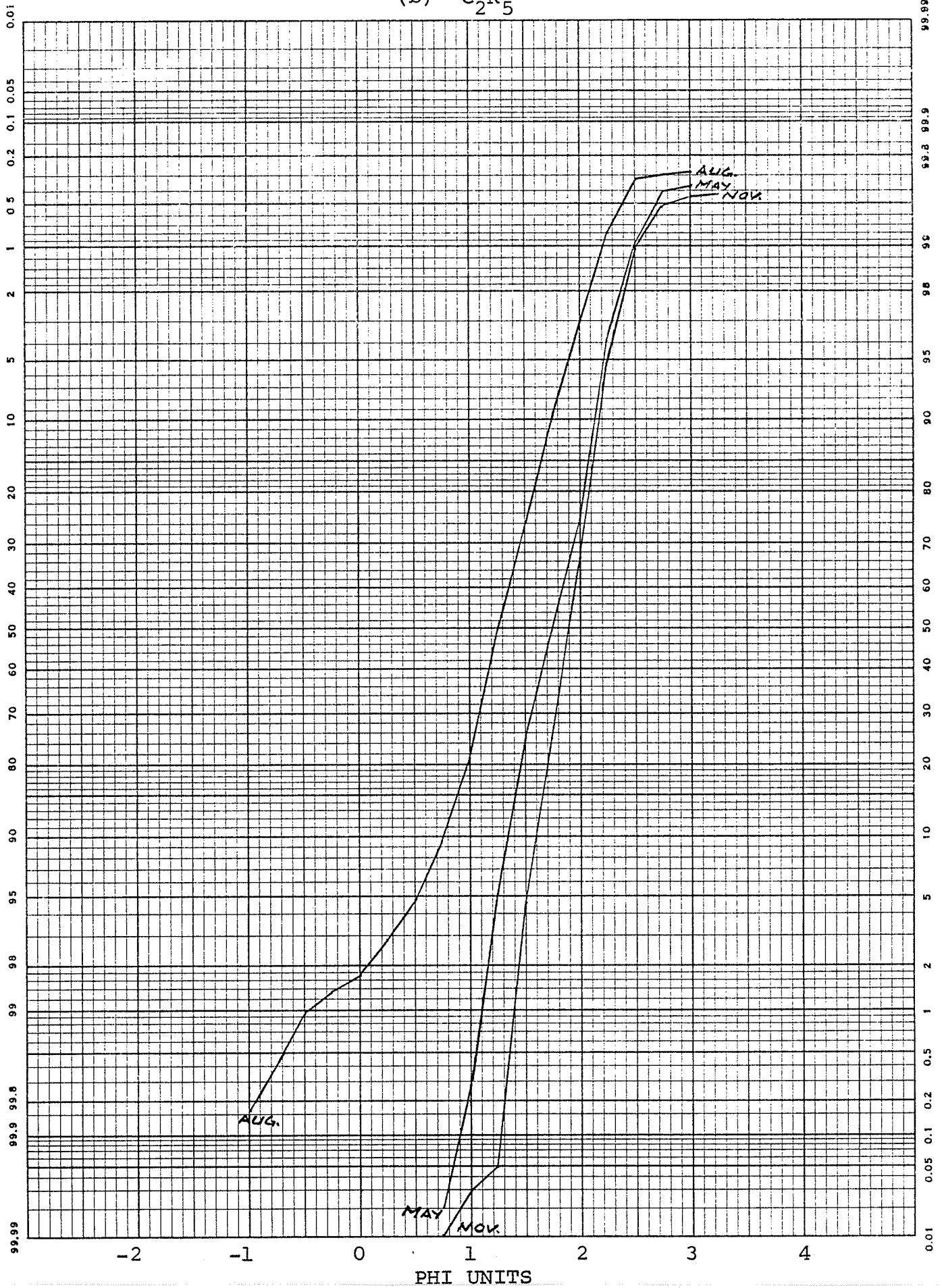


FIGURE 15 CONTINUED
(c) C_3R_3

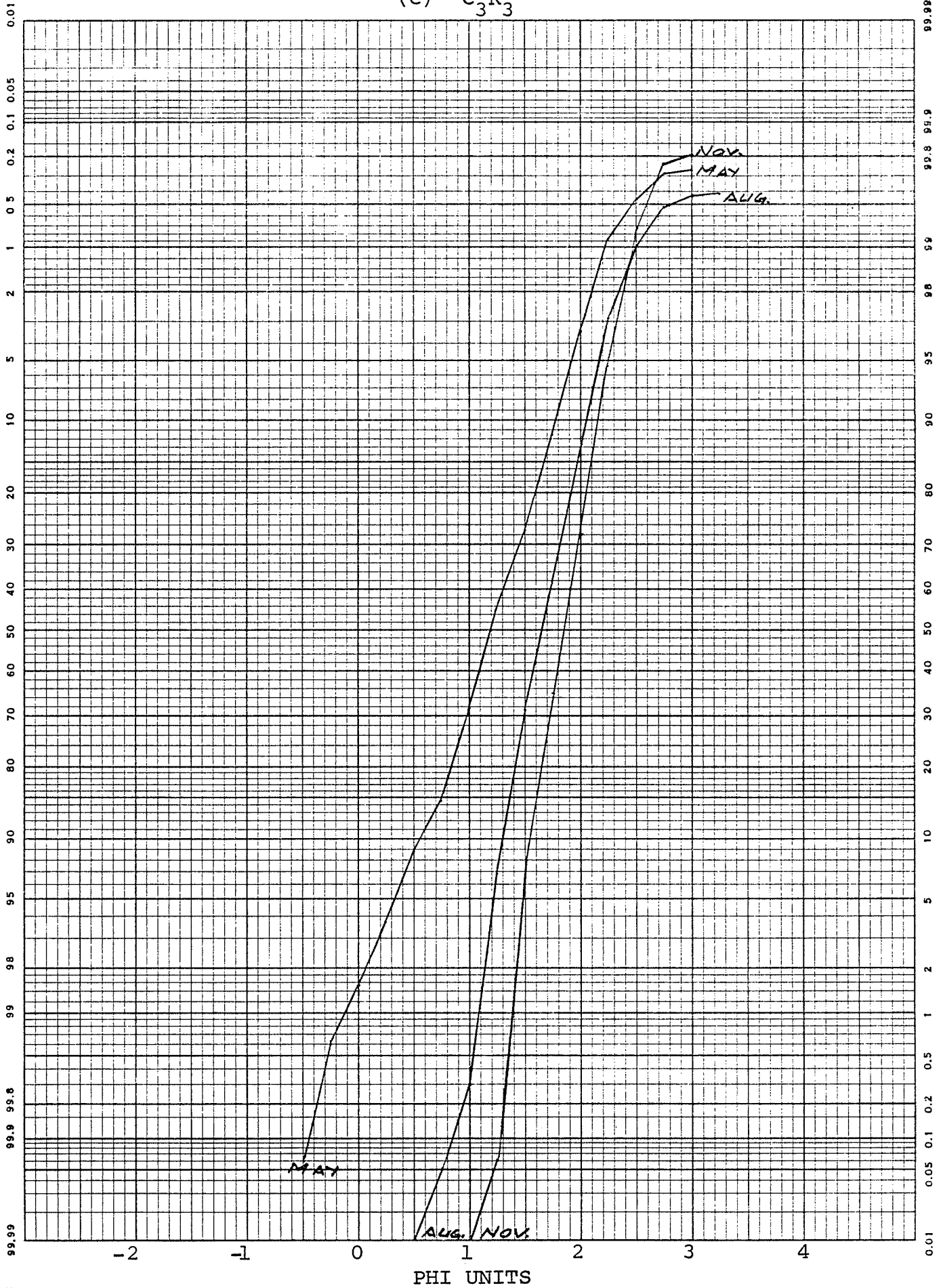


FIGURE 15 CONTINUED

(d) C_4R_4

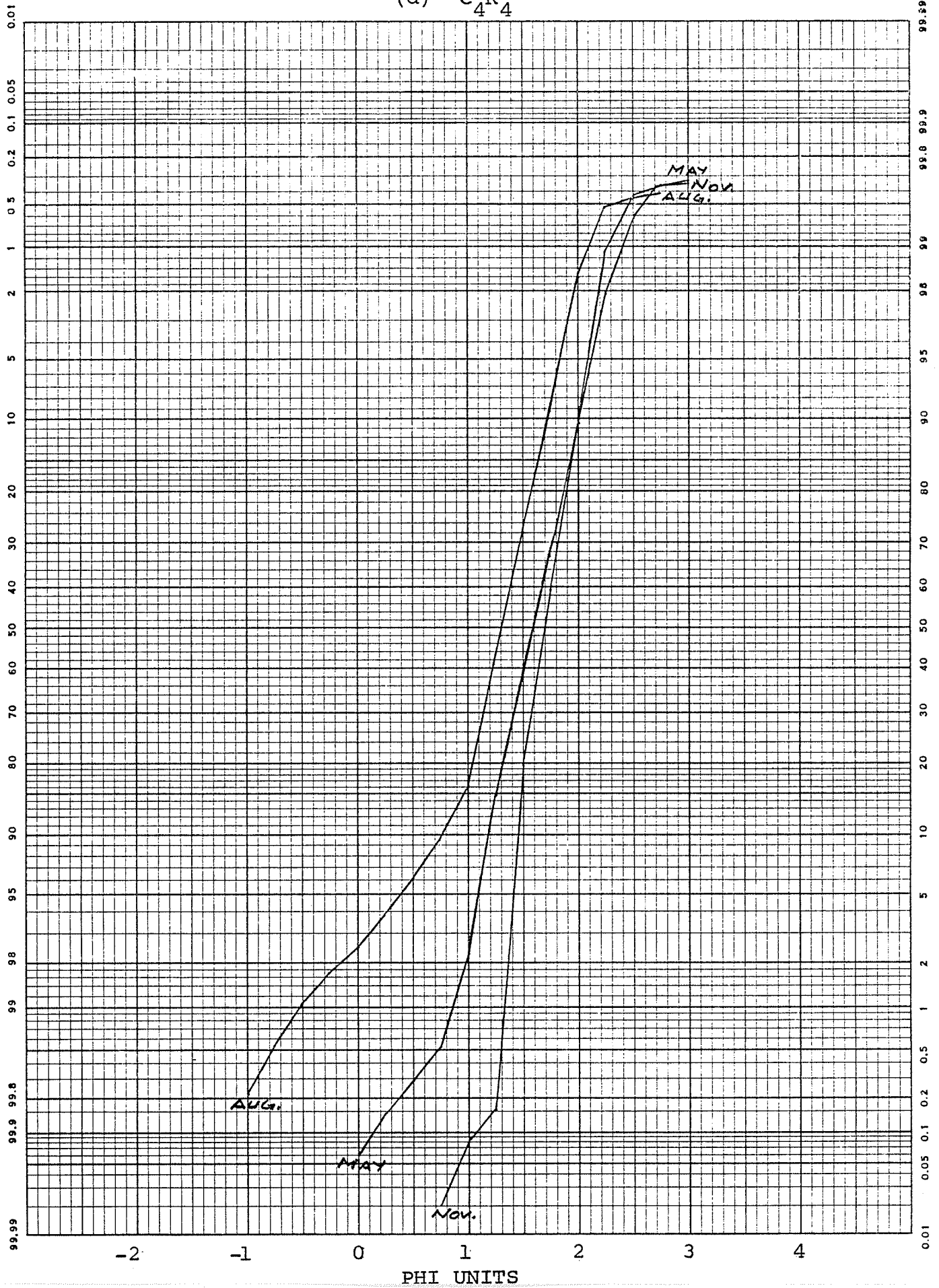
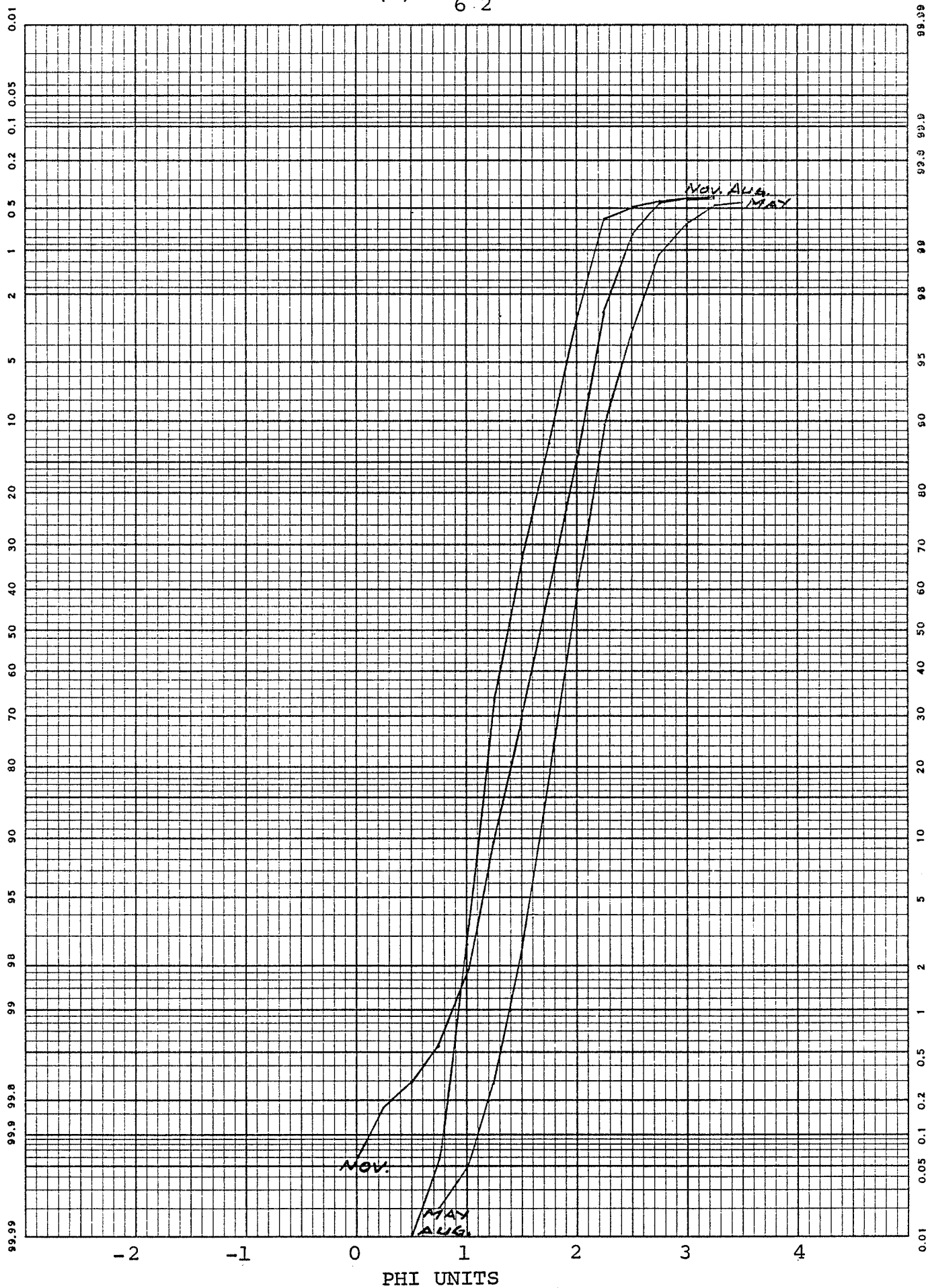


FIGURE 15 CONTINUED

(e) C_{6R_2}



Most of the quartz grains vary from sub-rounded to rounded, displaying frosted and pitted surfaces. Approximately 25 percent of the quartz grains have either angular or subangular faces, showing fresh fractures and vitreous lustre. Solohub and Klován (1970) state that:

"The frosted grains are identical to those comprising Ordovician Winnipeg Sandstone. The vitreous grains along with the feldspars and heavy minerals, are components of the glacial material."

(Solohub and Kovan 1970, p.86).

Figure 16 and 17 show the typically even distribution of grain size in a vertical and a horizontal plan (May 1974, C₂R₂).

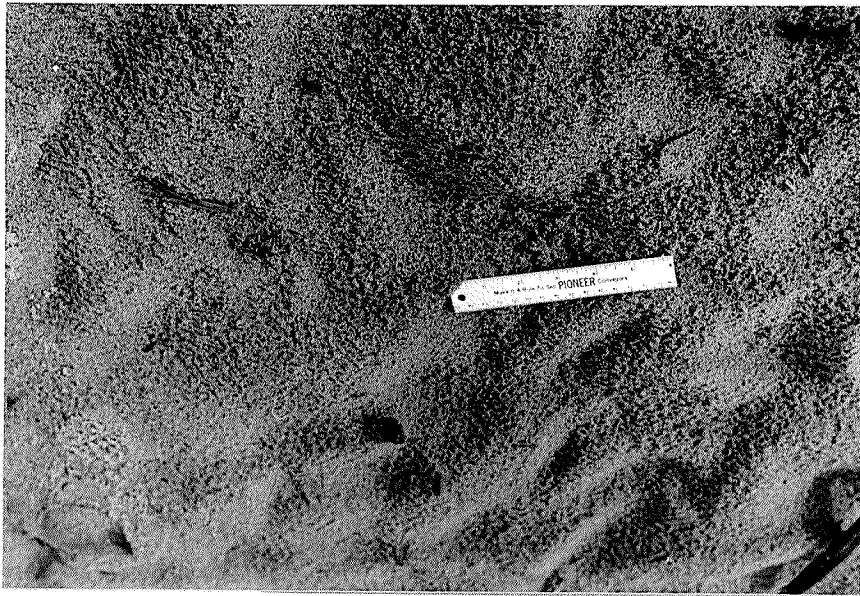


Figure 16

Grain Size - Horizontal Plane - May, 1974, C_2R_2

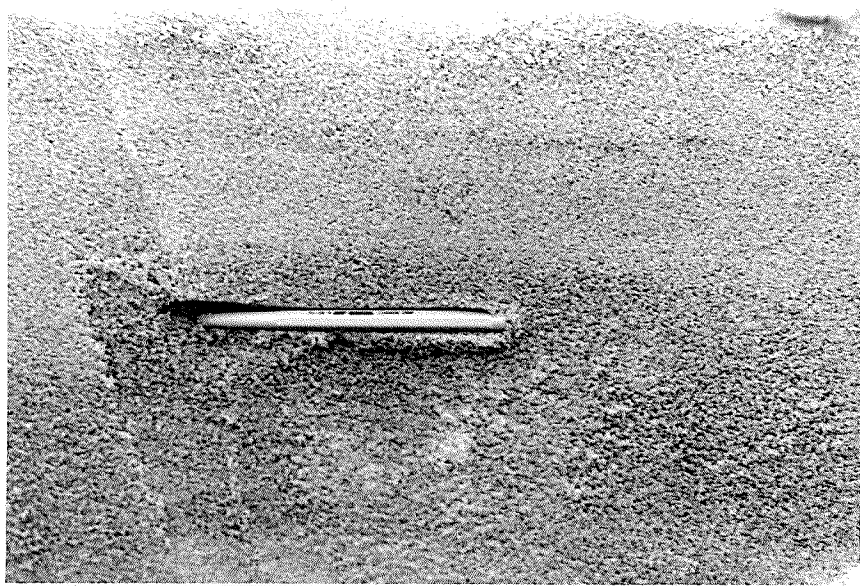


Figure 17

Grain Size - Vertical Plane - May, 1974 C_2R_2

Chapter IVINTERPRETATION AND DISCUSSION OF THE RESULTSIntroduction - Profiles

During the open water season, sand movement is continuously resulting in an endlessly changing beach. Changes in profiles and volumetric changes, accompanying the shifting of sand, are expressions of this phenomenon.

Shoreline erosion and recession have always been serious problems on Lake Winnipeg.

"An associated problem is the movement of sand away from beach areas, particularly during high water periods, causing a deterioration of these areas".

(Lake Winnipeg, Churchill and Nelson Rivers Study Board, 1971-75 p. 5-12)

It is therefore possible to postulate that changes in lake level, should cause a corresponding shift in beach profile. A rise in lake level should be accompanied by an upward and landward movement of the beach profile to maintain a constant position relative to the water level. Each profile can now be considered a representative part of a particular segment of a beach.

According to the Study Board (1971-75) the shore lines of the southern basin of Lake Winnipeg were eroding from one to two feet per year, the extreme values varying from

zero to twenty-five feet per year.

Seasonal Changes in Morphology

North Beach

The period of study covered three distinct parts of the year. The first survey was taken on May 29, 1974 soon after break-up, the second during the mid-season on August 29, 1974, and the third on November 29, 1974 before freeze-up.

In May, the beach showed the effects of the recent ice and snow cover, displaying overlying debris. Tsang (1973) mentions that ice piling along shore lines of Lake Simcoe is often a spectacular phenomenon, causing damages to shoreline properties. Figure 18 and Figure 19 may be used for illustration as they show debris from ice action in May and a relatively clean beach in November.

At the time of the May 29, 1974 survey, lake level was at 717.33 feet. All columns, except #6 of North Beach, delineate profiles having a well developed berm. Column six was sheltered from the action of water because of a higher elevation and a continuous belt of shrubs and bushes along the edge of the water. The shoreline was crescent shaped (Figure 20). Columns one and two show erosional aspects due to wave action of the lagoon which was about



Figure 18

Accumulation of Debris (North Beach) During
May Survey



Figure 19

Area Clean of Debris (North Beach) During
November Survey



Figure 20

Berm and Crescent Shaped Shoreline of

North Beach (May 29, 1974).

Vegetation Cover at Extreme End of Area

twenty feet away from the sample area at that time. The first portion of the beach, starting from row five leading up to the berm exhibited a slope of approximately three degrees. The backshore area of the beach was almost flat or had a gentle slope measuring up to $1\frac{1}{2}$ degrees.

The results of the survey in August indicate that the beach had undergone a period of flattening and lowering of the profiles. The lake level rose from 717.33 feet in May 29 to 717.62 feet on August 29, 1974. During this time segment, erosion appears to be the dominant feature. Most of the profiles were lowered by approximately six inches and all of the early berm had been destroyed. Some deposition took place as well, however, especially in row five for columns one to four. Columns five and six do not display any deposition, only erosion throughout their respective profiles. The beginning of a new berm and an associated landward movement of the beach had taken place (Figure 21). One reason for the berm not being fully developed may have been the high winds which had occurred on August 25 and 26, 1974 (Table 1b). A large portion of the sample area was inundated by the lake except column six, which was not covered by water at the time of the survey (Figure 22). The total erosion in volume of sand during the period from



Figure 21

Landward Movement of North Beach (August 29, 1974).

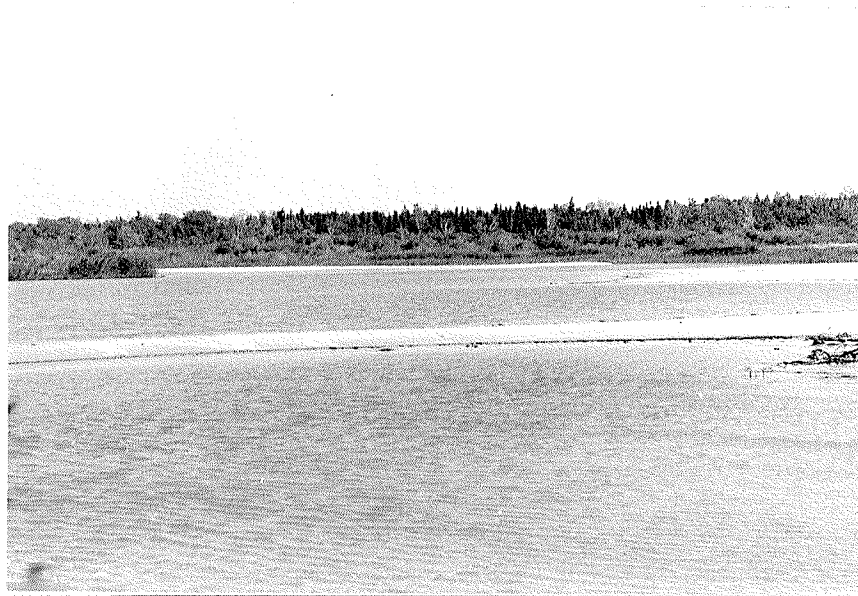


Figure 22

Column 6 of North Beach Not Covered by Water
In Spite of High Water Level (August 29, 1974).

the May to the August survey amounted to 168.38 cubic yards.

The last survey on November 29, by which time the lake level had fallen from 717.62 to 715.70 feet, showed a definite period of erosion during which the whole beach was lowered extensively. In some areas this amounted to a difference of one foot in elevation, corresponding to a drop in the lake level and indicating a definite lakeward migration of the foreshore. The shoreline exhibited several small baylets and the beginning of a definite berm (Figures 23 and 24). The change in volume of sand from August to November amounted to a decrease of 809.18 cubic yards, giving a total net change of 977.56 cubic yards for the entire season.

South Beach

The shore line along the South Beach did not at any time exhibit any formation of baylets or crescents. Throughout the season this beach displayed a relatively straight shore line. The reason for this occurrence can probably be traced to the absence of vegetation in the immediate sample area (Figures 25 and 26).

In May the area was covered with some debris. However, the amount was not as great as in the more sheltered North Beach. The berm in the foreshore area was not as pronounced as in the North Beach (Figure 27); the reason for this might



Figure 23

Beginning of Gentle Slope and Berm
North Beach. (November 29, 1974).



Figure 24

Baylets, North Beach. (November 29, 1974).



Figure 25

South Beach, (May 29, 1974).



Figure 26

South Beach, (November 29, 1974).



Figure 27

Debris, and Berm of South Beach (May 29, 1974)

be the comparatively steeper slope of the backshore area. The foreshore slope varied from 1° to $1\frac{1}{2}^{\circ}$ whereas the backshore slope had a variation of between 5° and 6° with column 5 having the steepest slope.

The August survey indicates that during the period between May and August heavy erosion had taken place. The area most affected was between rows one, two and three, another display of landward extension of the foreshore area. The entire area between rows two and five became the foreshore and exhibited a gentle slope, not unlike the foreshore slope during the May survey. The backshore area (row one and two) maintained approximately the same steep slope as before. The South Beach berm during this survey was even less developed than the North Beach during the same period. The decrease in volume during this period was 569.38 cubic yards.

The survey in November shows a similar extensive erosion as in the North Beach area. The shore line however remained relatively straight, and instead of baylets this area had the beginning of a small off-shore bar. The change in volume was 345.83 cubic yards, amounting to a total loss of 915.21 cubic yards for the season.

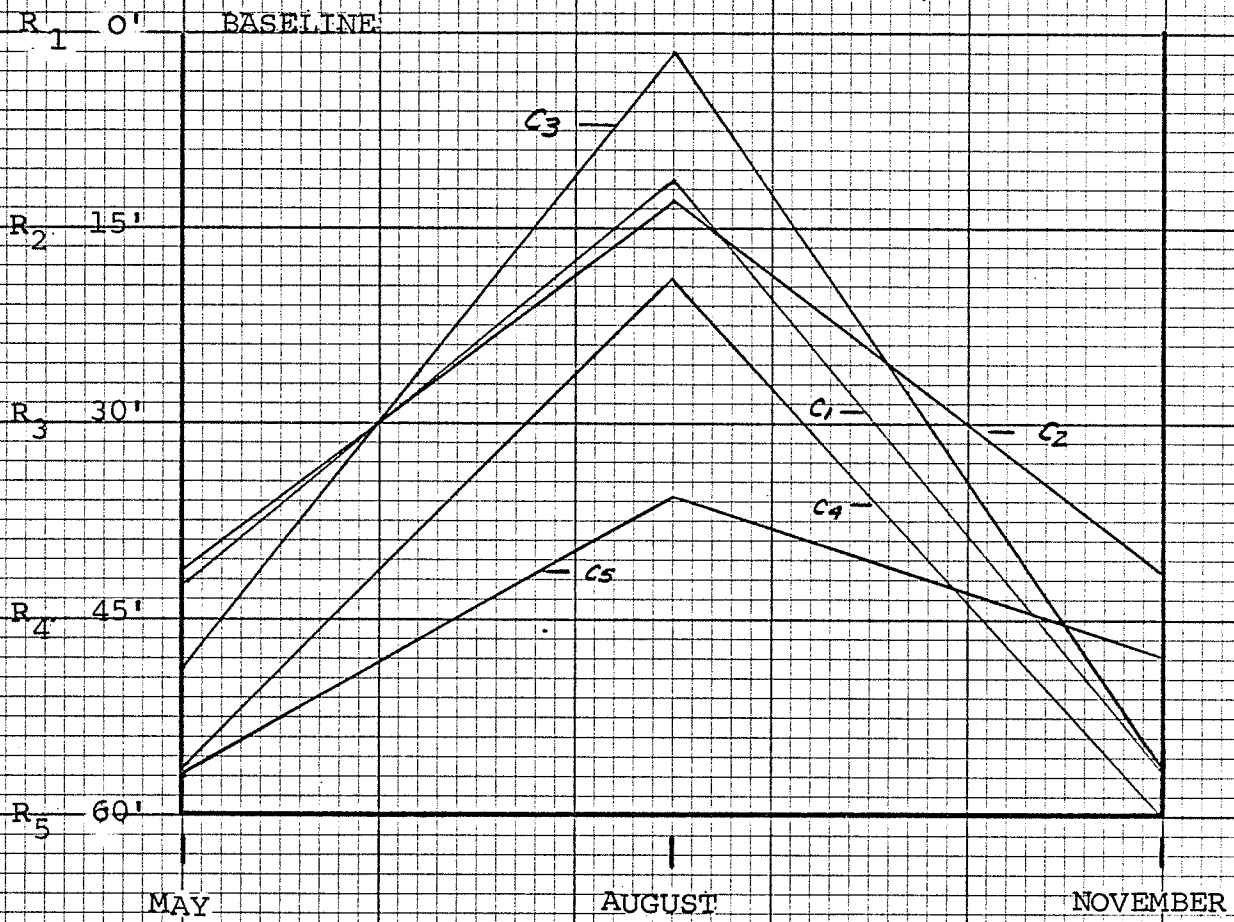
Both beaches illustrate the migration of their respective segments, either landward with high water level, or lakeward with lower lake levels. On inspecting the width of the beaches at various times (Figures 28 and 29), it can be clearly seen that the North Beach has a more erratic displacement whereas the South Beach is more geometrical or even. As both areas are subjected to the same rise in water level, the difference in morphology must be due to the original difference in topography. The South beach had a comparatively steeper slope, propagating a similar retrogradation and progradation of the area; not allowing any changes along the shore line. The shore line remained relatively straight throughout the season. The North beach, with its flatter topography, except column six which had a higher elevation, experienced more extensive flooding and the smoothing and flattening of the area was more pronounced. The shore line in November showed formation of various baylets throughout its length and the beginning of a crescent in the area of column five and six.

Introduction - Grain Size Parameters

Various measures are used to describe grain size distribution of sediment. Textural characteristics can be

FIGURE 28

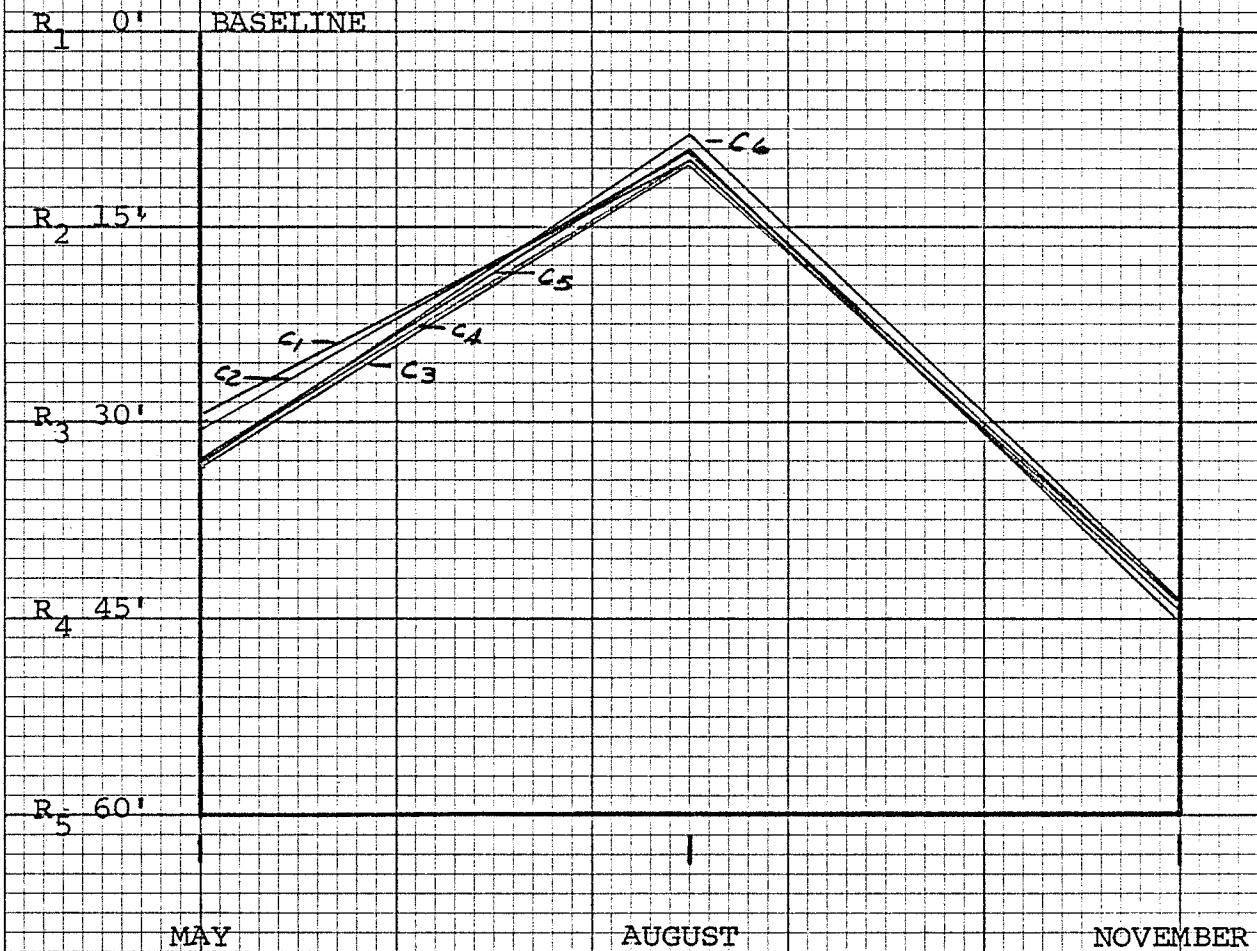
SUPERIMPOSED FLUCTUATIONS OF BEACH
 WIDTH WITH INCREASE IN WATER LEVEL
 ALONG COLUMNS - NORTH BEACH FOR MAY,
 AUGUST AND NOVEMBER



NOTE: C₆ WAS NOT INUNDATED BY WATER
 DURING ANY SURVEY.

FIGURE 29

SUPERIMPOSED FLUCTUATIONS OF BEACH
WIDTH WITH INCREASE IN WATER LEVEL
ALONG COLUMNS - SOUTH BEACH FOR MAY,
AUGUST AND NOVEMBER



NOTE: VERY MINOR DEVIATION OF STRAIGHTNESS
OF SHORELINE THROUGHOUT THE SEASON,
SOUTH BEACH WAS MORE EXPOSED TO LONG
SHORE DRIFT, HAD NO VEGETATION COVER
WITHIN SAMPLE AREA

identified through the use of descriptive statistics. In this study, Folk's four statistical parameters are adopted: the mean, standard deviation, skewness and kurtosis. The mean and standard deviation are in phi units, skewness and kurtosis are dimensionless statistics describing the symmetrical characteristics of a frequency curve. More complete information on different moment measures is detailed in Appendix A.

Krumbein (1938) gave an extensive treatment of specific concepts such as the log-normality of grain size distribution. Doeglas (1946) showed that grain size distribution followed an arithmetic probability. His analyses yielded an empirical classification of curve shapes which he related to specific environments.

Inman (1949) delineated the relationship between texture and process; defining three basic modes of transport, surface creep, saltation and suspension. In 1952, Inman further recommended the use of five parameters for statistical computation. These were: the mean diameter, standard deviation, kurtosis, and two measures of skewness. Folk and Ward (1957), Friedman (1961), and Visher (1969), have contributed significant studies using the same statistical parameters, except that only one measure of skewness was used. Solohub and Klovan (1970) stated that if it can be

assumed that the mean grain size is a measure of energy at the time of deposition, then the mean grain size delineates different environmental conditions.

Grain Size Parameters

North Beach

The mean size of all the samples taken in May ranges from 0.542 ϕ at C₂R₃, to 1.951 ϕ at C₄R₂; the difference amounting to 1.409 ϕ , indicating a fairly uniform grain size distribution for this area. Row three was an exception, C₂, C₃, C₅ and C₆ having readings under 1 ϕ delineating very coarse sand. However, the average mean size of row three (1.115 ϕ) is still within the medium sand range. The reason for this difference could be due to residual material from ice action of the previous winter. Furthermore, row three is located beyond the berm away from the waterline facilitating trapping of larger particles.

Samples from the August survey reflect a slightly larger difference in mean particle size, the range was between 1.387 ϕ at C₁R₂ and 2.56 ϕ at C₆R₄ amounting to a difference of 0.769 ϕ with all ϕ values greater than one. Most of the sample area was covered by water accounting for the absence of larger particle size.

The results of the November survey had very similar values to that of August. The difference was 0.791 ϕ

with a value of 1.286 ϕ at C₃R₂ to 2.077 ϕ at C₆R₃. All sample points along C₁ to C₅ except row five were above water level. Column six was above water throughout the season.

Sahu (1964) states that the mean particle size reflects the average kinetic energy of a process. Hjulstrom (1939) assumes a direct relationship between grains coarser than the range of 0.3 to 0.6 millimeters and the energy needed to move those particles. Materials finer than the range of 0.3 to 0.6 needed an increase of threshold velocity to initiate particle movement with decreasing particle size.

The study area illustrates a similar effect, with the rows parallel to the lakeshore showing the trend of advance and retreat of the lake level, and therefore the rise and fall of kinetic energy, by exhibiting a gradual decrease of fines away from the water line. Samples from August and November indicate the uniformity of the processes involved, and are lacking the larger sand particles which were found in May, satisfying the different geologic processes which must have acted on this area during the winter months.

The standard deviation, or sorting, depicts the variation of the kinetic energy about the mean energy level (Sahu, 1964): larger sorting coefficients reflect greater

variations in energy level. A large sorting value reflects a poorly sorted sediment, whereas a small range of sediment size denotes well sorted material.

The values for the standard deviations (except for one isolated sample unit) exhibited a small range for the whole area over the entire season. The low and high values for May, August, and November respectively are, 0.223 ϕ at C_5R_2 , 0.262 ϕ at C_4R_3 , 0.243 ϕ at C_5R_4 , 0.909 ϕ at C_3R_3 , 0.608 ϕ at C_2R_3 , and 0.559 ϕ at C_2R_1 . C_3R_3 is the only unit which is moderately sorted. All others range from very well sorted to moderately well sorted. It seems that beach sands tend to be raked back and forth by the continual motion of wave swash and this continuous reworking results in good sorting.

Skewness is a measure of the asymmetry of a sediment frequency distribution, it relates the position of the mean relative to the median (Dubois, 1972). In a positively or fine skewed sediment, the mean is offset from the median towards the fine tail, whereas in a negatively or coarse skewed sediment, the mean is offset from the median towards the coarse tail of the distribution. Friedman (1961) postulates the skewness for dune sand is generally positive and that medium sands which are subjected to high

wave energies will be negatively skewed because the fine clasts will have been winnowed out. Therefore the mean of a sediment distribution is offset from the median toward the coarse fraction generating a negatively skewed value (Dubois, 1972).

From the first survey in May, twenty three out of thirty values show negative skewness, the maximum range being between 0.171 to - 0.287. In August, twenty six values were negatively skewed with a maximum range of 0.186 to - 0.328. The November survey also had only four values exhibiting positive skewness. The range was between 0.112 to - 0.3117. August and November survey results tend slightly more to a beach environment than the May values. The positively skewed values may be due to some modification by the wind; some however were found below lake level. Stephenson (1970) suggests that skewness seems to be related to environmental conditions and energy. He claims that negative skewness can be associated with areas of erosion while positive skewness is indicative of deposition. Although all of the region shows aspects of erosion, some deposition may have occurred during any segment of time accounting for the positively skewed values.

Dubois (1972) states that the specific geologic meaning of kurtosis has not been fully determined. According

to Folk (1957) kurtosis as used by sedimentationists measures the ratio of the sorting in the extremes of the distribution compared with the sorting in the central part. If the central portion is better sorted than the tails, then the curve is leptokurtic or excessively peaked. If the tails are better sorted then the curve is deficiently peaked or platykurtic. If the curve is normal, or mesokurtic, the K_G value is from 0.90 to 1.11 (Folk, 1968). On working out average values along rows, it can be seen that all samples fall within the mesokurtic category - again pointing to the uniformity of the environmental material and environmental processes.

South Beach

The mean size of the samples collected during the May survey had no values under 1.0 ϕ as was the case at the North beach. On examining the average values of all the rows, it is also row three which has the lowest value. The average mean value of the North beach of row three was 1.115 ϕ and the South beach 1.366 ϕ . The reason for slightly larger particles in both areas may be that they delineate particular processes occurring during the dormant season. Wave energy may be strong enough to carry larger particles to this limit and deposition will take place, whereas the swash is lacking the energy required for transportation of

larger particles.

The August mean averages indicate larger particles for all rows, whereas the November mean averages show an increase of smaller particle size. This phenomenon may be due to the fact that the shoreline of the South beach was relatively straight, lacking any vegetation cover and therefore more exposed to environmental processes. With a higher water level during August a higher kinetic energy would substantiate movement of the larger particle size; during November a lower lake level would constitute a lower energy level and subsequently a smaller particle size.

During the November survey, the South beach does not exhibit as clearly the gradual decrease of fines away from the water line as does the North beach - especially row four which has a higher average mean value than any other row during this time segment. On examining elevations and profiles, it can be seen that row four delineates the boundary of water environment and shows the beginning of a berm which could cause trapping of slightly larger particles.

The values for the standard deviation for the South beach again indicate the uniformity of the area; most of the values for the season range from very well sorted to well sorted, having σ values from 0.35 σ or less to 0.50 σ . Three exceptions were found during the May survey at C₃R₂

(0.5434), C_1R_4 (0.540 ϕ) and at C_1R_5 (0.626 ϕ). As well, three other values found in the August survey were in the moderately well sorted to moderately sorted category, being (0.797 ϕ) at C_1R_3 , (0.557 ϕ) at C_1R_5 and (0.617 ϕ) at C_5R_5 .

Ball (1972) states that scalloping may account for deviation of values in similar areas and this would account for the higher values in May; however, the larger material in August is hard to explain as all stations in questions were covered by water and therefore acted upon by the same processes.

The South beach had twelve samples in May, sixteen samples in August and five samples in November positively skewed, indicating the more exposed aspect of the beach. It is clearly not a dune environment, but rather a beach which has been slightly modified by wind action.

Similar average values for kurtosis are exhibited by the South beach as the North beach. Of fifteen readings only one is leptokurtic, the remainder are mesokurtic. The high values are from the August survey and are found in row five, which was covered by more than one foot of water which caused less disturbance of the material.

Visher (1969) found that 50-99% of beach material between 0.5 ϕ and 4.25 ϕ was a result of saltation activity. The range was determined by the Coarse Truncation and Fine

Truncation points on a cumulative frequency curve. 90% of the samples from the study area were found to be in the range from 1 σ to 3 σ showing similarities in the cumulative frequency curves (Figures 14 and 15), thus appearing to bear out Visher's findings.

Chapter VCONCLUSION

Coastal studies of marine beaches have shown that changes in morphology are manifestly associated with wind regimes. Increased storms during the winter season cause erosion of beaches. In summer when moderate winds prevail reconstruction of the beach takes place. Dubois (1972) further states that not all marine beaches undergo a seasonal change. Deviations from the cycle are functions of beach orientation relative to wind direction, and of the absence of major seasonal climatic variation.

Seasonal changes in morphology of limnic beaches respond to seasonal fluctuations in lake levels. In this study it was found that the rise and fall of lake level was indeed the dominant variable affecting the study area. Winds and waves must also bring about changes for they can be related directly to the elevation of lake level.

Moisture content and consequent saturation of the sand, which was not measured, should also be considered as a factor altering the beach profile. The rate of erosion, or lowering of the area due to heavy saturation of the material, must have been excessive, but can only be inferred from the amount of erosion over the whole season. Vegetation cover, slope, and the original topography of the

area will influence subsequent erosion or deposition.

Lake level for the entire season (1974) was markedly higher than the mean level of 713.22 feet (Figure 5b) denoting an atypical increase in water level. The Technical Report (1971-75) states that the average relationship between water level increase and loss of beach area was estimated to be a twenty percent loss of beach width for each one foot increase in water levels above the mean level.

Throughout the season South and North beaches underwent a heavy cycle of erosion. Deposition or build up of the beach area had not manifested itself clearly by November, 1974; therefore the seasonal cycle of dynamic equilibrium normally associated with limnic lakes had not been fully completed at this particular segment of time. As the lake level rose during the season, a marked landward movement of the beach was noted. While the lake level increased by 0.29 feet, the loss in width was approximately twenty to twenty-five feet. In November the reverse occurred. With the drop in lake level, a definite lakeward movement of the foreshore took place. The drop in lake level between August and November amounted to 1.13 feet, expanding the beach by approximately thirty feet.

"The effect of lake level on sand beaches is very complex. Neither the available data for Lake Winnipeg nor the present state of the art of analysis is adequate to offer more than some very general comments on the subject."

(Program for Regulation of Lake Winnipeg, 1972, p.11).

On comparing the grain-size parameters of the two beaches it can be deduced that the two areas may be viewed as a single beach unit delineating the uniform nature of source sediments. The foreshore bottom, the foreshore and backshore have similar parameters throughout the season with some isolated exceptions. The original physiography was slightly different from one region to the other, accounting for the sequential change in topography for the respective areas.

Plotting mean size against the distance from the base line emphasizes the uniformity of grain size (Figures 30 to 31), and on plotting standard deviation against mean size a definite clustering of samples is revealed, indicating the same environmental processes (Figures 32 to 33), (Wong, P.P., 1971). These results coincide with Dubois' findings.

FIGURE 30
 PLOT OF MEAN SIZE AGAINST DISTANCE FROM
 BASELINE. NORTH BEACH

• MAY
 + AUGUST
 x NOVEMBER

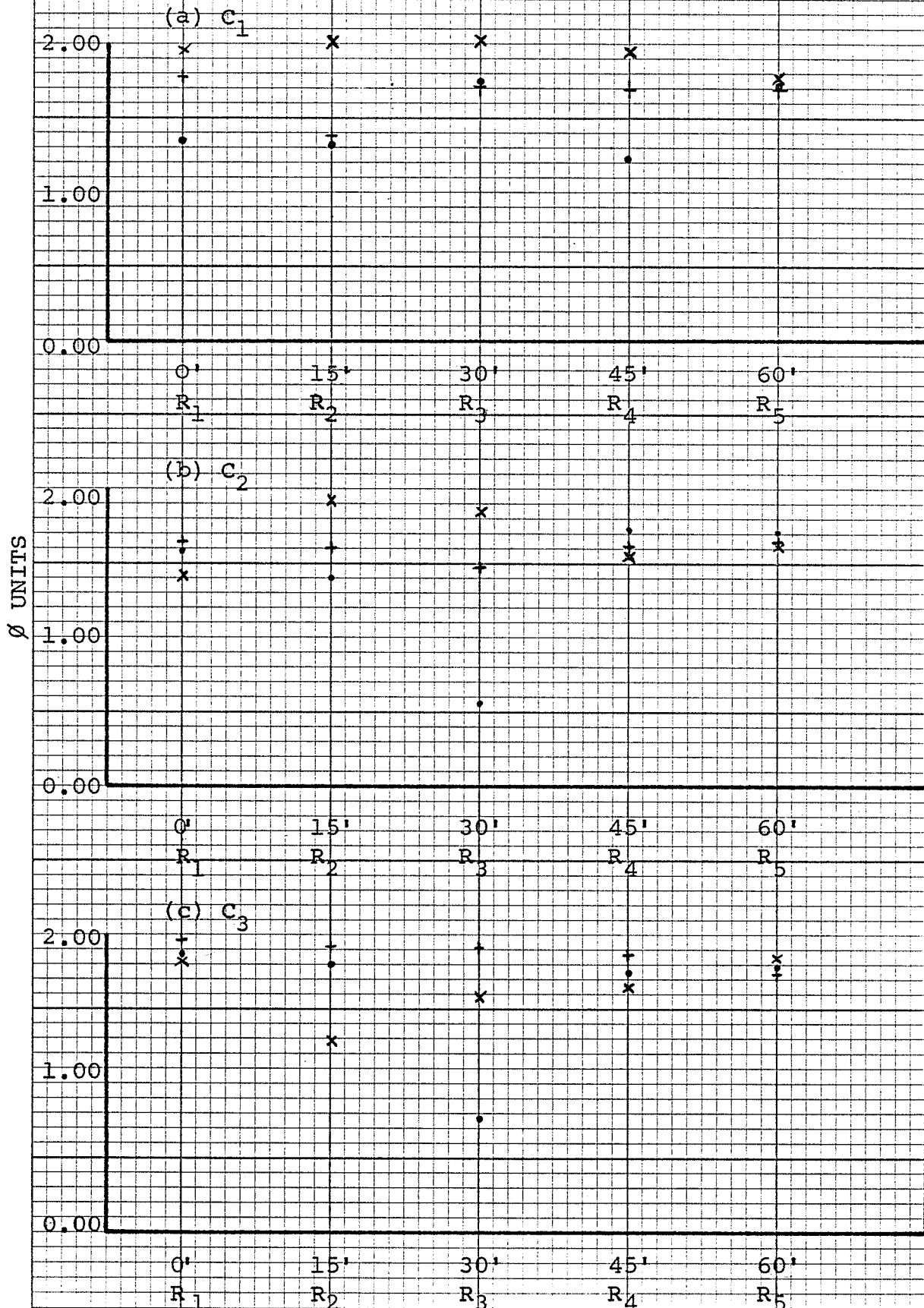


FIGURE 30 CONTINUED

Ø UNITS

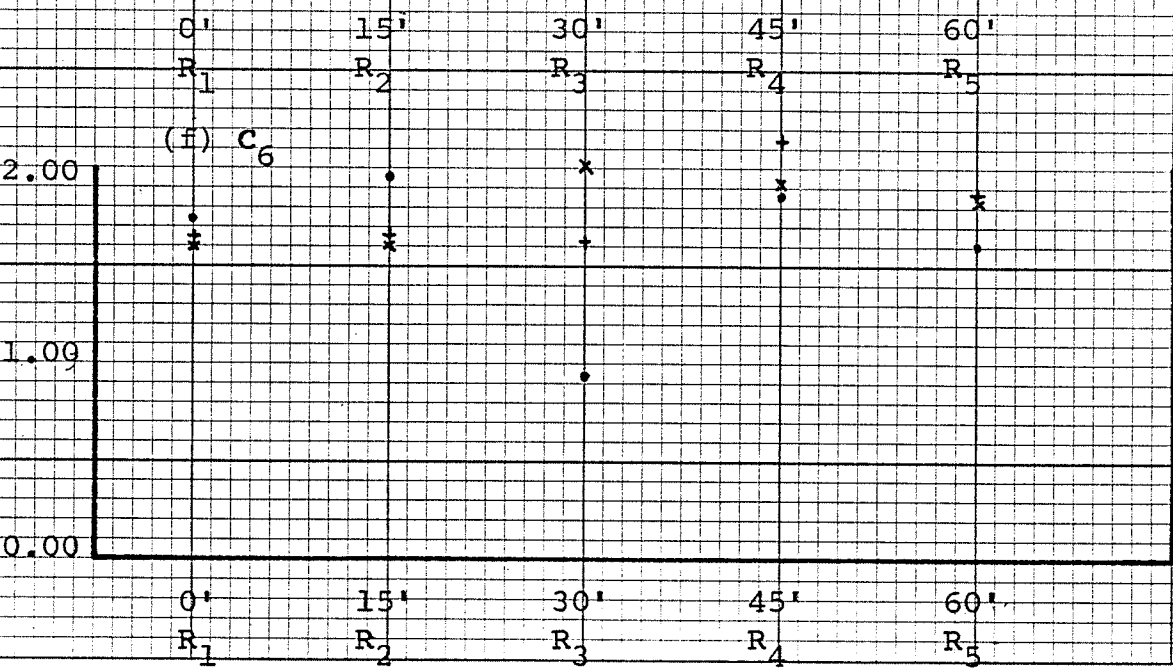
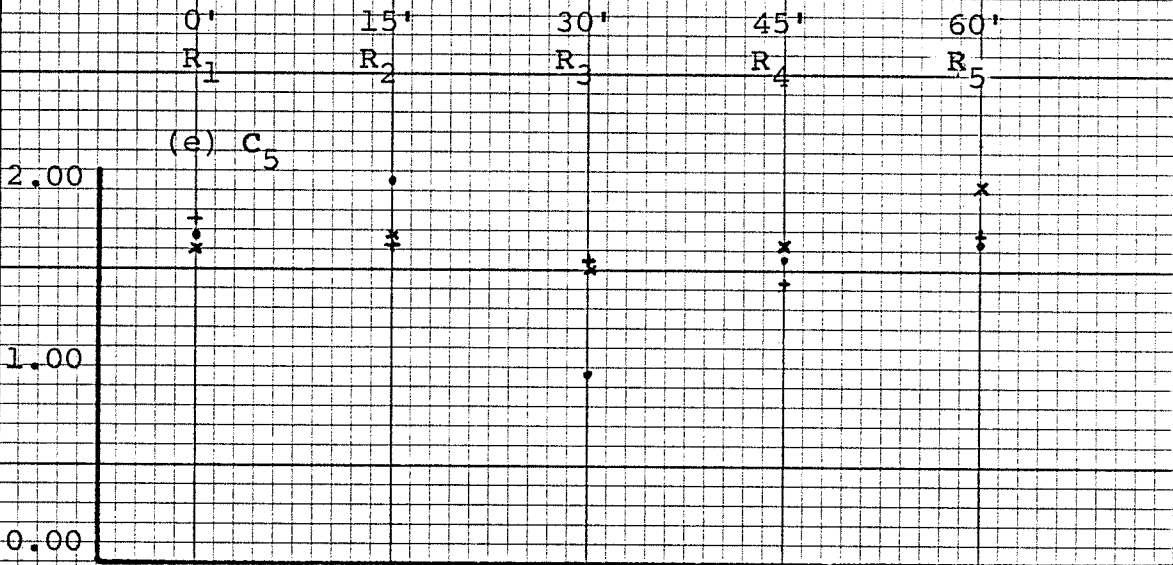
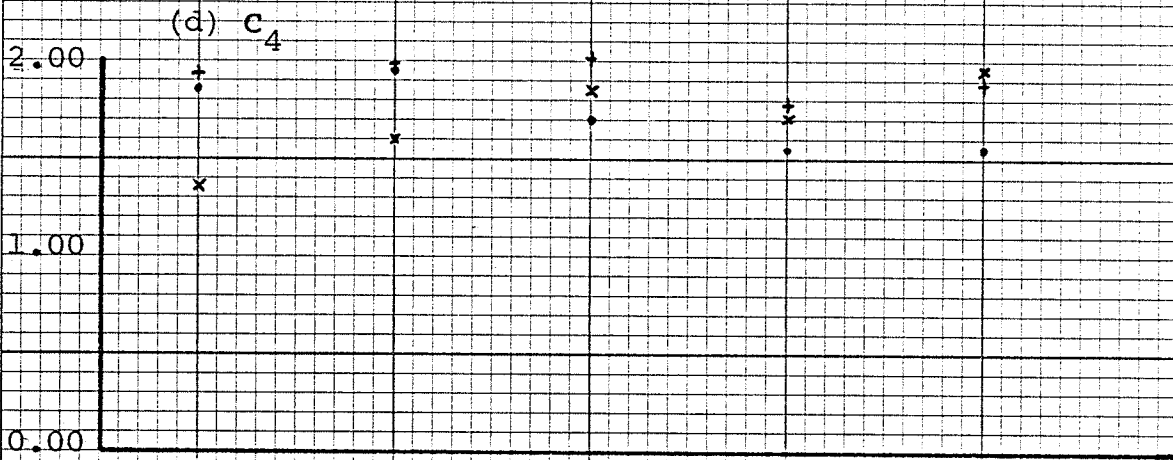


FIGURE 31
 PLOT OF MEAN SIZE AGAINST DISTANCE FROM
 BASELINE. SOUTH BEACH

• MAY
 + AUGUST
 x NOVEMBER

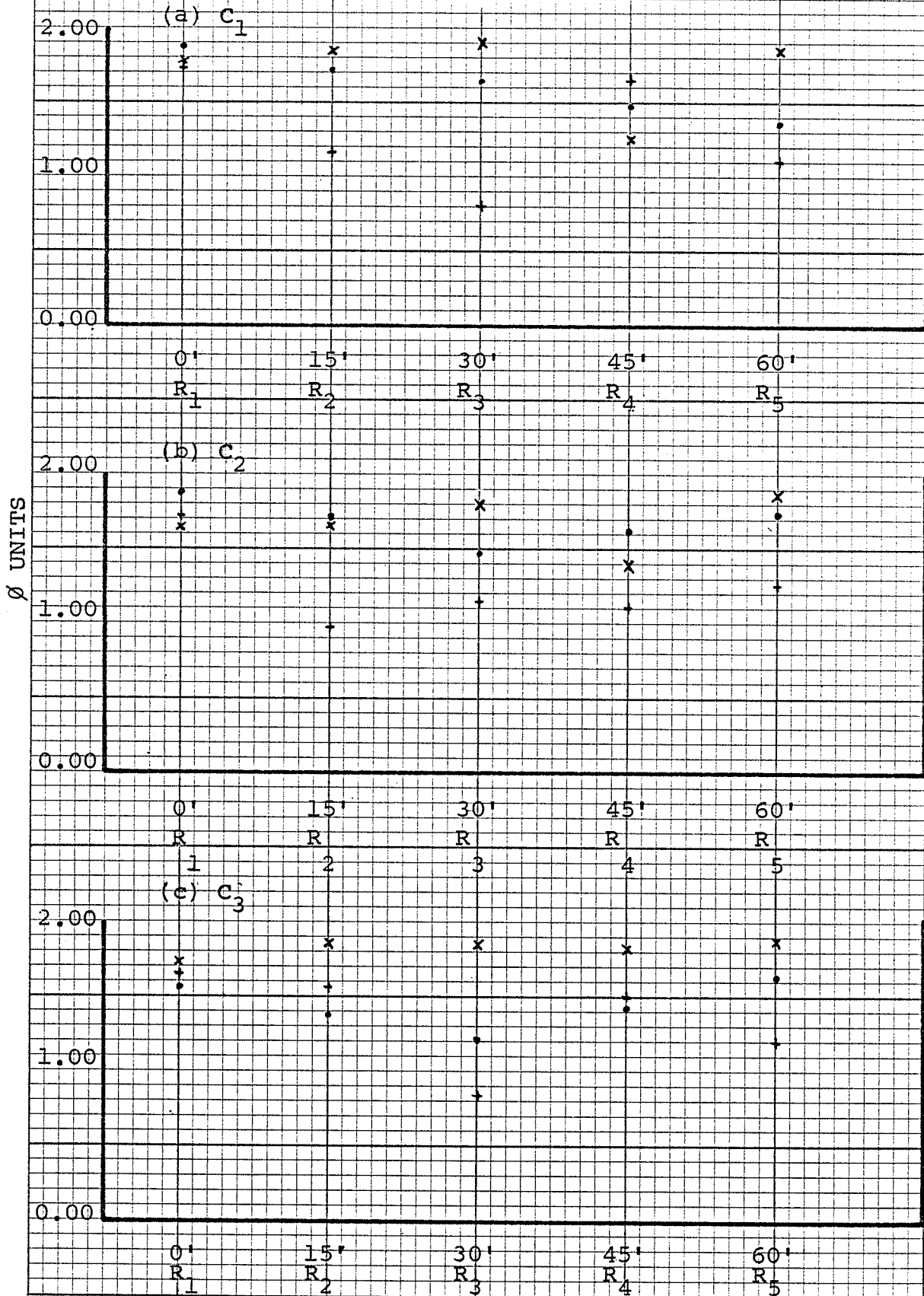


FIGURE 31 CONTINUED

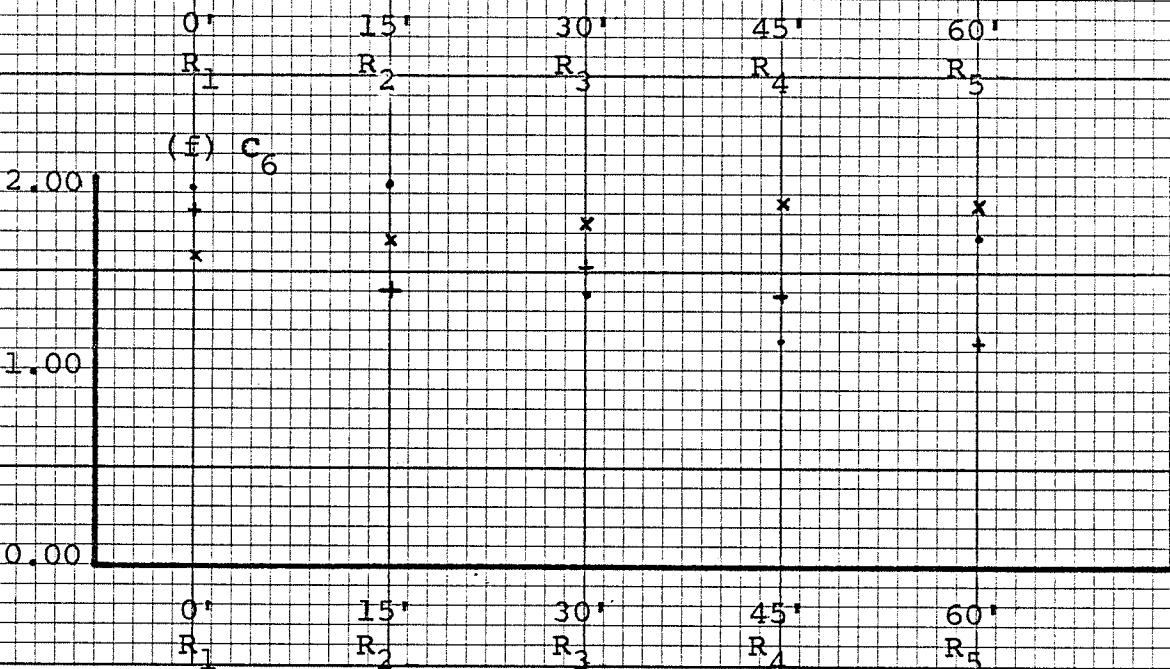
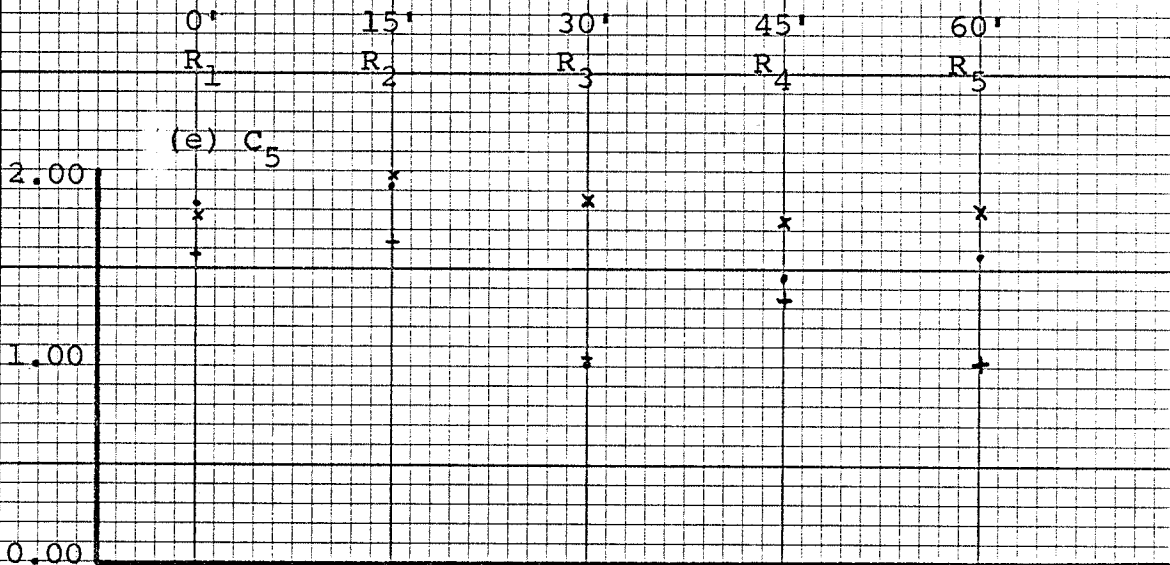
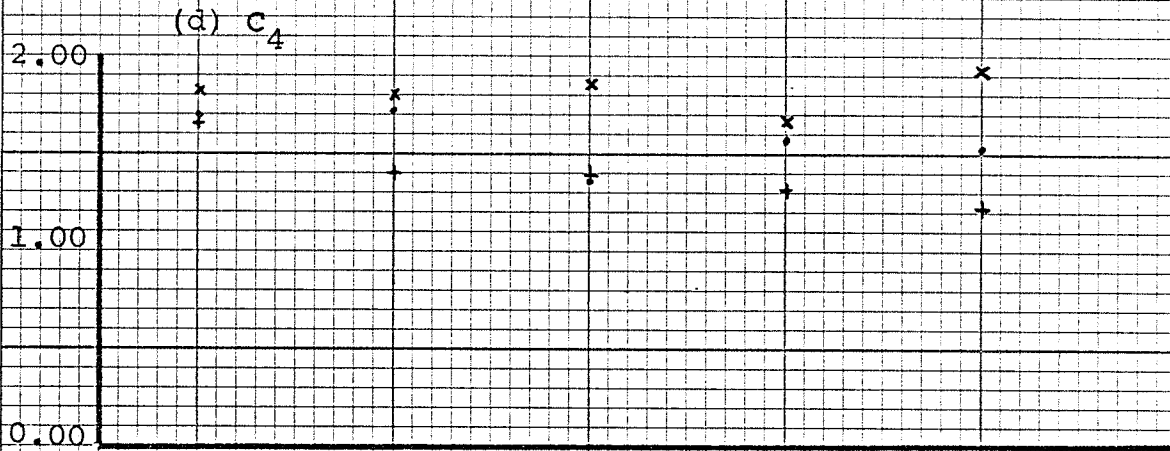


FIGURE 32
 PLOT OF MEAN SIZE AGAINST STANDARD DEVIATION
 NORTH BEACH

• MAY
 + AUGUST
 x NOVEMBER

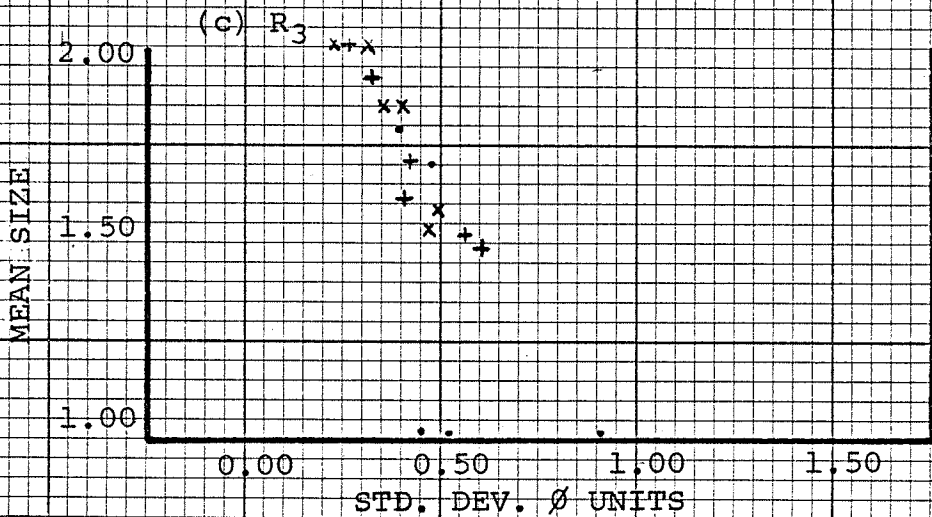
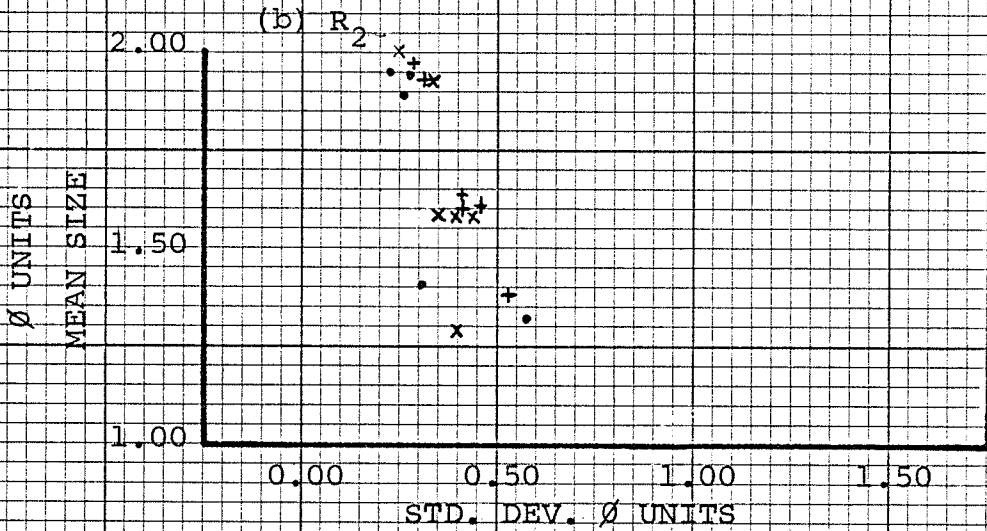
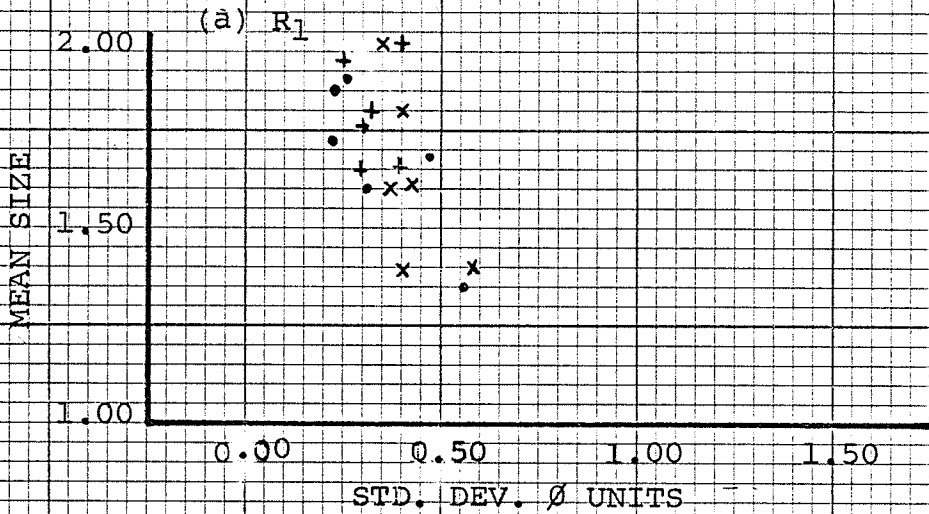


FIGURE 32 CONTINUED

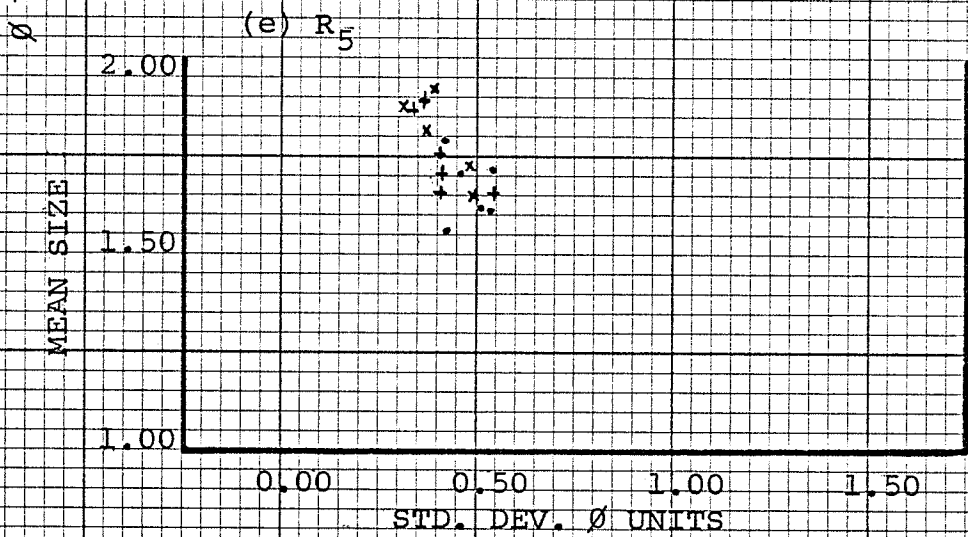
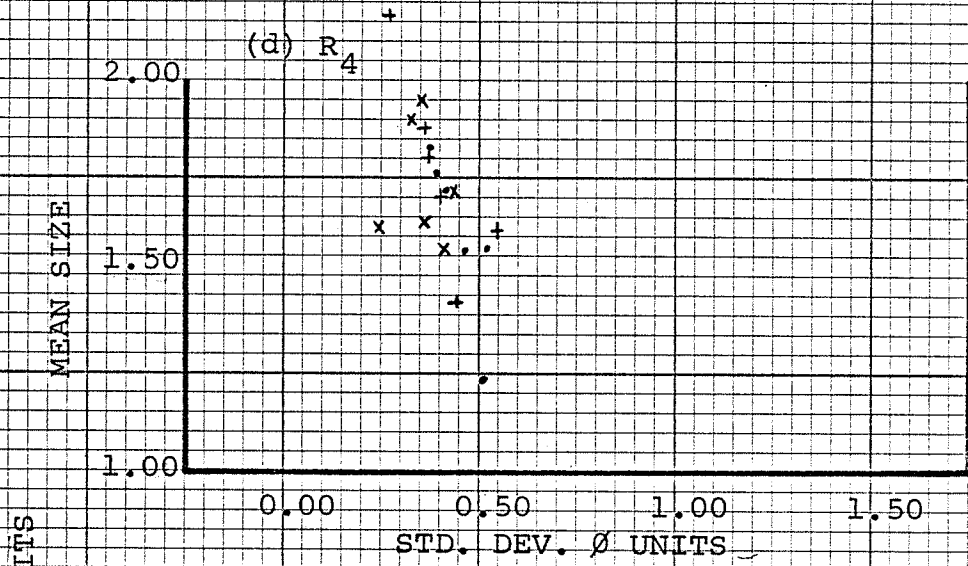


FIGURE 33
 PLOT OF MEAN SIZE AGAINST
 STANDARD DEVIATION SOUTH BEACH

• MAY
 + AUGUST
 x NOVEMBER

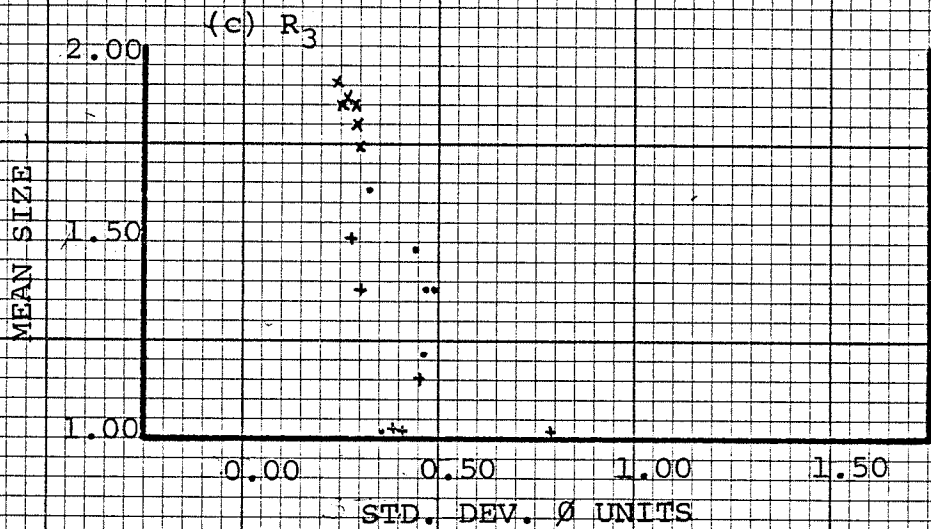
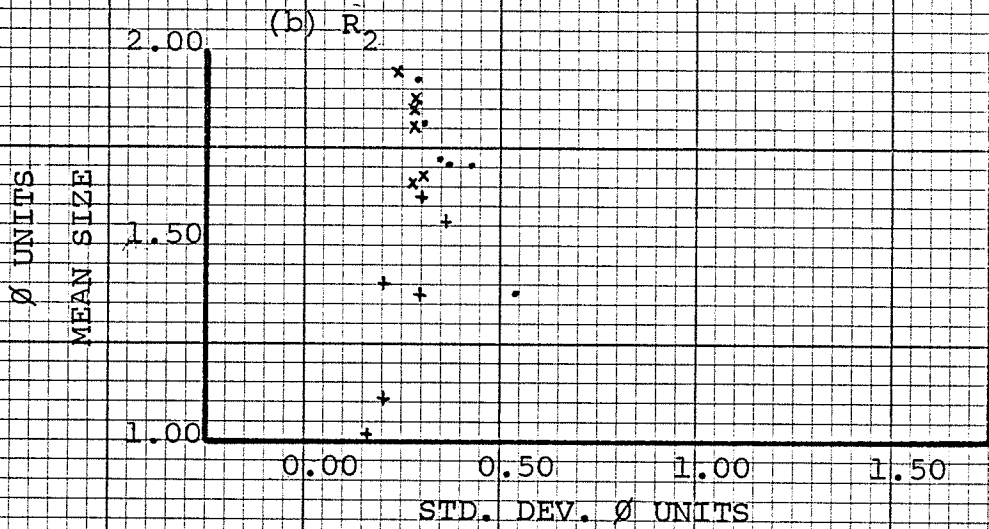
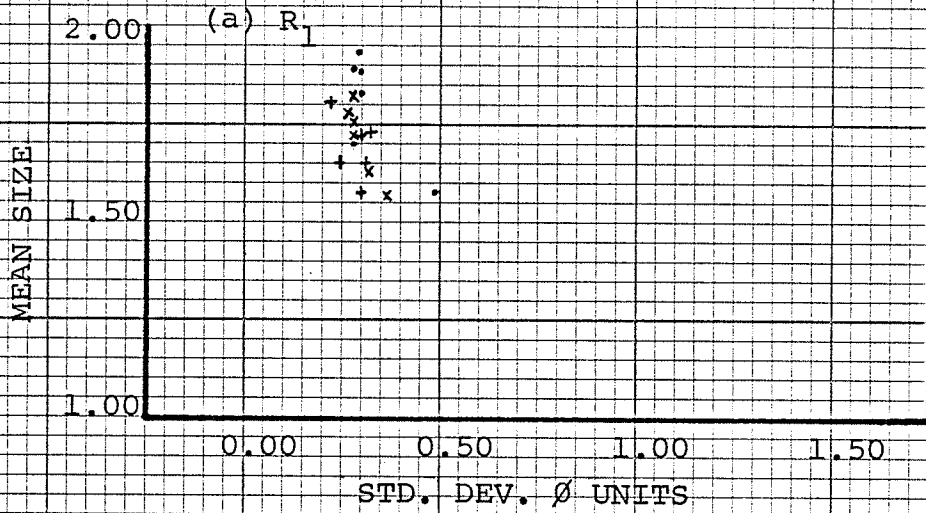
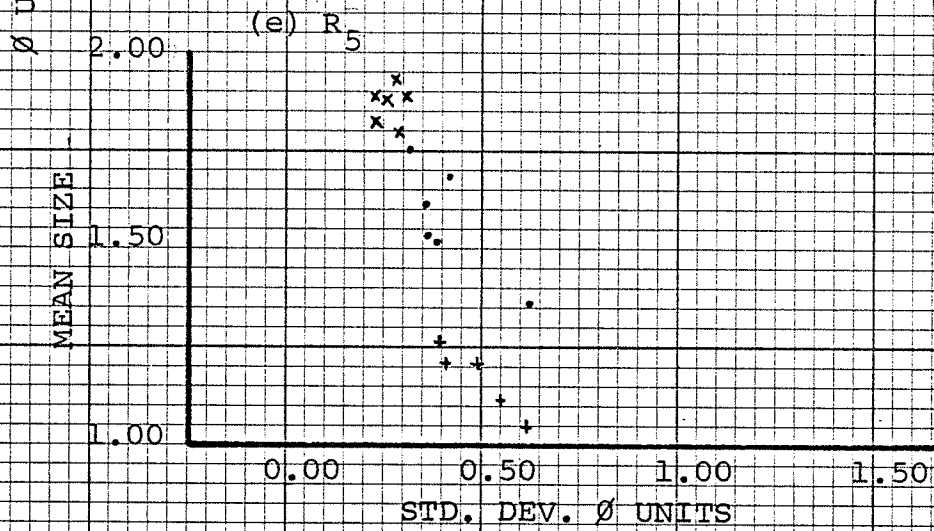
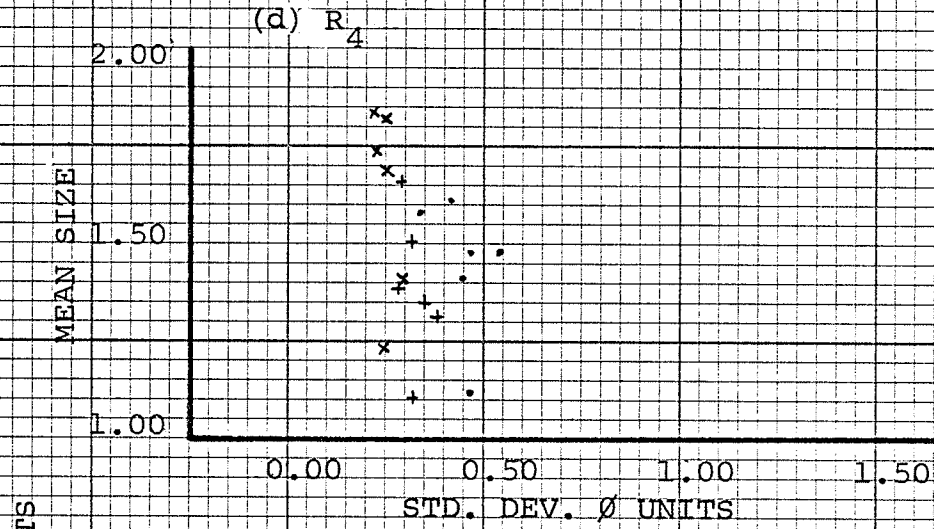


FIGURE 33 CONTINUED



"No surficial sediments showed variations both in mean particle size and in sorting between the study periods. Judged on the averages of mean grain size and sorting, the textural properties of the nearshore sediments appear to be constant throughout the year."

(Dubois, 1972, p.78).

It was hoped that a relationship between the changes in morphology and the changes in sedimentation might be established. These changes could not be correlated, however, due to the very uniformity of source material and processes acting on the study area.

Suggestions for further study to acquire a more comprehensive knowledge of seasonal changes of a beach, would include monitoring of an area over a period of several years. This may confirm seasonal patterns during typical years, and deviations from those patterns during atypical years, and establish whether these atypical years disrupt the dynamic cycle of equilibrium.

Furthermore, observation on a daily basis, particularly after catastrophic events, would add much data to the study of beach movement. For such a study, a permanently manned station, set aside from public use, could be envisioned. An undertaking of this nature is, however, outside the scope of this paper.

As a final conclusion for this study, it can be stated that changes in morphology of a beach are specific to locality, source material and variables governing the rise and fall of lake level. Sandy beaches should be considered as unique entities which differ from region to region, reaffirming the view that they are dynamic and variable landforms.

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APPENDIX A

A Review of Grain Size Parameters as per Computer Program.

Introduction

One of the problems in sedimentology is the ability to give a definite conclusion with respect to its environment and measurable sediment parameters. McBride for example, states:

"Grain size analyses are made for one or more of the following reasons:

- 1) To describe samples in terms of statistical measures.
- 2) To determine the agent (wind, river, turbidity, current, etc.) of transportation and deposition.
- 3) To correlate samples from similar depositional environments or stratigraphic units.
- 4) To determine the process (suspension, traction, saltation, etc.) of final deposition.
- 5) To determine the environment of depositional channel, floor plain, beach, dune, neritic marine, etc."

(McBride, 1971 p. 109).

The mean is a measure to determine overall average size. The mean is affected by every grain in the distribution and therefore should be an indicator of the environmental process.

The standard deviation (σ_1) is used to find an acceptable method of sorting measure. As in graphic approximation to the mean, the more of a curve that is used for a sorting measure, the more accurate the measure will be.

Skewness measures the degree of asymmetry and whether a curve has an asymmetrical tail on the left or right. Friedman (1961) states that the skewness for dune sand is generally positive, while the skewness for beach sand is negative. Furthermore, Stephenson (1970) suggests that skewness seems to be related to environmental conditions and environment energy. He claims that negative skewness has been related to areas of erosion while positive skewness is indicative of deposition.

Kurtosis measures the ratio between the spread in the central part of the distribution and the spread in the tails of the curve.

Folk (1968) states:

"If the central portion is better sorted than the tails, the curve is said to be excessively peaked or leptokurtic; if the tails are better sorted than the central portion, the curve is deficiently or flat peaked and platykurtic."

(Folk 1968, p. 48).

Statistical Parameters and Formulae.

Folk and Ward (1957)

1) Mean size (M_z)

$$= \frac{\emptyset 16 + \emptyset 50 + \emptyset 84}{3}$$

2) Inclusive graphic standard deviation (σ_I)

$$= \frac{\emptyset 84 - \emptyset 16}{4} + \frac{\emptyset 95 - \emptyset 5}{66}$$

3) Skewness (SK_I)

$$= \frac{\emptyset 84 + \emptyset 16 - 2 \emptyset 50}{2 (\emptyset 84 - \emptyset 16)} + \frac{\emptyset 95 + \emptyset 5 - 2 \emptyset 50}{2 (\emptyset 95 - \emptyset 5)}$$

4) Kurtosis (KG)

$$= \frac{\emptyset 95 - \emptyset 5}{2.44 (\emptyset 75 - \emptyset 25)}$$

Inman (1952)

1) Mean

$$= \frac{\emptyset 16 + \emptyset 84}{2}$$

2) Standard Deviation (σ_I)

$$= \frac{\emptyset 84 - \emptyset 16}{2}$$

3) Skewness (SK_I)

The first measure of skewness is used for the central part

$$= \frac{\emptyset 16 + \emptyset 84 - 2 \emptyset 50}{\emptyset 84 - \emptyset 16}$$

and the second measure of skewness is used for the tails,

$$= \frac{\emptyset 5 + \emptyset 95 - 2 \emptyset 50}{\emptyset 84 - \emptyset 16}$$

4) Kurtosis (KG)

$$= \frac{(\phi 95 - \phi 84 - \phi 16)}{\phi 84 - \phi 16}$$

Friedman (1961) Moment Measure

Mean ($\bar{x}\phi$) - First moment

$$\bar{x}\phi = 1/100 \sum f_m \phi$$

where $\bar{x}\phi$ is the mean grain size (phi units) f is the grade size frequency, and $m\phi$ is the mid point of each grain size (phi units).

Skewness (3ϕ) Third moment

$$3\phi = (1/100) \phi^{-3} \sum f (m\phi - \bar{x}\phi)^3$$

where 3ϕ is the skewness and ϕ is the standard deviation (phi units) and which is expressed by

$$\phi\phi = (\sum f) (m\phi - \bar{x}\phi) 2/100 \frac{1}{2}$$

Kurtosis (4ϕ) - Fourth moment

$$4\phi = (1/100) \phi^{-4} \sum f (m\phi - \bar{x}\phi)^4$$

where 4ϕ is the kurtosis.

APPENDIX B

ROW 3

\emptyset	C_1			C_2			C_3		
	May	Aug.	Nov.	May	Aug.	Nov.	May	Aug.	Nov.
-1.00	0.00	0.00	0.00	0.08	0.50	0.00	0.52	0.00	0.00
-0.75	0.02	0.04	0.00	0.20	0.62	0.00	4.07	0.04	0.00
-0.50	0.22	0.16	0.00	1.00	0.71	0.00	5.78	0.06	0.00
-0.25	0.33	0.20	0.00	3.34	0.86	0.00	6.39	0.10	0.07
-0.00	0.43	0.28	0.07	5.27	1.02	0.06	6.40	0.12	0.30
0.25	0.46	0.32	0.08	13.91	1.76	0.06	7.40	0.14	0.44
0.50	0.50	0.54	0.16	22.22	2.67	0.13	7.76	0.17	0.96
0.75	0.71	1.07	0.36	21.39	4.35	0.27	7.00	0.22	2.18
1.00	1.94	3.09	0.76	18.80	7.34	0.86	9.50	0.64	8.88
1.25	6.01	9.45	1.76	8.14	12.89	4.10	12.61	2.21	13.03
1.50	12.69	13.42	3.53	2.21	13.76	9.20	9.60	5.75	15.62
1.75	21.18	18.76	8.95	1.62	14.27	20.37	9.24	13.99	17.98
2.00	29.65	26.09	26.38	0.93	20.39	32.12	8.47	33.49	19.97
2.25	18.79	18.46	36.52	0.37	14.09	23.19	3.88	30.95	13.12
2.50	5.25	5.28	16.29	0.05	3.31	6.54	0.74	8.98	4.67
2.75	1.22	1.77	4.12	0.02	0.82	2.11	0.13	2.36	1.82
3.00	0.16	0.46	0.59	0.00	0.13	0.44	0.02	0.35	0.49
3.25	0.03	0.08	0.05	0.00	0.04	0.05	0.00	0.07	0.07
3.50	0.01	0.04	0.01	0.00	0.02	0.01	0.00	0.04	0.03
3.75	0.00	0.01	0.00	0.00	0.01	0.00	0.00	0.02	0.01
4.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.01	0.00
	C_4			C_5			C_6		
-1.00	0.02	0.06	0.00	0.00	0.25	0.00	0.09	0.07	0.00
-0.75	0.10	0.08	0.00	0.10	0.69	0.00	0.13	0.14	0.00
-0.50	0.32	0.16	0.06	0.11	0.75	0.09	0.43	0.19	0.00
-0.25	0.41	0.20	0.10	0.69	0.99	0.31	1.01	0.25	0.00
0.00	0.57	0.22	0.15	2.06	1.27	0.89	2.69	0.30	0.00
0.25	0.60	0.24	0.19	4.19	1.41	1.06	4.45	0.35	0.00
0.50	1.01	0.26	0.37	8.81	1.69	1.83	8.65	0.47	0.01
0.75	1.57	0.28	0.61	13.84	2.27	2.69	15.86	0.79	0.02
1.00	4.24	0.37	0.93	24.38	6.69	5.10	25.17	2.83	0.03
1.25	9.04	0.81	5.52	21.45	9.45	12.79	20.56	14.50	0.21
1.50	12.20	2.47	9.65	8.69	13.00	18.86	8.29	19.72	0.90
1.75	16.60	9.54	15.82	5.87	22.97	22.03	4.92	19.84	5.12
2.00	25.14	31.33	29.87	5.33	22.51	19.72	4.50	22.65	29.63
2.25	21.13	40.63	23.96	3.36	11.87	10.54	2.34	13.45	43.36
2.50	5.60	11.39	8.12	0.80	2.91	2.40	0.41	3.34	15.86
2.75	1.13	1.57	2.73	0.19	0.65	0.87	0.09	0.64	3.95
3.00	0.16	0.12	1.01	0.05	0.14	0.28	0.01	0.06	0.53
3.25	0.04	0.01	0.29	0.01	0.02	0.06	0.00	0.01	0.03
3.50	0.02	0.00	0.15	0.00	0.00	0.02	0.00	0.00	0.01
3.75	0.00	0.00	0.09	0.00	0.00	0.01	0.00	0.00	0.00
4.00	0.00	0.00	0.02	0.00	0.00	0.00	0.00	0.00	0.00

APPENDIX C

Kurtosis South Beach

	c_1	c_2	c_3	c_4	c_5	c_6
<u>Row 1</u>						
May	1.034	1.088	1.002	1.009	1.013	1.116
Aug.	0.974	0.913	0.938	0.997	0.936	0.965
Nov.	0.963	0.956	0.945	1.002	1.014	0.910
<u>Row 2</u>						
May	1.383	1.065	0.941	0.974	1.051	1.096
Aug.	1.495	0.776	0.903	1.356	0.934	0.899
Nov.	0.947	0.932	1.019	1.019	1.011	0.978
<u>Row 3</u>						
May	0.983	1.072	1.089	1.059	1.487	0.860
Aug.	0.989	1.248	1.038	0.962	1.160	0.966
Nov.	0.998	1.072	0.949	1.006	0.969	0.942
<u>Row 4</u>						
May	1.111	1.023	1.185	0.932	0.926	1.047
Aug.	0.860	1.393	0.944	1.367	1.009	1.040
Nov.	1.099	0.978	1.024	0.996	1.013	1.037
<u>Row 5</u>						
May	1.068	0.783	0.913	0.994	0.935	1.038
Aug.	1.311	1.215	1.348	1.414	1.587	1.352
Nov.	1.029	0.999	1.017	1.077	0.976	0.977

Skewness South Beach

	c_1	c_2	c_3	c_4	c_5	c_6
<u>Row 1</u>						
May	-0.038	-0.047	-0.253	0.135	0.035	0.076
Aug.	-0.057	-0.007	0.022	0.052	0.228	0.018
Nov.	-0.086	-0.024	-0.009	-0.046	-0.030	0.011
<u>Row 2</u>						
May	-0.376	-0.192	-0.128	-0.088	0.072	0.092
Aug.	-0.226	0.146	0.215	0.113	0.069	0.169
Nov.	-0.095	-0.043	-0.072	-0.057	0.042	-0.058
<u>Row 3</u>						
May	0.010	-0.067	-0.016	-0.119	0.193	0.239
Aug.	-0.369	-0.015	-0.190	0.076	-0.012	0.082
Nov.	-0.081	-0.085	-0.009	-0.127	-0.041	-0.069
<u>Row 4</u>						
May	-0.263	-0.150	-0.053	0.044	0.132	0.507
Aug.	0.074	0.191	-0.020	-0.137	0.033	-0.002
Nov.	0.156	0.211	-0.052	0.041	-0.049	-0.023
<u>Row 5</u>						
May	-0.246	-0.038	-0.063	-0.001	0.098	-0.128
Aug.	-0.185	0.021	-0.100	-0.115	-0.273	-0.222
Nov.	-0.057	-0.016	-0.074	-0.044	-0.058	-0.118

Standard Deviation σ Units South Beach

	c_1	c_2	c_3	c_4	c_5	c_6
<u>Row 1</u>						
May	0.275	0.294	0.497	0.276	0.307	0.298
Aug.	0.334	0.305	0.310	0.242	0.311	0.227
Nov.	0.297	0.337	0.296	0.283	0.272	0.365
<u>Row 2</u>						
May	0.436	0.366	0.543	0.349	0.247	0.258
Aug.	0.204	0.169	0.361	0.203	0.306	0.289
Nov.	0.280	0.301	0.288	0.270	0.239	0.329
<u>Row 3</u>						
May	0.327	0.448	0.472	0.467	0.352	0.463
Aug.	0.797	0.457	0.429	0.306	0.386	0.280
Nov.	0.244	0.256	0.273	0.278	0.257	0.306
<u>Row 4</u>						
May	0.540	0.413	0.453	0.336	0.374	0.476
Aug.	0.297	0.321	0.311	0.388	0.354	0.280
Nov.	0.248	0.292	0.250	0.255	0.236	0.223
<u>Row 5</u>						
May	0.626	0.333	0.355	0.383	0.357	0.435
Aug.	0.557	0.399	0.496	0.423	0.617	0.536
Nov.	0.251	0.242	0.307	0.285	0.289	0.311

Mean Size ϕ Units South Beach

	C_1	C_2	C_3	C_4	C_5	C_6
<u>Row 1</u>						
May	1.887	1.885	1.557	1.702	1.836	1.943
Aug.	1.735	1.724	1.656	1.654	1.582	1.807
Nov.	1.776	1.640	1.718	1.829	1.788	1.575
<u>Row 2</u>						
May	1.701	1.707	1.383	1.712	1.925	1.950
Aug.	1.183	0.984	1.573	1.400	1.639	1.395
Nov.	1.855	1.684	1.862	1.807	1.947	1.661
<u>Row 3</u>						
May	1.646	1.485	1.214	1.375	1.086	1.393
Aug.	0.806	1.159	0.831	1.393	1.044	1.518
Nov.	1.907	1.808	1.850	1.864	1.853	1.748
<u>Row 4</u>						
May	1.479	1.614	1.412	1.591	1.474	1.130
Aug.	1.666	1.107	1.505	1.324	1.353	1.394
Nov.	1.246	1.402	1.830	1.697	1.747	1.847
<u>Row 5</u>						
May	1.360	1.752	1.622	1.524	1.575	1.690
Aug.	1.115	1.267	1.207	1.216	1.051	1.137
Nov.	1.879	1.895	1.809	1.943	1.803	1.842

Kurtosis North Beach

	c_1	c_2	c_3	c_4	c_5	c_6
<u>Row 1</u>						
May	1.095	0.928	0.951	1.098	0.860	0.947
Aug.	0.960	1.000	1.307	0.959	1.005	0.970
Nov.	1.097	1.032	1.091	0.992	0.894	1.005
<u>Row 2</u>						
May	1.133	0.874	0.986	1.119	1.032	1.011
Aug.	1.034	0.913	1.204	1.048	1.102	0.983
Nov.	1.084	1.153	1.633	0.944	0.926	1.017
<u>Row 3</u>						
May	1.111	1.085	0.813	0.998	1.311	1.314
Aug.	1.014	1.025	1.217	1.123	1.265	0.860
Nov.	1.196	1.080	0.923	1.113	1.100	1.047
<u>Row 4</u>						
May	0.929	1.096	1.055	0.955	0.969	0.991
Aug.	0.980	1.048	1.156	1.092	0.860	1.077
Nov.	1.130	0.953	0.926	1.009	1.164	1.148
<u>Row 5</u>						
May	1.028	1.105	1.070	0.956	0.966	1.138
Aug.	0.985	1.111	1.092	1.184	0.972	1.075
Nov.	0.956	1.100	1.077	1.388	1.021	1.100

Skewness North Beach

	C_1	C_2	C_3	C_4	C_5	C_6
<u>Row 1</u>						
May	-0.249	-0.019	0.006	-0.026	-0.269	-0.118
Aug.	-0.065	0.186	-0.224	-0.008	-0.028	-0.115
Nov.	-0.081	-0.131	-0.073	-0.062	-0.027	-0.197
<u>Row 2</u>						
May	-0.257	0.171	-0.016	-0.026	0.050	0.019
Aug.	-0.051	-0.110	-0.090	0.070	-0.223	-0.106
Nov.	-0.071	-0.050	0.408	-0.086	-0.053	-0.102
<u>Row 3</u>						
May	-0.154	-0.008	-0.170	-0.264	0.088	0.043
Aug.	-0.153	-0.254	-0.113	-0.127	-0.328	-0.048
Nov.	-0.118	-0.061	-0.067	-0.130	-0.149	0.008
<u>Row 4</u>						
May	0.065	-0.270	-0.241	-0.146	-0.182	-0.113
Aug.	-0.217	-0.227	-0.151	-0.227	0.110	0.115
Nov.	-0.150	-0.120	0.055	-0.085	0.112	-0.091
<u>Row 5</u>						
May	-0.217	-0.287	-0.243	-0.116	-0.197	-0.256
Aug.	-0.219	-0.296	-0.224	-0.171	-0.149	-0.104
Nov.	-0.244	-0.317	-0.089	-0.147	0.064	-0.086

Standard Deviation σ Units North Beach

	C_1	C_2	C_3	C_4	C_5	C_6
<u>Row 1</u>						
May	0.556	0.319	0.254	0.245	0.468	0.379
Aug.	0.325	0.295	0.402	0.236	0.296	0.393
Nov.	0.371	0.559	0.413	0.419	0.364	0.417
<u>Row 2</u>						
May	0.587	0.306	0.283	0.273	0.223	0.225
Aug.	0.527	0.416	0.312	0.279	0.460	0.416
Nov.	0.251	0.323	0.396	0.359	0.361	0.392
<u>Row 3</u>						
May	0.396	0.451	0.909	0.473	0.528	0.518
Aug.	0.440	0.608	0.323	0.262	0.560	0.406
Nov.	0.316	0.349	0.499	0.397	0.476	0.246
<u>Row 4</u>						
May	0.503	0.411	0.391	0.460	0.536	0.373
Aug.	0.405	0.549	0.364	0.370	0.445	0.268
Nov.	0.334	0.412	0.371	0.416	0.243	0.327
<u>Row 5</u>						
May	0.477	0.549	0.427	0.426	0.505	0.540
Aug.	0.413	0.542	0.413	0.370	0.403	0.338
Nov.	0.480	0.496	0.365	0.393	0.273	0.328

TABLE 5 - GRAIN SIZE PARAMETERS

Mean Size ϕ Units North Beach

	C_1	C_2	C_3	C_4	C_5	C_6
<u>Row 1</u>						
May	1.353	1.593	1.898	1.857	1.680	1.720
Aug.	1.797	1.656	1.969	1.938	1.774	1.665
Nov.	1.968	1.407	1.818	1.371	1.600	1.612
<u>Row 2</u>						
May	1.313	1.407	1.805	1.951	1.948	1.949
Aug.	1.387	1.602	1.936	1.957	1.616	1.662
Nov.	2.018	1.928	1.286	1.594	1.650	1.630
<u>Row 3</u>						
May	1.759	0.542	0.760	1.706	0.982	0.924
Aug.	1.727	1.492	1.926	2.004	1.541	1.615
Nov.	2.045	1.853	1.587	1.853	1.543	2.077
<u>Row 4</u>						
May	1.230	1.729	1.755	1.557	1.566	1.827
Aug.	1.700	1.611	1.878	1.795	1.430	2.156
Nov.	1.953	1.574	1.640	1.739	1.634	1.911
<u>Row 5</u>						
May	1.702	1.714	1.795	1.555	1.623	1.602
Aug.	1.709	1.664	1.777	1.899	1.660	1.878
Nov.	1.721	1.659	1.822	1.926	1.926	1.877

APPENDIX D

BY GREG McMILLAN, A GRADUATE STUDENT, DEPT. OF GEOGRAPHY, THE UNIVERSITY OF MANITOBA FROM A PROGRAM WRITTEN BY W.C. ISOPHORDING, OF THE DEPT. OF GEOLOGY, THE UNIVERSITY OF SOUTH ALABAMA, MOBILE.

C
C
C

-PROGRAM LISTING-

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1 CHARACTER*2 TITLE(12)
2 DIMENSION PHI(2,99),PCT(99),WT(99),DF(4),DEL(99)
3 NREAD=5
4 NWRIT=6
5 IRR=0
6 1000 READ (NREAD,1001) J,(TITLE(I),I=1,12),PHI(1,1),TOTWT,XMQ
7 IF (J.EQ.0) GO TO 10000
8 1001 FORMAT (I2,2X,12A2,F7.2,F7.2,F7.2)
9 JA=J-1
10 READ (NREAD,1002) (DEL(N),N=1,J)
11 1002 FORMAT (8F10.2)
12 DO 5 N=2,J
13 PHI(1,N)=PHI(1,N-1)+DEL(N)
14 5 CONTINUE
15 READ (NREAD,1002) (WT(N),N=1,J)
16 I=1
17 K=1
18 WTTOT=0.0
19 DO 3 N=1,J
20 PCT(N)=100.*(WT(N)/TOTWT)
21 WTTOT=WTTOT+WT(N)
22 3 CONTINUE
23 WTDIF=TOTWT-WTTOT
24 PHI(2,1)=PCT(1)
25 DO 4 N=2,J
26 PHI(2,N)=PCT(N)+PHI(2,N-1)
27 4 CONTINUE
28 WRITE (NWRIT,601) (TITLE(I),I=1,12)
29 601 FORMAT ('1',12A2)
30 WRITE (NWRIT,417)
31 N=1
32 IF(PHI(2,J)-84.0)100,310,310
33 310 IF(PHI(2,J)-102.0)612,920,920
34 920 WRITE (NWRIT,319)
35 319 FORMAT (' ','**ERROR** SIEVE CONTENTS GREATER THAN TOTAL SAMPLE WE
*IGHT,SUGGEST YOU CHECK YOUR INPUT DATA.')
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36 IRR=IRR+1
37 999 GO TO (1000,70,71,72),IRR
38 70 WRITE (NWRIT,970)
39 GO TO 1000
40 71 WRITE (NWRIT,971)
41 GO TO 1000
42 72 WRITE (NWRIT,972)
43 IRR=IRR-1
44 GO TO 1000
45 612 IF(PHI(2,1)-5.0)613,613,614
46 614 PHI5=5.*DEL(N)/PHI(2,1)+PHI(1,N)
47 IF (PHI(2,1)-16.0)9,9,615
48 615 PHI16=16.*DEL(N)/PHI(2,1)+PHI(1,1)
49 GO TO 13
50 613 IF(PHI(2,N+1)-5.0)6,6,8
51 6 N=N+1
52 GO TO 612
53 8 PHI5 =DEL(N)*((5.0-PHI(2,N))/(PHI(2,N+1)-PHI(2,N)))+PHI(1,N)
54 9 IF(PHI(2,N+1)-16.)11,11,12

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57 14 N=N+1
60 GO TO 13
61 15 PHI25=DEL(N)*((25.0-PHI(2,N))/(PHI(2,N+1)-PHI(2,N)))+PHI(1,N)
62 16 IF(PHI(2,N+1)-50.)17,17,18
63 17 N=N+1
64 GO TO 16
65 18 PHI50=DEL(N)*((50.0-PHI(2,N))/(PHI(2,N+1)-PHI(2,N)))+PHI(1,N)
66 19 IF(PHI(2,N+1)-75.)20,20,21
67 20 N=N+1
68 GO TO 19
69 21 PHI75=DEL(N)*((75.0-PHI(2,N))/(PHI(2,N+1)-PHI(2,N)))+PHI(1,N)
70 22 IF(PHI(2,N+1)-84.)23,23,24
71 23 N=N+1
72 IF(N-JA)22,22,100
73 24 PHI84=DEL(N)*((84.0-PHI(2,N))/(PHI(2,N+1)-PHI(2,N)))+PHI(1,N)
74 25 IF(PHI(2,N+1)-95.)26,26,27
75 26 N=N+1
76 IF(N-JA)25,25,625
77 27 PHI95=DEL(N)*((95.0-PHI(2,N))/(PHI(2,N+1)-PHI(2,N)))+PHI(1,N)
78 EMZ=.333*(PHI16+PHI50+PHI84)
79 SIGI=.25*(PHI84-PHI16)+.1515*(PHI95-PHI5)
80 SNAP=.5*((PHI5+PHI95-2.*PHI50)/(PHI95-PHI5))
81 SKI=.5*((PHI16+PHI84-2.*PHI50)/(PHI84-PHI16))+SNAP
82 AL2FI=(PHI95+PHI5-2.*PHI50)/(PHI84-PHI16)
83 CAYGP=.4098*((PHI95-PHI5)/(PHI75-PHI25))
84 28 CONTINUE
85 EMPHI=.5*(PHI16+PHI84)
86 SIGFI=.5*(PHI84-PHI16)
87 ALFI=(PHI84+PHI16-2.*PHI50)/(PHI84-PHI16)
88 GO TO (701,702),I
89 701 CAYGI=(PHI16-PHI5+PHI95-PHI84)/(PHI84-PHI16)
90 702 CONTINUE
91 C THIS POINT INDICATES COMPLETION OF FOLK, INMAN. START MOMENTIME
92 GO TO (29,29,200),K
93 29 CONTINUE
94 C IF(PHI(2,J)-99.5)300,30,30
95 C M=1 SETS SWITCH 1 OFF
96 C M=2 SETS SWITCH 1 ON
97 30 M=2
98 GO TO (320,31),M
99 31 CONTINUE
100 DO 35 L=1,4
101 DF(L)=0
102 35 CONTINUE
103 DO 40 N=1,JA
104 IF (PCT(N)) 40,40,41
105 41 A=PHI(1,N)-.5*DEL(N)-XMQ
106 IF (A) 42,40,42
107 42 D=A*A
108 E=A*D
109 DF(1)=PCT(N)*A+DF(1)
110 DF(2)=PCT(N)*D+DF(2)
111 DF(3)=PCT(N)*E+DF(3)
112 DF(4)=PCT(N)*E*A+DF(4)
113 40 CONTINUE
114 DO 45 L=1,4
115 DF(L)=.01*DF(L)
116 45 CONTINUE
117 XBAR=XMQ+DF(1)
118 B=DF(1)*DF(1)
119 SSQD=DF(2)-B
120 S=SQRT(SSQD)

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122 C THIS COMPLETES THE CALCULATION OF MOMENT MEASURE
      C GO TO 400
      C THE FOLLOWING SECTION PROVIDES OUTPUT FORMS AND SELECTS FOLKS
      C TEXTURAL GROUPING TERMS
123 102 FORMAT(33H CALCULATION OF INMANS STATISTICS)
124 103 FORMAT(32H CALCULATION OF FOLKS STATISTICS)
125 104 FORMAT(41H CALCULATION OF MOMENT MEASURE STATISTICS)
126 105 FORMAT(1X,'INSUFFICIENT DATA')
127 106 FORMAT(28H MOMENT MEASURE NOT COMPUTED)
128 107 FORMAT(48H DATA FOR DRAWING A FREQUENCY DISTRIBUTION CURVE)
129 109 FORMAT(17H VERY WELL SORTED)
130 111 FORMAT(1X,'MODERATELY SORTED')
131 1313 FORMAT(1X,'MODERATELY WELL SORTED')
132 110 FORMAT(1X,'WELL SORTED')
133 112 FORMAT(14H POORLY SORTED)
134 113 FORMAT(19H VERY POORLY SORTED)
135 114 FORMAT(24H EXTREMELY POORLY SORTED)
136 121 FORMAT(21H STRONGLY FINE SKEWED)
137 122 FORMAT(12H FINE SKEWED)
138 123 FORMAT(17H NEAR SYMMETRICAL)
139 124 FORMAT(14H COARSE SKEWED)
140 125 FORMAT(23H STRONGLY COARSE SKEWED)
141 131 FORMAT(17H VERY PLATYKURTIC)
142 720 FORMAT(1X,'KG (INMAN)='F7.3,' KURTOSIS VALUE')
143 132 FORMAT(12H PLATYKURTIC)
144 133 FORMAT(11H MESOKURTIC)
145 134 FORMAT(12H LEPTOKURTIC)
146 135 FORMAT(17H VERY LEPTOKURTIC)
147 136 FORMAT(22H EXTREMELY LEPTOKURTIC)
148 141 FORMAT(7H GRAVEL)
149 142 FORMAT(13H SANDY GRAVEL)
150 143 FORMAT(19H MUDDY SANDY GRAVEL)
151 144 FORMAT(13H MUDDY GRAVEL)
152 145 FORMAT(14H GRAVELLY SAND)
153 146 FORMAT(20H GRAVELLY MUDDY SAND)
154 147 FORMAT(13H GRAVELLY MUD)
155 148 FORMAT(23H SLIGHTLY GRAVELLY SAND)
156 149 FORMAT(29H SLIGHTLY GRAVELLY MUDDY SAND)
157 150 FORMAT(28H SLIGHTLY GRAVELLY SANDY MUD)
158 151 FORMAT(22H SLIGHTLY GRAVELLY MUD)
159 160 FORMAT(5H SAND)
160 161 FORMAT(12H CLAYEY SAND)
161 162 FORMAT(11H MUDDY SAND)
162 163 FORMAT(11H SILTY SAND)
163 164 FORMAT(11H SANDY CLAY)
164 165 FORMAT(10H SANDY MUD)
165 166 FORMAT(11H SANDY SILT)
166 167 FORMAT(5H CLAY)
167 168 FORMAT(4H MUD)
168 169 FORMAT(5H SILT)
      C THIS LIST PROVIDES ALPHA OUTPUT TO BE USED IN THE FOLLOWING
      C DECISION NETWORK BASED UPON FOLKS TEXTURAL TRIANGLE DIAGRAMS
      C FOLK, JOUR GEOL. V0162, P345-351, JULY 1954
169 275 FORMAT(38H DATA IS TOO OPENENDED FOR CALCULATION)
      C THE FOLLOWING PROCEEDURE NAMES THE SAMPLE ACCORDING TO FOLKS
      C TEXTURAL GROUPING
170 277 GRPCT=0.0001
171 N=1
172 801 IF(PHI(1,N)+1.) 802,803,804
173 802 GRPCT=GRPCT+PCT(N)
174 805 IF(N-J)805,806,806
175 805 N=N+1

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181 807 GRPCT=GRPCT+PCT(N)-SPCT
182 807 IF(N-J)808,809,809
183 808 N=N+1
184 810 IF(PHI(1,N)-4.)810,811,812
185 810 SPCT=SPCT+PCT(N)
186 807 GO TO 807
187 811 SPCT=SPCT+PCT(N)
188 809 GO TO 809
189 812 SIPCT=PCT(N)*(PHI(1,N)-4.)/DEL(N)
190 812 SPCT=SPCT+PCT(N)-SIPCT
191 813 GO TO 813
192 809 SIPCT=0.0001
193 813 IF(N-J)814,815,815
194 814 N=N+1
195 816 IF(PHI(1,N)-8.)816,817,818
196 816 SIPCT=SIPCT+PCT(N)
197 813 GO TO 813
198 817 SIPCT=SIPCT+PCT(N)
199 815 CLPCT=0.0001
200 819 GO TO 819
201 818 CLPCT=PCT(N)*(PHI(1,N)-8.)/DEL(N)
202 819 SIPCT=SIPCT+PCT(N)-CLPCT
203 819 IF(N-J)820,215,822
204 822 WRITE (NWRIT,823)
205 823 FORMAT(23H POSSIBLE MACHINE ERROR)
206 820 CLPCT=CLPCT+PHI(2,J)-PHI(2,N)
207 215 CONTINUE
208 417 WRITE (NWRIT,417)
209 218 WRITE (NWRIT,218)
210 250,250,217 IF(GRPCT-.1)250,250,217
211 217 EMPCT=SIPCT+CLPCT
212 220,219,219 IF(GRPCT-80.)220,219,219
213 219 WRITE (NWRIT,141)
214 500 GO TO 500
215 220,221,221 IF(GRPCT-30.)222,221,221
216 221 IF(EMPCT-SPCT)223,224,224
217 224 WRITE (NWRIT,144)
218 500 GO TO 500
219 223 IF((SPCT/EMPCT)-9.)225,225,226
220 225 WRITE (NWRIT,143)
221 500 GO TO 500
222 226 WRITE (NWRIT,142)
223 500 GO TO 500
224 222,228,228 IF(GRPCT-5.)227,228,228
225 228 IF(EMPCT-SPCT)230,229,229
226 229 WRITE (NWRIT,147)
227 500 GO TO 500
228 230,231,232 IF((SPCT/EMPCT)-9.)231,231,232
229 231 WRITE (NWRIT,146)
230 232 WRITE (NWRIT,145)
231 500 GO TO 500
232 227,234,234 IF(EMPCT-SPCT)233,234,234
233 233 IF((SPCT/EMPCT)-9.)235,235,236
234 235 WRITE (NWRIT,149)
235 500 GO TO 500
236 236 WRITE (NWRIT,148)
237 500 GO TO 500
238 234,237,238 IF((EMPCT/SPCT)-9.)237,237,238
239 237 WRITE (NWRIT,150)
240 500 GO TO 500
241 238 WRITE (NWRIT,151)

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247 253 IF(SPCT-50.)255,256,256
248 254 IF((SIPCT/CLPCT)-2.)261,261,262
249 256 WRITE (NWRIT,161)
250 GO TO 500
251 255 IF(SPCT-10.)259,260,260
252 259 WRITE (NWRIT,167)
253 GO TO 500
254 260 WRITE (NWRIT,164)
255 GO TO 500
256 261 IF(SPCT-50.)263,264,264
257 264 WRITE (NWRIT,162)
258 GO TO 500
259 263 IF(SPCT-10.)265,266,266
260 265 WRITE (NWRIT,168)
261 GO TO 500
262 266 WRITE (NWRIT,165)
263 GO TO 500
264 262 IF(SPCT-50.)267,268,268
265 268 WRITE (NWRIT,163)
266 GO TO 500
267 267 IF(SPCT-10.)269,270,270
268 269 WRITE (NWRIT,169)
269 GO TO 500
270 270 WRITE (NWRIT,166)
271 GO TO 500
272 100 WRITE (NWRIT,275)
273 IRR=IRR+1
274 GO TO 999
275 200 WRITE (NWRIT,276)
276 K=3
277 GO TO 301
278 300 WRITE (NWRIT,105)
279 320 WRITE (NWRIT,106)
280 K=2
281 GO TO 301
282 400 K=1
283 GO TO 301
284 625 K=3
285 I=2
286 GO TO 28
C M = 1 SETS SWITCH 2 OFF
C M = 2 SETS SWITCH 2 ON
287 301 M=2
288 GO TO (660,661),M
289 660 CONTINUE
290 661 GO TO (302,303,304),K
291 302 WRITE (NWRIT,417)
292 WRITE (NWRIT,104)
293 WRITE (NWRIT,402) XBAR
294 WRITE (NWRIT,416) SSQD
295 WRITE (NWRIT,401) S
296 WRITE (NWRIT,403) SK
297 WRITE (NWRIT,404) CTSIS
298 WRITE (NWRIT,415) EM3,EM4
299 GO TO 600
300 303 WRITE (NWRIT,417)
301 WRITE (NWRIT,103)
302 WRITE (NWRIT,406) EMZ
303 WRITE (NWRIT,405) SIGI
304 WRITE (NWRIT,403) SKI
305 WRITE (NWRIT,404) CAYGP

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311      WRITE (NWRIT,102)
312      WRITE (NWRIT,408) EMPHI
313      WRITE (NWRIT,409) SIGFI
314      WRITE (NWRIT,403) ALFI
315      GO TO (703,704),I
316  703  WRITE (NWRIT,720) CAYGI
317      WRITE (NWRIT,411) AL2FI
318  704  CONTINUE
319      GO TO 600
320  500  CONTINUE
321      IF(ES-4.)502,502,501
322  501  WRITE (NWRIT,114)
323      GO TO 520
324  502  IF(ES-2.)504,504,503
325  503  WRITE (NWRIT,113)
326      GO TO 520
327  504  IF(ES-1.)506,506,505
328  505  WRITE (NWRIT,112)
329      GO TO 520
330  506  IF (ES-.71)1508,1508,507
331  507  WRITE (NWRIT,111)
332      GO TO 520
333  1508 IF (ES-.50) 508,508,1507
334  1507 WRITE (NWRIT,1313)
335      GO TO 520
336  508  IF(ES-.35)510,510,509
337  509  WRITE (NWRIT,110)
338      GO TO 520
339  510  WRITE (NWRIT,109)
340  520  GO TO (705,540),I
341  705  CONTINUE
342      IF(X1-3.)522,522,521
343  521  WRITE (NWRIT,136)
344      GO TO 540
345  522  IF(X1-1.5)524,524,523
346  523  WRITE (NWRIT,135)
347      GO TO 540
348  524  IF(X1-1.11)526,526,525
349  525  WRITE (NWRIT,134)
350      GO TO 540
351  526  IF(X1-.90)528,528,527
352  527  WRITE (NWRIT,133)
353      GO TO 540
354  528  IF(X1-.67)530,530,529
355  529  WRITE (NWRIT,132)
356      GO TO 540
357  530  WRITE (NWRIT,131)
358  540  CONTINUE
359      IF((ESK+1.)-1.30)542,542,541
360  541  WRITE (NWRIT,121)
361      GO TO 600
362  542  IF((ESK+1.)-1.1)544,544,543
363  543  WRITE (NWRIT,122)
364      GO TO 600
365  544  IF((ESK+1.)-.9)546,546,545
366  545  WRITE (NWRIT,123)
367      GO TO 600
368  546  IF((ESK+1.)-.70)548,548,547
369  547  WRITE (NWRIT,124)
370      GO TO 600
371  548  WRITE (NWRIT,125)

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C      M = 1 SETS SWITCH 4 OFF
      M = 2 SETS SWITCH 4 ON
376      M=2
377      GO TO (1000,751),M
378      751  WRITE (NWRIT,417)
379      WRITE (NWRIT,1005)
380      1005  FORMAT (' ', 'PHI ', 2X, ' SIEVE ', 2X, '% SIEVE ', 2X, ' CUM. % ', /' '
      *, ' VALUE ', 2X, ' WEIGHTS ', 2X, ' WEIGHTS ', 2X, ' WEIGHTS ')
381      DO 650 N=1,J
382      WRITE (NWRIT,412) PHI(1,N),WT(N),PCT(N),PHI(2,N)
383      650   CONTINUE
384      700   GO TO 1000
385      218   FORMAT(27H FOLKS TEXTURAL DESCRIPTION)
386      401   FORMAT(1X, 'STANDARD DEVIATION=',F7.3)
387      402   FORMAT(1X, 'MEAN=',F8.3)
388      403   FORMAT(1X, 'SKEWNESS=',F7.3)
389      404   FORMAT(1X, 'KURTOSIS=',F7.3)
390      405   FORMAT(1X, 'SORTING=',F6.3)
391      406   FORMAT(1X, 'MZ=',F6.3, ' MEAN DIAMETER IN PHI UNITS')
392      408   FORMAT(1X, 'PHI=',F7.3, ' MEAN DIAMETER IN PHI UNITS')
393      409   FORMAT(1X, 'SIGMA PHI=',F6.3, ' SORTING VALUE')
394      411   FORMAT(1X, 'ALPHA TWO PHI=',F6.3)
395      412   FORMAT(' ',F7.2,2X,F7.2,2X,F7.2,2X,F7.2)
396      415   FORMAT(1X, 'THIRD MOMENT =',E12.5, ' FOURTH MOMENT =',E12.5)
397      416   FORMAT(1X, 'VARIANCE =',E12.5)
398      417   FORMAT(1HS)
399      276   FORMAT(49H DATA IS TOO OPENENDED FOR FOLK OR MOMENT MEASURE)
400      970   FORMAT(57H TWO ERRORS,CUT DOWN ON THE COFFEE BREAKS AND GET TO WOR
      1K)
401      971   FORMAT(75H THREE ERRORS,ARE YOU TRYING TO THINK OR IS SOMEONE BURN
      1ING AN OLD OVERSHOE)
402      972   FORMAT(34H YOU STUPID CLOD YOU GOOFED AGAIN)
403      1000  WRITE (NWRIT,1006)
404      1006  FORMAT('1')
405      STOP
406      END

```

\$ENTRY

C IBM 370 COMPUTER. THE PROGRAM AS FOLLOWS HAS BEEN SLIGHTLY MODIFIED
 C BY GREG McMILLAN, A GRADUATE STUDENT, DEPT. OF GEOGRAPHY, THE UNIVER-
 C SITY OF MANITOBA FROM A PROGRAM WRITTEN BY W.C. ISOPHORDING, OF THE
 C DEPT. OF GEOLOGY, THE UNIVERSITY OF SOUTH ALABAMA, MOBILE.

C
 C
 C

-PROGRAM LISTING-

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1 CHARACTER*2 TITLE(12)
2 DIMENSION PHI(2,99),PCT(99),WT(99),DF(4),DEL(99)
3 NREAD=5
4 NWRIT=6
5 IRR=0
6 1000 READ (NREAD,1001) J,(TITLE(I),I=1,12),PHI(1,1),TOTWT,XMQ
7 IF (J.EQ.0) GO TO 10000
8 1001 FORMAT (I2,2X,12A2,F7.2,F7.2,F7.2)
9 JA=J-1
10 READ (NREAD,1002) (DEL(N),N=1,J)
11 1002 FORMAT (8F10.2)
12 DO 5 N=2,J
13 PHI(1,N)=PHI(1,N-1)+DEL(N)
14 5 CONTINUE
15 READ (NREAD,1002) (WT(N),N=1,J)
16 I=1
17 K=1
18 WTTOT=0.0
19 DO 3 N=1,J
20 PCT(N)=100.*(WT(N)/TOTWT)
21 WTTOT=WTTOT+WT(N)
22 3 CONTINUE
23 WTDIF=TOTWT-WTTOT
24 PHI(2,1)=PCT(1)
25 DO 4 N=2,J
26 PHI(2,N)=PCT(N)+PHI(2,N-1)
27 4 CONTINUE
28 WRITE (NWRIT,601) (TITLE(I),I=1,12)
29 601 FORMAT ('1',12A2)
30 WRITE (NWRIT,417)
31 N=1
32 IF (PHI(2,J)-84.0)100,310,310
33 310 IF (PHI(2,J)-102.0)612,920,920
34 920 WRITE (NWRIT,319)
35 319 FORMAT (' ', '**ERROR** SIEVE CONTENTS GREATER THAN TOTAL SAMPLE WE
*IGHT, SUGGEST YOU CHECK YOUR INPUT DATA. ')
36 IRR=IRR+1
37 999 GO TO (1000,70,71,72),IRR
38 70 WRITE (NWRIT,970)
39 GO TO 1000
40 71 WRITE (NWRIT,971)
41 GO TO 1000
42 72 WRITE (NWRIT,972)
43 IRR=IRR-1
44 GO TO 1000
45 612 IF (PHI(2,1)-5.0)613,613,614
46 614 PHI5=5.*DEL(N)/PHI(2,1)+PHI(1,N)
47 IF (PHI(2,1)-16.0)9,9,615
48 615 PHI16=16.*DEL(N)/PHI(2,1)+PHI(1,1)
49 GO TO 13
50 613 IF (PHI(2,N+1)-5.0)6,6,8
51 6 N=N+1
52 GO TO 612
53 8 PHI5 =DEL(N)*((5.0-PHI(2,N))/(PHI(2,N+1)-PHI(2,N)))+PHI(1,N)
54 9 IF (PHI(2,N+1)-16.)11,11,12

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54 14 N=N+1
55 14 GO TO 13
61 15 PHI25=DEL(N)*((25.0-PHI(2,N))/(PHI(2,N+1)-PHI(2,N))+PHI(1,N)
62 16 IF(PHI(2,N+1)-50.)17,17,18
63 17 N=N+1
64 17 GO TO 16
65 18 PHI50=DEL(N)*((50.0-PHI(2,N))/(PHI(2,N+1)-PHI(2,N))+PHI(1,N)
66 19 IF(PHI(2,N+1)-75.)20,20,21
67 20 N=N+1
68 20 GO TO 19
69 21 PHI75=DEL(N)*((75.0-PHI(2,N))/(PHI(2,N+1)-PHI(2,N))+PHI(1,N)
70 22 IF(PHI(2,N+1)-84.)23,23,24
71 23 N=N+1
72 23 IF(N-JA)22,22,100
73 24 PHI84=DEL(N)*((84.0-PHI(2,N))/(PHI(2,N+1)-PHI(2,N))+PHI(1,N)
74 25 IF(PHI(2,N+1)-95.)26,26,27
75 26 N=N+1
76 26 IF(N-JA)25,25,625
77 27 PHI95=DEL(N)*((95.0-PHI(2,N))/(PHI(2,N+1)-PHI(2,N))+PHI(1,N)
78 EMZ=.333*(PHI16+PHI50+PHI84)
79 SIGI=.25*(PHI84-PHI16)+.1515*(PHI95-PHI5)
80 SNAP=.5*((PHI5+PHI95-2.*PHI50)/(PHI95-PHI5))
81 SKI=.5*((PHI16+PHI84-2.*PHI50)/(PHI84-PHI16))+SNAP
82 AL2FI=(PHI95+PHI5-2.*PHI50)/(PHI84-PHI16)
83 CAYGP=.4098*((PHI95-PHI5)/(PHI75-PHI25))
84 28 CONTINUE
85 EMPHI=.5*(PHI16+PHI84)
86 SIGFI=.5*(PHI84-PHI16)
87 ALFI=(PHI84+PHI16-2.*PHI50)/(PHI84-PHI16)
88 GO TO (701,702),I
89 701 CAYGI=(PHI16-PHI5+PHI95-PHI84)/(PHI84-PHI16)
90 702 CONTINUE
91 C THIS POINT INDICATES COMPLETION OF FOLK, INMAN. START MOMENTIME
92 GO TO (29,29,200),K
93 29 CONTINUE
94 C IF(PHI(2,J)-99.5)300,30,30
95 C M=1 SETS SWITCH 1 OFF
96 C M=2 SETS SWITCH 1 ON
97 30 M=2
98 GO TO (320,31),M
99 31 CONTINUE
100 DO 35 L=1,4
101 DF(L)=0
102 CONTINUE
103 DO 40 N=1,JA
104 IF (PCT(N)) 40,40,41
105 A=PHI(1,N)-.5*DEL(N)-XMQ
106 IF (A) 42,40,42
107 42 D=A*A
108 E=A*D
109 DF(1)=PCT(N)*A+DF(1)
110 DF(2)=PCT(N)*D+DF(2)
111 DF(3)=PCT(N)*E+DF(3)
112 DF(4)=PCT(N)*E*A+DF(4)
113 40 CONTINUE
114 DO 45 L=1,4
115 DF(L)=.01*DF(L)
116 CONTINUE
117 XBAR=XMQ+DF(1)
118 B=DF(1)*DF(1)
119 SSQD=DF(2)-B
120 S=SQRT(SSQD)

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122 C THIS COMPLETES THE CALCULATION OF MOMENT MEASURE
GO TO 400
C THE FOLLOWING SECTION PROVIDES OUTPUT FORMS AND SELECTS FOLKS
C TEXTURAL GROUPING TERMS
123 102 FORMAT(33H CALCULATION OF INMANS STATISTICS)
124 103 FORMAT(32H CALCULATION OF FOLKS STATISTICS)
125 104 FORMAT(41H CALCULATION OF MOMENT MEASURE STATISTICS)
126 105 FORMAT(1X,'INSUFFICIENT DATA')
127 106 FORMAT(28H MOMENT MEASURE NOT COMPUTED)
128 107 FORMAT(48H DATA FOR DRAWING A FREQUENCY DISTRIBUTION CURVE)
129 109 FORMAT(17H VERY WELL SORTED)
130 111 FORMAT(1X,'MODERATELY SORTED')
131 1313 FORMAT(1X,'MODERATELY WELL SORTED')
132 110 FORMAT(1X,'WELL SORTED')
133 112 FORMAT(14H POORLY SORTED)
134 113 FORMAT(19H VERY POORLY SORTED)
135 114 FORMAT(24H EXTREMELY POORLY SORTED)
136 121 FORMAT(21H STRONGLY FINE SKEWED)
137 122 FORMAT(12H FINE SKEWED)
138 123 FORMAT(17H NEAR SYMMETRICAL)
139 124 FORMAT(14H COARSE SKEWED)
140 125 FORMAT(23H STRONGLY COARSE SKEWED)
141 131 FORMAT(17H VERY PLATYKURTIC)
142 720 FORMAT(1X,'KG (INMAN)=' ,F7.3,' KURTOSIS VALUE')
143 132 FORMAT(12H PLATYKURTIC)
144 133 FORMAT(11H MESOKURTIC)
145 134 FORMAT(12H LEPTOKURTIC)
146 135 FORMAT(17H VERY LEPTOKURTIC)
147 136 FORMAT(22H EXTREMELY LEPTOKURTIC)
148 141 FORMAT( 7H GRAVEL)
149 142 FORMAT(13H SANDY GRAVEL)
150 143 FORMAT(19H MUDDY SANDY GRAVEL)
151 144 FORMAT(13H MUDDY GRAVEL)
152 145 FORMAT(14H GRAVELLY SAND)
153 146 FORMAT(20H GRAVELLY MUDDY SAND)
154 147 FORMAT(13H GRAVELLY MUD)
155 148 FORMAT(23H SLIGHTLY GRAVELLY SAND)
156 149 FORMAT(29H SLIGHTLY GRAVELLY MUDDY SAND)
157 150 FORMAT(28H SLIGHTLY GRAVELLY SANDY MUD)
158 151 FORMAT(22H SLIGHTLY GRAVELLY MUD)
159 160 FORMAT( 5H SAND)
160 161 FORMAT(12H CLAYEY SAND)
161 162 FORMAT(11H MUDDY SAND)
162 163 FORMAT(11H SILTY SAND)
163 164 FORMAT(11H SANDY CLAY)
164 165 FORMAT(10H SANDY MUD)
165 166 FORMAT(11H SANDY SILT)
166 167 FORMAT( 5H CLAY)
167 168 FORMAT( 4H MUD)
168 169 FORMAT( 5H SILT)
C THIS LIST PROVIDES ALPHA OUTPUT TO BE USED IN THE FOLLOWING
C DECISION NETWORK BASED UPON FOLKS TEXTURAL TRIANGLE DIAGRAMS
C FOLK,JOUR GEOL. V0162, P345-351, JULY1954
169 275 FORMAT(38H DATA IS TOO OPENENDED FOR CALCULATION)
C THE FOLLOWING PRUCEEDURE NAMES THE SAMPLE ACCORDING TO FOLKS
C TEXTURAL GROUPING
170 277 GRPCT=0.0001
171 N=1
172 801 IF(PHI(1,N)+1.) 802,803,804
173 802 GRPCT=GRPCT+PCT(N)
174 805 IF(N-J)805,806,806
175 805 N=N+1

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100      004      SPCT=PCT(N)*((PHI(1,N)+1.)/DEL(N))
181      GRPCT=GRPCT+PCT(N)-SPCT
182      807      IF(N-J)808,809,809
183      808      N=N+1
184      IF(PHI(1,N)-4.)810,811,812
185      810      SPCT=SPCT+PCT(N)
186      GO TO 807
187      811      SPCT=SPCT+PCT(N)
188      GO TO 809
189      812      SIPCT=PCT(N)*((PHI(1,N)-4.)/DEL(N))
190      SPCT=SPCT+PCT(N)-SIPCT
191      GO TO 813
192      809      SIPCT=0.0001
193      813      IF(N-J)814,815,815
194      814      N=N+1
195      IF(PHI(1,N)-8.)816,817,818
196      816      SIPCT=SIPCT+PCT(N)
197      GO TO 813
198      817      SIPCT=SIPCT+PCT(N)
199      815      CLPCT=0.0001
200      GO TO 819
201      818      CLPCT=PCT(N)*((PHI(1,N)-8.)/DEL(N))
202      SIPCT=SIPCT+PCT(N)-CLPCT
203      819      IF(N-J)820,215,822
204      822      WRITE (NWRIT,823)
205      823      FORMAT(23H POSSIBLE MACHINE ERROR)
206      820      CLPCT=CLPCT+PHI(2,J)-PHI(2,N)
207      215      CONTINUE
208      WRITE (NWRIT,417)
209      WRITE (NWRIT,218)
210      IF(GRPCT-.1)250,250,217
211      217      EMPCT=SIPCT+CLPCT
212      IF(GRPCT-80.)220,219,219
213      219      WRITE (NWRIT,141)
214      GO TO 500
215      220      IF(GRPCT-30.)222,221,221
216      221      IF(EMPCT-SPCT)223,224,224
217      224      WRITE (NWRIT,144)
218      GO TO 500
219      223      IF((SPCT/EMPCT)-9.)225,225,226
220      225      WRITE (NWRIT,143)
221      GO TO 500
222      226      WRITE (NWRIT,142)
223      GO TO 500
224      222      IF(GRPCT-5.)227,228,228
225      228      IF(EMPCT-SPCT)230,229,229
226      229      WRITE (NWRIT,147)
227      GO TO 500
228      230      IF((SPCT/EMPCT)-9.)231,231,232
229      231      WRITE (NWRIT,146)
230      232      WRITE (NWRIT,145)
231      GO TO 500
232      227      IF(EMPCT-SPCT)233,234,234
233      233      IF((SPCT/EMPCT)-9.)235,235,236
234      235      WRITE (NWRIT,149)
235      GO TO 500
236      236      WRITE (NWRIT,148)
237      GO TO 500
238      234      IF((EMPCT/SPCT)-9.)237,237,238
239      237      WRITE (NWRIT,150)
240      GO TO 500
241      238      WRITE (NWRIT,151)

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240 222 IF((CLPCT/SIPCT)-2.)254,253,253
247 253 IF(SPCT-50.)255,256,256
248 254 IF((SIPCT/CLPCT)-2.)261,261,262
249 256 WRITE (NWRIT,161)
250 GO TO 500
251 255 IF(SPCT-10.)259,260,260
252 259 WRITE (NWRIT,167)
253 GO TO 500
254 260 WRITE (NWRIT,164)
255 GO TO 500
256 261 IF(SPCT-50.)263,264,264
257 264 WRITE (NWRIT,162)
258 GO TO 500
259 263 IF(SPCT-10.)265,266,266
260 265 WRITE (NWRIT,168)
261 GO TO 500
262 266 WRITE (NWRIT,165)
263 GO TO 500
264 262 IF(SPCT-50.)267,268,268
265 268 WRITE (NWRIT,163)
266 GO TO 500
267 267 IF(SPCT-10.)269,270,270
268 269 WRITE (NWRIT,169)
269 GO TO 500
270 270 WRITE (NWRIT,166)
271 GO TO 500
272 100 WRITE (NWRIT,275)
273 IRR=IRR+1
274 GO TO 999
275 200 WRITE (NWRIT,276)
276 K=3
277 GO TO 301
278 300 WRITE (NWRIT,105)
279 320 WRITE (NWRIT,106)
280 K=2
281 GO TO 301
282 400 K=1
283 GO TO 301
284 625 K=3
285 I=2
286 GO TO 28
C M = 1 SETS SWITCH 2 OFF
C M = 2 SETS SWITCH 2 ON
287 301 M=2
288 GO TO (660,661),M
289 660 CONTINUE
290 661 GO TO (302,303,304),K
291 302 WRITE (NWRIT,417)
292 WRITE (NWRIT,104)
293 WRITE (NWRIT,402) XBAR
294 WRITE (NWRIT,416) SSQD
295 WRITE (NWRIT,401) S
296 WRITE (NWRIT,403) SK
297 WRITE (NWRIT,404) CTSIS
298 WRITE (NWRIT,415) EM3,EM4
299 GO TO 600
300 303 WRITE (NWRIT,417)
301 WRITE (NWRIT,103)
302 WRITE (NWRIT,406) EMZ
303 WRITE (NWRIT,405) SIGI
304 WRITE (NWRIT,403) SKI
305 WRITE (NWRIT,404) CAYGP

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310	304	WRITE (NWRIT,417)
311		WRITE (NWRIT,102)
312		WRITE (NWRIT,408) EMPHI
313		WRITE (NWRIT,409) SIGFI
314		WRITE (NWRIT,403) ALFI
315		GO TO (703,704),I
316	703	WRITE (NWRIT,720) CAYGI
317		WRITE (NWRIT,411) AL2FI
318	704	CONTINUE
319		GO TO 600
320	500	CONTINUE
321		IF(ES-4.)502,502,501
322	501	WRITE (NWRIT,114)
323		GO TO 520
324	502	IF(ES-2.)504,504,503
325	503	WRITE (NWRIT,113)
326		GO TO 520
327	504	IF(ES-1.)506,506,505
328	505	WRITE (NWRIT,112)
329		GO TO 520
330	506	IF(ES-.71)1508,1508,507
331	507	WRITE (NWRIT,111)
332		GO TO 520
333	1508	IF(ES-.50) 508,508,1507
334	1507	WRITE (NWRIT,1313)
335		GO TO 520
336	508	IF(ES-.35)510,510,509
337	509	WRITE (NWRIT,110)
338		GO TO 520
339	510	WRITE (NWRIT,109)
340	520	GO TO (705,540),I
341	705	CONTINUE
342		IF(X1-3.)522,522,521
343	521	WRITE (NWRIT,136)
344		GO TO 540
345	522	IF(X1-1.5)524,524,523
346	523	WRITE (NWRIT,135)
347		GO TO 540
348	524	IF(X1-1.11)526,526,525
349	525	WRITE (NWRIT,134)
350		GO TO 540
351	526	IF(X1-.90)528,528,527
352	527	WRITE (NWRIT,133)
353		GO TO 540
354	528	IF(X1-.67)530,530,529
355	529	WRITE (NWRIT,132)
356		GO TO 540
357	530	WRITE (NWRIT,131)
358	540	CONTINUE
359		IF((ESK+1.)-1.30)542,542,541
360	541	WRITE (NWRIT,121)
361		GO TO 600
362	542	IF((ESK+1.)-1.1)544,544,543
363	543	WRITE (NWRIT,122)
364		GO TO 600
365	544	IF((ESK+1.)-.9)546,546,545
366	545	WRITE (NWRIT,123)
367		GO TO 600
368	546	IF((ESK+1.)-.70)548,548,547
369	547	WRITE (NWRIT,124)
370		GO TO 600
371	548	WRITE (NWRIT,125)

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C      M = 1 SETS SWITCH 4 OFF
C      M = 2 SETS SWITCH 4 ON
M=2
376   GO TO (1000,751),M
377   WRITE (NWRIT,417)
378   751  WRITE (NWRIT,1005)
379   WRITE (NWRIT,1005)
380   1005  FORMAT (' ',PHI ',2X,' SIEVE ',2X,'% SIEVE',2X,' CUM. % ',/ '
*, ' VALUE ',2X,' WEIGHTS',2X,' WEIGHTS',2X,' WEIGHTS')
381   DO 650 N=1,J
382   WRITE (NWRIT,412) PHI(1,N),WT(N),PCT(N),PHI(2,N)
383   650   CONTINUE
384   700   GO TO 1000
385   218   FORMAT(27H FOLKS TEXTURAL DESCRIPTION)
386   401   FORMAT(1X,'STANDARD DEVIATION=',F7.3)
387   402   FORMAT(1X,'MEAN=',F8.3)
388   403   FORMAT(1X,'SKEWNESS=',F7.3)
389   404   FORMAT(1X,'KURTOSIS=',F7.3)
390   405   FORMAT(1X,'SORTING=',F6.3)
391   406   FORMAT(1X,'MZ=',F6.3,' MEAN DIAMETER IN PHI UNITS')
392   408   FORMAT(1X,'PHI=',F7.3,' MEAN DIAMETER IN PHI UNITS')
393   409   FORMAT(1X,'SIGMA PHI=',F6.3,' SORTING VALUE')
394   411   FORMAT(1X,'ALPHA TWO PHI=',F6.3)
395   412   FORMAT(' ',F7.2,2X,F7.2,2X,F7.2,2X,F7.2)
396   415   FORMAT(1X,'THIRD MOMENT =',E12.5,' FOURTH MOMENT =',E12.5)
397   416   FORMAT(1X,'VARIANCE =',E12.5)
398   417   FORMAT(1HS)
399   276   FORMAT(49H DATA IS TOO OPENENDED FOR FOLK OR MOMENT MEASURE)
400   970   FORMAT(57H TWO ERRORS,CUT DOWN ON THE COFFEE BREAKS AND GET TO WOR
1K)
401   971   FORMAT(75H THREE ERRORS,ARE YOU TRYING TO THINK OR IS SOMEONE BURN
1ING AN OLD OVERSHOE)
402   972   FORMAT(34H YOU STUPID CLOD YOU GOOFED AGAIN)
403   10000 WRITE (NWRIT,1006)
404   1006  FORMAT('1')
405   STOP
406   END

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\$ENTRY